The Origin of Gold in South Africa

Ancient rivers filled with gold, a spectacular upwelling of magma and a colossal meteor impact combined to make the Witwatersrand basin a very special place

Jason Kirk, Joaquin Ruiz, John Chesley and Spencer Titley

rammed shoulder to shoulder with a couple dozen burly miners in an all-too-small elevator, we are about to descend into one of South Africa's largest gold mines. Three men, who must weigh well over 100 kilograms each, lock elbows and heave backwards in unison, forcing us visiting geologists off our feet and farther into the lift. The gate is forced shut and, as we plummet down the billiondollar shaft, our ears pop and swirling debris finds its way into our eyes and noses. Two minutes and two vertical kilometers later, we tumble out into oppressive heat and humidity. Walking through enormous tunnels, dodging ore carts and monstrous mining trucks, we reach a series of ski lifts that take us a kilometer deeper. After a long, muddy hike, we near a passage that slopes up at a 20-degree angle and

Jason Kirk is a Ph.D. candidate in the department of geosciences at the University of Arizona. His research interests include the source and timing of gold mineralization as well as large-scale crustmantle processes. Joaquin Ruiz is professor of geosciences and dean of the College of Science at the University of Arizona. His long-term interests have been the evolution of the crust and mantle. His group at the Unversity of Arizona has extensively used the Re-Os isotopic system to study the geochemical evolution of Earth's systems. John Chesley is a senior research associate in the department of geosciences. His research interests include detailed understanding of the timing and formation of ore deposits, and the use of isotopic and elemental systems to trace the source and pathways of surficial processes. Spence Titley is professor of geosciences in economic geology. He has studied ores of many types, especially gold and copper, in several regions around the world. His current focus is on problems of metallogenesis and the relationships of crust and metal endowment. Kirk's address: Geosciences Department, 1040 East Fourth Street, University of Arizona, Tucson, AZ 85721. Internet: jkirk@geo.arizona.edu

is less than a meter high. Hunched over, we make our way toward the sound of pneumatic drills and water jets, where workers are blasting away at the working face of the mine. More than three kilometers of rock, weighing thousands of tons, lies above our heads. The sedimentary rocks being mined are uninspiring to look at from arm's length—they are what geologists call conglomerates, being composed mostly of rounded pebbles (here consisting predominantly of quartz) cemented together. But under a small magnifying lens, hundreds of small specks of gold appear. The rocks are chock-full of it. Nearly half the gold ever mined in human history has come from these conglomerates in South Africa's Witwatersrand basin.

Small-scale gold prospecting and mining began in this area in the early 1850s, but the first mother lode wasn't discovered until 1885. Two itinerant prospectors, George Walker and George Harrison, stumbled on surface outcrops of gold-rich conglomerate on an old farm-land that is now near the center of Johannesburg. In what must be one of the biggest financial blunders in history, both men quickly sold their claims for the equivalent of a few hundred dollars. Today, the gold fields in the region are worth many billions.

The Witwatersrand basin, which covers an area about the size of West Virginia, contains almost as much gold as the rest of the Earth's surface combined. There is, of course, enormous interest in the origin of these deposits, and more than a hundred years of mining and scientific research have revealed a complex history for these gold-bearing conglomerates. But the chronology of geologic events is im-

perfectly understood and the ideas about how all this gold came to be here are quite controversial.

Our research group at the University of Arizona has recently taken significant steps toward answering questions about the origin of the Witwatersrand gold. Here we discuss the research and the implications of our results for the geologic history of South Africa.

Gold Records

There are two major hypotheses for the origin of gold within the Witwatersrand basin—the "placer" model and the "hydrothermal" model. Both concepts date back more than 100 years, and each has traded places several times with the other as the favorite among scientists. Determining which of these theories is correct not only concerns earth scientists who wish to unlock the geologic past, but it also has great economic significance for mining companies. The exploratory strategies for gold within the Witwatersrand basin and other parts of the world are continually being modified according to current scientific models.

Everyone agrees that the sediments of the Witwatersrand were originally carried in by a system of braided rivers that eroded material from the surrounding highlands and deposited clay, sand and gravel at the edge of an inland sea (or possibly a great lake). As the rivers emptied into this vast body of water, heavier sediments, such as large quartz pebbles and heavy minerals, settled first, building gravel-rich deltas close to the shoreline, whereas sand and clay were carried farther out to greater depths. Over millions of years, fluctuations in sea level continued to change the position of the river-

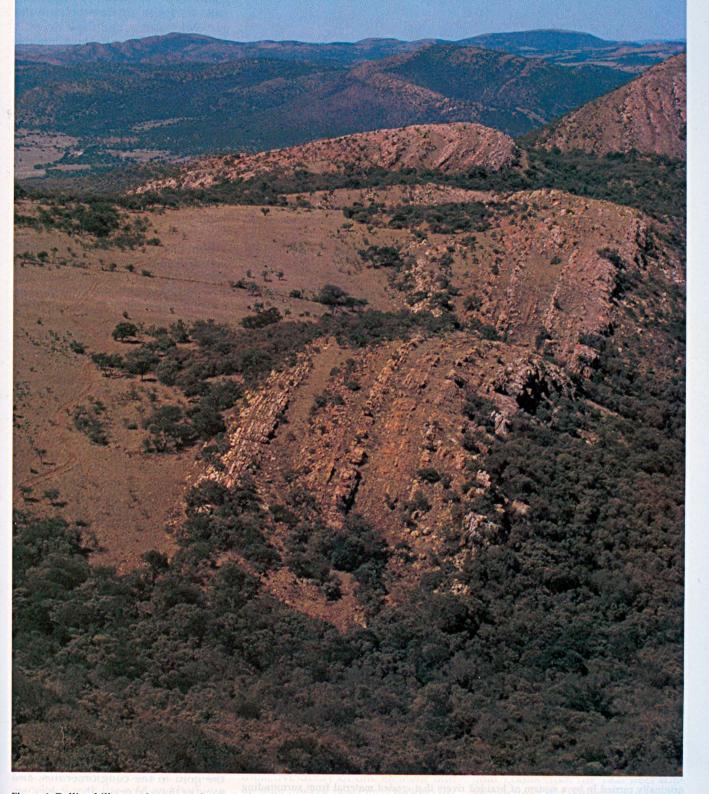


Figure 1. Rolling hills near the center of the Witwatersrand basin show highly fractured sandstone beds turned nearly vertical, or even overturned in places. They are part of the Vredefort dome, which was produced by the impact of a large meteor about two billion years ago. The collision tilted gold-bearing sedimentary layers and so helped partially preserve these rocks from erosion until the present day. The basin is host to a number of peculiar geologic structures, which may help to explain the mineralogical riches in the region. (Except where noted, all photographs courtesy of the authors.)



Figure 2. Witwatersrand basin, which contains nearly half the world's gold, covers an area the size of West Virginia in the northeastern part of South Africa.

sea interface causing the deltaic gravels to be covered by sand and clay layers, which were in turn covered by other gravels, and later more sand and clay. This sequence grew many kilometers thick and was then overlaid by large eruptions of lava (called *flood basalts*) and more sediment. The weight of these layers provided the heat and pressure necessary to transform the unconsolidated sediments into coherent sedimentary rocks.



Figure 3. Conglomerate rock, consisting of quartz pebbles cemented together, contains most of the gold in the sedimentary rocks of the Witwatersrand basin. The sediments in the basin were originally carried in by a system of braided rivers that eroded material from surrounding highlands and deposited them at the edge of a large inland body of water. The source of the gold within the sediments is hotly disputed (see Figure 4). This conglomerate layer, the Kimberly Reef, is about 1.5 kilometers beneath the surface within the Evander goldfields. At this particular location the ore contains about one kilogram of gold for every metric ton of rock, which is extremely high, even for the Witwatersrand basin.

The placer model holds that these rivers carried small grains of gold and rounded pyrite ("fool's gold") into the basin. Because of their high density, the gold and pyrite fell out of suspension with the larger quartz pebbles in the gravel-rich deltas, and these deposits were eventually transformed into the conglomerates being mined today.

The hydrothermal model states that the sediments that washed into the basin contained very little or no gold. Instead, gold-rich hot fluids emanating from deep within the Earth's crust, and traveling along faults and fractures, added gold to the basin long after the sediments consolidated into rock. The gold precipitated from these fluids along chemically favorable horizons within the basin, corresponding to the layers of conglomerate.

Both theories agree that most of the gold *appears* to be hydrothermal—it is concentrated in small fractures and around pyrite and carbon within the conglomerates. So, on the face of it, the hydrothermal camp seems to have a fairly strong case. But, as we shall see, all that glitters may not be hydrothermal, and the real answer may require some further geological sleuthing.

Indeed, dozens of scientific papers in the past two decades have offered numerous lines of evidence for (or against) each of the two models. One important observation is that gold is confined almost exclusively to the conglomerates. Supporters of the placer model argue that this correspondence shows that the gold was deposited under the transition from high fluid energy to low, which caused the gravel of the conglomerates to accumulate beneath the river deltas. Supporters of the hydrothermal model counter that the conglomerates fracture more readily than other rocks under the stress of tectonic forces, and the resulting cracks would therefore provide the best conduits for gold-bearing fluids. In this view, carbon and iron in the conglomerates change the local oxidation state of the fluid and act as precipitation sites, bringing the gold out of solution.

Another interesting observation is that much of the pyrite associated with the gold in the conglomerates, and some of the gold grains themselves, are rounded. In the placer model, rounded pyrite and gold result from abrasion during stream transport and wind action during deposition. In the hydrothermal model, dissolved sulfur in the hydrothermal fluids would react with rounded iron-oxide mineral grains (magnetite), replacing the oxygen in the minerals with sulfur, and creating rounded pyrite. Most proponents of this model dispute the existence of rounded gold grains.

Because observations such as these can accommodate either model, a "smoking gun" is needed to choose between the two theories. One possibility is to determine when the gold was mineralized. If the gold grains are older than their host conglomerate, then they must have come from a source that predated the sedimentation. In this view river waters eroded the gold from older source terrains and transported it, along with other sediments, into the basin—the placer model. If the gold grains are younger than their host rocks, then hot groundwater must have added them after the conglomerates were deposited—the hydrothermal model.

The test sounds simple, but gold mineralization has been notoriously difficult to date directly. Previous attempts have relied on dating minerals that often coexist with gold, such as mica, pyrite or uraninite. These ages are used as proxies for the age of the gold but may in fact date events millions of years before or after the gold was actually formed.

Using other minerals to date the gold has been especially problematic in the Witwatersrand basin. Some materials associated with the Witwatersrand gold give ages older than the host conglomerates whereas others give younger ages. Pyrite is a good example; it is intimately associated with gold in the Witwatersrand conglomerates and mining geologists often associate large abundances of pyrite with high-grade gold. We have determined that the ages of rounded, compact pyrite grains are older than the host conglomerates, supporting ages determined by other workers and the supposition that they were rounded by stream transport. Cubic crystals of pyrite, which are almost certainly hydrothermal in origin, give less precise but younger ages than the conglomerates. Both types of pyrites are spatially associated with the gold and both can be used to support either model of gold deposition.

Gold grains are difficult to date because they are composed primarily of

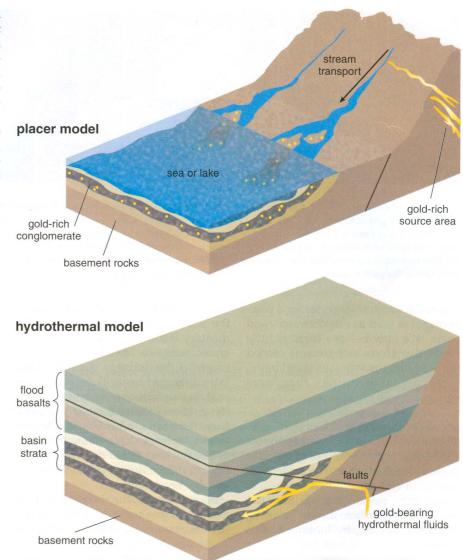


Figure 4. Two major theories—the placer model and the hydrothermal model—attempt to account for the origin of gold in the Witwatersrand basin. The placer model holds that gold was eroded from a pre-existing source and transported into the basin with other sediments forming the conglomerate strata. Wind and wave action further concentrated the gold near an ancient shoreline. In contrast, the hydrothermal model argues that hot fluids from deep within the Earth's crust carried the gold along faults and fractures within the basin conglomerates long after they were consolidated into rock. In the placer model the gold was mineralized before it was deposited in the basin sediments. In the hydrothermal model the gold was mineralized after the sediments were transformed into rocks. The authors tested these models by directly dating the age of the gold and the conglomerate rock (see Figure 5).

elemental gold and minor amounts of silver and mercury and even lesser amounts of bismuth, selenium, platinum group elements and other metals, such as rhenium. Most of these elements are isotopically stable, so dating techniques that rely on the radioactive decay of one element into another are not possible. The lone exception is an isotope of rhenium, rhenium-187 (187Re), which radioactively decays over time into Osmium-187 (187Os) at a known rate. New analytical techniques now allow measurements of the extremely small amounts of rhenium and osmium found in gold, making it possible to determine its age directly.

We recently employed this method to determine a very precise age for gold grains from the Vaal Reef conglomerate of the Witwatersrand basin. It turns out that the gold minerals are 3.01 billion years old—significantly older than the host conglomerates, which are 2.76 to 2.89 billion years old.

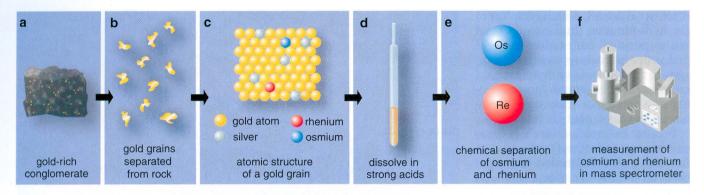


Figure 5. Gold can be dated by measuring the radioactive isotope rhenium–187, which decays into osmium–187 at a known rate. First the gold grains must be physically separated from the conglomerate (*a and b*). Trace amounts of rhenium and osmium, which were trapped within the gold when it was mineralized, are removed by dissolving the grains in strong acids (*c and d*). The rhenium and osmium atoms are chemically separated from each other (*e*), and the relative amounts of rhenium–187 and osmium–187 are independently measured in a mass spectrometer (*f*). This technique revealed that the Witwatersrand gold was mineralized before the basin sediments were deposited, which supports the placer model of the gold's origin.

The result supports theories for a placer origin of gold in the Witwatersrand basin. We now believe there is little doubt that rivers and streams carried the gold into the Witwatersrand basin, probably in quantities that were unique in geologic history.

Searching for Eldorado

This leaves us with the question of where this vast amount of gold came from in the first place. For contemporary gold-rich stream sediments it is sometimes possible to follow the stream back to where the gold is being eroded. Likewise, by looking at features such as the size and orientation of the pebbles and orientations of sedimentary features within the conglomerates, scientists have been able to reconstruct the drainage patterns of the Witwatersrand basin. These studies reveal that ancient river systems brought the gold and the sediments primarily from the north and the west. Despite years of intense exploration, however, geologists have failed to locate the figurative mountain of gold at these primordial headwaters.

As it happens, the rhenium and osmium isotopes may also help identify the source of the gold—whether it

comes from the Earth's crust or from below that, somewhere in the mantle. The method we used to determine the age of the gold gives two additional pieces of information: the initial composition of the osmium isotopes and the concentration of rhenium and osmium in the gold. For ease of measurement and comparison, ¹⁸⁷Re and ¹⁸⁷Os (which is the daughter isotope produced by the decay of ¹⁸⁷Re) are referenced to a stable isotope of osmium, ¹⁸⁸Os. The more rhenium a rock or mineral contains initially, and the older it is, the higher the resulting ¹⁸⁷Os/¹⁸⁸Os ratio. Therefore, we can find an age for the formation of the gold by measuring the ¹⁸⁷Re/¹⁸⁸Os and the ¹⁸⁷Os/¹⁸⁸Os ratios in the gold today. We can also calculate the ¹⁸⁷Os/¹⁸⁸Os ratio for when the gold was formedthe so-called initial Os isotopic ratio, $^{187}\mathrm{Os}/^{188}\mathrm{Os}_{i}.$ The $^{187}\mathrm{Os}/^{188}\mathrm{Os}_{i}$ ratio at the age of formation can then be compared to the ¹⁸⁷Os/¹⁸⁸Os ratio of different crustal rocks and the mantle of the same age.

It turns out that the mantle has relatively low amounts of rhenium compared with osmium, whereas the crust generally has higher amounts. This is because crustal rocks are the products of partial melting of the mantle (and potentially re-melting of previously formed crust) and rhenium goes more readily into the melt. So as crust evolves, it develops ¹⁸⁷Os/¹⁸⁸Os ratios much greater than the mantle over the same time frame. In a few tens of millions of years, the ¹⁸⁷Os/¹⁸⁸Os ratio of the mantle and the crust diverge rapidly. Most crustal rocks develop elevated ¹⁸⁷Os/¹⁸⁸Os ratios quickly,



Figure 6. Grains of rounded pyrite (brass-colored ovoids), which are often associated with gold grains (not visible here), sit atop a quartz pebble (grayish white) in this sample from the Witwatersrand conglomerate rocks. According to the placer model, the rounded pyrite and some of the gold grains were shaped by abrasion during stream transport. The rounded grains are harder to explain in the hydrothermal model.

whereas the ¹⁸⁷Os/¹⁸⁸Os ratios of the much more voluminous mantle change very little. Thus gold that originated from the mantle will have a very different osmium "fingerprint" compared with gold derived from crustal rocks.

The ¹⁸⁷Os/¹⁸⁸Os ratio of the threebillion-year-old gold from the Witwatersrand basin is the same as that of the Earth's mantle three billion years ago. It has long been recognized that episodes of metamorphism caused by various tectonic events have led to infiltration of hydothermal fluids throughout the basin. These tectonic events mobilized fluids from within the continental crust between 2.7 and 2.0 billion years ago. If these hydrothermal fluids, which originated in the crust, had deposited the Witwatersrand gold, then osmium in these fluids and gold that was precipitated from the fluids should contain elevated ¹⁸⁷Os/¹⁸⁸Os ratios, much as the crustal rocks themselves. But the Witwatersrand gold has low ¹⁸⁷Os/¹⁸⁸Os values, much like that of the three-billion-yearold mantle, suggesting that Witwatersrand's gold was not originally derived from normal crust. Instead, it originated directly from the mantle or from a particular class of rocks called komatiites, which are rich in magnesium and sulfur and are made from upper mantle that was melted at very high temperatures.

Furthermore, the mineralized Witwatersrand gold has very high concentrations of both rhenium and osmium relative to younger conglomerate-hosted gold deposits, hydrothermal deposits and average concentrations in the continental crust. Gold from the Witwatersrand basin has rhenium and osmium concentrations that show a very clear affinity with mantle samples and with komatiites. Komatiites were formed almost exclusively in the Archaean Era—2.5 billion years ago and older—and are found predominantly in the ancient centers of the continents. Even though komatiites are crustal rocks in the strict sense, the high-temperature conditions associated with their genesis also causes a high proportion of the mantle to melt, and so imparts mantle-like characteristics to the komatiites. This includes qualities such as relatively high proportions of gold, other platinum group elements and osmium with mantle-like ¹⁸⁷Os/¹⁸⁸Os ratios.

The sediments found in the Witwa-

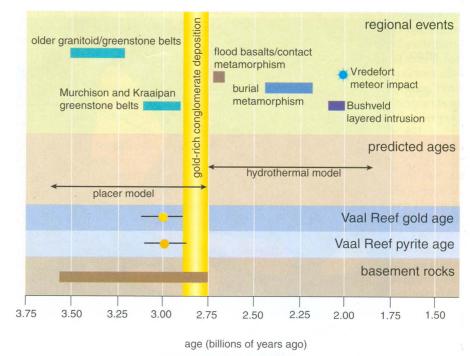


Figure 7. Geologic events in the Witwatersrand basin provide a historical context for the deposition of the gold in the region. Radioisotope-dating techniques reveal that gold and pyrite from the Vaal Reef deposits in the basin are older than the conglomerate rocks in which they are located, which is consistent with the placer model (*see Figure 5*). Several kilometers of lava and other rocks (such as the flood basalts), progressively buried the gold-rich conglomerates, producing enough pressure and heat to transform the sediments into rock. The Vredefort meteor impact, which may have helped preserve the rocks from later erosion (*see Figure 1*), took place about 2 billion years ago.

tersrand basin are made up of minerals that have long been recognized to originate from granite-greenstone belts—terrains made up of greenstone, a metamorphosed basalt or komatiite, and intruded by granite domes. The

nature of the sediments and the mantle-like osmium concentration and composition of the gold, make the gold-bearing komatiites our favored source for the Witwatersrand gold.

There are two areas that might serve

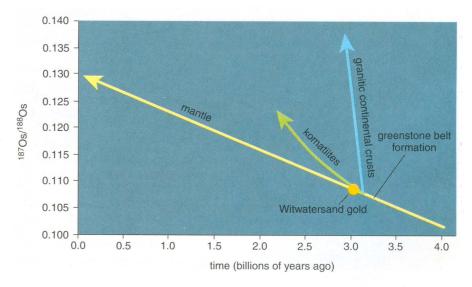


Figure 8. Osmium composition—the ratio of ¹⁸⁷Os to ¹⁸⁸Os—evolves at different rates in different types of rocks. The osmium ratio of the 3-billion-year-old Witwatersrand gold is similar to that of similarly aged mantle and rocks known as komatiites, suggesting that they may be the original source of the gold that was eroded into the basin.

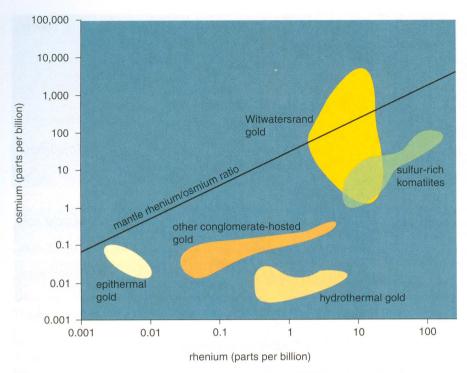


Figure 9. Concentrations of osmium and rhenium in the Witwatersrand gold are similar to those of the mantle and of sulfur-rich komatiites. Hydrothermal gold and gold from deposits in other parts of the world tend to have lower concentrations of osmium. These results further support a placer model in which the Witwatersrand gold is eroded from highlands containing mantle rocks or komatiites.

as the source of these komatiites: the Kraaipan granite-greenstone belt and the Murchison granite-greenstone belt. These belts are found to the west and the north of the Witwatersrand basin—exactly where reconstructions of the river drainage patterns suggest they

should be. Moreover, many of the rocks in these belts are approximately the same age as the Witwatersrand gold, about three billion years. We are currently analyzing gold from within the Kraaipan and Murchison rocks and this should soon determine whether it

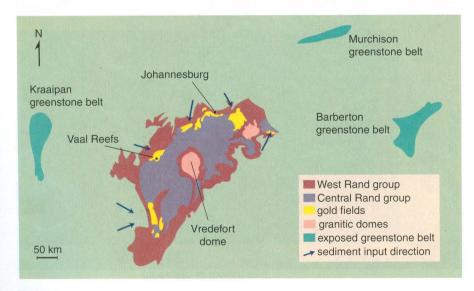


Figure 10. Orientations of elongated cobbles and sedimentary features within the gold-bearing conglomerates in the Witwatersrand basin show the direction (*arrows*) of ancient river flow, which is consistent with the hypothesis that the gold was eroded from old greenstone belts—containing komatiites and mantle rocks—to the north and west of the basin.

is also the same age and composition as the Witwatersrand gold.

Will a close age correspondence start a new gold rush to these terrains of South Africa? This seems unlikely. The rocks of the Kraaipan and Murchison belts contain only slightly elevated concentrations of gold relative to normal crust, and they are not rich enough to be of much economic interest on their own. The low concentrations of gold in these granite-greenstone belts suggest that the gold-rich parts have already been eroded away or that younger rocks cover them or that the Witwatersrand's depositional processes (wind and wave action) concentrated the available gold.

There are other places in the world—for example, in Jacobina, Brazil, and Blind River, Canada—with conglomerate formations that are almost identical to the Witwatersrand conglomerates, except that they are younger and have much smaller quantities of gold. Did these other deposits simply lack the gold-rich source terrains that fed the Witwatersrand basin? The source rocks for these younger conglomerate deposits are also granitegreenstone terrains, but they are hundreds of millions of years younger than the Witwatersrand source rocks. The Earth's mantle loses heat exponentially and so younger greenstone terrains form from melting much smaller proportions of the solid mantle at lower temperatures. A higher percentage of mantle melting may imply that more gold can go into the melt. Is the richness of the Witwatersrand source rocks simply a result of their age? We don't know the answers yet.

A Golden Age

The new evidence from our rhenium and osmium analyses and the work of many others provides a clearer picture of the history of gold mineralization in the Witwatersrand basin. Scientists now know that volcanic eruptions and granitic intrusions produced the nuclei of the South African continental crust—such as the Barberton granitegreenstone belt—over three and a half billion years ago. These terrains provided a foundation on which other volcanic arcs and plateaus were progressively accreted through the action of plate tectonics. The Kraaipan greenstone belt and the Murchison greenstone belt were two of the terrains (approximately 3.1 to 2.7 billion years

old) that were plastered onto the northern and western portions of the continental nucleus. The region of the mantle that fed these terrains may have been extremely rich in gold or the melting processes that generated the komatiites may have been exceptionally efficient at extracting gold out of the mantle rock. After the Kraaipan and Murchison greenstone belts attached themselves to the continental nucleus, the crust became stable enough to form one of the world's first large sedimentary basins. Waters flowing from the high relief of these terrains carried gold into the neighboring inland sea.

By placing age constraints on the mineralization of the Witwatersrand basin, and by confirming the sedimentary nature of the gold, our work adds information to current debate regarding studies of the nature of the Archaean atmosphere—the era when the sediment was transported and deposited. The presence of uranium minerals and pyrite in the sediments suggests that the atmosphere was very different 2.9 to 2.7 billion years ago. In today's environment, minerals such as pyrite and uraninite are unstable because of atmospheric oxygen and do not survive long in surficial waters. Extremely low concentrations of oxygen in the atmosphere are required for pyrite and uranium minerals to be stable in surficial waters. This supports current theories based on Precambrian iron deposits, oxidized sandstones and ancient soils, which suggest that oxygen was not abundant in the Earth's atmosphere in the Early Archaean and did not reach present-day levels until much later.

Approximately 2.7 billion years ago, the supply of sediments into the basin ended as voluminous flood basalts covered the Witwatersrand basin. More than seven kilometers of sedimentary rocks interspersed with goldrich conglomerates were buried beneath several more kilometers of lavas and other rocks. This progressive burial eventually produced enough pressure and heat to squeeze water out of the sediments—compacting, cooking and cementing the sediments into rock. The fluids released by burial may have dissolved some of the originally sedimentary gold and moved it over very short distances, while leaving some of the gold untouched. This is evident by the presence of both gold grains with

rounded disk-like shapes (which are sedimentary) and gold grains with irregular dendritic shapes (which are hydrothermal).

Partial dissolution of the gold and local re-precipitation not only explains the two different morphologies of gold, but also helps explains why the rhenium-osmium clock wasn't reset by this local mobilization. Rhenium and osmium are generally very insoluble and therefore, hydrothermal fluids contain very low concentrations and transport only small amounts of these elements. Witwatersrand gold on the other hand has orders of magnitude higher concentrations of rhenium and osmium relative to hydrothermal fluids and crustal rocks. Therefore almost all of the rhenium and osmium would stay with the undissolved gold and it would therefore retain the three-billion-year age. The age of the gold that was locally mobilized would be reset and be much younger than that of the undissolved gold. However, because of the extreme differences in osmium concentration between the placer gold and the mobilized gold, the rheniumosmium clock would not be affected and all of the gold would appear to have the three-billion-year-old age.

Two other events that transpired well after the gold was deposited may help explain why South Africa is so rich in valuable minerals. The first was a large upwelling of magma from the mantle about 100 kilometers northeast of the Witwatersrand basin. which resulted in the Bushveld igneous complex, the world's largest source of platinum and chromium. This event also triggered fluids to move through the conglomerates of the Witwatersrand basin once again and may have moved some of the gold very short distances.

The second event took place about 30 million years after the emplacement of the Bushveld: A huge meteor slammed into the sedimentary rocks at the center of the Witwatersrand basin. This collision was larger than the dinosaur-killing impact that produced the Chicxulub crater in Mexico. The South African blast brought up rocks from the lower crust and upper mantle and tilted the layers of the basin. The tilting of the gold-bearing sedimentary layers may have helped partially preserve these rocks from later erosion until the present day.

Since those two events the continental crust of South Africa has been geologically stable. The gold-rich conglomerates waited undisturbed for billions of years until prospectors discovered the tip of a gold iceberg, now being explored by tens of thousands of mine workers laboring in the deepest holes on Earth.

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