

The formation of the giant planets: early or late?

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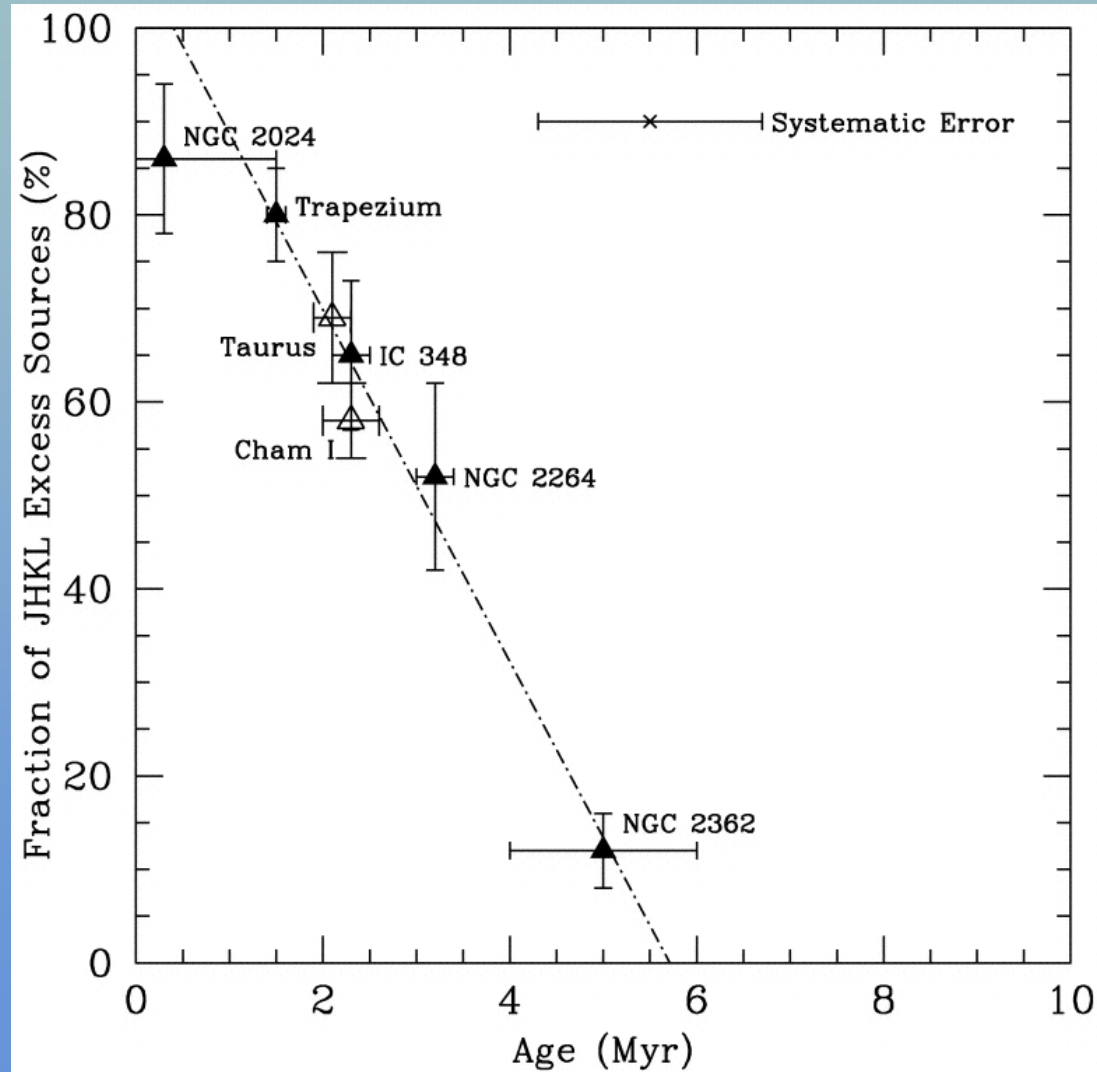
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Basic facts

- Giant planets are made of hydrogen & helium
- Gaseous disks seem to disappear quickly
- ~7% of solar type stars seem to have giant planetary companions

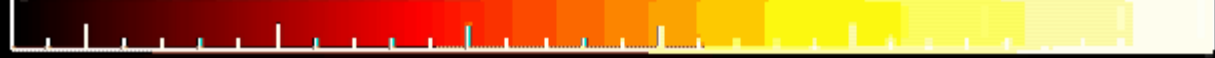


The dissipation of gaseous disks...

- Viscous dissipation: rapid inside 10 AU, but very inefficient for $r > 100 \text{ AU}$.
- Photoevaporation by the central star: efficient close to the star only ($< 10 \text{ AU}$)
- Photoevaporation by external sources: efficient far from the central star ($> 20 \text{ AU}$), but requires massive stars
- Tidal stripping: efficient at $r > 100 \text{ AU}$, but requires dense stellar population

Dimensions: 82500. AU

Time: 197220. yr



-1.5 -1.0 -0.5 0.0 0.5 1.0

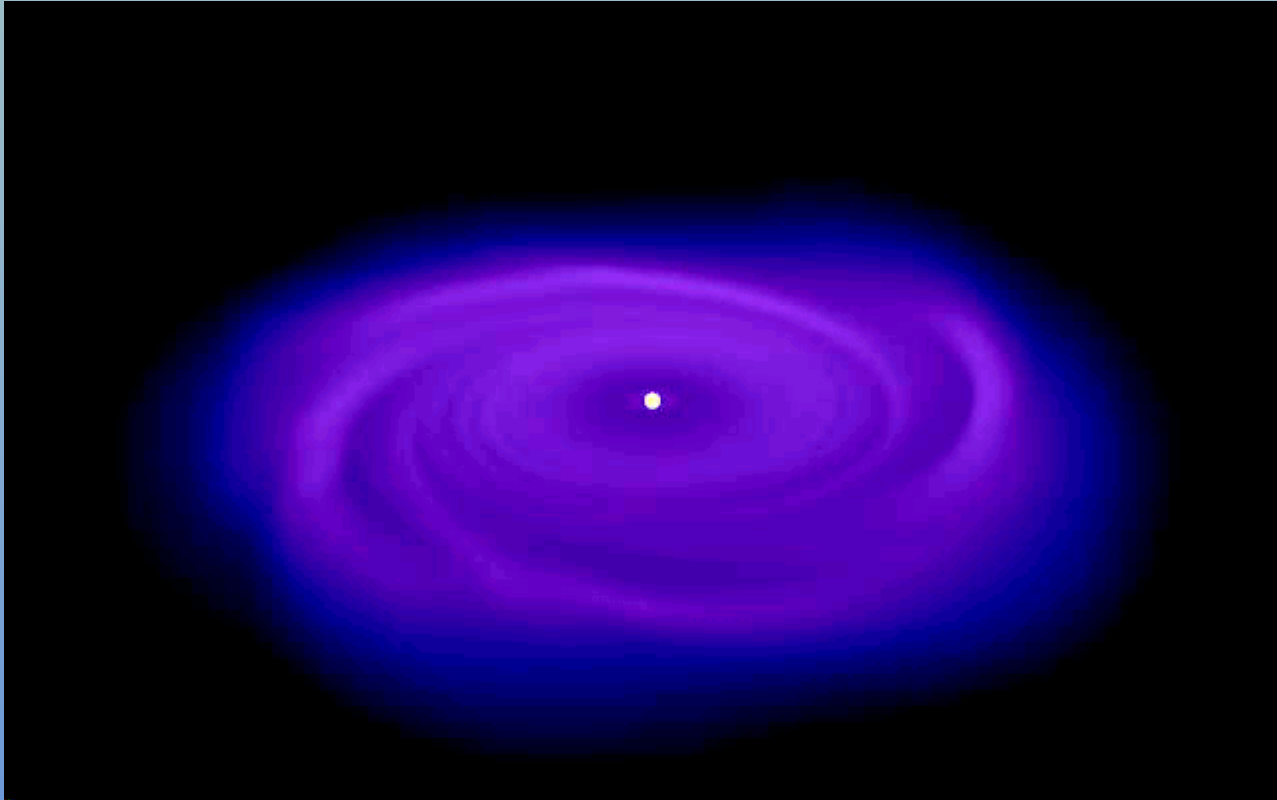
Log Column Density [g/cm^2]

Matthew Bate

Basic questions

- Do giant planets form rapidly ($t < 3\text{Myr}$)?
 - NO? Implies giant planets are RARE
 - YES? Implies sometimes they don't form, or they disappear
- Can we grow them fast enough?
 - With a direct gravitational instability: YES
 - By core accretion: probably YES
- Did they form early (in a massive disk), or late (in a disk with a mass of a few M_{jup})?

A quick formation mode: direct gravitational instability

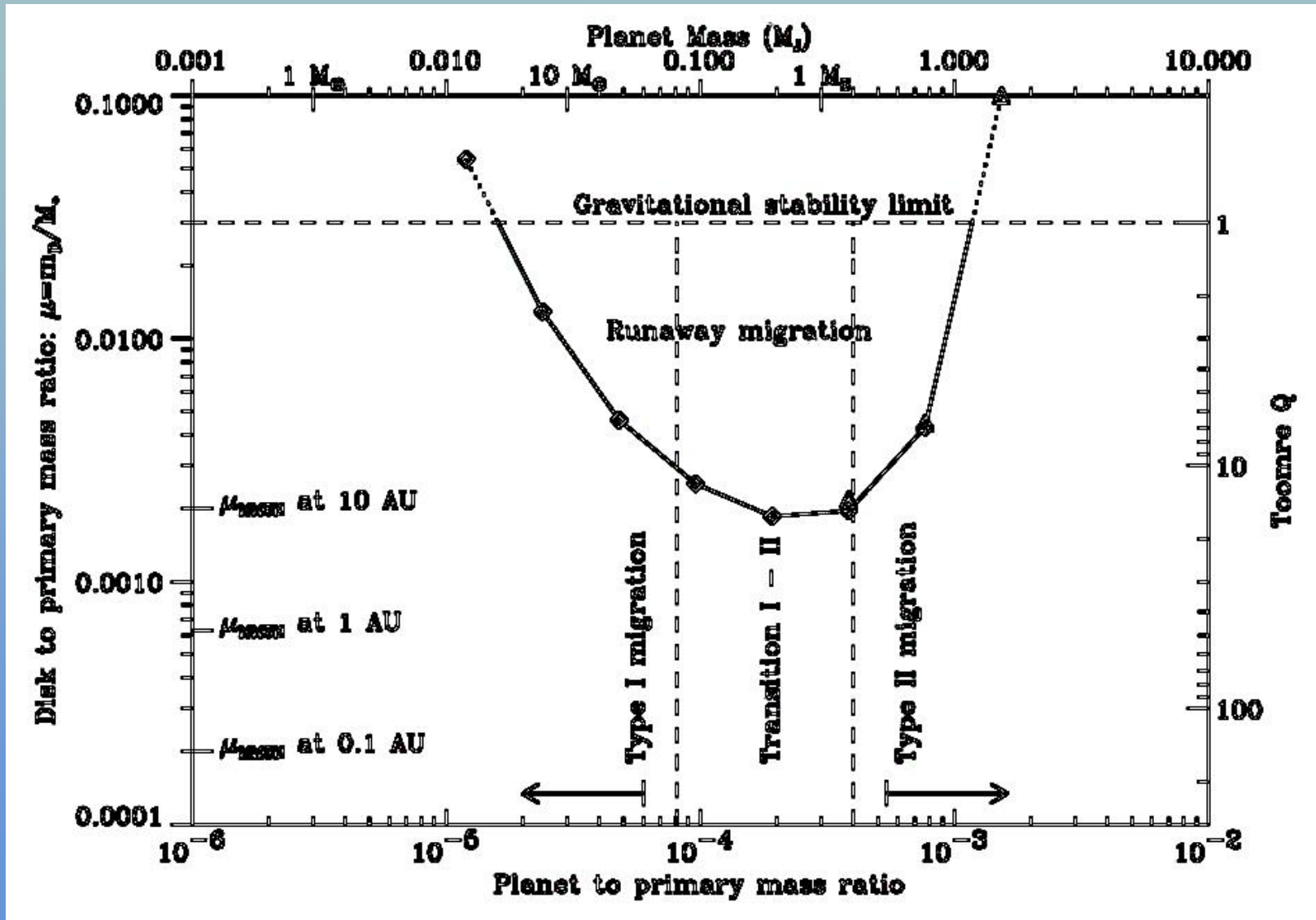


Mayer et al. 2002

A quick formation mode: direct gravitational instability

- Requires a massive & cold nebula ($Q < 2$)
- Simulations: thermodynamics?
 - Two independent methods (SPH, Hydrodynamics)
- Formation timescale (a few 100 years!) is very short compared to the timescale over which the disk is formed by collapse (at least 10^5 years)
- Giant planets would then form early, in a still massive accretion disk

A critical problem: Type III migration



Another problem: halting planetary growth

- Gap opening doesn't appear to be capable of halting planetary growth
 - particularly critical for the gas instability formation scenario
 - possible timing with nebula dissipation requires fine tuning!

Mass flow through a gap

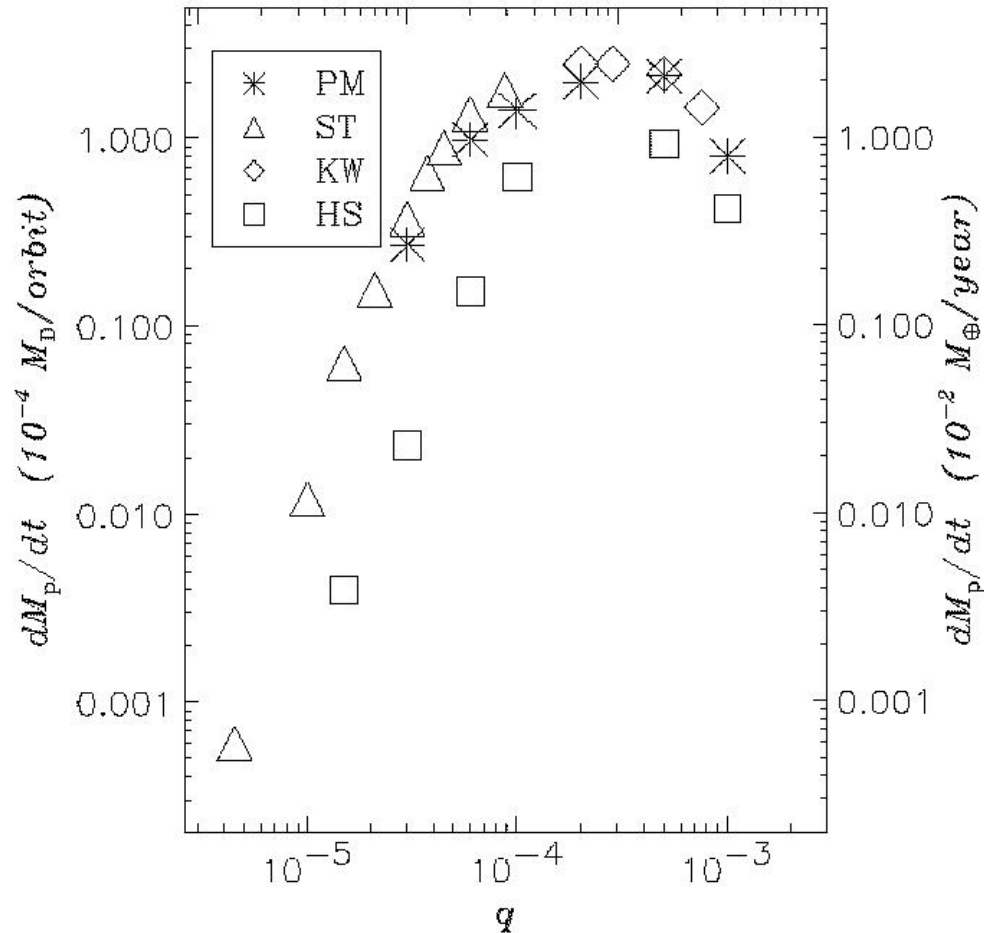
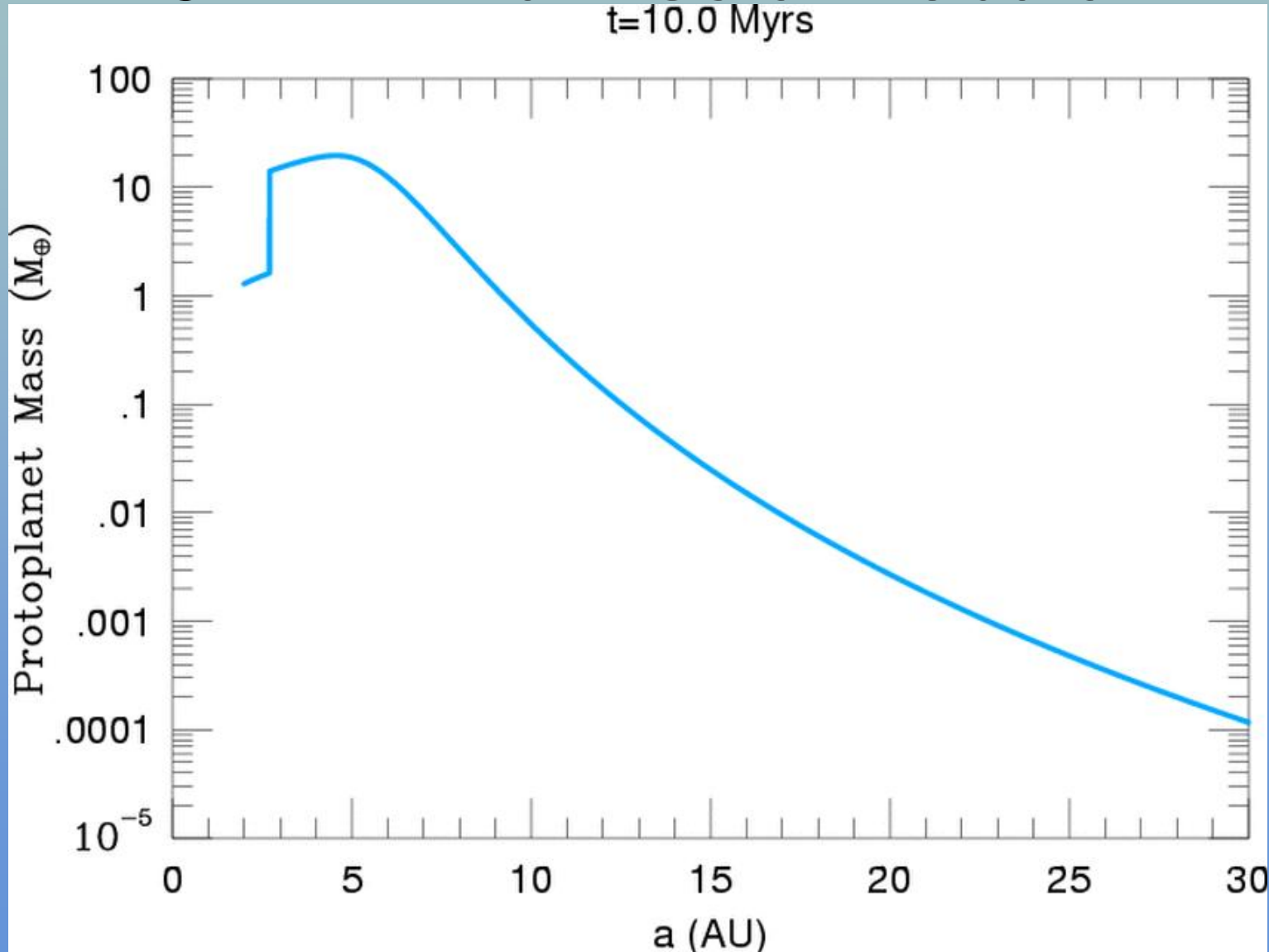


Fig. 7.— Planet's accretion rate as function of the normalized planet mass q . Different symbols stand for the different forms of gravitational potential Φ_p adopted in the computations. Apart from those models run with the homogeneous sphere potential Φ_p^{HS} (see eq. [6]), mass is removed from a volume, centered on the planet, with radius $\kappa_{\text{ac}} = 0.1 R_H$ (see Table 3 for some details concerning the simulation with $M_p = 1.5 M_\oplus$). Models for which $\Phi_p = \Phi_p^{\text{HS}}$ have κ_{ac} equal to $0.2 R_H$ if $M_p > 20 M_\oplus$ and to $0.15 R_H$ if $M_p = 20 M_\oplus$; otherwise κ_{ac} is set to $0.1 R_H$.

A slow formation mode: concurrent accretion of solids and gas

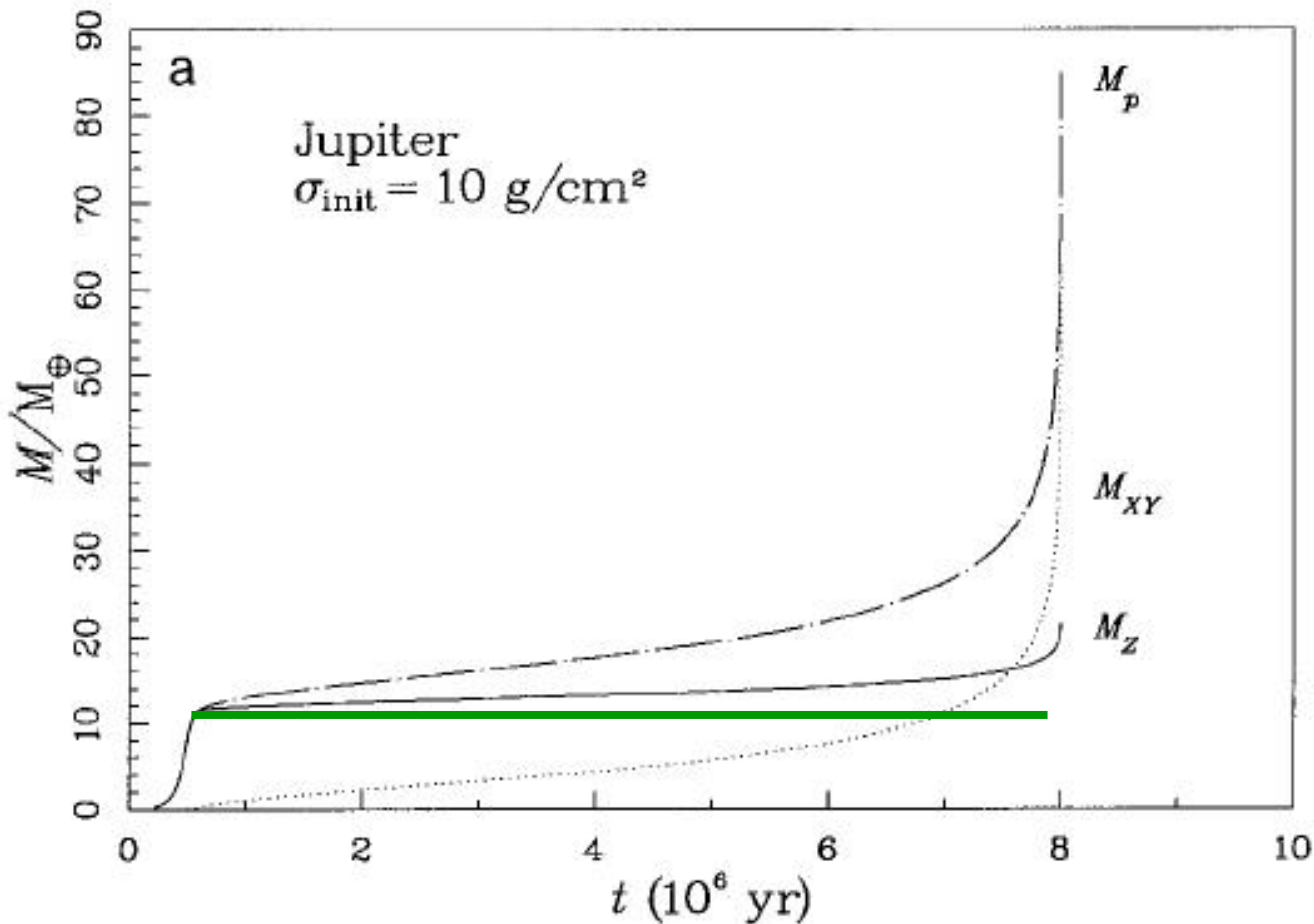
- Accretion of solids by runaway growth
 - Not always easy: migration of solids & excitation of planetesimals can dramatically affect the growth
- Slow capture of the surrounding hydrogen and helium as soon as $M_{\text{core}} > \sim 5 M_{\text{Earth}}$
- Rapid capture of the envelope as soon as $M > \sim 30 M_{\text{Earth}}$
- Same process for Jupiter, Saturn, Uranus, Neptune and terrestrial planets

Runaway growth: the end product in a 5x minimum solar nebula



Forming Jupiter by accretion

- Concurrent accretion of solids and gas (Pollack et al. 1996)
- Final core mass depends on surface density in the nebula



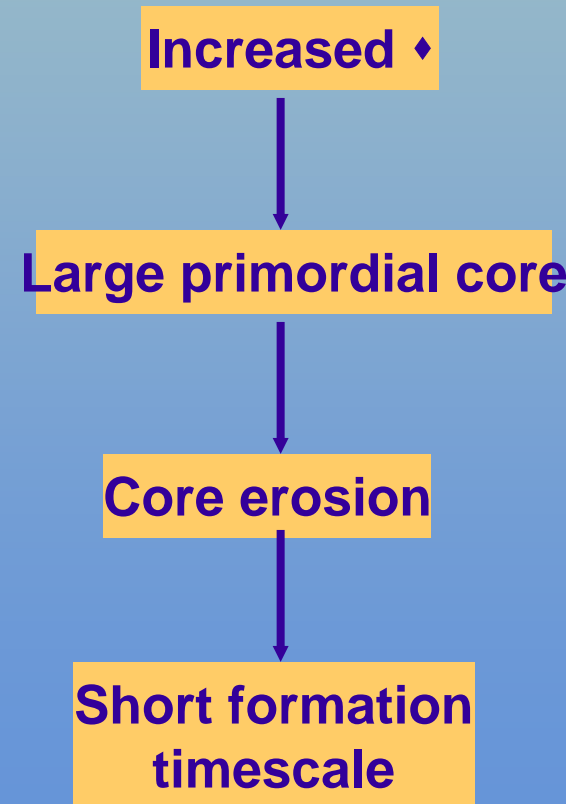
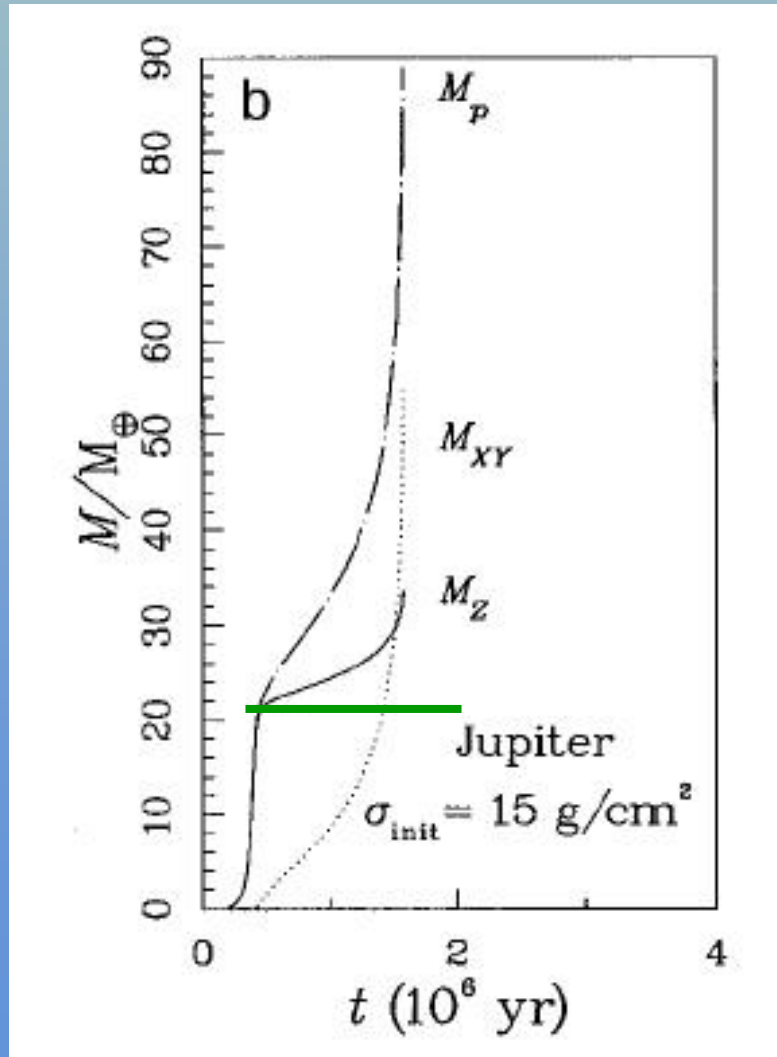
Reduced ♦

Small core

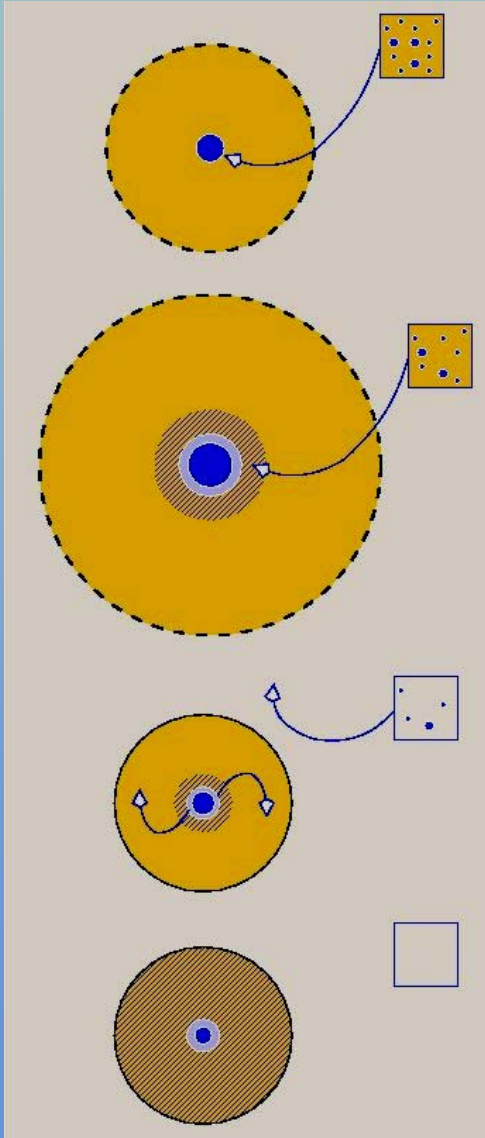
Long formation timescale

Possible rapid formation of Jupiter by accretion & core erosion

- Concurrent accretion of solids and gas (Pollack et al. 1996)



Delivering planetesimals (& water) to the giant planets



- Core accretion: planetesimals are delivered onto the central core.
- Core accretion: planetesimals cannot reach the core intact. (Podolak et al. 1988; Pollack et al. 1996)
- Envelope capture: accretion efficiency drops (Guillot & Gladman 2000): core erosion?
- Present: enriched atmospheres.

Core erosion

Energy needed to redistribute a small core of mass m_{core} in a planet of total mass M and radius R :

$$\Delta E_{\text{grav}} = -\varpi \frac{GM}{R} \Delta m_{\text{core}}$$

Erosion mass flux (fraction of the convective energy in the first convective cell):

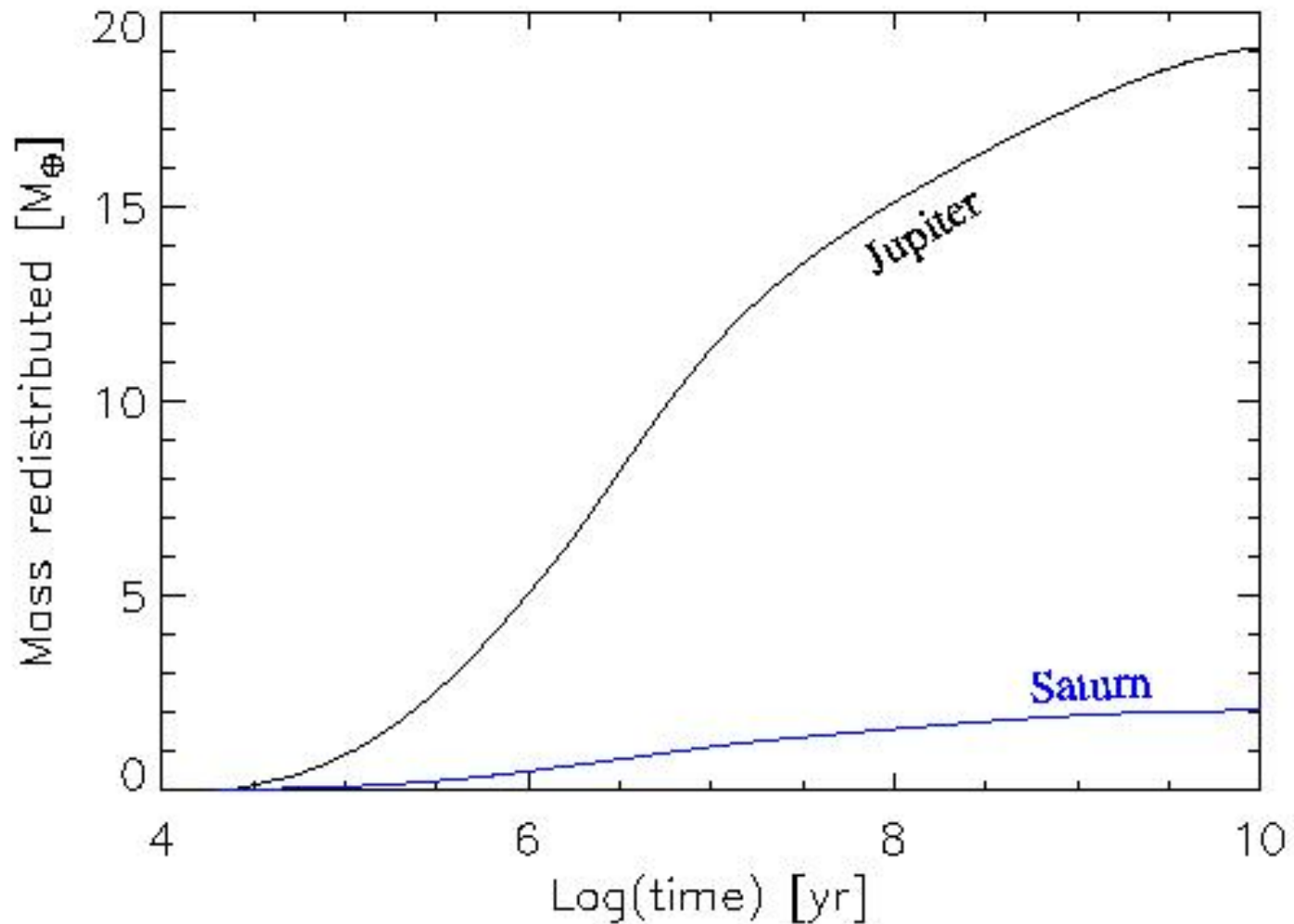
$$\dot{m}_{\text{core}} = -\frac{\chi}{\varpi} \frac{RL_1}{GM}$$

$$\varpi \approx 3/10$$

$$\chi \approx \sqrt{D/K} \approx 0.1$$

molecular diffusivity $D \sim 10^{-3} - 10^{-4} \text{ cm}^2\text{s}^{-1}$
thermal diffusivity $K \sim 0.1 \text{ cm}^2\text{s}^{-1}$

Core erosion: Jupiter & Saturn

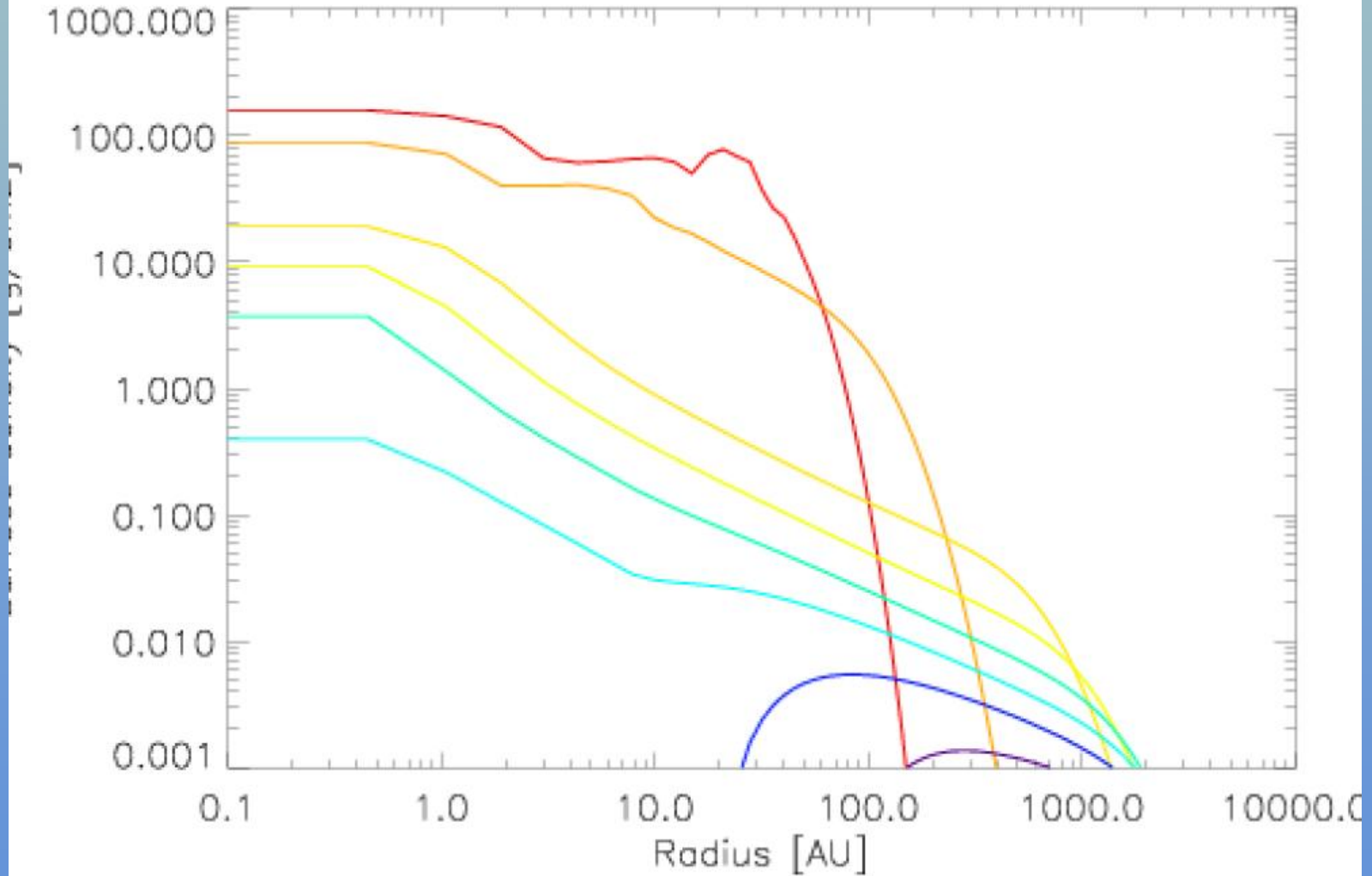


A “Nice” formation scenario

- Assume collapse of molecular cloud in solid rotation
- Assume constant turbulent alpha (here $\alpha=0.03$)
- Begin forming the cores late in nebula evolution when type I migration is suppressed
- Account for nebula photoevaporation
- Giant planets swallow the remains of the nebula (gas + solids)
- Don't bother with details (accretion, planetesimal growth, vertical structure, etc.)

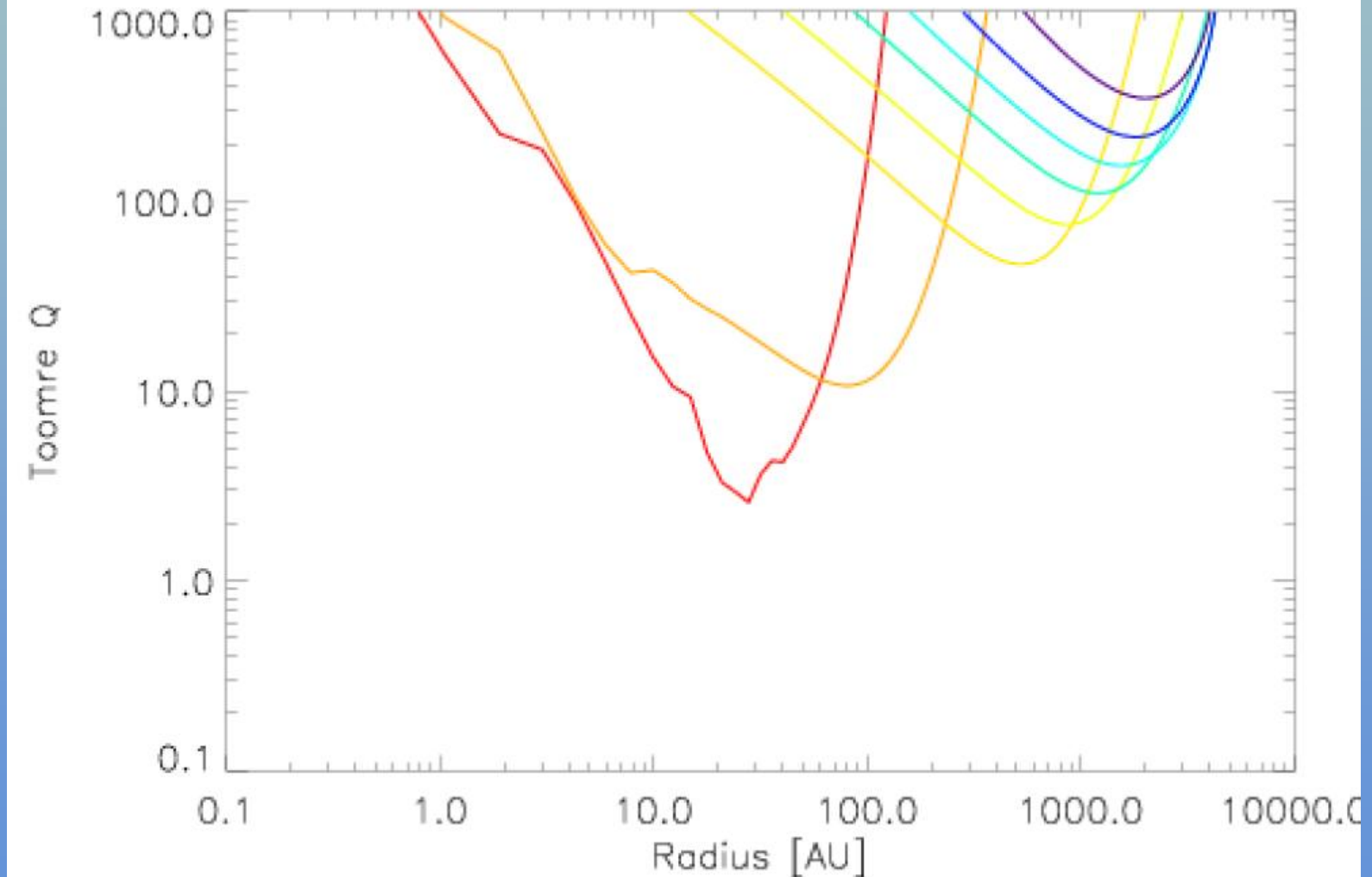
Nebula evolution: surface densities

0.5, 1, 2, 4, 6, 8, 10, 12 Myr

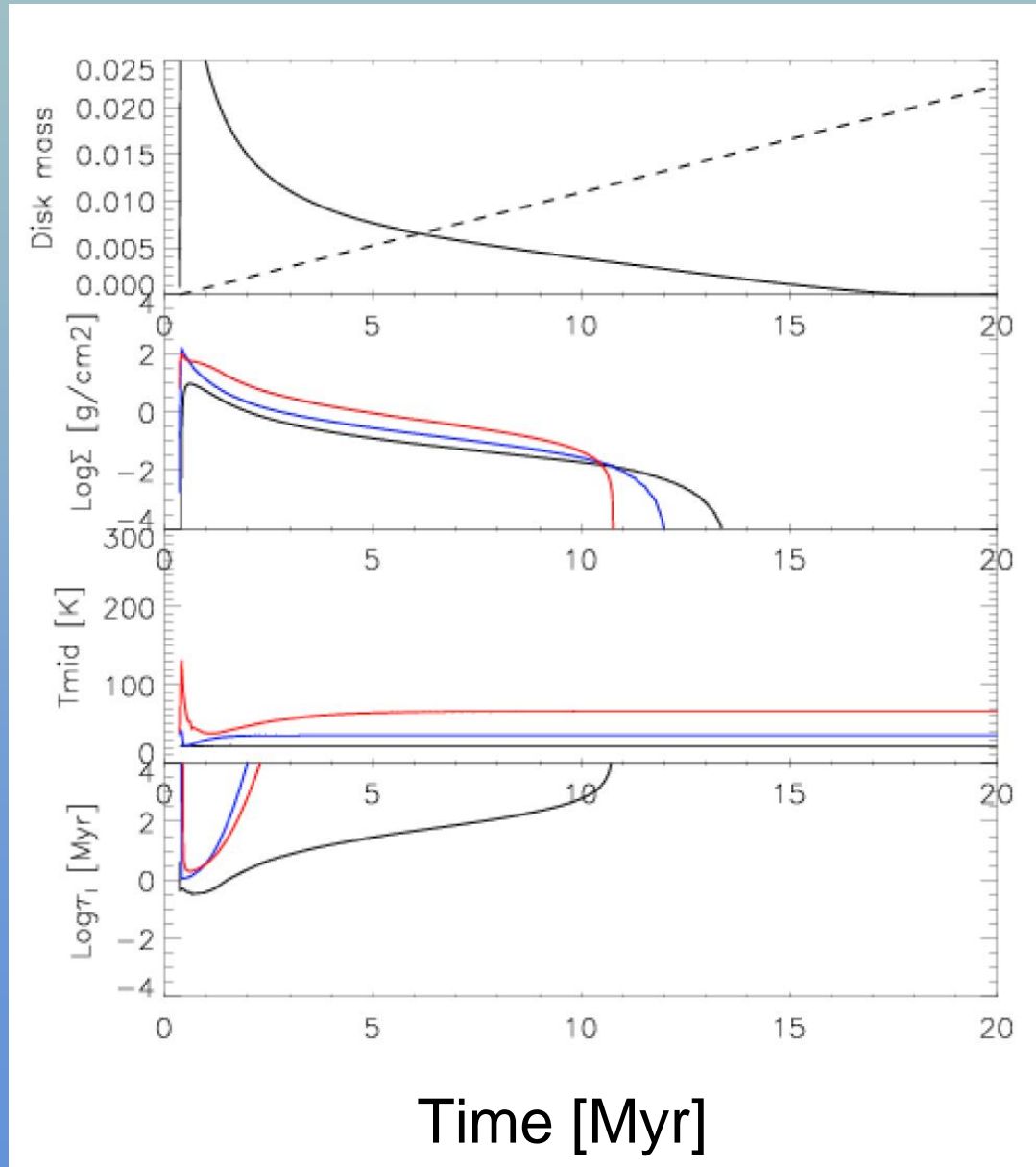


Nebula evolution: Toomre parameter

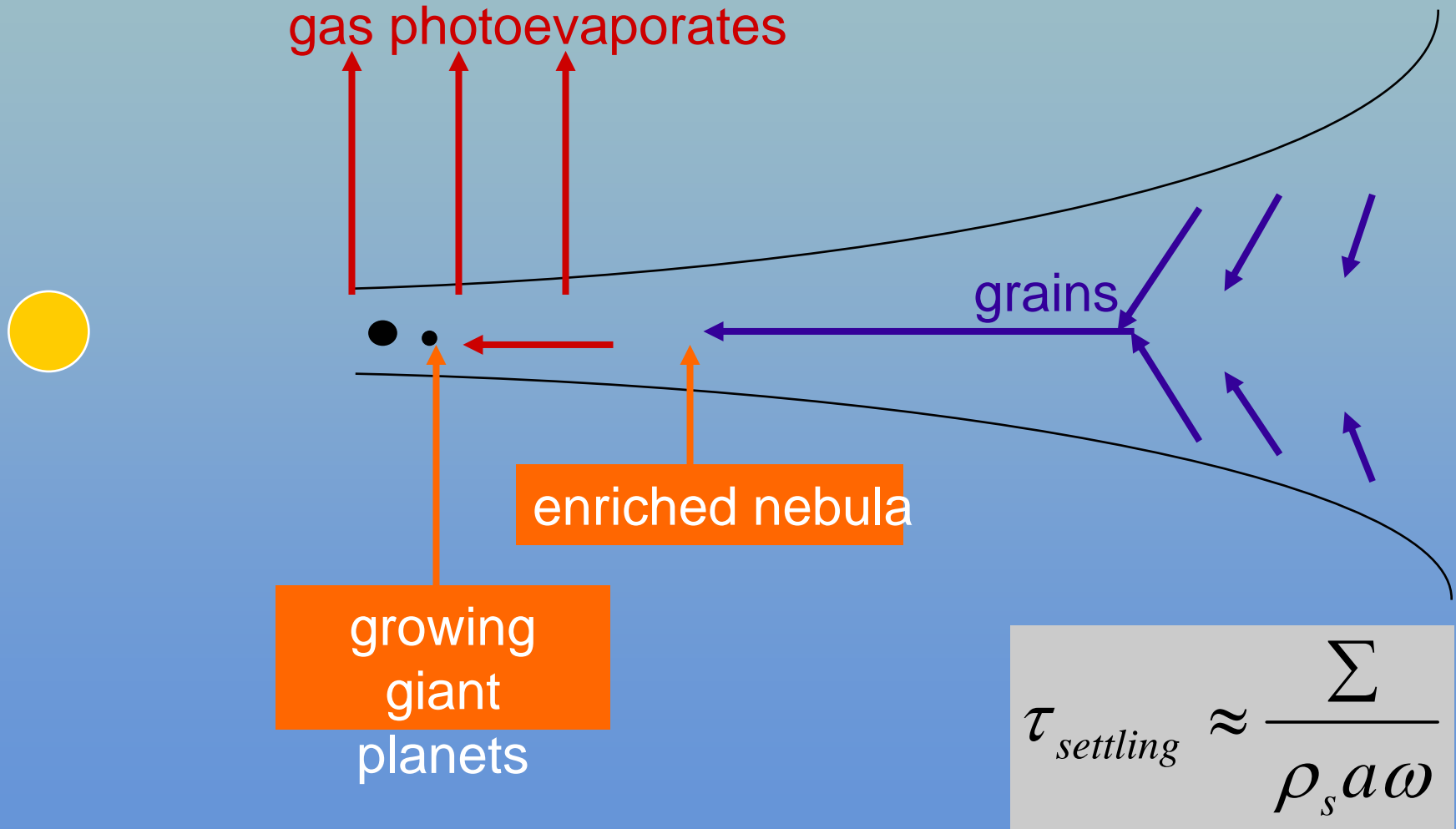
0.5, 1, 2, 4, 6, 8, 10, 12 Myr



Nebula evolution: time dependence



Global picture



Photoevaporation

- Leads to hydrodynamic escape (mean free path $\ll H$)
- Assumed surface density loss (hydrogen):

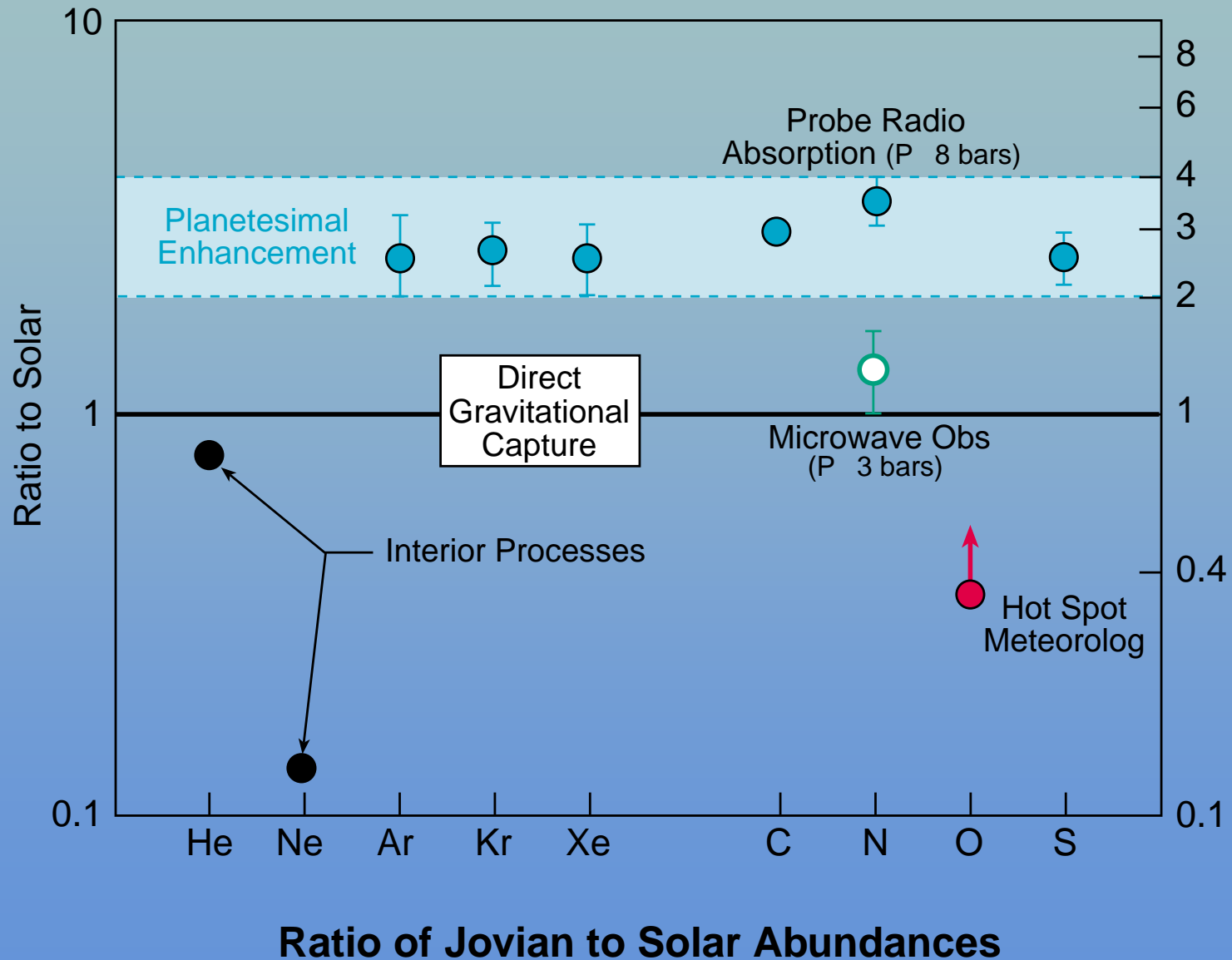
$$\dot{\Sigma}_{UV} \approx 2 \times 10^{-12} \sqrt{\phi_{41}} R_{g,AU} r_{AU}^{-5/2} \text{ g.cm}^{-2} .\text{s}^{-1}$$

- Simple approach (Hunten et al. 1986): elements escape except above a critical mass:

$$\frac{m_c}{m_1} = 1 + \frac{kT \dot{\Sigma}_{UV}}{bg_0 X_1 m_1^2} \approx 20000 @ 10AU$$

- Atomic argon escapes, not grains

The Galileo probe results



Delivering argon to Jupiter

- Ar condenses at very low temperatures (~30K)
- Present in 3 times solar abundance in Jupiter's atmosphere
 - requires decoupling from H,He delivery
- Possible clathration (Gautier et al. 2001, Hersant et al. 2003)
 - implies late delivery
 - requires a number of free cages \Rightarrow high H₂O abundance
 - requires that most of the mass is in small particules
- Direct condensation?
 - requires low temperatures: TBI (work with L. Abe)

Argon: clathration

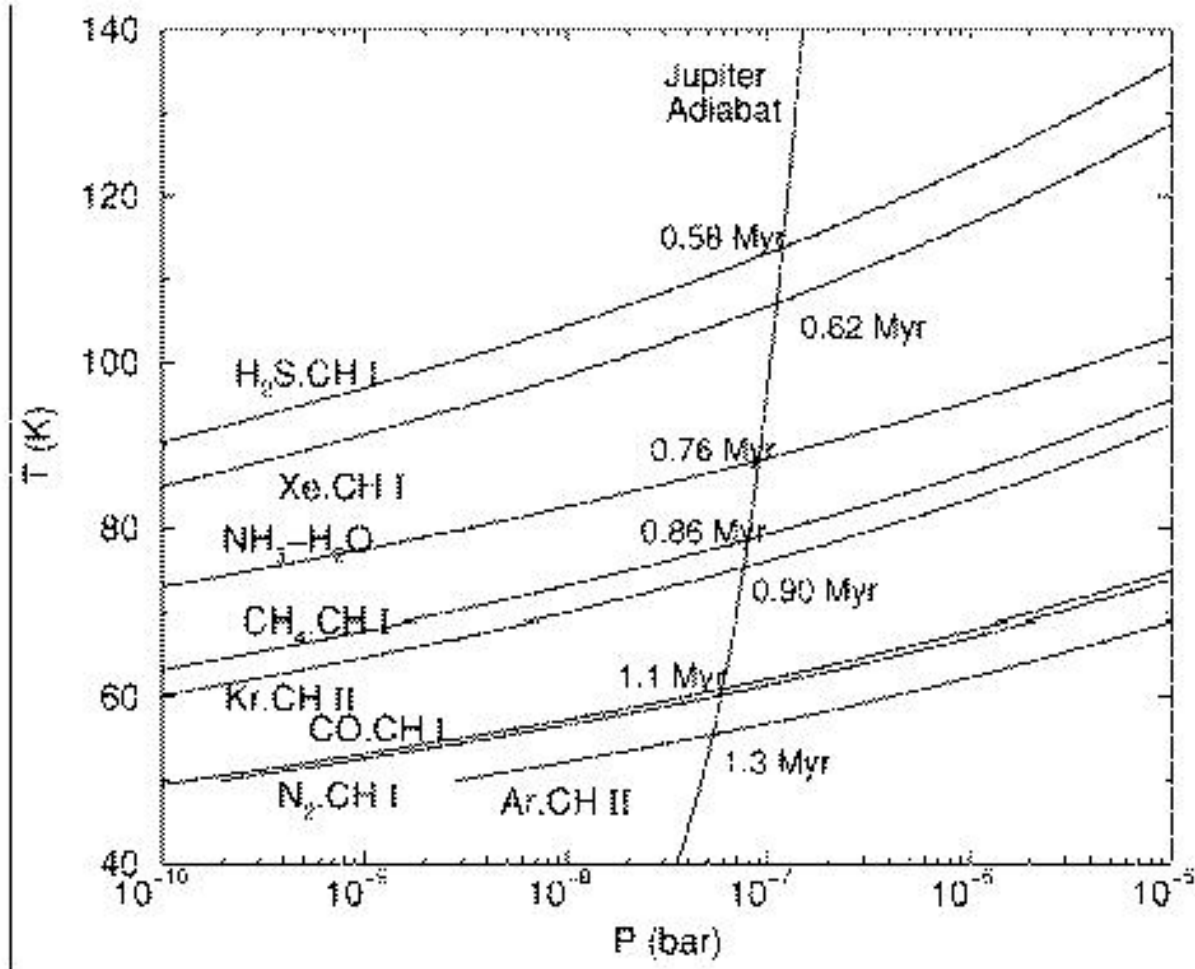


FIG. 2.—Curves of stability of clathrate hydrates of H₂S, Xe, CH₄, CO, N₂, and Ar, together with the Jupiter adiabat from the nominal nebula of the model of Hersant et al. (2001). CH I and CH II refer to clathrate hydrates of type I and type II, respectively (see text for more details). The condensation curve of the NH₃-H₂O hydrate is also shown.

In conclusion...

- A slow formation of Jupiter agrees with models of formation of the galilean moons (Stevenson, Canup & Ward)
- Qualitative explanation of the star metallicity/radial velocimetry planets correlation (eg. Gonzalez, Santos et al.)
- Possibility of gravitational instability of the dust disk (gas depletion)

but

- Terrestrial planets (X-wind?)
- Growth of planetesimals?
 - Runaway growth in a low surface density nebula yields small cores
 - requires migration to grow
- Timing of growth?
 - Jupiter, Saturn, Neptune, Uranus...and the nebula dissipates
- Gas accretion onto the planets?
 - Requires slow accretion (isotropic) for 10 Me cores
- Gravitational interactions?