

Contributions of past and present human generations to committed warming caused by carbon dioxide

Pierre Friedlingstein*[†] and Susan Solomon[‡]

*Institut Pierre Simon Laplace/Laboratoire des Sciences du Climat et de l'Environnement, Commissariat à l'Energie Atomique-Saclay, L'Orme des Merisiers, 91191 Gif sur Yvette, France; and [‡]National Oceanic and Atmospheric Administration, Boulder, CO 80305

Contributed by Susan Solomon, June 7, 2005

We developed a highly simplified approach to estimate the contributions of the past and present human generations to the increase of atmospheric CO₂ and associated global average temperature increases. For each human generation of adopted 25-year length, we use simplified emission test cases to estimate the committed warming passed to successive children, grandchildren, and later generations. We estimate that the last and the current generation contributed approximately two thirds of the present-day CO₂-induced warming. Because of the long time scale required for removal of CO₂ from the atmosphere as well as the time delays characteristic of physical responses of the climate system, global mean temperatures are expected to increase by several tenths of a degree for at least the next 20 years even if CO₂ emissions were immediately cut to zero; that is, there is a commitment to additional CO₂-induced warming even in the absence of emissions. If the rate of increase of CO₂ emissions were to continue up to 2025 and then were cut to zero, a temperature increase of $\approx 1.3^\circ\text{C}$ compared to preindustrial conditions would still occur in 2100, whereas a constant-CO₂-emissions scenario after 2025 would more than double the 2100 warming. These calculations illustrate the manner in which each generation inherits substantial climate change caused by CO₂ emissions that occurred previously, particularly those of their parents, and shows that current CO₂ emissions will contribute significantly to the climate change of future generations.

climate | greenhouse | radiative forcing

Changes in greenhouse gases and aerosols can alter Earth's energy budget, representing a radiative forcing that in turn leads to climate-system responses. Some of these responses can be rapid, as evidenced by changes in surface temperature and water vapor within months after major volcanic eruptions (1, 2). In contrast, the mixed layer of the ocean displays a response time scale of the order of decades, whereas the deep ocean and the great polar ice sheets change over centuries or even millennia. The thermal inertia of ocean and ice is thus among the key factors that cause lags between increases in concentrations of greenhouse gases and observed atmospheric temperature changes. Such lags have been variously referred to as unrealized warming, residual warming, or committed warming (3–7); here we use the term “committed warming.” Regardless of the nomenclature used, these and other studies have highlighted the fact that simple physical considerations linked to time scales for heat transport within the Earth system imply that radiative forcing leads not only to warming at a given time but also to additional future warming even if the concentrations are stabilized. Commitments to warming are also accompanied by commitments to future sea-level rise over time scales of many centuries or even millennia (6–8), but in this article we confine our attention to committed warming only.

Previous studies have shown that a range of forcing agents is contributing to modern climate change (e.g., see refs. 9–11). It is well known that the time scale of commitment depends not only on the climate-system response time but also on the atmospheric lifetime of the radiative forcing agent in question

(gas or aerosol) (e.g., see ref. 12). Relatively short-lived species such as aerosols, tropospheric ozone, and methane are associated with shorter commitments than longer-lived substances such as carbon dioxide or perfluorocarbons (13, 14). We do not attempt to examine the range of forcing agents that are believed to be contributing to climate change in this article, nor do we examine how feedbacks influence the climate changes caused by forcing.

The purpose of this article is to probe in a qualitative fashion the nature and scope of the warming commitments associated with increases in the dominant greenhouse gas, carbon dioxide, by using a highly simplified approach. We will show how multiple time scales that characterize the removal of this gas from the atmosphere affect the warming commitments associated with its emission. Because the time scale of these warming commitments may extend over decades to centuries, the commitments can be passed from one generation to another and are the focus of this study.

We use 25 years as the approximate time scale for a human generation, noting that this may be somewhat lower than the time scale sometimes used to refer to certain types of generations of infrastructure (e.g., power plants, factories, etc.) and technologies. Our goal is to explore physical responses rather than to discuss the plausibility of possible future emissions scenarios. It should be noted that changes in atmospheric concentrations of carbon dioxide are well established for the 20th century, whereas the connections of concentrations to emissions are subject to greater uncertainties relating to sources and the details of processes that remove carbon dioxide from the atmosphere (e.g., see ref. 15). To probe the warming commitments undertaken by several past generations as well as the present one, we examine highly simplified test cases per generation by using (i) zero emissions, (ii) constant emissions, and (iii) increasing emissions derived from observations. Because of the long lifetime of carbon dioxide and its slow removal rate compared to current input emissions, the concentration of CO₂ will continue to increase even for constant emissions. Similarly, CO₂ concentrations do not decrease rapidly even in the complete absence of emissions. We will first show the implications of these different emission cases on the future evolution of atmospheric CO₂. Then, we will show that these cases illustrate how present and past changes in carbon dioxide concentrations affect the climate experienced by each generation, their children, their grandchildren, and later generations.

Description of the Approach

It is well known from atmospheric and ice-core data that the concentration of atmospheric carbon dioxide has changed during past ice ages through many hundreds of thousands of years. It has also increased dramatically in the industrial era at rates that are unprecedented compared to those of at least the previous half-million years (see Fig. 1). Today this carbon dioxide increase contributes $\approx 1.5 \text{ W/m}^2$ to warming of Earth's climate, which is $\approx 60\%$ of the estimated total forcing caused by all known radiative warming agents assessed by the Intergovernmental

[†]To whom correspondence should be addressed. E-mail: pierre.friedlingstein@cea.fr.

© 2005 by The National Academy of Sciences of the USA

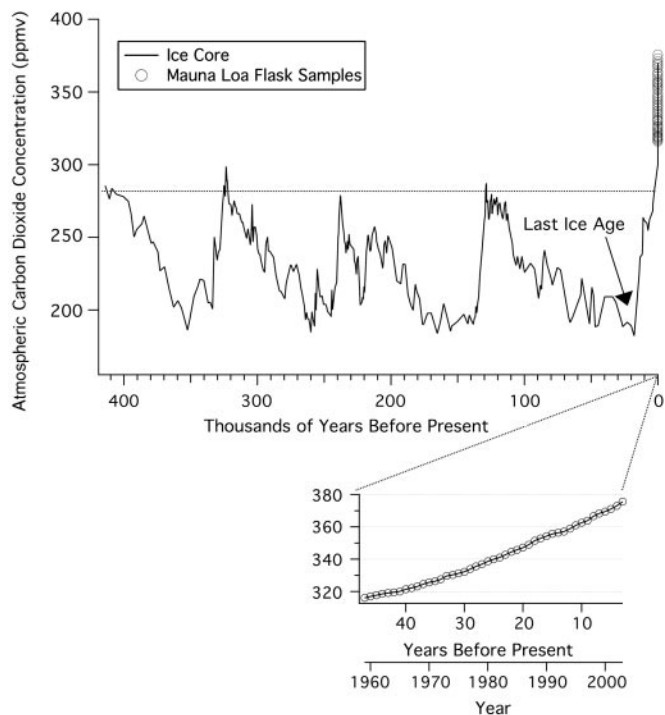


Fig. 1. Observed carbon dioxide concentrations versus time based on ice-core records spanning >400,000 years and flask air samples spanning almost 50 years. Note the different time scales in the figure and *Inset*.

Panel on Climate Change as of 2001 (12). Only carbon dioxide is considered in this article.

We divided the past 150 years into six 25-year-long generations as illustrated in Fig. 2. For each generation we examined the effects of the unrealistic cases of zero emissions (i.e., emissions are cut to zero at the end of that generation up to 2100) and constant emissions (for which emissions are set to their value at the end of that generation up to year 2100). A stepwise increment for each generation up to the year 2000 was also evaluated based on emission data from the Marland *et al.* (16) estimate for

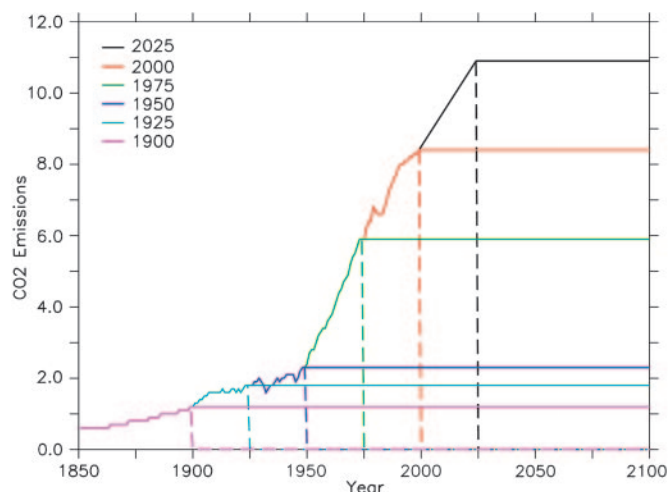


Fig. 2. Annual emissions for the zero-emissions (dashed line) and constant-emissions (solid lines) cases along with historical changes within each 25-year generation (gigatons of carbon per year) up to 2000. The 2000–2025 emissions case was obtained assuming linear increases of emission that are the same as those for the 1975–2000 period.

fossil-fuel CO₂ emissions and from Houghton and Hackler (17) for land-use-change CO₂ emissions.

For the current 2000–2025 generation, we also impose a simple test case assuming increases of emission that are the same as those for 1975–2000 (i.e., an additional 0.1 gigatons of carbon per year). Then in 2025 we again impose zero emissions as well as constant emissions for 2025–2100 for the present generation (Fig. 1). All calculations were performed for the 1850–2100 time period. Calculations including zero emissions were extended to 2200 to indicate the decay of CO₂ over longer time periods that would occur in the absence of emissions.

We first calculate the atmospheric CO₂ concentrations resulting from each of these emission time series by using a pulse-response approach. Atmospheric CO₂ is linked to removal by several different earth-system processes involving the biosphere and oceans. Some CO₂ is removed relatively rapidly, over a time scale of decades, whereas a portion (≈15%) of the added CO₂ is expected to remain in the atmosphere even after a millennium, implying some long-ranging commitments over time scales of human history (18, 19). The atmospheric CO₂ change at time *t* for a CO₂ pulse at time *t*₀ is given by the response function of Joos *et al.* (20, 21):

$$dCO_2(t) = dCO_2(t_0) [b_0 + b_1(t - t_0) + b_2(t - t_0)^2] / [b_0 + b_3(t - t_0) + b_4(t - t_0)^2 + b_5(t - t_0)^3], \quad [1]$$

where $dCO_2(t_0) = \text{emi}(t_0)/2.12$, with $\text{emi}(t_0)$ being the CO₂ emission at time *t*₀.

The actual change in CO₂ is therefore the integral of dCO_2 over past years. Uncertainties in this function are related to differences between different carbon-cycle models as well as dependencies on the absolute abundance of CO₂ (especially for larger perturbations relative to preindustrial conditions) and are on the order of ±15% for the limited range of CO₂ concentrations considered here (22, 23)

When forced with the fossil-fuel and land-use CO₂ emissions for the 1850–2000 period, these equations simulate an atmospheric CO₂ level that is consistent with ice-core data for the 1850–1950 period but slightly overestimate (up to 10 ppm) the CO₂ over the last few decades. We note that there is a large uncertainty in the sources of CO₂ from deforestation. Recent satellite-based estimates show tropical-deforestation emissions being a factor of 2 lower than the Houghton and coworkers (24, 25) estimate, which alone could explain the discrepancy in our simulated CO₂ calculations. Also, as mentioned above, there is uncertainty in the CO₂ pulse-response functions. Finally, such a response function makes the implicit assumption that the CO₂ sinks are only driven by atmospheric CO₂. This assumption may not be realistic for the land uptake, in which other processes such as climate change, nitrogen deposition, or changes in land management are also likely to play a role in the land carbon sink (e.g., see refs. 15 and 26).

The estimate of the corresponding climate change for a given CO₂ time series is made by using a temperature-change pulse function. The temperature change δT at time *t* caused by an atmospheric CO₂ change of δCO_2 at time *t*₀ is assumed to be given by:

$$\delta T(t) = \frac{\alpha}{\tau} \delta CO_2(t_0) \exp(-t/\tau), \quad [2]$$

where α is the climate sensitivity and τ is the response time of the climate system to a CO₂ pulse. The actual temperature change is therefore the integral of δT over past years. These parameters are chosen to broadly match recent modeling studies (e.g., see ref. 7). The assumed climate sensitivity adopted here for a doubling of

emissions strategy had been adopted. The rate of increase in CO₂ emission was largest for the 1950–1975 generation (before the first oil crisis). Keeping emissions constant at the 1975 level would lead to an atmospheric CO₂ level of 465 ppm by 2100, or ≈100 ppm more than what would be reached if emissions had been capped at the 1950 level (Table 1). Capping the emissions at the year 2000 level would add 60 ppm in the atmosphere, whereas doing the same by 2025 would add another ≈50 ppm. We also note that the difference in estimated CO₂ by 2100 if the current and future 21st-century generations were to adopt a zero-emissions versus a constant-emissions strategy is almost 200 ppm (573 versus 382 ppm). Thus, Fig. 3 and these examples show how current and near-future actions on CO₂ emissions are expected to have a very large impact on atmospheric CO₂. In the next section we show how these actions will affect the climate experienced by the next several generations.

Estimated Temperature Changes. Fig. 4 shows estimated temperature changes due to the calculated carbon dioxide time series presented in Fig. 3. In terms of contributions to current warming caused by carbon dioxide increases, Fig. 4 illustrates that the past two generations (1950–1975 and 1975–2000) are believed to share much of the responsibility for the present-day CO₂-induced warming. Currently, globally averaged warming since 1850 is estimated to be $0.6 \pm 0.2^\circ\text{C}$, reflecting contributions to warming and cooling from CO₂ and other radiative forcing agents not considered here (8). Thus, the currently observed warming is comparable to the values shown for 2000 in Fig. 4. The temperature change caused by the first 4 generations from 1980 to 1950 represents ≈0.27°C, or approximately one third of the total estimated CO₂-induced warming in 2000. The 1950–1975 generation added ≈0.23°C, and the 1975–2000 generation added another 0.25°C.

Because of the time response of the carbon cycle shown above along with the response times of the climate system (particularly linked to the lags between oceanic and atmospheric temperature), if emissions were set to zero by 2000, temperature would still keep increasing by another several tenths of a degree Celsius for another ≈30 years (Fig. 4), although atmospheric CO₂ would display some decrease (Fig. 3). Thus, the comparison between Figs. 3 and 4, particularly the shapes of the curves after cessation of emissions in the two cases, illustrates the character and duration of the warming commitment. Different climate models suggest quantitative differences in the magnitude and duration of the existing warming commitment, believed to be linked largely to differences in representations of ocean/atmosphere interactions, but the general behavior shown in Fig. 4 is robust across models (e.g., see refs. 27 and 28).

Stated differently, Fig. 4 underscores the character of the warming commitment: today's observed climate change represents a portion rather than the total of what is expected to occur in the future even if emissions of CO₂ are reduced dramatically. Although the warming commitment implies that near-term actions do not have an immediate effect on CO₂-induced climate change, Fig. 4 shows that they do have an affect within a few decades (i.e., on the time scale of the next assumed generation). This change occurs because CO₂ removal is characterized by multiple time scales, i.e., a portion of the emitted CO₂ is removed on time scales of decades.

Thus, the difference between cutting emissions to zero in 2025 and continuing constant emissions is ≈0.5°C of added warming by approximately 2050. Because of additional removal of CO₂ on longer time scales, the difference reaches 1.6°C by 2100. That is, cutting the emissions of CO₂ to zero in 2025 leads to a total warming of 1.3°C by 2100, with a peak value of 1.4°C around 2050, whereas capping emissions at their 2025 level leads to a warming of ≈1.9°C by 2050 and 2.9°C by 2100. Thus, the actions of each generation influence the climate passed on to their children and later generations, but each inherits much of their own climate. Fig. 4 illustrates how this also applied for the last generation. Constant emissions after 2000 lead to a warming of 1.7°C higher than with zero emissions by 2100.

Summary

This article has probed some simple ways in which time scales influence how carbon dioxide emissions lead to future commitments to a warmer climate. There is an existing and substantial commitment to additional CO₂-induced climate change that is expected to dominate the global average temperature changes in the next several decades. Stated differently, the current human generation is already committed to greater CO₂-induced warming in decades to come than that observed today.

Warming commitments are linked to several factors. One is the commitment of the physical climate system to additional changes in response to the radiative forcing linked to increased concentrations of carbon dioxide. The physical constraints on how each generation passes climate change to the next (even if emissions had been constant or zero) are tied to the time responses of carbon dioxide as well as those of the climate system. In the absence of other changes (such as changes in shorter-lived radiative forcing agents, including human-made contributions to aerosols, methane, or tropospheric ozone, or natural forcings such as volcanic cooling), global mean temperatures are expected to increase by several tenths of a degree Celsius in coming decades as a result of this commitment.

Fig. 2 shows that estimated CO₂ emissions have increased with each passing generation during the 20th century. Although not explicitly probed here, it is evident that these increases in emissions are linked to a different type of commitment: that to energy sources and practices. We emphasize that the zero- and constant-emissions cases assumed here are not realistic and serve only to elucidate the physical-science aspects of the warming commitment.

The simple calculations shown here illustrate that each past generation has inherited substantial climate change caused by carbon dioxide increases that occurred previously. Although substantial quantitative uncertainties are associated with the highly simplified approach taken in this article, the results are qualitatively robust and linked to basic physical considerations. They serve to illustrate how each human generation to date has added to carbon dioxide concentrations, thereby passing increasing warming commitments on to its children, grandchildren, and later descendants. Thus, today's carbon dioxide emissions can be expected to contribute significantly to the climate to be experienced by future generations.

P.F. thanks J. L. Dufresne for fruitful discussions on CO₂ radiative forcing.

1. Soden, B. J., Wetherald, R. T., Stenchikov, G. L. & Robock, A. (2002) *Science* **296**, 727–730.
2. Forster, P. M. d. & Collins, M. (2004) *Clim. Dyn.* **23**, 207–214.
3. Hansen, J., Russel, G., Laci, A., Fung, I., Rind, D. & Stone, P. (1985) *Science* **229**, 857–859.
4. Hansen, J., Nazarenko, L., Ruedy, R., Sato, M., Willis, J., DelGenio, A., Koch, D., Laci, A., Lo, K., Menon, S., et al. (2005) *Science* **308**, 1431–1435.
5. Wetherald, R. T., Stouffer, R. J. & Dixon, K. W. (2001) *Geophys. Res. Lett.* **28**, 1535–1538.
6. Wigley, T. (2005) *Science* **307**, 1766–1769.

7. Meehl, G. A., Washington, W. M., Collins, W. D., Arblaster, J. M., Hu, A., Buja, L. E., Strand, W. G. & Teng, H. (2005) *Science* **307**, 1769–1772.
8. Houghton J. T., Ding, Y., Griggs, D. J., Noguer, M., van der Linden, P. J., Dai, X., Maskell, K. & Johnson, C. A., eds. (2001) *Climate Change 2001: The Scientific Basis: Contributions of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change* (Cambridge Univ. Press, Cambridge, U.K.).
9. Hansen, J., Sato, M., Laci, A., Ruedy, R., Tegen, I. & Matthews, E. (1998) *Proc. Natl. Acad. Sci. USA* **95**, 12753–12758.
10. Boucher, O. & Haywood, J. (2001) *Clim. Dyn.* **18**, 297–302.

11. Myhre, G., Myhre, A. & Stordal, F. (2001) *Atmos. Environ.* **35**, 2361–2373.
12. Rawaswamy V, Boucher, O., Haigh, J., Hauglustaine, D., Haywood, J., Myhre, G., Nakajima, T., Shi, G. Y. & Solomon, S. (2001) in *Climate Change 2001: The Scientific Basis: Contributions of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change*, eds. Houghton, J. T., Ding, Y., Griggs, D. J., Noguer, M., van der Linden, P. J., Dai, X., Maskell, K. & Johnson, C. A. (Cambridge Univ. Press, Cambridge, U.K.), pp. 349–416.
13. Hansen, J., Sato, M. & Ruedy, R. (1997) *J. Geophys. Res.* **102**, 6831–6864.
14. Hansen, J., Sato, M., Nazarenko, L., Ruedy, R., Lacis, A., Koch, D., Tegen, I., Hall, T., Shindell, D., Santer, B. *et al.* (2002) *J. Geophys. Res.* **107**, 4347, 10.1029/2001JD001143.
15. Prentice, I. C., Farquhar, G. D., Fasham, M. J. R., Goulden, M. L., Heimann, M., Jaramillo, V. J., Kheshgi, H. S., Le Quéré, C., Scholes, R. J. & Wallace D. W. R. (2001) in *Climate Change 2001: The Scientific Basis: Contributions of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change*, eds. Houghton, J. T., Ding, Y., Griggs, D. J., Noguer, M., van der Linden, P. J., Dai, X., Maskell, K. & Johnson, C. A. (Cambridge Univ. Press, Cambridge, U.K.), pp. 349–416.
16. Marland, G., Boden, T. A., Andres, R. J., Brenkert, A. L. & Johnson, C. A. (1999) in *Trends Online: A Compendium of Data on Global Change* (Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, U.S. Department of Energy, Oak Ridge, TN).
17. Houghton, R. A. & Hackler, J. L. (2000) in *Trends: A Compendium of Data on Global Change* (Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, U.S. Department of Energy, Oak Ridge, TN).
18. Archer, D. E., Kheshgi, H. & Maier-Reimer, E. (1997) *Geophys. Res. Lett.* **24**, 405–408.
19. Maier-Reimer, E. & Hasselmann, K. (1987) *Clim. Dyn.* **2**, 63–90.
20. Siegenthaler, U. & Joos, F. (1992) *Tellus* **44B**, 186–207.
21. Joos, F., Bruno, M., Fink, R., Stocker, T. F., Siegenthaler, U., LeQuéré, C. & Sarmiento, J. L. (1996) *Tellus* **48B**, 397–417.
22. Albritton, D. L., Derwent, R. G., Isaksen, I. S. A. & Wuebbles, D. J. (1994) in *Climate Change 1994: Radiative Forcing of Climate Change*, eds. Houghton, J. T., Meira, L. G., Bruce, J., Lee, H., Callendar, B. A., Haites, E., Harris, N. & Maskell, K. (Cambridge Univ. Press, Cambridge, U.K.), pp. 205–231.
23. Schimel, D., Enting, I. G., Heimann, M., Wigley, T. M. L., Raynaud, D., Alves, D. & Siegenthaler, U. (1994) in *Climate Change 1994: Radiative Forcing of Climate Change*, eds. Houghton, J. T., Meira, L. G., Bruce, J., Lee, H., Callendar, B. A., Haites, E., Harris, N. & Maskell, K. (Cambridge Univ. Press, Cambridge, U.K.), pp. 35–72.
24. Achard, F., Eva, H. D., Stibig, H.-J., Mayaux, P., Gallego, J., Richards, T. & Malingreau, J.-P. (2002) *Science* **29**, 999–1002.
25. DeFries, R. S., Houghton, R. A., Hansen, M. C., Field, C. B., Skole, R. & Townshend, J. (2002) *Proc. Natl. Acad. Sci. USA* **99**, 14256–14261.
26. Schimel, D. S., House, J. I., Hibbard, K. A., Bousquet, P., Ciais, P., Peylin, P., Braswell, B. H., Apps, M. J., Baker, D., Bondeau, A., *et al.* (2001) *Nature* **414**, 169–172.
27. Raper, S. C., Gregory, J. M. & Stouffer, R. J. (2002) *J. Clim.* **15**, 124–130.
28. Stott, P. A. & Kettleborough, J. (2002) *Nature* **416**, 723–726.