

## JOINT TYPHOON WARNING CENTER GUAM, MARIANA ISLANDS

$\mathcal{F R O N T T}$ COVER: The digitized image (center of Lower square) of surface wind speed shows Typhoon Kelly (19W) (circular pattern at the top right) and the Pfilippine Islands (black sfapes at Gottom left). The technique to develop these surface wind speed fields is currently under development. The surface wind speed field algorithm uses the polarized 19 (horizontal), 22 (vertical) and 37 (vertical and horizontal) gigafiertz ( $\mathcal{G H z}$ ) channels of the Defense Meteorological Satelfite Program's new special sensor, the microwave imager (SSM/I), which is a four-channel passive microwave radiometer.

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To: Distribution
Subj: PROMULGATION OF 1987 ANNUAL TROPICAL CYCLONE REPORT
Ref: (a) USCINCPACINST 3140.1S (NOTAL)

1. The Annual Tropical Cyclone Report for 1987 is promulgated in accordance with the provisions of reference (a).
2. The 1987 tropical cyclone season marked the beginning of a new era in tropical cyclone forecasting. Despite an unusually active season, forecasters made a mid-season transition from an aircraft based reconnaissance system to one based mostly on satellites, while recording the lowest track errors in the center's history.
3. The initial release of the Joint Typhoon Warning Center Automation Program hardware and software package also arrived on Guam in 1987. This program has already proved very successful and promises to be one of the most significant advances in the operational forecasting of tropical cyclones.
4. Despite the tremendous added pressures of the 1987 season, the staff has pulled together and done an outstanding job in publishing this document six months ahead of last year. I hope you find this report a valuable contribution to your library.


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## FOREWORD

The Annual Tropical Cyclone Report is prepared by the staff of the Joint Typhoon Warning Center (JTWC), a combined USAF/USN organization operating under the command of the Commanding Officer, U.S. Naval Oceanography Command Center/Joint Typhoon Warning Center, Guam. JTWC was established in April 1959 when USCINCPAC directed USCINCPACFLT to provide a single tropical cyclone warning center for the western North Pacific region. The operations of JTWC are guided by CINCPACINST 3140.1 (series).

The mission of the Joint Typhoon Warning Center is multi-faceted and includes:

1. Continuous monitoring of all tropical weather activity in the northem and southern hemispheres, from 180 degrees longitude westward to the east coast of Africa, and the prompt issuance of appropriate advisories and alerts when tropical cyclone development is anticipated.
2. Issuing warnings on all significant tropical cyclones in the above area of responsibility.
3. Determination of reconnaissance requirements for tropical cyclone surveillance and assignment of appropriate priorities.
4. Post-storm analysis of all significant tropical cyclones occurring within the western North Pacific and North Indian Oceans, which includes an in-depth analysis of tropical cyclones of note and all typhoons.
5. Cooperation with the Naval Environmental Prediction Research Facility (NEPRF), Monterey, California, on the operational evaluation of tropical cyclone models and forecast aids, and the development of new techniques to support operational forecast scenarios.

Satellite imagery used throughout this report represents data obtained by the tropical cyclone satellite surveillance network. The personnel of Detachment 1, 1WW, collocated with JTWC at Nimitz Hill, Guam, coordinate the satellite acquisitions and tropical cyclone surveillance with the following units:

Det 4, 20WS, Hickam AFB, Hawaii
Det 5, 20WS, Clark AB, RP
Det 8, 20WS, Kadena AB, Japan
Det 15, 30WS, Osan AB, Korea
Air Force Global Weather Central, Offutt AFB, Nebraska

In addition, the Naval Oceanography Command Detachment, Diego Garcia, and Defense Meteorological Satellite Program (DMSP) equipped U.S. Navy aircraft carriers have been instrumental in providing vital satellite position fixes of tropical cyclones in the Indian Ocean.

Should JTWC become incapacitated, the Alternate Joint Typhoon Warning Center (AJTWC) located at the U.S. Naval Western Oceanography Center, Pearl Harbor, Hawaii, assumes warning responsibilities. Assistance in determining satellite reconnaissance requirements, and in obtaining the resultant data, is provided by Det 4, 20WS Hickam AFB, Hawaii.

Changes to this year's publication include: statistical verification for individual wamings for the North Indian Ocean and the southem hemisphere are provided. Again, as last year, raw fix data files previously printed in Annex A, plus the raw warning, forecast and best track data, will be available, upon request (the requested data will be copied onto 5.25 inch "floppy" diskettes provided by the requestor); and, with reference to best track philosophy, a conscious effort has been made to extend the post-warning best tracks to provide better verification for the 48 - and 72 -hour forecasts.

A special thanks is extended to the men and women of: 27th Information Systems Squadron, Operating Location C and the Operations section of the Naval Oceanography Command Center, Guam for their continuing support by providing high quality real-time satellite imagery; Marine Corps Air Station, Futenma, Japan for their satellite fix support; the Pacific Fleet Audio-Visual Center, Guam for their assistance in the reproduction of satellite data for this report; to the Navy Publications and Printing Service Branch Office, Guam; the Royal Observatory Hong Kong for supporting synoptic data on Super Typhoon Lynn (20W); the Central Weather Bureau, Taiwan for radar scope photographs of Typhoons Vernon (06W), Alex (08W) and Gerald (14W); Dr. Bob Abbey of the Office of Naval Research for his technical support to this publication; Mr. Michael Fiorino at NEPRF for his software conversion for the statistical programs; Mr. S.D. Rice, manager of Mobil Oil Micronesia, Inc. for his damage photos of Ulithi Atoll after Typhoon Orchid (01W); Dr. Greg Holland for sharing the ship's log account of Typhoon Lynn (20W); and Captain K. W. Reese (USAF) for the reconnaissance photograph of Typhoon Wynne (07W)

Note: Appendix IV contains information on how to obtain past issues of the Annual Tropical Cyclone Report (titled Annual Typhoon Report prior to 1980).

## TABLE OF CONTENTS


3. North Indian Ocean Tropical Cyclones ..... 138
INDIVIDUAL TROPICAL CYCLONES
TROPICAL CYCLONE AUTHOR PAGE TROPICAL CYCLONE AUTHOR PAGE
TC 01B DREKSLER 140 TC 05B STEWART ..... 148
TC 02B SCOVIL ..... 144 TC 07A MUNDELL ..... 152
TC 04B FATJO 146 TC 08B FATJO ..... 154
CHAPTER IV SUMMARY OF SOUTH PACIFIC AND
SOUTH INDIAN OCEAN TROPICAL CYCLONES

1. General ..... 157
2. South Pacific and South Indian Ocean Tropical Cyclones ..... 160
CHAPTER V SUMMARY OF FORECAST VERIFICATION
3. Annual Forecast Verification ..... 163
4. Comparison of Objective Techniques ..... 173
CHAPTER VI TROPICAL CYCLONE SUPPORT SUMMARY
5. Naval Environmental Research Prediction Facility ..... 179
6. Joint Typhoon Warning Center ..... 181
ANNEX A TROPICAL CYCLONE TRACK AND FIX DATA 1. General ..... 183
7. Warning Verification Statistics ..... 183
APPENDICES I. Definitions ..... 209
II. Names of Tropical Cyclones ..... 210
III. References ..... 211
IV. Past Annual Tropical Cyclone Reports ..... 213

## CONTRACTIONS

| ABIO | Significant Tropical <br> Weather Advisory for the <br> Indian Ocean | CM | Centimeter | INJAH | North Indian Ocean <br> Component of TYAN |
| :--- | :--- | :--- | :--- | :--- | :--- |
| ABPW | Significant Tropical <br> Weather Advisory for the <br> Western Pacific Ocean | CPA | Cyclops Objective Steering <br> Model Output Statistics | INST | Instruction |


| NET | Near-Equatorial Trough | SLP | Sea-Level Pressure | TUTT | Tropical Upper- <br> Tropospheric Trough |
| :--- | :--- | :--- | :--- | :--- | :--- |
| NM | Nautical Mile(s) | SRP | Selective Reconnaissance <br> Program | ULAC | Upper-Level Anticyclone |
| NOAA | National Oceanic and <br> Atmospheric | STNRY | Stationary | ULCC | Upper-Level Circulation |
|  | Administration | SST | Sea Surface Temperature |  | Center |

## CHAPTER I - OPERATIONAL PROCEDURES

## 1. GENERAL

The Joint Typhoon Warning Center (JTWC) provides a variety of routine services to the organizations within its area of responsibility, including:
a. Significant Tropical Weather Advisories: issued daily, these products describe all tropical disturbances and assess their potential for further development during the advisory period;
b. Tropical Cyclone Formation Alerts: issued when synoptic, satellite and/or aircraft reconnaissance data indicate development of a significant tropical cyclone in a specified area is likely;
c. Tropical Cyclone Warnings: issued periodically throughout each day for significant tropical cyclones, giving forecasts of position and intensity of the system; and
d. Prognostic Reasoning Messages: issued twice daily for tropical storms and typhoons in the western North Pacific; these messages discuss the rationale behind the most recent JTWC warnings.

The recipients of the services of JTWC essentially determine the content of JTWC's products according to their ever changing requirements. Therefore, the spectrum of routine services is subject to change from year to year. Such changes are usually the result of deliberations held at the Annual Tropical Cyclone Conference.

## 2. DATA SOURCES

## a. COMPUTER PRODUCTS:

A standard array of synoptic-scale computer analyses and prognostic charts are available from the Fleet Numerical Oceanography Center (FLENUMOCEANCEN) at Monterey, California. These products are provided to JTWC via the Naval Environmental Data Network (NEDN).

## b. CONVENTIONAL DATA:

This data set is comprised of land-based and shipboard surface and upper-air observations taken at, or near, synoptic times, cloud-motion winds derived twice daily from satellite data, and enroute meteorological observations from commercial and military aircraft (AIREPS) within six hours of synoptic times. Conventional data charts are prepared daily at 0000 Z and 1200 Z using computer- and hand-plotted data for the surface/gradient and 200 mb (upper-tropospheric) levels. In addition to these analyses, charts at the $925,850,700$, 500 and 400 mb levels are computer-plotted from rawinsonde/pibal observations at the 12 hour synoptic times.

## c. AIRCRAFT RECONNAISSANCE:

Data provided by aircraft weather reconnaissance are invaluable for locating the position of the center of developing systems and essential for the accurate determination of:

- maximum surface and flight-level wind
- minimum sea-level pressure
- horizontal surface and flight-level wind distribution
- eye/center temperature and dew point

In addition, wind and pressure-height data at the $500 \mathrm{and} /$ or 400 mb levels, provided by the aircraft while enroute to, or from fix missions, or during dedicated synoptic-scale flights, provide a valuable supplement to the all too sparse data fields of JTWC's area of responsibility. A more detailed discussion of aircraft weather reconnaissance is presented in Chapter II.

## d. SATELLITE RECONNAISSANCE:

Meteorological satellite data obtained from the Defense Meteorological Satellite Program (DMSP) and National Oceanic and Atmospheric Administration (NOAA) spacecraft played a major role in the early
detection and tracking of tropical cyclones in 1987. A discussion of the role of these programs is presented in Chapter II.

## e. RADAR RECONNAISSANCE:

During 1987, as in previous years, landbased radar coverage was utilized extensively when available. Once a tropical cyclone moved within the range of land-based radar sites, their reports were essential for determination of small-scale movement. Use of radar reports during 1987 is discussed in Chapter II.

## f. DRIFTING METEOROLOGICAL BUOYS:

JTWC received wind speed, sea-level pressure, sea surface temperature and air temperature reports from six drifting meteorological buoys deployed by the U. S. Navy beginning in the middle of June 1987. One line of three buoys was deployed along 7 degrees North Latitude from south of Guam eastward toward the Marshall Islands. Another set of three was deployed along 11 degrees North Latitude from southwest of Guam eastward through the Caroline Islands. The buoys performed flawlessly throughout most of the western North Pacific tropical cyclone season. At the end of the year, four buoys continued to operate, one no longer transmitted data and annother was apparently taken to Tandag City, Mindanao, R.P., where it continued to transmit. The three northernmost buoys tracked basically westward and covered 25 to 35 degrees of longitude. The southern buoys drifted more slowly and erratically. One buoy either snagged its drouge on a submerged reef east of Woleai Atoll or became trapped in an eddy within the island/reef chain.

JTWC received at least one position update from each buoy per day and up to eight meteorological data updates per buoy per day. Buoy data were consistent with the data from other conventional sources to the extent that they was considered to be, in most cases, more reliable and more accurate than ship reports and some island reporting stations. As a backup and position check, JTWC also received buoy data, on a delay basis, over the AWN (Manop header SSVX6 LFPW).

An expanded buoy network for the 1988 tropical cyclone season is being planned.
a. JTWC currently has access to three primary communications circuits.

## 3. COMMUNICATIONS

(1) The Automated Digital Network (AUTODIN) is used for dissemination of warnings, alerts and other related bulletins to Department of Defense installations. These messages are relayed for further transmission over U.S. Navy Fleet Broadcasts, and U.S. Coast Guard CW (continuous wave Morse Code) and voice broadcasts. Inbound message traffic for JTWC is received via AUTODIN addressed to NAVOCEANCOMCEN GQ or DET 1, 1WW NIMITZ HILL GQ.
(2) The Air Force Automated Weather Network (AWN) provides weather data to JTWC through a dedicated circuit from the Automated Digital Weather Switch (ADWS) at Hickam AFB, Hawaii. The ADWS selects and routes the large volume meteorological reports necessary to satisfy JTWC requirements for the right data at the right time. Weather bulletins prepared by JTWC are inserted into the AWN circuit via the Naval Environmental Display Station (NEDS) through the Nimitz Hill Naval Telecommunications Center (NTCC) of the Naval Communications Area Master Station Western Pacific.
(3) The Naval Environmental Data Network (NEDN) is the communications link with the computers at FLENUMOCEANCEN. JTWC is able to receive environmental data from FLENUMOCEANCEN and provide data directly to the computers to execute numerical techniques.
b. NEDS has been the backbone of the JTWC communications system for several years. Currently, JTWC is undergoing an upgrade that will make use of microcomputer technology and automate much of the work that goes into message preparation and transmission. This will decrease the work load on the NEDS and allow JTWC to interface directly with NTCC for AWN and AUTODIN messages.

## 4. ANALYSES

A composite surface/gradient-level ( 3000 ft ( 914 m )) manual analysis of the JTWC area of responsibility is accomplished daily on the 0000 Z and 1200 Z conventional data. Analysis of the wind field using streamlines is stressed for tropical and subtropical regions. Analysis of the pressure field outside the tropics is accomplished routinely by the Naval Oceanography Command Center Operations watch team and is used by JTWC in conjunction with their analysis of the tropical wind fields.

A composite upper-tropospheric manual streamline analysis is accomplished daily utilizing rawinsonde data from 300 mb through 100 mb , winds obtained from satellitederived cloud motion analysis, and AIREPS (taken plus or minus three hours of chart valid time) at or above 31,000 feet $(9,449 \mathrm{~m})$. Wind and height data are used to generate a representative analysis of tropical cyclone outflow patterns, mid-latitude steering currents, and features that may influence tropical cyclone intensity. All charts are hand-plotted in the tropics to provide all available data as soon as possible to the Typhoon Duty Officer (TDO). These charts are augmented by computerplotted charts for the final analysis.

Computer-plotted charts for the 925 , $850,700,500$ and 400 mb levels are available for streamline and/or height-change analysis from the 0000 Z and 1200 Z data base. Additional sectional charts at intermediate synoptic times and auxiliary charts, such as station-time plot diagrams and pressure-change charts, are also analyzed during periods of significant tropical cyclone activity.

## 5. FORECAST AIDS

The following objective techniques were employed in tropical cyclone forecasting during 1987 (a description of these techniques is presented in Chapter V):

## a. MOVEMENT

(1) 12-HOUR EXTRAPOLATION
(2) CLIMATOLOGY
(3) COSMOS (Model Output Statistics)
(4) CSUM (Colorado State University Model)
(5) OTCM (Dynamic Model)
(6) TAPT (Empirical)
(7) HPAC (Half Persistence - Half Climatology Blend)
(8) TYAN78 (Analog)
b. INTENSITY
(1) CLIMATOLOGY
(2) DVORAK (Empirical)
(3) THETA -E (Empirical)
c. WIND RADIUS (Analytical)

## 6. FORECAST PROCEDURES

## a. INITIAL POSITIONING

The warning position is the best estimate of the center of the surface circulation at synoptic time. It is estimated from an analysis of all fix information received up to one and one-half hours after synoptic time. This analysis is based on a semi-objective weighting of fix information based on the historical accuracy of the fix platform and the meteorological features used for the fix. The interpolated warning position reduces the weighting of any single fix and results in a more consistent movement and a warning position that is more representative of the larger-scale circulation. If the fix data are not available due to reconnaissance platform malfunction or communication problems, synoptic data or extrapolation from previous fixes are used.

## b. TRACK FORECASTING

A preliminary forecast track is developed based on an evaluation of the rationale behind the previous warning and the guidance given by the most recent set of objective techniques and numerical prognoses. This preliminary track is then subjectively modified based on the following considerations:
(1) The prospects for recurvature or erratic movement are evaluated. This determination is based primarily on the present and forecast positions and amplitudes of the middle-tropospheric, mid-latitude troughs and ridges as depicted on the latest upper-air analysis and numerical forecasts.
(2) Determination of the best steering level is partly influenced by the maturity and vertical extent of the tropical cyclone. For mature tropical cyclones located south of the subtropical ridge axis, forecast changes in speed of movement are closely correlated with anticipated changes in the intensity or relative position of the ridge. When steering currents are relatively weak, the tendency for tropical cyclones to move northward due to internal forces is an important consideration.
(3) Over the 12 - to 72 -hour (12- to 48 -hour in the southern hemisphere) forecast period, speed of movement during the early forecast period is usually biased towards persistence, while the later forecast periods are biased towards objective techniques. When a tropical cyclone moves poleward, and toward the mid-latitude steering currents, speed of movement becomes increasingly more biased toward a selective group of objective techniques capable of estimating acceleration.
(4) The proximity of the tropical cyclone to other tropical cyclones is closely evaluated to determine if there is a possibility of binary interaction.

A final check is made against climatology to determine whether the forecast track is reasonable. If the forecast deviates greatly from one of the climatological tracks, the forecast rationale may be reappraised.

## c. INTENSITY FORECASTING

For this parameter, heavy reliance is placed on intensity trends from aircraft reconnaissance reports when available, wind and pressure data from ships and land stations in the vicinity of the tropical cyclone, the Dvorak satellite empirical model and climatology. An evaluation of the entire synoptic situation is made, including the location of major troughs and ridges, the position and intensity of any nearby tropical upper-tropospheric troughs (TUTTs), the vertical and horizontal extent of the tropical cyclone's circulation and the extent of the associated upper-level outflow pattern. An essential element affecting each intensity forecast is the accompanying forecast track and the environmental influences along that track, such as terrain, vertical wind shear, and the existence of an extratropical environment.

## d. WIND RADII FORECASTING

Once the forecast intensities have been derived, the horizontal distribution of surface winds (winds greater than $30-, 50-$, and $100-\mathrm{kt}$ ) is determined. The most recent wind radii and associated asymmetrics are deduced from all available surface wind observations and reconnaissance aircraft reports. Based on the current surface wind distribution, preliminary estimates of future wind radii are provided by an empirically derived objective technique (Holland, 1980). These estimates may be subjectively modified based upon the anticipated interaction of the tropical cyclone's circulation with forecast locations of large-scale wind regimes and significant land masses. Other factors including the tropical cyclone's speed of movement and possible extratropical transition are also considered.

## 7. WARNINGS

Tropical cyclone warnings are issued when a closed circulation is evident and maximum sustained winds are forecast to increase to 34 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) within 48 -hours, or if the tropical cyclone is in such a position that life or property may be endangered within $72-$ hours. Warnings may also be issued in other situations if it is determined that there is a need
to alert military or civil interests to threatening tropical weather conditions.

Each tropical cyclone warning is numbered sequentially and includes the following information: the position of the surface center; estimate of the position accuracy and the supporting reconnaissance (fix) platforms; the direction and speed of movement during the past six hours (past 12 -hours in the southern hemisphere); the intensity and radial extent of over $30-, 50$-, and 100 -knot surface winds, when applicable. At forecast intervals of $12-, 24-, 48-$, and 72 -hours (12-, 24-, and 48hours in the southern hemisphere), information on the tropical cyclone's anticipated position, intensity and wind radii are also provided. Vectors indicating the mean direction and mean speed between forecast positions are also included in all warnings.

Warnings in the western North Pacific and North Indian Oceans are issued every six hours valid at standard times; 0000Z, 0600Z, 1200 Z and 1800 Z (every 12-hours; 0000Z, 1200 Z or $0600 \mathrm{Z}, 1800 \mathrm{Z}$ in the southern hemisphere). All warnings are released to the communications network no earlier than synoptic time and no later than synoptic time plus two and one-half hours so that recipients will have a reasonable expectation of having all warnings "in hand" by synoptic time plus three hours $(0300 \mathrm{Z}, 0900 \mathrm{Z}, 1500 \mathrm{Z}$ and 2100 Z ).

Warning forecast positions are later verified against the corresponding "best track" positions (obtained during detailed post-storm analysis to determine the actual path and intensity of the cyclone). A summary of the verification results for 1987 is present in Chapter V.

## 8. PROGNOSTIC REASONING MESSAGES

For tropical storms and typhoons in the western North Pacific Ocean, prognostic reasoning messages are transmitted following the 0000 Z and 1200 Z warnings, or whenever the previous forecast reasoning is no longer valid. This plain language message is intended to provide meteorologists with the reasoning behind the latest forecast.

In addition to this message, prognostic reasoning information applicable to all customers is provided in the remarks section of warnings when significant forecast changes are made or when deemed appropriate by the TDO.

## 9. TROPICAL CYCLONE FORMATION ALERTS

Tropical Cyclone Formation Alerts (TCFAs) are issued whenever interpretation of satellite imagery and other meteorological data indicate that the formation of a significant tropical cyclone is likely. These formation alerts will specify a valid period not to exceed twenty-four hours and must either be cancelled, reissued, or superseded by a tropical cyclone warning prior to the expiration of the valid time.

## 10. SIGNIFICANT TROPICAL WEATHER ADVISORIES

This product contains a general, nontechnical description of all tropical disturbances in JTWC's area of responsibility (AOR) and an assessment of their potential for further (tropical cyclone) development. In addition, all tropical cyclones in warning status are briefly discussed. Two separate messages are issued daily and are valid for a 24 -hour period. The Significant Tropical Weather Advisory for the western Pacific Ocean (ABPW PGTW) covers the area east of 100 degrees East Longitude to the dateline and is issued by 0600 Z . The Significant Tropical Weather Advisory for the Indian Ocean (ABIO PGTW) covers the area west of 100 degrees East Longitude to the coast of Africa and is issued by 1800 Z . It is reissued whenever the situation warrants. For each suspect area, the words "poor", "fair", and "good" are used to describe the potential for development. "Poor" will be used to describe a tropical disturbance in which meteorological conditions are currently unfavorable for development; "fair" will be used to describe a tropical disturbance in which the meteorological conditions are favorable for development but significant development has not commenced; and "good" will be used to describe the potential for development of a tropical disturbance covered by a TCFA.

## CHAPTER II - RECONNAISSANCE AND FIXES

## 1. GENERAL

The Joint Typhoon Warning Center depends on reconnaissance to provide necessary, accurate, and timely meteorological information in support of each warning. JTWC relies primarily on three reconnaissance platforms: aircraft, satellite, and radar. In data rich areas, synoptic data are also used to supplement the above. Optimum utilization of all available reconnaissance resources is obtained through the Selective Reconnaissance Program (SRP); various factors are considered in selecting a specific reconnaissance platform including capabilities and limitations, and the tropical cyclone's threat to life and property both afloat and ashore. A summary of reconnaissance fixes received during 1987 is included in Section 6 of this chapter.

## 2. RECONNAISSANCE AVAILABILITY

## a. Aircraft

Aircraft weather reconnaissance for JTWC was performed by the 54th Weather Reconnaissance Squadron (54th WRS) located at Andersen Air Force Base, Guam. Due to budgetary decisions, 1987 was the final year for dedicated weather reconnaissance in the western North Pacific. The 54th WRS was deactivated effective 1 October 1987. The phaseout of aircraft and personnel began well before the actual deactivation of the squadron and effected aircraft availability from the very beginning of the tropical cyclone season. Only four aircraft were on station at the start of the year, three of which were storm-capable. One storm-capable aircraft was transferred to Keesler Air Force Base, Mississippi on 15 July leaving just two capable airframes to fly reconnaissance missions up to the date of deactivation. The shortage of both aircraft and personnel significantly limited the number of reconnaissance missions that the 54th WRS was able to fly throughout the season until closure. The JTWC aircraft reconnaissance requirements were provided daily to the Tropical Cyclone Aircraft Reconnaissance Coordinator (TCARC). The TCARC then married the tasking from

JTWC with the available airframes from the 54th WRS.

As in the previous years, aircraft reconnaissance provided direct measurements of standard pressure-level heights, temperatures, flight-level winds, sea-level pressures, estimated surface winds and numerous additional parameters. The meteorological data were gathered by the Aerial Reconnaissance Weather Officer and dropsonde operators from Detachment 3, 1st Weather Wing who flew with the 54th WRS. These data provided the Typhoon Duty Officer with indications of changing tropical cyclone characteristics, radii of associated winds and current tropical cyclone position and intensity. Another important aspect was the availability of the data for research on tropical cyclone analysis and forecasting.
b. Satellite

Satellite fixes from.USAF/USN ground sites and USN ships provide day and night coverage in JTWC's area of responsibility. Interpretation of this satellite imagery provides tropical cyclone positions and estimates of current and forecast intensities through the Dvorak technique.

## c. Radar

Land-based radar provides positioning data on well-developed tropical cyclones when in the proximity (usually within 175 nm ( 324 km )) of the radar sites in the Philippines, Taiwan, Hong Kong, Japan, South Korea, Kwajalein, and Guam.

## d. Synoptic

JTWC also determines tropical cyclone positions based on the analysis of the surface/gradient-level synoptic data. These positions were helpful in situations where the vertical structure of the tropical cyclone was weak or accurate surface positions from aircraft or satellite were not available.

TABLE 2-1.
AIRCRAFT RECONNAISSANCE EFFECTIVENESS
MISSIONS
FIXES
INVESTS
SYNOPTIC TRACKS

| TASKED | COMPLETED | MISSED | PERCENT |
| :---: | :---: | :---: | :---: |
| 68 | 57 | 11 | 82.98 |
| 20 | 16 | 4 | $76.6 \%$ |
| 8 | 7 | 1 | $87.5 \%$ |

MISSION EFFECTIVENESS GRADING

FIX MISSIONS TASKED
SATISFACTORY DEGRADED ( BUT SATISFACTORY ) UNSATISEACTORY

| TOTAL | PERCENT |
| :---: | :---: |
| 68 | $-81.0 \%$ |
| 55 | $8.8 \%$ |
| 6 | $19.0 \%$ |


|  |  | LEVIED VS. MISSED FIXES |  |
| ---: | ---: | ---: | ---: |
|  |  | LEVIED | MISSED | PERCENT

## 3. AIRCRAFT RECONNAISSANCE SUMMARY

During 1987, JTWC levied requirements for 68 vortex fixes and 20 investigative missions of which only 1 was flown into a disturbance which did not develop. In addition to the levied fixes, 54 intermediate fixes were obtained. Two airborne radar fixes were provided from C-141 aircraft of opportunity missions which are not included in the statistics below. Eight synoptic track missions were requested, seven of which were completed. The synoptic tracks provide mid-level steering flow information. The average position error for the combined fixes during the 1987 season was 12 nm ( 22 km ).

Aircraft reconnaissance effectiveness for the 1987 season is summarized in Table 2-1. The grading criteria is based on the Mission Effectiveness Grading (MEG) system which was developed and employed for the first time in 1986. This system grades the performance of each mission as satisfactory, degraded but satisfactory, unsatisfactory or missed. A mission could be degraded if certain critical weather parameters were not obtained such as temperature, dew point, minimum sea-level pressure, flight-level height in meters, etc. If the required time constraints between the primary and intermediate fixes were not met, the mission could still be deemed satisfactory but degraded.

## 4. SATELLITE RECONNAISSANCE SUMMARY

The Air Force provides satellite reconnaissance support to JTWC through a tropical cyclone satellite surveillance network consisting of both tactical and centralized facilities. Tactical DMSP sites monitoring DMSP, NOAA and geostationary satellite data are located at Nimitz Hill, Guam; Clark AB, Republic of the Philippines; Kadena AB, Okinawa, Japan; Osan AB, Republic of Korea; and Hickam AFB, Hawaii. These sites provide a combined coverage that includes most of JTWC's area of responsibility in the western North Pacific from near the dateline westward to the Malay Peninsula. For the remainder of its AOR, JTWC relies on the Air Force Global Weather Central (AFGWC) to provide coverage using stored satellite data. The Naval Oceanography Command Detachment, Diego Garcia, provides NOAA polar orbiting coverage in the central Indian Ocean as a supplement to this support. U.S. Navy ships equipped for direct readout also provide supplementary support.

AFGWC, located at Offutt AFB, Nebraska, is the centralized member of the tropical cyclone satellite surveillance network. In support of JTWC, AFGWC processes stored imagery from DMSP and NOAA spacecraft. Imagery recorded onboard the spacecraft as they pass over the earth is later down-linked to AFGWC via a network of command readout sites and communication satellites. This enables AFGWC to obtain the coverage necessary to fix all tropical systems of interest to JTWC. AFGWC has the primary responsibility to provide tropical cyclone surveillance over the entire Indian Ocean, southwest Pacific, and the area near the dateline. Additionally, AFGWC can be tasked to provide tropical cyclone positions in the entire western North Pacific as backup to coverage routinely available in that region.

The hub of the network is Detachment 1 , First Weather Wing (Det 1, 1WW), colocated with JTWC on Nimitz Hill, Guam. Based on available satellite coverage, Det 1, 1WW is responsible for coordinating satellite reconnaissance requirements with JTWC and tasking
the individual network sites for the necessary tropical cyclone fixes, intensity estimates and forecast intensities. When a particular fix is important to the development of JTWC's next tropical cyclone warning, two sites are tasked to fix the tropical cyclone from the same satellite pass. This "dual-site" concept provides the necessary redundancy to virtually guarantee JTWC an accurate satellite fix on the tropical cyclone.

The network provides JTWC with several products and services. The main service is one of monitoring its AOR for indications of tropical cyclone development. If an area exhibits the potential for development, JTWC is notified. Once JTWC issues either a Tropical Cyclone Formation Alert or warning, the network is tasked to provide three products: tropical cyclone positions, intensity estimates and forecast intensities. Each satellite tropical cyclone position is assigned a Position Code Number (PCN) to indicate the accuracy of the fix position. This is determined by the availability of visible landmarks in the image for precise gridding, and the degree of organization of the tropical cyclone's cloud system (Table 2-2).

TABLE 2-2. POSITION CODE NUMBERS (PCN)

PCN METHOD OF CENTER DETERMINATION/GRIDDING

EYE/GEOGRAPHY
2 EYE/EPHEMERIS
WELL-DEFINED CIRCULATION CENIER/GEOGRAPHY
4 WELL-DEFINED CIRCULATION CENTER/EPHEMERIS
POORLY DEFINED CIRCULATION CENTER/GEOGRAPHY
POORLY DEFINED CIRCULATION CENTER/EPHEMERIS

During 1987, Detachment 1, First Weather Wing increased the number of estimates of the tropical cyclone's current intensity from two to four per day once a Tropical Cyclone Formation Alert or tropical cyclone warning was issued. Intensity estimates and 24 -hour intensity forecasts were made using the Dvorak technique (NOAA Technical Report

$\mathcal{F}$ igure 2-1. Dvorak code for communicating estimates of current and forecast intensity derived from satellite data. In the example, the current 'T-number' is 3.5, but the current intensity is 4.5 (equivalent to 77 Kt ( 40 $\mathrm{m} / \mathrm{sec}$ )). The cloud system fias weakened by $1.5^{\circ} \mathrm{T}$. numbers" since the previous evaluation conducted 24. fours earlier. The plus ( + ) symbol indicates an expected reversal of the weakening trend or very little further weakening of the tropical cyclone during the next 24. four period.
NESDIS 11) for both visual and enhanced infrared imagery (Figure 2-1).

Figure 2-2 shows the status of operational polar orbiting spacecraft. Three Defense Meteorological Satellite Program (DMSP) spacecraft were operational in 1987. The 19543 (F8) satellite was launched in June as a replacement for the aging 17540 (F6) spacecraft. The imaging instrument on the 18541 (F7) spacecraft failed on 17 October, which left only one DMSP spacecraft for support during the remainder of the tropical cyclone season. The special passive sensor, microwave imager (SSM/I) on the F8 spacecraft performed well until overheating forced the sensor to be temporarily shut down on 3


December. The NOAA 9 and NOAA 10 spacecraft performed well throughout the year.

On 16 August with the loss of dedicated aircraft reconnaissance, data from the satellite reconnaissance network became the primary input to warnings and best tracks in the western North Pacific. This heightened emphasis on satellite data resulted in an increase from 60 percent (in 1986) to 88 percent of warnings based on satellite.

During 1987, the satellite reconnaissance network provided JTWC with a record total of 2,835 satellite fixes on 25 tropical cyclones in the western North Pacific Ocean. In addition, 311 fixes were made on tropical cyclones in the North Indian Ocean, more than eight times the total for 1986. For the southern hemisphere, 1,192 satellite fixes were provided. A comparison of those fixes in the JTWC area of responsibility and their corresponding JTWC best track is shown in Tables 2-3A and 2-3B. (Note: Those fixes which were out-of-limits when compared with the best track are not included.)

The relationship between tropical cyclone "T-number", maximum surface wind speed and minimum sea-level pressure is outlined in Table 2-4. Table 2-5A, B and C address the verification of satellite-derived intensity estimates for developing, weakening and all cases of tropical cyclones, respectively. In each table the first column states the "Tnumber" in parentheses and expected current and forecast intensity. The verifying average intensities from the current and 24 -hour best tracks are included to the right in the second and third columns, respectively.

## 5. RADAR RECONNAISSANCE SUMMARY

Fifteen of the twenty-five significant tropical cyclones in the western North Pacific during 1987 passed within range of land-based radar with sufficient cloud pattern organization to be fixed. The land-based radar fixes that were obtained and transmitted to JTWC totaled 806. Only one radar fix was obtained by reconnaissance aircraft.

TABLE 2-3A. MEAN DEVIATION (NM) OF ALL SATELLITE DERIVED TROPICAL CYCLONE POSITIONS FROM THE JTWC BEST TRACK POSITIONS IN THE WESTERN NORTH PACIFIC AND NORTH INDIAN OCEANS. NUMBER OF CASES (IN PARENTHESES).

WESTERN NORTH PACIFIC OCEAN 1977-1986 AVERAGE 1987

| PCN | (ALL SITES) | (ALL SITES) |  |  |
| :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |
| 1 | 14.2 | $(1689)$ | 14.9 | $(182)$ |
| 2 | 16.3 | $(2118)$ | 13.0 | $(511)$ |
| 3 | 21.3 | $(2410)$ | 21.4 | $(219)$ |
| 4 | 23.9 | $(1546)$ | 18.7 | $(576)$ |
| 5 | 37.8 | $(4456)$ | 32.6 | $(195)$ |
| 6 | 39.5 | $(4222)$ | 34.6 | $(1048)$ |
| $1 \& 2$ | 15.4 | $(3807)$ | 13.5 | $(693)$ |
| $3 \& 4$ | 22.3 | $(3956)$ | $19.5(795)$ |  |
| $5 \& 6$ | 38.6 | $(8678)$ | 34.6 | $(1243)$ |
|  |  |  |  |  |
| TOTAL | 29.3 | $(16441)$ | 24.2 | $(2731)$ |

NORTH INDIAN OCEAN


TABLE 2-3B
MEAN DEVIATION (NM) OF ALL SATELLITE-DERIVED TROPICAL CYCLONE POSITIONS IN THE SOUTH

PACIFIC AND SOUTH INDIAN OCEANS. NUMBER OF CASES ARE IN PARENTHESES.

|  | 1985 - 1986 AVERAGE |  | 1987 |  |
| :---: | :---: | :---: | :---: | :---: |
| PCN |  |  | (ALL | SITES) |
| 1 | 17.6 | ( 68) | 14.5 | ( 14) |
| 2 | 15.5 | ( 312) | 17.4 | ( 130) |
| 3 | 33.7 | ( 97) | 40.4 | ( 15) |
| 4 | 27.2 | ( 301) | 26.5 | ( 107) |
| 5 | 46.8 | ( 399) | 28.8 | ( 75) |
| 6 | 38.1 | (2152) | 32.9 | ( 786) |
| 1 ¢ 2 | 15.9 | ( 380) | 17.3 | ( 144) |
| $3 \times 4$ | 28.8 | ( 398) | 28.2 | ( 122) |
| $5 \& 6$ | 39.5 | (2551) | 32.6 | ( 861) |
| TOTALS | 35.5 | (3329) | 30.1 | (1127) |

TABLE 2-4. MAXIMUM SUSTAINED WIND SPEED (KT) AS A FUNCTION OF DVORAK CI FI (CURRENT AND FORECAST INTENSITY) NUMBER AND MINIMUM SEA-LEVEL PRESSURE (MSLP)

| TROPICAL CYCLONE INTENSITY NUMBER | $\begin{array}{r} \text { WIND } \\ \text { SPEED } \end{array}$ | (NW | MSLP <br> PACIFIC |
| :---: | :---: | :---: | :---: |
| 0.0 | <25 |  | ---- |
| 0.5 | 25 |  | ---- |
| 1.0 | 25 |  | ---- |
| 1.5 | 25 |  | ---- |
| 2.0 | 30 |  | 1000 |
| 2.5 | 35 |  | 997 |
| 3.0 | 45 |  | 991 |
| 3.5 | 55 |  | 984 |
| 4.0 | 65 |  | 976 |
| 4.5 | 77 |  | 966 |
| 5.0 | 90 |  | 954 |
| 5.5 | 102 |  | 941 |
| 6.0 | 115 |  | 927 |
| 6.5 | 127 |  | 914 |
| 7.0 | 140 |  | 898 |
| 7.5 | 155 |  | 879 |
| 8.0 | 170 |  | 858 |


| TABLE 2-5A. |  | DEVELOPING STAGE |  | TABLE 2-5B. |  | WEAKENING STAGE |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| CURRENT OR FORECAST INTENSITY* |  | VERIFYING | VERIFYING | CUR | OR | VERIFYING | VERIFYING |
|  |  | AVERAGE BT | AVE 24HR BT |  | AST | AVERAGE BT | AVE 24 HR BT |
|  |  | INTENSITY | INTENSITY | INT | ITY* | INTENSITY | INTENSITY |
| (T \#) |  | KT | KT | (T \#) |  | KT | KT |
| (0.0) | <25 | --- | - | (0.0) | $<25$ | --- | --- |
| (1.0) | 25 | 22 | 28 | (1.0) | 25 | 19 | 14 |
| (1.5) | 25 | 25 | 31 | (1.5) | 25 | 27 | 22 |
| (2.0) | 30 | 30 | 37 | (2.0) | 30 | 30 | 24 |
| (2.5) | 35 | 35 | 47 | (2.5) | 35 | 38 | 30 |
| (3.0) | 45 | 47 | 65 | (3.0) | 45 | 43 | 31 |
| (3.5) | 55 | 57 | 75 | (3.5) | 55 | 57 | 40 |
| (4.0) | 65 | 65 | 80 | (4.0) | 65 | 65 | 50 |
| (4.5) | 77 | 75 | 92 | (4.5) | 77 | 77 | 53 |
| (5.0) | 90 | 88 | 110 | (5.0) | 90 | 88 | 70 |
| (5.5) | 102 | 102 | 110 | (5.5) | 102 | 98 | 75 |
| (6.0) | 115 | 115 | 122 | (6.0) | 115 | 113 | 90 |
| (6.5) | 127 | 127 | 123 | (6.5) | 127 | 123 | 108 |
| (7.0) | 140 | 138 | 115 | (7.0) | 140 | 133 | 114 |
| (7.5) | 155 | --- | --- | (7.5) | 155 | --- | --- |
| (8.0) | 170 | --- | --- | (8.0) | 170 | --- | --- |
| * DVORAK, 1984 |  |  |  | * D | аK, |  |  |


| TABLE 2-5c. |  | ALl CASES |  |
| :---: | :---: | :---: | :---: |
| CURRENT OR FORECAST INTENSITY* |  | VERIFYING | VERIFYING |
|  |  | AVERAGE BT | AVE 24HR BT |
|  |  | INTENSITY | INTENSITY |
| ( T \#) | KT | KT | KT |
| (0.0) | <25 | --- | --- |
| (1.0) | 25 | 22 | 26 |
| (1.5) | 25 | 25 | 29 |
| (2.0) | 30 | 29 | 33 |
| (2.5) | 35 | 36 | 41 |
| (3.0) | 45 | 46 | 55 |
| (3.5) | 55 | 57 | 59 |
| (4.0) | 65 | 65 | 65 |
| (4.5) | 77 | 76 | 76 |
| (5.0) | 90 | 88 | 88 |
| (5.5) | 102 | 99 | 88 |
| (6.0) | 115 | 114 | 101 |
| (6.5) | 127 | 125 | 114 |
| (7.0) | 140 | 135 | 114 |
| (7.5) | 155 | --- | --- |
| (8.0) | 170 | --- | --- |
| * DVORAK, 1984 |  |  |  |

The WMO radar code defines three categories of accuracy: good (within 10 km ( 5 nm ), fair (within $10-30 \mathrm{~km}(5-16 \mathrm{~nm})$ ), and poor (within $30-50 \mathrm{~km}(16-27 \mathrm{~nm})$ ). Of the 807 radar fixes coded in this manner; 309 were good, 190 were fair, and 308 were poor. Compared to JTWC's best track, the mean vector deviation for land-based radar sites was $19 \mathrm{~nm}(35 \mathrm{~km})$. Excellent support through timely and accurate radar fix positioning allowed JTWC to track and forecast tropical cyclone movement through even the most difficult erratic tracks.

## 6. TROPICAL CYCLONE FIX DATA

As in previous years, no radar reports were received on North Indian Ocean tropical cyclones.

A total of 3,754 fixes on twenty-five western North Pacific tropical cyclones and 311 fixes on eight North Indian Ocean tropical cyclones were received at JTWC. Table 2-6A, Fix Platform Summary, delineates the number of fixes per platform for each individual tropical cyclone. Season totals and percentages are also indicated. (Table 2-6B provides the same information for the South Pacific and South Indian Oceans.)

\begin{tabular}{|c|c|c|c|c|c|}
\hline TABLE 2-6A. \& \multicolumn{5}{|l|}{FIX PLATFORM SUMMARY FOR 1987} <br>
\hline WESTERN NORTH PACIFIC \& AIRCRAFT \& SAtELLite \& RADAR \& SYNOPTIC \& TOTAL <br>
\hline TY ORCHID (01W) \& 17 \& 100 \& 0 \& 0 \& 117 <br>
\hline TS PERCY (02W) \& 4 \& 60 \& 0 \& 0 \& 64 <br>
\hline TS RUTH (03W) \& 0 \& 41 \& 20 \& 0 \& 61 <br>
\hline TS SPERRY (04W) \& 12 \& 82 \& 8 \& 0 \& 102 <br>
\hline Sty thelma (05W) \& 24 \& 141 \& 72 \& 0 \& 237 <br>
\hline TS VERNON (06W) \& 11 \& 97 \& 27 \& 0 \& 135 <br>
\hline TY WYNNE (07W) \& 21 \& 198 \& 41 \& 0 \& 260 <br>
\hline TY ALEX (08W) \& 1 \& 100 \& 77 \& 0 \& 178 <br>
\hline STY BETTY (09W) \& 13 \& 144 \& 71 \& 0 \& 228 <br>
\hline TY CARY (10W) \& 9 \& 181 \& 72 \& 0 \& 262 <br>
\hline STY DINAH (11W) \& 0 \& 159 \& 106 \& 0 \& 265 <br>
\hline TS ED (12W) \& 0 \& 68 \& 0 \& 0 \& 68 <br>
\hline TY FREDA (13W) \& 0 \& 176 \& 29 \& 0 \& 205 <br>
\hline TY GERALD (14W) \& 0 \& 119 \& 81 \& 0 \& 200 <br>
\hline STY HOLLY (15W) \& 0 \& 151 \& 0 \& 0 \& 151 <br>
\hline TY IAN (16W) \& 0 \& 138 \& 5 \& 0 \& 143 <br>
\hline TD 17W (17W) \& 0 \& 30 \& 0 \& 0 \& 30 <br>
\hline TY PEKE (02C) \& 0 \& 131 \& 0 \& 0 \& 131 <br>
\hline TS JUNE (18W) \& 0 \& 43 \& 0 \& 0 \& 43 <br>
\hline TY KELLY (19W) \& 0 \& 111 \& 63 \& 0 \& 174 <br>
\hline STY LYNN (20W) \& 0 \& 159 \& 56 \& 0 \& 215 <br>
\hline TS MAURY (21W) \& 0 \& 95 \& 0 \& 0 \& 95 <br>
\hline STY NINA (22W) \& 0 \& 176 \& 79 \& 0 \& 255 <br>
\hline TS OGDEN (23W) \& 0 \& 17 \& 0 \& 0 \& 17 <br>
\hline TY PHYLLIS (24W) \& 0 \& 118 \& 0 \& 0 \& 118 <br>
\hline TOTALS \& 1,12 \& 2835 \& 807 \& 0 \& 3754 <br>
\hline \% OF TOTAL NR OF FIXES \& 3.0\% \& 75.5\% \& 21.5\% \& 0.0\% \& 100.0\% <br>
\hline \multirow[t]{9}{*}{NORTH

INDIAN OCEAN
TC} \& SATELLIte \& SYNOPTIC \& T \& als \& <br>
\hline \& 59 \& 0 \& \& 59 \& <br>
\hline \& 59 \& 0 \& \& 59 \& <br>
\hline \& 38 \& 0 \& \& 38 \& <br>
\hline \& 15 \& 0 \& \& 15 \& <br>
\hline \& 43 \& 0 \& \& 43 \& <br>
\hline \& 32 \& 0 \& \& 32 \& <br>
\hline \& 16 \& 0 \& \& 16 \& <br>
\hline \& 49 \& 0 \& \& 49 \& <br>
\hline totals \& 311 \& 0 \& \& 311 \& <br>
\hline \% OF TOTAL
NR OF FIXES \& 100.0\% \& 0.0\% \& \& .0\% \& <br>
\hline
\end{tabular}

TABLE 2-6B. FIX PLATFORM SUMMARY FOR 1987


## CHAPTER III - SUMMARY OF WESTERN NORTH PACIFIC AND NORTH INDIAN OCEAN TROPICAL CYCLONES

## 1. GENERAL

During the calendar year 1987, JTWC issued warnings on 25 different significant tropical cyclones in the western North Pacific -six super typhoons, 12 typhoons, six tropical storms and one tropical depression. This includes one typhoon, Peke (02C), which initially developed in the central North Pacific (Table 3-1). The total number of western North Pacific tropical cyclones is lower than the climatological mean of 30.7, and two tropical cyclones below the 1986 total (Table 3-2). A record-setting eight significant tropical cyclones (all were of tropical storm intensity) developed in the North Indian Ocean. This is twice the climatological mean of four. Therefore, during 1987, JTWC issued warnings on a total of 33 northern hemisphere tropical cyclones.

During 1987 in the western North Pacific there were 139 "warning days". (A warning day is defined as a day during which JTWC was issuing warnings on at least one tropical cyclone. A "two-cyclone day" refers to a day when there were warnings issued on two different tropical cyclones simultaneously, a "three-cyclone day" -- three tropical cyclones at one time, and so on...). Considering only the western North Pacific, there were 30 twocyclone days, 10 three-cyclone days and no four or five-cyclone days (Table 3-3). When North Indian Ocean tropical cyclones are included, there were 156 warning days, 38 two-cyclone days, 11 three-cyclone days and no four- or five-cyclone days. Thus, JTWC was in warning status 42.7 percent of the year; it was in a multiple-cyclone situation (that is, warning on two or more tropical cyclones) for 38 days or about 10.4 percent of the year.

JTWC issued 668 warnings on the 25 western North Pacific tropical cyclones (two warnings from January 1st on Typhoon Norris (26W) of 1986 are included in the total) and 83 warnings on the eight North Indian Ocean tropical cyclones, for a grand total of 751 warnings. There were thirty-one initial Tropical Cyclone Formation Alerts (TCFAs) issued for western North Pacific and eleven for the North

Indian Ocean. Twenty-four western North Pacific and seven North Indian Ocean tropical cyclones developed subsequent to the issuance of a TCFA. Three of the western North Pacific tropical cyclones regenerated during their lifetime and each was covered by a TCFA. Typhoon Peke (02C) was passed to JTWC while in warning status and thus no TCFA was required (Table 3-4). For the western North Pacific, the false alarm rate was 24 percent and the mean lead time (to issuance of the first warning) was 13.5 hours. For the North Indian Ocean, the false alarm rate was 18 percent, with a mean lead time of 10.1 hours. One system (Tropical Cyclone 03A) was warned on without the benefit of a preceding TCFA.

## 2. WESTERN NORTH PACIFIC TROPICAL CYCLONES

The distinguishing features of 1987 in the western North Pacific were the low number of total tropical cyclones (25), the large number of super typhoons (6) and the number of "midgets" (4).

## JANUARY THROUGH JUNE

The activity began in early in January with Typhoon Orchid (01W). Orchid (01W) was an unusually small system which transited the wintertime western North Pacific before being sheared apart by the northeast monsoon east of the Philippine Islands. The island of Ulithi experienced $100 \mathrm{kt}(51 \mathrm{~m} / \mathrm{sec})$ winds and extensive damage when Orchid (01W) passed directly overhead. Tropical Storm Percy (02W) was the only significant tropical cyclone in April. It struggled to get started, but tenaciously resisted dissipation until it passed over the island of Luzon. Tropical Storm Ruth (03W) was a short-lived system which developed southeast of Hong Kong and eventually dissipated over southern China. Typhoon Sperry (04W) was the second tropical cyclone to reach typhoon intensity and also the second "midget" typhoon (Orchid (01W) was the first) of 1987. It was also the first to enter

the mid-latitude westerlies and recurve toward the northeast.

## JULY

Super Typhoon Thelma (05W) was the first of four significant tropical cyclones to develop in July and the first super typhoon of 1987. Forecasting the timing and location of recurvature presented a problem for JTWC. Dynamic forecast aids did indicate recurvature, but much sooner than was ultimately observed. After damaging northern Luzon and causing the
evacuation of aircraft from Okinawa, Japan, Thelma (05W) slammed into Korea causing $\$ 124$ million in damages and the loss of several hundred lives. Typhoon Vernon (06W) followed closely on the heels of Super Typhoon Thelma (05W). It was weak and disorganized throughout most of its life. As a result, initial positioning problems arose in the Philippine Sea due to differences between real-time fix information from radar, satellite and aircraft. Typhoon Wynne (07W) was the third "midget" typhoon of the season. It tracked along a constant bearing of 294 degrees for four days
and maintained a visible eye for six days. Meteorologists on Kwajalein Atoll provided radar fix information which was instrumental in relocating Wynne (07W) early on. Wynne (07W) caused extensive damage as it passed through the northern Mariana Islands ${ }^{1}$. Typhoon Alex ( 08 W ) was the final tropical cyclone in July. Together with Wynne (07W) it formed the first multiple tropical cyclone situation during 1987. Alex (08W) brushed the eastern coast of Taiwan before making landfall on the coast of mainland China. Although damage to Taiwan and mainland China was relatively light, the remnants of Alex (08W) enhanced a band of precipitation that stalled over Korea, and as a consequence, 12 inches ( 308 mm ) of rain fell in 24 -hours causing major flooding and loss of life.

## AUGUST

Super Typhoon Betty (09W) was the second super typhoon of the season and had the lowest reported minimum sea-level pressure ( 891 mb ) up to that time. It explosively deepened just prior to making landfall in the Philippine Islands, where 20 people were killed and 60,000 left homeless. Typhoon Cary (10W), together with Betty (09W) and Dinah (11W) formed the first three-storm situation for 1987. The last scheduled western North Pacific aircraft reconnaissance mission was flown on Cary (10W) on the 15 th of August. Cary (10W) eventually made landfall on the coast of northern Vietnam and ultimately dissipated over Burma. Super Typhoon Dinah (11W) was the most destructive typhoon to strike Okinawa and the southern islands of Japan in the past 20 years. Throughout its life, JTWC consistently forecast recurvature and accelerations towards the northeast through the Sea of Japan, as a result, Dinah's (11W) forecast track errors were smaller than average. Tropical Storm Ed (12W) was a very difficult tropical cyclone to locate and forecast due to fluctuations in its intensity, speed and track direction and its poorly defined

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cloud signature. Drifting buoy reports of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ were key to the decision to issue the first warning on Ed (12W).

## SEPTEMBER

Typhoon Freda (13W) was the first of seven tropical cyclones to develop during September and was the middle (geographically) of a three-storm situation (the other tropical cyclones being Gerald (14W) and Holly (15W)). This was the second three-storm situation of the year. Freda (13W) was unusual because it traversed less than 10 degrees of longitude, but 25 degrees of latitude. Freda's (13W) thirteen day life span and 50 warnings were records for 1987. Typhoon Gerald (14W) was unique in that it matured within the monsoon trough, but did not detach from it. The most distinctive feature was its unusually large eye. Super Typhoon Holly (15W) was the third tropical cyclone to develop from the active monsoon trough which also spawned Freda (13W) and Gerald (14W). Although very intense, it had a very uneventful life as it recurved far to the east of Japan. After a six day respite in tropical cyclone activity, Typhoon Ian (16W) developed about 330 nm ( 611 km ) to the east-northeast of Guam. Andersen

Air Force Base Weather was able to provide several radar fixes of Ian (16W) as it passed to the north of Guam. It eventually recurved and transitioned to a sub-tropical system north of 25 degrees North Latitude. Tropical Depression 17W developed north of the Marshall Islands at about the same time as Ian. Tropical Depression 17W did not reach tropical storm strength because it was suppressed by the combined effects of the outflow from Ian (16W) and a mid-level short-wave trough to the north. Typhoon Peke (02C) was the first hurricane to form in the central North Pacific and cross to the western North Pacific in the past twenty years. It meandered basically northnorthwestward and eventually dissipated just west of the dateline. Tropical Storm June (18W) was the third tropical cyclone of the final three-storm situation during 1987. Throughout its life, June's (18W) upper-level outflow was restricted by the strong outflow from Ian.


## OCTOBER

Typhoon Kelly (19W) was the first of only two significant tropical cyclones to occur during October. Kelly (19W) developed when the monsoon trough re-established itself in lowlatitudes after it had been displaced to a position about 25 degrees North Latitude the previous week due to Ian (16W), June (18W) and Peke (02C). Super Typhoon Lynn (20W) was the fifth super typhoon of the year and the third to produce winds of at least $140 \mathrm{kt}(72 \mathrm{~m} / \mathrm{sec})$. It attained a minimum sea-level pressure of 898 mb . At one point, Lynn (20W) appeared to be headed straight for Guam. Fortunately, a last minute jog toward the north spared the island from a direct hit. Saipan, which is north of Guam, received gusts to $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$, however. Lynn (20W) eventually passed south Taiwan, enhanced convection and increased

| BLE 3-4 | FORMATION ALERT SUMMARY WESTERN NORTH PACIFIC |  |  |  |
| :---: | :---: | :---: | :---: | :---: |
|  | $\begin{gathered} \text { NUMBER } \\ \text { op } \\ \text { TCFAS } \\ \hline \end{gathered}$ | NUMBER OF SYSTEMS WARNED ON | TOTAL NUMBER OF SYSTEMS <br> systems | $\begin{gathered} \text { FALSE } \\ \text { ALARM } \\ \text { RATE } \\ \hline \end{gathered}$ |
| 1975 | 34 | 25 | 25 | $26 \%$ |
| 1976 | 34 | 25 | 25 | 268 |
| 1977 | 26 | 20 | 21 | 238 |
| 1978 | 32 | 27 | 32 | 168 |
| 1979 | 27 | 23 | 28 | 158 |
| 1980 | 37 | 28 | 28 | 248 |
| 1981 | 29 | 28 | 29 | 37 |
| 1982 | 36 | 26 | 28 | 288 |
| 1983 | 31 | 25 | 25 | 198 |
| 1984 | 37 | 30 | 30 | 19\% |
| 1985 | 39 | 26 | 27 | 338 |
| 1986 | 38 | 27 | 27 | 298 |
| 1987 | 31 | 24 | 25 | 238 |
| (1975-1987) <br> average | 33.1 | 25.7 | 26.9 | $21.8 \%$ |
| CASES | 431 | 334 | 350 |  |

wind speeds to the north that caused the deaths of 42 people. Over 68 inches ( 1744 mm ) of rain fell on Taipei over a two day period due to Lynn (20W). Tropical Storm Maury (21W) was a relatively weak, but persistent, tropical cyclone which formed southeast of Guam and tracked basically westward across the Philippine Sea, through the Philippine Islands and into the South China Sea.

## NOVEMBER THROUGH DECEMBER

Super Typhoon Nina (22W) was the sixth, and last, super typhoon. It also proved to be the most intense and destructive tropical cyclone of the year. Nina was interesting from a meteorological point of view because it unexpectedly intensified while still accelerating toward the west. It devastated Truk, killing five and injuring 38, before moving on to the Philippine Sea. Once there Nina (22W) explosively intensified, to $145 \mathrm{kt}(75 \mathrm{~m} / \mathrm{sec})$, prior to making landfall. An estimated 658 people were killed on southern Luzon, making it the most destructive typhoon to hit the Philippine Islands in 20 years. Tropical Storm Ogden (23W) was another minimal tropical storm which developed in the South China Sea and moved westward before making landfall on the Vietnam coast, north of Cam Rahn Bay. Typhoon Phyllis (24W), the fourth "midget" of 1987, was the last tropical cyclone of the 1987 season. Phyllis (24W) formed southeast of Guam and initially followed what appeared to be a broad recurvature track, passing to the southwest of Guam. Unfortunately, it weakened, moved toward the west-southwest and then explosively intensified (to 100 kt ( 51 $\mathrm{m} / \mathrm{sec}$ )) before striking the island of Samar in the central Philippine Islands and moving into the South China Sea.




## TYPHOON ORCHID (01W)

Typhoon Orchid (01W), the first tropical cyclone of 1987 , was an unusually small system. It transited across the wintertime western North Pacific before being sheared apart by the northeast monsoon east of the Philippines.

The disturbance that eventually developed into Typhoon Orchid was first detected at 0000 Z on January 3rd as a small area of persistent convection in the nearequatorial trough near the dateline. It was first mentioned on the Significant Tropical Weather


Figure 3-01-1. The 110000Z Ianuary aircraft reconnaissance fix mission of Typhoon Orcfid (01W) near maximum intensity. Note the tight gradient of surface winds to the south of the vortex center.


Figure 3-01-2. Damage to the Outer-Isfand Sciool Located on the north side of Fafafop Isfand on the qufitfi gtoll. gll these buildings sustained some damage wfiife others (dornitories and classrooms) vere totally destroyed (Photo courtesy of Mobi( Oi( Miceroresia, Inc.).

Advisory (ABPW PGTW) at 0306002. Over the next four to five days, this area drifted toward the west and slowly increased in organization and convection until a small ragged central dense overcast (CDO) formed and upper-level outfiow improved. A Tropical Cyclone Formation Alert followed at 072130Z and a daylight hours aircraft reconnaissance investigative mission was tasked for the next day. No surface data was available near the system; however, at 080419Z Dvorak satellite
intensity analysis of the disturbance estimated maximum sustained surface winds of 25 to 30 kt ( 13 to $15 \mathrm{~m} / \mathrm{sec}$ ). This prompted JTWC to issue the first warning on Tropical Depression 01W at 081200Z. The next morning, the aircraft investigative mission reported maximum sustained surface winds of 45 kt ( 23 $\mathrm{m} / \mathrm{sec}$ ). This prompted the upgrade of the system on the third warning (at 090000Z) to Tropical Storm Orchid (01W).


Figure 3-01-3. Corrugated sheet roofing embedded in a coconut log on Fafatop Island on Ulithi Atolf. This building material becomes a deadly ofject to life and property when airborne (Photo courtesy of Mobil OiL Micronesia, Inc.).

There were two unusual aspects of Typhoon Orchid (01W). The first was its small radius of maximum winds. For example, at its peak intensity of $95 \mathrm{kt}(49 \mathrm{~m} / \mathrm{sec})$ at 110000 Z , the radii of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ winds were only 45 $\mathrm{nm}(83 \mathrm{~km})$ in the south semicircle and 140 nm ( 259 km ) in the northwest semicircle (see Figure 3-01-1). The larger wind radius in the northwest semicircle was due partly to interaction with high pressure ridging to the north and the motion vector addition to the
winds in the west-northwest, or right front, quadrant. Typhoon Orchid (01W), near maximum intensity, passed directly over the island of Ulithi (WMO 91203). As a result, Ulithi reported surface winds near 100 kt ( 51 $\mathrm{m} / \mathrm{sec}$ ) and sustained extensive damage (Figures 3-01-2 and 3-01-3). A few hours later, however, Typhoon Orchid passed about 45 nm $(83 \mathrm{~km})$ north of the island of Yap (WMO 91413), where surface winds of only 20 to 25 kt ( 10 to $13 \mathrm{~m} / \mathrm{sec}$ ) and fair skies were reported.


Figure 3-01-4. Enfanced infrared (EIR) imagery of Typhoon Orchid (01W) near maximum intensity. $\mathfrak{K}$ (ote the small well-defined eye ( $102259 Z$ January NOAA infrated imagery).

The second unusual aspect of Typhoon Orchid was the fact that during the two days from 101200 Z to 121200 Z when Orchid was the most intense (Figure 3-01-4), the Dvorak intensity estimates were 10 to 20 kt ( 5 to $10 \mathrm{~m} / \mathrm{sec}$ ) higher than the intensity reported by aircraft reconnaissance (Figure 3-01-5).

After reaching maximum intensity, Orchid continued moving northwestward. Between 110000 Z and 140000 Z (when the last warning was issued), Orchid came under the influence of the strong wintertime low-level
northeast monsoonal flow. This, coupled with 200 mb westerly flow aloft, set up a strong vertical shearing environment in which the upper portion of Orchid was displaced toward the east. Once the central convection stripped away, the remaining surface circulation was then steered by the low-level northeast monsoonal flow toward the southwest before dissipating over water. This wintertime shearing situation was a common factor in the end of the last five significant tropical cyclones in November and December of 1986.


Figure 3-01-5. PLot of intensities obtained for Typfioon Orckid (01W) 6 y aircraft reconnaissance vortex fixes and Dvorak analysis of satellite imagery. Also plotted are the Final Best Track intensities for comparison. Note the figher satelfite intensities especially around the time of maximum intensity.


## TROPICAL STORM PERCY (O2W)

Percy was the only significant western North Pacific tropical disturbance during April. The vortex struggled to get started, only achieved minimal tropical storm intensity and tenaciously resisted dissipation.

During the first week of April, brisk 30 $\mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ northeasterly trades clashed with a low-latitude westerly surge associated with a tropical disturbance in the southern hemisphere. This created an area of cyclonic rotation in the low-level wind field and slightly lower pressures in the eastern Caroline Islands. The resulting
convection was first noted on the Significant Tropical Weather Advisory (ABPW PGTW) for 030600 Z April. The slow formation of a cloud system led to the issuance of a Tropical Cyclone Formation Alert (TCFA) at 061800 Z . Maximum sustained surface winds, at that time, were estimated to be 25 to 30 kt ( 13 to 15 $\mathrm{m} / \mathrm{sec}$ ). The TCFA was reissued at 071800 Z because the area had increased in organization, although the overall convection decreased. By 081800 Z , the mid- to upper-level winds over the system increased and the second TCFA was cancelled.


Figure 3-02-1. Plot of aircraft reconnaissance data from the third mission into Tropical Storm Percy (02W) shows increasing surface winds near the center.


Figure 3-02-2. Tropical Storm Percy (02W), while southwest of Guam, showing an exposed low-Level circulation center ( $110022 Z$ April DMSSP visual imagery).

However, the weakened disturbance continued to drift slowly westward for the next three days and possessed fair potential for significant development. By $101407 Z$ the convection and organization had again improved, and the vertical wind shear on the system decreased sufficiently to justify the issuance of a third TCFA. Satellite intensity analysis (Dvorak, 1984) estimated maximum surface winds of $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$ at that time, however, an aircraft daylight investigative mission flown on the morning of the 10th produced unexpected results.

Enroute to the circulation, the aircraft reported gradually increasing flight-level winds ( $1500 \mathrm{ft}(457 \mathrm{~m}$ )) and observed surface winds from $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec}$ ) near Guam to $40 \mathrm{kt}(21$ $\mathrm{m} / \mathrm{sec}$ ). Upon reaching the expected location of the low-level cyclonic circulation, the Aerial Recon- naissance Weather Officer (ARWO) reported flight-level winds of $56 \mathrm{kt}(29 \mathrm{~m} / \mathrm{sec})$ and surface winds of $50 \mathrm{kt}(26 \mathrm{~m} / \mathrm{sec})$ at 102325 Z . The center location was consistent with the increasing surface winds encountered enroute (Figure 3-02-1); however, the magnitude of the winds were inconsistent with
the minimum sea-level pressures (MSLPs) enroute ( 1010 mb to 1003 mb ) and at the circulation center ( 1001 mb ). Dvorak satellite intensity analysis at 110022 Z estimated 25 kt ( $13 \mathrm{~m} / \mathrm{sec}$ ) surface winds (Figure 3-02-2). This was more consistent with the extrapolated MSLPs. According to Atkinson and Holliday (1977), an environmental MSLP of near 1012 mb , together with maximum sustained surface winds of $50 \mathrm{kt}(26 \mathrm{~m} / \mathrm{sec})$ usually implies a central MSLP of about 987 mb . After the aircraft reconnaissance flight-level wind observations, were double-checked, Tropical Depression 02W was upgraded to a tropical storm.

Aircraft reconnaissance was available again at 112100 Z . Flight-level and surface winds were much lighter than observed 24hours earlier. Values ranged from 20 to 32 kt ( 10 to $16 \mathrm{~m} / \mathrm{sec}$ ) at flight level enroute to the early fix at 112120 Z , and 15 to 20 kt ( 8 to 10 $\mathrm{m} / \mathrm{sec}$ ) prior to the primary fix at 112354 Z . The Dvorak satellite intensity analysis at 122304 Z estimated $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec}$ ) maximum sustained surface winds. As a result, the final warning on Tropical Storm Percy (02W) followed at 130000 Z. Percy's circulation persisted as an exposed low-level center through the 15th, with the remaining convection located well to the northeast and southwest (Figure 3-02-3). The residual low-level eddy, that remained, finally dissipated near northern Luzon on the 19th of April.


Figure 3-02-3. Tropical Storm Percy (02W) in the Pfilippine Sea just Gefore the final warning was issued (122341Z April DMSSP visual imagery).


## TROPICAL STORM RUTH (03W)



Figure 3-03-1. Tropical Storm Rutf was a short-fived tropical cyclone. Only six warnings were issued on the system before it moved infand and dissipated over southern China. It began as a monsoon depression 240 nm ( 444 Km ) southeast of $\mathcal{H}$ (ong Kong over the South China Sea. Eafly on 17 Iune, convection consolidated into convective bands prompting its mention on the Significant Tropical Weather $\mathfrak{A d v i s o r y}(\mathcal{A B P W}$ PGT'W) at 0600 Z as faving good potential for development. Ms a result, IT'WC issued a Tropical Cyclone Formation Alert, valid at the same time, because satelfite imagery indicated upper-level anticyclonic outflow was becoming established. ITWC issued the first warning on Tropical Depression $03 \mathcal{W}$ at $180000 Z$, after symoptic reports indicated surface pressures in the area had dropped significantly overnight from 1001 mb to 995 mb . The initial forecast tracks indicated the system would move nortiwestward, but subsequent forecasts gradually shifted the track further west as the subtropical ridge east of Ruth began ridging slowly westward across southern China. The system was upgraded to tropical storm intensity at $181800 Z$ based upon a $\mathcal{D}$ vorak intensity estimate of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ maximum sustained surface winds associated with convective bands which were wrapped halfway around the center (see image above). Ruth was dowingraded to a tropical depression on the fifth warning as it interacted with the southern coast of China. Hong Kong (WMO 45005) radar reports were excellent and proved instrumental in accurately tracking this tropical cyclone for a day before it made landfalf. Ruth dissipated witfin eighteen fours of moving infand, causing fittle damage and no known deatfs (181138Z gune DMSSP infrared imagery).


Typhoon Sperry was the second tropical cyclone to reach typhoon intensity and also the second "midget" typhoon in 1987. It was also the season's first to enter the midlatitude westerlies and recurve toward the northeast.

The tropical disturbance that eventually developed into Typhoon Sperry was first detected by synoptic data on 24 June as a broad, weak surface circulation in the western extension of the monsoon trough 200 nm ( 370 km ) to the northwest of the island of Truk in the eastern Caroline Islands. The convection in this area appeared to be random. At the same
time, a second area of disorganized convection was developing 210 nm ( 389 km ) east of the island of Enewetak in the Marshalls. To the north and east, a Tropical Upper-Tropospheric Trough (TUTT) extended from Wake Island southwestward to just northeast of Guam. The broad subtropical ridge dominated the low-level flow pattern in the northwest Pacific. Although the two convective areas consolidated on 25 June, the resultant disturbance still struggled for two more days before reaching tropical storm intensity. The most probable cause for this slow intensification was the close proximity of a TUTT low (Sadler, 1979) to the northeast. This low aloft, in conjunction with the lower


Figure 3-04-1. Visual sateflite imagery showing the tropical disturbance that would later develop into Typhoon Sperry. Note the Low-Level circulation center
displaced to the nortfieast of the main convection (252347Z June DMSP visual imagery).


Figure 3-04-2. Typfioon Sperry with an intensity of $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ just prior to peaking (280047Z gune DMSS visual imagery).
tropospheric subtropical ridge, created an area of strong vertical wind shear (Figure 3-04-1).

The low-level disturbance drifted westnorthwestward for the next several days. At that point, the separation between the TUTT low, or cell, and the low-level tropical disturbance to the southwest remained static. However, an interesting change occurred aloft. By 251200 Z a plume of dense cirrus, associated with a $55 \mathrm{kt}(28 \mathrm{~m} / \mathrm{sec})$ wind maximum entering the western side of the upper cold low from the north, moved southward. Within eighteenhours the cirrus plume had plunged into the southwest portion of the TUTT cell. The cell responded. The circulation within the core of the upper low tightened up and became more symmetrical. This, in turn, reduced the vertical wind shear across the system and as a result, the central convection started to increase within the low-level disturbance again. Earlier (at 241200Z), the Navy Operational Global Atmospheric Prediction System upper-air prognoses, had correctly forecast this lessening of vertical shear. The new convection was initially mentioned on the 260600 Z Significant

Tropical Weather Advisory (ABPW PGTW).
Based upon satellite intensity analysis (Dvorak, 1984) of satellite imagery between 1500 Z and 2100 Z on the 26 th , analysts of Detachment 1, 1st Weather Wing estimated that the disturbance had $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ surface winds, based on more organized and intense convection. The satellite reconnaissance inputs prompted the issuance of a Tropical Cyclone Formation Alert (TCFA) at 262230Z. An aircraft reconnaissance investigative mission was requested for the following day. At the time of the TCFA, synoptic data was not available near the center of the disturbance. However, surface data on the periphery of the disturbance implied that at least a 10 kt (5 $\mathrm{m} / \mathrm{sec}$ ) low-level circulation was present. The only reported stronger wind was the gradientlevel ( $3000 \mathrm{ft}(914 \mathrm{~m}$ )) report at Yap (WMO 91413 ), which increased from $10 \mathrm{kt}(5 \mathrm{~m} / \mathrm{sec})$ at 261200 Z to $15 \mathrm{kt}(8 \mathrm{~m} / \mathrm{sec})$ at 270000 Z as the disturbance passed northeast of the island on the 26th.

The first warning on Tropical Storm


Figure 3-04-3. Sperry interacting with a frontal boundary south of Iapan (302345Z June DMSSP visual imagery).

Sperry was issued on the 27 th , valid at 0000 Z , after visual satellite imagery showed that a central dense overcast and 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) maximum sustained surface winds were present. Aircraft reconnaissance later in the day located a 1001 mb circulation center with 40 kt ( 21 $\mathrm{m} / \mathrm{sec}$ ) maximum sustained surface winds, extending out to $50 \mathrm{~nm}(93 \mathrm{~km}$ ) southeast of the center. Initial forecasts called for Sperry to follow an around-the-ridge scenario and recurve. This forecast philosophy proved to be correct.

Sperry attained typhoon intensity 24 hours later at about 280000Z. The Aerial Reconnaissance Weather Officer reported Sperry as very compact, with $70 \mathrm{kt}(36 \mathrm{~m} / \mathrm{sec})$ maximum sustained surface winds surrounding a small, circular $15 \mathrm{~nm}(28 \mathrm{~km}$ ) diameter eye. The eye was open to the north and had a minimum sea-level pressure of 983 mb . Sperry developed a ragged eye while moving northwestward under the influence of the midlevel steering flow around the western periphery of the subtropical ridge. Its intensity peaked at 75 kt ( $39 \mathrm{~m} / \mathrm{sec}$ ) between 281200 Z and 281800Z (Figure 3-04-2). This set the stage for Typhoon Sperry's final phase.

By 290000 Z , with a frontal boundary and associated mid-latitude trough moving eastward across southern Japan, a recurvature scenario appeared most probable. JTWC incorporated this into the warnings and called for recurvature in 48-hours. Sperry came under the influence of the mid-latitude westerlies and recurved passing 175 nm ( 324 km ) to the east of the island of Okinawa in the Ryukyu Island chain.

After recurvature, Sperry started a gradual acceleration toward the northeast. By 1800 Z on the 30th, the intense central convection became displaced south-southwest of the low-level circulation center. A steady decrease in cloud organization and intensity followed. Figure 3-04-3 shows the proximity of the frontal boundary and effect of the strong vertical wind shear on the remaining convection. The final warning was issued as Sperry transitioned to extratropical at 010600 Z .

After completing extratropical transition, the low-level circulation drifted eastward embedded in the frontal boundary. There were no reports of lives lost or damage to shipping due to Typhoon Sperry.


## SUPER TYPHOON THELMA (05W)

Thelma was the first of four significant tropical cyclones to develop in July and the first super typhoon of 1987. Forecasting the timing and location of recurvature presented a problem for JTWC. After recurvature, Thelma slammed into Korea causing extensive damage and the loss of many lives.

As a tropical disturbance, Thelma's initial intensification was slow, but once the system became organized it developed at very near the normal Dvorak rate (Dvorak, 1984) of one "T-number" per day from 25 to 130 kt (13 to $67 \mathrm{~m} / \mathrm{sec}$ ). Thelma originated in the monsoon trough as a broad area of convection with slight curvature. Dvorak analysis estimated an intensity of $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$, while synoptic data indicated a cyclonic surface circulation was present along with upper-level divergence. As a result, the area was mentioned on the Significant Tropical Weather Advisory (ABPW PGTW) at 060600Z. Over the next
eight hours, the amount of convection and its organization increased. In addition, an aircraft reconnaissance investigative mission early on the 7th was able to close off the low-level circulation center and found a minimum sealevel pressure (MSLP) of 1003 mb . They also reported maximum sustained surface winds of $20 \mathrm{kt}(10 \mathrm{~m} / \mathrm{sec})$. At 070300 Z , JTWC issued a Tropical Cyclone Formation Alert.

The first warning on Tropical Depression 05W was issued at 071800 Z when the system demonstrated a steady increase in convection and organization, and satellite intensity analysis estimated $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ sustained surface winds. The forecast philosophy called for movement toward the north for 24 -hours through a weakness in the 700 mb ridge. The ridge was then expected to strengthen and drive the system toward the west. This did occur, but at speeds nearly triple those forecast.


Figure 3.05-1. PLot of OTCM guidance. The OTCM, IT'WC's primary dymamic aid, repeatedly indicated recurvature.

Initially, Thelma did not develop as quickly as expected. Once the first warning had been issued on Tropical Depression 05W, the system became broader and less organized. Aircraft reconnaissance scheduled for 080000 Z was unable to close off a surface center. Thirteen hours later, the poorly organized system passed about $60 \mathrm{~nm}(111 \mathrm{~km}$ ) to the north of Guam. Finally at 090000 Z (on warning number six), the system was upgraded to tropical storm intensity. The upgrade was based on aircraft reconnaissance data at 090029 Z which reported a MSLP of 996 mb and maximum sustained surface winds of 50 kt ( 26 $\mathrm{m} / \mathrm{sec}$ ). (Post-analysis indicated that the intensification had most probably occurred 12hours earlier.)

JTWC's primary aid, the One-Way Interactive Tropical Cyclone Model (OTCM), preferred a northwesterly track or hinted at recurvature in the 48 - to 72 -hour time frame beginning with the guidance for warning number 3 (080600Z July) (see Figure 3-05-1). Recurvature forecasts started with warning
number 3, valid at 080600Z. Although Thelma continued tracking in a westward direction, JTWC mistakenly continued to forecast recurvature for the next 30 -hours (spanning six warnings).

At 091200Z, Thelma began developing a banding eye. Warning number 10 , valid at 100000Z, upgraded the system to a typhoon. This action was based on the aircraft reconnaissance data at 092138 Z and 100011 Z that indicated an extrapolated MSLP of 974 mb and estimated maximum sustained surface winds of $80 \mathrm{kt}(41 \mathrm{~m} / \mathrm{sec})$. Typhoon Thelma reached its maximum intensity at 111200 Z , after a 36 -hour pressure fall of 66 mb (and a 12 -hour pressure fall of 25 mb ) down to 911 mb . During this time, Dvorak intensity estimates kept pace from approximately $77 \mathrm{kt} \mathrm{( } 40 \mathrm{~m} / \mathrm{sec}$ ) to approximately $127 \mathrm{kt}(65 \mathrm{~m} / \mathrm{sec})$. At 111200Z, Thelma became the season's first super typhoon. Afterward, infrared satellite imagery indicated a warming of the cloud tops which indicated that Thelma had peaked in intensity. Satellite


Figure 3.05-2. Plot of statistical aids (CSUM and COSMOS), dynamic numerical aid (OTCM), and persistence (XITRP) along with the final best track at 120000 Z , at the point of the abrupt track change towand the north.


Figure 3-05-3. Visual satelfite imagery sfiowing Typhoon Thelma after reacking maximum intensity. Note how the upper-Level outflow hias become restricted to the north (110632Z Iufy NOAA visual imagery).
imagery also indicated the system's upper-level outflow had become restricted to the north (see Figure 3-05-3). Aircraft reconnaissance at 112353 Z found that the eye was open to the north and was becoming elliptical.

Typhoon Thelma began a sharp turn toward the north at 120000 Z . Earlier, the dynamic forecast aid OTCM had repeatedly forecast movement toward the north or northwest (see Figure 3-05-2), but the typhoon continued to track westward. By 121200Z, Thelma was heading just west of north and the OTCM guidance was on track.

Even though Thelma's abrupt course change occurred 300 nm ( 556 km ) east of northern Luzon, heavy rains and high seas resulted in at least twelve fatalities in the

Philippine Islands. The northerly track took the typhoon west of the island of Okinawa, Japan, and resulted in the evacuation of military aircraft. Commercial airlines also interrupted service, which stranded thousands of air travelers as Thelma passed by.

Finally Typhoon Thelma slammed into South Korea, where widespread flooding caused death and destruction. Floods from Thelma covered thousands of houses, ruptured reservoirs, and destroyed roads, railroad tracks and embankments. News coverage from Korea reported that Thelma killed at least 123 people with 212 additional people listed as missing. The missing were largely seamen and fisherman, who were caught offshore. Officials estimated losses at more than $\$ 124$ million from damaged or destroyed houses, crops, and water craft.


## TYPHOON VERNON (06W)

Typhoon Vernon, the second of four significant tropical cyclones to develop in July followed closely on the heels of Super Typhoon Thelma (05W). It was a weak and disorganized system throughout most of its lifetime. As such, initial positioning problems arose in the Philippine Sea due to differences between realtime fix information from radar, satellite and aircraft.

The initial tropical disturbance was first detected in the near-equatorial trough near the island of Truk in the eastern Caroline Islands at 140000 Z July and was subsequently listed on the Significant Tropicar Weather Advisory (ABPW PGTW) as having poor potential for
development into a significant tropical cyclone. However, six hours later a low-level circulation was apparent on visual satellite imagery. The satellite intensity estimate (Dvorak, 1984) was 25 kt ( $13 \mathrm{~m} / \mathrm{sec}$ ) and a Tropical Cyclone Formation Alert (TCFA) was issued at 140830Z. Convective activity did not increase appreciably for the next two days, but the TCFA was reissued twice due to its persistence. On 161800 Z , the first warning was issued for Tropical Depression 06W, based on an estimate of maximum sustained surface winds of 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) from satellite imagery. The initial forecasts were based on a persistent westward trend with higher than normal speeds of 17 kt ( $32 \mathrm{~km} / \mathrm{hr}$ ).


Figure 3-06-1. The area of intense convection as seen on enfanced infrared satellite imagery. This feature was used to fix Vernon during most of the period between 170000 Z and the relocation at 181200Z (180042Z July $\mathcal{D M E S P}$ infrated imagery).


Figure 3-06-2. The exposed low-fevel circulation center whicf was mistakenly thought to be a secondary circulation can be seen in the center of the above image. This Low-Level circulation was not identified as Vernon's main circulation center until $181200 Z$ ( $180042 Z$ Iuly DMSSP visual imagery).

Between 170600 Z and 181200 Z , satellite fixes and radar reports indicated Vernon had continued to move westward toward the central Philippine Islands. During this time, the most intense area of curved convection remained just east of the Philippines (see Figure 3-06-1) and appeared to be the dominant feature. However, there was also a low-level circulation center northeast of the deep convection which was initially believed to be a secondary circulation center. Figure 3-062 shows this exposed low-level circulation center about 180 nm ( 333 km ) northeast of the primary mass of convection at 180042 Z .

After the last successful fix at 170224 Z July, keeping track of Vernon's weak low-level circulation center became increasingly more difficult. To compound the problem, the radar fixes at two different sites in the Philippine Islands (WMO 98558 and WMO 98447) reinforced the satellite analysis which continued to fix on the main convective mass that was moving towards southem Luzon (see Figure 3-06-3). As a result, Vernon's low-level circulation was not recognized until an exposed low-level circulation was identified by the satellite analyst at 181200 Z . Immediately thereafter, Vernon was relocated approximately


Figure 3-06-3. Typhoon Vernon's best track during the period when the system was relocated. Note that no aircraft fixes were made between $170224 Z$ and $181325 Z$ Iufy. Miso note the preponderance of radar fixes and satelfite fixes which indicated westward movement when the system was actually moving nortifwestward (dashed (ines). These fixes could not be used in the final best track.
$145 \mathrm{~nm}(269 \mathrm{~km})$ to the north-northeast of the original 180600 Z position. This relocation was subsequently verified at 181325 Z by the first aerial reconnaissance fix mission in nearly 36 hours.

The Aerial Weather Reconnaissance Officer (ARWO) on the fix mission reported passing a probable vortex center on the inbound leg of the primary fix mission as the aircraft was heading toward the fly-to-point given by the Typhoon Duty Officer. After consulting the

Typhoon Duty Officer, it was decided that the satellite and land radar fix position should be investigated. Once there, the ARWO reported rising heights at the 850 mb level and no lowlevel vortex. The Typhoon Duty Officer then concluded there was only one circulation center, vice multiple vortices. The aircraft crew was then requested to return to the vortex they had passed earlier and investigate it (see Figure 3-$06-4$ ). The ARWO subsequently located the vortex and reported a center height of 1363 m at 850 mb .


Figure 3-06-4. Plot of the 181325Z Iuly aircraft fix mission which determined the fow-fevel circulation center associated with Vernon.

After the relocation, radar reports from Guiuan Airport (WMO 98558) continued to fix Vernon $170 \mathrm{~nm}(315 \mathrm{~km})$ south of the aircraft verified position. From synoptic data, there is no evidence that a distinct circulation ever
existed separate from Vernon's exposed lowlevel center. There is also no evidence that the central Philippine Islands ever experienced any significant winds associated with Vernon.


Figure 3.06-5. Typhoon Vernon at its peak intensity of $65 \mathrm{Kt}(33 \mathrm{~m} / \mathrm{sec})(200143 Z \mathrm{guly}$ DMSP visual imagery).

Vernon began to track steadily northeastward toward Taiwan. It reached minimal typhoon intensity at 191200 Z (see Figure 3-06-5). At that point positioning by satellite was no longer a problem due to the better defined central features. On 21 July,


Figure 3-06-6. The spiralling rainband of Vernon as seen by radar from Hualien, Taiwan (WMO 46699) at 201500Z Iuly (Photograph courtesy of Central Weather Bureau, Taipei, Taizuar)..
Typhoon Vernon began to interact with the terrain of Taiwan as it skirted the eastern shore and rapidly weakened to a tropical depression (Figure 3-06-6). Vernon dissipated in the East China Sea on 22 July after passing the northern tip of Taiwan.


## TYPHOON WYNNE (07W)

Typhoon Wynne was the fifth typhoon in the western North Pacific in 1987 and was of interest due to several factors. Early communication with meteorologists from Kwajalein Atoll (WMO 91366) proved instrumental in relocating Wynne, using radar fixes during its formative stages. The system developed into the third "midget" typhoon of the year and maintained a visible eye for six days. Wynne tracked along a constant 294 degree bearing for four consecutive days, during which time, it crossed the northern Mariana Islands, causing extensive damage to the islands of Alamagan and Agrihan.

Wynne appeared as an amorphous, but persistent, mass of cloud in the maximum cloud


Figure 3-07-1. Wynne at the tropical storm stage of development about $200 \mathrm{~nm}(370 \mathrm{~km})$ west-nortfrwest of the Kwajalein $\mathcal{A t o l l}$. Note the relatively cloudfree ring surrounding the small bright CDO (222300Z Iuly $\mathcal{D M S S}$ visual imagery).
zone east of the dateline and was first mentioned on the 200600 Z July Significant Tropical Weather Advisory (ABPW PGTW). Analysis of the sparse synoptic data indicated convergence enhancing cross-equatorial lowlevel flow into the system in the horizontal with moderate wind shear in the vertical.

Wynne moved westward and continued to improve in convective organization. Satellite intensity analysis (Dvorak, 1984) of the welldefined spiral cloud bands at 210000 Z estimated $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ surface winds and 45 $\mathrm{kt}(23 \mathrm{~m} / \mathrm{sec}$ ) surface winds were forecast for the next day. Based on this information, a Tropical Cyclone Formation Alert (TCFA) followed at 210430 Z . Through the 20th, Wynne's track remained westward in response to the synoptic-scale flow south of the subtropical ridge axis. On 21 July, however, satellite reconnaissance fix positions indicated cloud system center movement towards the northwest. Due to this track change the alert area was redefined at 212030 Z and the TCFA was reissued.

Discussions on 22 July between the Typhoon Duty Officer (TDO) and meteorologists on the Kwajalein Atoll, Marshall Islands, provided invaluable positioning information. Kwajalein was receiving light winds and radar showed the main convection associated with the tropical cyclone to be well to the north of their location. The result was a $120 \mathrm{~nm}(222 \mathrm{~km})$ northward relocation of the 221200 Z warning position from its expected location. By the end of the day, Wynne had separated from the maximum cloud zone and drawn down into a small bright central dense overcast (CDO) (Figure 3-07-1).

An eye first became visible on satellite data at 240000 Z . From that point onward (a period of six days), the system was characterized by a small eye. The eye diameter changed slightly from $12 \mathrm{~nm}(22 \mathrm{~km})$ to 18 to 22 nm ( 33 to 41 km ) in diameter. Typical of a smaller than normal system, it had smaller than average $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ wind radii. Aircraft reconnaissance revealed this anamoly.

Figure 3-07.2. Plot of aircraft reconnaissance data from $252134 Z$ to 260125Z Iuly, showing the surface and 700 mb flight-level wind distribution around Typhoon Wynne. Note the greater extent of the wind radii in the northeastern semicircle.

Figure 3-07-3. Typfioon Wynne two fours before crossing the northerm Marianas and near its closest point of approach to Guam. With the low morning sun off the right side of the picture, differences in cloud top keigfits are accentuated by shadowing. In this image an apparent 'stadium' effect can Ge seen; the larger upper-Level inner eye wall Goundary slopes downward to the concentric smalfer low-level eye (252015Z July DSMSP visual imagery).


Figure 3.07-4. Midget Typfioon Wynne near maximum intensity. Note the well-defined 15 $n m(28 \mathrm{~km})$ eye ( $261815 Z$ Iuly $\mathcal{N} \mathcal{O A R}$ infrared imagery).

Wynne's wind radii were nearly twice as large in the northeast semicircle as elsewhere (Figure 3-07-2). This appears to be related to the pressure gradient between Wynne and the subtropical anticyclone to the north. Figure 3-07-3 shows Typhoon Wynne two hours before it passed directly over the northern Marianas island of Alamagan ( 240 nm ( 444 km ) northnortheast of Guam). At approximately the same time ( 240000 Z through 271200 Z ) Wynne followed an almost straight track along a mean 294 degree bearing. While on this course and mean speed of 16 kt ( $30 \mathrm{~km} / \mathrm{hr}$ ), the typhoon attained its maximum intensity of 125 kt ( 64 $\mathrm{m} / \mathrm{sec}$ ) on the 26th (Figure 3-07-4).

An interesting aspect of Wynne's travel across the western North Pacific was that it maintained a brisk forward speed of movement, even through recurvature, where it slowed only slightly to $10 \mathrm{kt}(19 \mathrm{~km} / \mathrm{hr}$ ). Typically, a larger decrease in forward speed is expected as a system passes through the area of weaker steering flow at the break in the subtropical ridge axis.

As Wynne rounded the western end of the mid-level subtropical ridge, it began to experience increasing vertical shear from the north. The exposed low-level cyclonic circulation became visible at 291500 Z (Figure 3-07-5). Even with this unfavorable environment in the vertical, there were strong winds associated with the system for the next two days. Aircraft reconnaissance at 291111 Z found 700 mb winds of $76 \mathrm{kt}(39 \mathrm{~m} / \mathrm{sec})$.

Wynne was downgraded from typhoon to tropical storm intensity on the 32nd warning (valid at 300000 Z ) after a satellite intensity estimate of $50 \mathrm{kt}(26 \mathrm{~m} / \mathrm{sec})$ was attained. Subsequent aircraft reconnaissance at 292157 Z and 300027 Z also reported maximum 700 mb flight-level winds of $60 \mathrm{kt}(31 \mathrm{~m} / \mathrm{sec})$.

Wynne continued slowly weakening as it moved eastward, south and southeast of the main Japanese island of Honshu. Its forward speed increased as a result of stronger mid-level westerly flow. At the same time, Wynne began entraining cooler, drier air from the north. As a consequence, extratropical transition was complete at 010000 Z August.


Figure 3-07-5. Whnne, at tropical storm intensity 130 $n m(241 \mathrm{~km})$ east of the gapanese island of Kyushu. Low-Level cloudiness defines the exposed circulation center. Of interest, the bright and dark patches on the ocean's surface to the east of the system are the result of sun-glint. These patches indicate the areas of relatively smooth ocean surface with less wind waves which are usually the result of lighter surface winds near the axis of the fower-tropospheric subtropical ridge. In this case the ridge axis runs east-to-west near 22 degrees North Latitude. Understanding the location and atmospfieric processes associated with this ridge are vitally important to tropical cyclone forecasting (292359Z July $\mathcal{D M} S P$ visual imagery).

In retrospect, the islands of Alamagan and Agrihan suffered the only recorded major damage due to Wynne's passage. Their crops were 90 to 100 percent destroyed and all coconut trees were downed. Fortunately no lives were lost. Except for this head-on meeting between Wynne and these islands, no synoptic data revealed the potent punch of this midget typhoon. Only direct aircraft measurement and indirect satellite reconnaissance recorded the wind intensities because of the system's small size.


## TYPHOON ALEX (08W)

Typhoon Alex was the fourth and final tropical cyclone to develop during the month of July, and combined with Typhoon Wynne ( 07 W ) to form the first multiple-storm situation of the 1987 western North Pacific tropical cyclone season. Wynne ( 07 W ) passed through the Marshall Islands and intensified to tropical
storm intensity as Alex showed initial signs of development on July 22nd. Six days later, on the 28th, Wynne ( 07 W ) began to slowly recurve south of Japan as Alex dissipated over the eastern China coast. The closest the two systems came to one another was 740 nm (1370 km ) late on the 28th.


Figure 3-08-1. Morning view of the tropical disturbance in the Philippine Sea which would develop into Typfioon Alex. Convective banding is evident in the low. Level cloud fines (220102Z July DMSP visual imagery).


Figure 3-08-2. Synoptic surface/gradient-level streanfine analysis of 230000Z Iuly data shows a broad cyclonic circulation in the Philippine Sea witf an estimated minimum sea-level pressure of 1000 mb and winds of $30 \mathrm{Kt}(15 \mathrm{~m} / \mathrm{sec})$. ( $\mathfrak{N o t e}$ : drifting buoy wind speeds (in Kt$)$ enclosed in 6 oxes.)

Alex developed in the western end of an active monsoon trough which stretched east-to-west 2400 nm ( 4445 km ) (south of 10 degrees North Latitude) from the dateline across the Marshall and Caroline Islands. Late on the 21 st , routine analysis of satellite imagery indicated a tropical disturbance persisting in an area of poorly organized convection 200 nm $(370 \mathrm{~km})$ to the southwest of Guam. This area was noted on the Significant Tropical Weather Advisory (ABPW PGTW) at 220600 Z due to its persistence and indications of convective banding in the low-level cloud lines visible on visual imagery that morning (Figure 3-08-1).

Over the next twelve hours, the convection increased and upper-level organization improved rapidly. Infrared satellite imagery at 221800 Z indicated a central core of heavy convection had developed. Surface winds were estimated at $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$ based on the Enhanced Infrared (EIR) technique (Dvorak, 1984). As a result, JTWC promptly issued a Tropical Cyclone Formation Alert (TCFA) at 221930 Z even though synoptic data indicated only a broad surface circulation with an estimated minimum sea-level pressure of 1005 mb .

Satellite intensity analysis at 230000 Z estimated surface winds of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ associated with this disturbance. A 30 kt (15 $\mathrm{m} / \mathrm{sec}$ ) ship observation north of the disturbance for this same time provided some ground truth to the Dvorak estimate (see Figures 3-08-2 and 3-08-3). Based on these data, JTWC immediately issued the first warning on Tropical

Depression 08W. Six hours later, on the second warning, Alex was upgraded to tropical storm intensity based on increased organization that became evident on satellite imagery at 230600 Z . Within 12 -hours a well-defined convective band could be seen on satellite imagery wrapping into the center.


Figure 3.08-3. Visual satelfite imagery near the time of the first warning on Tropical Depression 08W. See the 230000 Z Iuly synoptic surface/gradient-Cevel streamfine analysis in Figure 3-08-2 for comparison (230041Z Iufy DMSP visual imagery).


Figure 3-08-4. Initial OTCM 72-four guidance for ALex indicated the system would remain south of the subtropical ridge and move across the Philippine Islands. Afex's best track is also shown for comparison.


Figure 3-08-5. Initial $\mathcal{H}$ PAC 72-fiour guidance for $\mathcal{A l e x}$ agreed with the OTCM in keeping the system south of the subtropical ridge and moving it across the Philippine Istands. Alex's best track is also shown for comparison.

The main forecast problem occurred early, during the first two days of the system's lifetime. Alex was forecast to track across the Philippine Islands on warnings one through five. The primary guidance came from two forecast aids -- the One-Way Interactive Tropical Cyclone Model (OTCM) and the Half Climatology and Persistence Model (HPAC). Figures 3-08-4 and 3-08-5 show the guidance received for the first six warnings from the OTCM and HPAC, respectively. They incorrectly suggested Alex would remain south of the strong subtropical ridge, move across the Philippine Islands and then turn northward towards mainland China. JTWC forecasters determined the OTCM and HPAC guidance was flawed and, on the sixth warning, relocated Alex further north after several satellite fixes indicated it was moving towards the northwest rather than the west-northwest. Unfortunately, beginning at 240900 Z , there was increased
scatter in the satellite fixes as a cirrus canopy developed over the center. This left JTWC forecasters with no clear-cut indication of exactly where Alex's low-level center was. A solitary aircraft radar fix was obtained at 240916 Z which provided some close in information, however a trained Aerial Weather Reconnaissance Officer was not onboard the flight and the meteorological accuracy of the position was suspect. Figure $3-08-6$ shows a satellite image prior to the time of the aircraft fix. Notice the exposed low-level center is displaced slightly northeast of the heaviest convection. The radar site at Guiuan (WMO 98558) in the Philippine Islands fixed this area of heavy convection and added to the uncertainty as to where the actual location of Alex's center was.

Forecast guidance for the next five warnings indicated Alex should track through

Figure 3-08-6. Morning view of Alex. The exposed low-level center is displaced slightly northeast of the heaviest convection (240021Z Iuly $\mathfrak{\wedge}$ (ОЯA visual imagery).


59
the Luzon Strait and make landfall over mainland China to the west of Taiwan. JTWC forecasts for this time period ( 240600 Z through 250600Z) reflected this guidance. Also during this period, Alex continued to slowly intensify. Between 241500Z and 241800Z, it developed an eye. This eye was first implied by a warm spot in the central cloud mass on the nighttime infrared imagery (see Figure 3-08-7).

At 1200 Z on the 25 th, Alex reached its maximum intensity of $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ and was upgraded to typhoon status. At that time, Alex was 120 nm ( 222 km ) east of the northeast tip of Luzon. Forecast guidance at 251200 Z changed significantly, suggesting a more northward movement, which would take Alex east of Taiwan vice through the Luzon Strait. The reason for this change in computer forecast guidance appears to be twofold. First, a surface frontal boundary stalled across the eastern coast of Asia, and second, a large break developed between the upper-level subtropical ridge south of Japan and the Siberian High.

Alex remained at minimal typhoon intensity for another 30-hours and then began to slowly weaken. It was then steered toward the north by the low-level southerly flow east of the stalled front, which caused it to brush the eastern portion of Taiwan (Figure 3-08-8) and pass within $30 \mathrm{~nm}(56 \mathrm{~km}$ ) of the capital city of Taipei.

Shortly after passing Taipei, Alex was drawn slightly westward by the lee effect of its interaction with Taiwan's mountainous terrain. This caused Alex to make landfall on the China coast near the city of Wenzhou, 200 nm ( 370 km ) south of Shanghai. The system then moved inland and dissipated as a significant tropical cyclone. Figure 3-08-9 shows Alex with respect to Wynne for this same time period. Near 281800Z, the remnants of Alex, with its residual vorticity and moisture, once again moved over water but did not regenerate into a significant tropical cyclone. It did, however, add to the band of precipitation that had stalled over Korea and, as a consequence, over 12 inches ( 300 mm ) of rain fell within 24 -hours. This deluge triggered major flooding, landslides and loss of life. In contrast, the damage to Taiwan and China was minor.


Figure 3-08-7. An implied eye appears as a warm (white) spot in the central cloud mass (dark gray) (241837Z Jufy 고겨 enfuanced infrared imagery).


Figure 3-08-8. The tightly curved rainband and eye wall of Typhoon $\mathfrak{A l e x}$ as seen by radar from $\mathcal{H}$ ualien, Taiwan (WMO 46699) at 261400Z July (Thotograph courtesy of Central Weather Bureau, Taipei,


Figure 3-08-9. Typhoons $\mathcal{A}$ fex and Wynne (07W) appeared together on this thresfolded infrared satelfite image (Note: coldest cloud tops appear black). Alex had just moved infand over the eastern coast of China and Wynne was still on a northwestward track, heading toward OKinawa, Iapan (271341Z Iuly DSSSP inverted infrared imagery).


## SUPER TYPHOON BETTY (09W)

Super Typhoon Betty was the first of two tropical cyclones to hit Vietnam during the month of August. Betty was also the second super typhoon (intensity equal to or greater than $130 \mathrm{kt}(67 \mathrm{~m} / \mathrm{sec})$ ) of the 1987 western North Pacific tropical cyclone season and had the lowest reported minimum sea-level pressure ( 891 mb ). It intensified (deepened) explosively (Holliday and Thompson, 1979) prior to making landfall in the Philippine Islands. Other distinguishing characteristics were the large size of the area of intense convection, the small radius of maximum wind and the associated strong low-level southwest monsoonal inflow. Also of note was the large radius of gale force winds in Betty's northwest semicircle, due to the enhancement of surface winds by a strong pressure gradient between the tropical cyclone and the subtropical ridge.

After Typhoon Alex (08W), which had developed in the low-level southwest monsoon
trough, dissipated on the 28 th of July, the midlevel subtropical ridge again became wellestablished over the western North Pacific. Coincident with Alex's (08W) movement toward the north was the replacement of the strong low-level southwest monsoonal flow over the South China Sea by the ridge.

Betty was first detected on the 7th of August as a tropical disturbance embedded in the monsoon trough, which extended from the Marshall Islands westward to the Philippine Islands. Satellite intensity estimates (Dvorak, 1984) showed surface winds of $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$ when the disturbance was $65 \mathrm{~nm}(120 \mathrm{~km})$ north-northwest of the island of Belau in the western Caroline Islands: The system cloudiness developed rapidly early on the 8th prompting JTWC to issue a Tropical Cyclone Formation Alert at 0300Z. Figure 3-09-1 shows the disturbance on the 8th of August exhibiting


Figure 3-09-1. Super Typfioon Betty as a tropical disturbance in the monsoon trough Signs of organized upper fevel outflow were present ( 0812572 August $\mathcal{D M S P}$ visual imagery).


Figure 3-09-2. Plot of Betty's minimum sea-level pressure and central 700 mb equivalent potential temperature during the period $082300 Z$ and $120000 Z$ August. Once the critical crossing of the surface pressure and Theta-E traces occurred (at point $\mathcal{A}$ ), explosive deepening was expected.
signs of organization in its upper-level outflow pattern. The system moved westward and reached tropical storm intensity on 9 August.

In the 37 -hour period between 100000 Z and 111300 Z , Betty's minimum sea-level pressure dropped from 985 mb to 892 mb , a decrease of 93 mb . This translates to a drop of approximately $2.5 \mathrm{mb} / \mathrm{hr}$ (sustained for at least 12-hours) or explosive intensification. JTWC uses a technique (Dunnavan, 1981), in which the 700 mb equivalent potential temperature, Theta-E, (a measure of the tropical cyclone's thermodynamic energy based on the central 700 mb temperature and dew point) and the
minimum sea-level pressure are compared to forecast explosive intensification. This technique forecasts intensification to below 925 mb whenever the plots of minimum sea-level pressure and Theta-E intersect near the critical values of 950 mb and 360 degrees Kelvin, both values being statistical means derived from analysis of past intense tropical cyclones. Figure 3-09-2 is a plot of Betty's minimum sealevel pressure and Theta-E during the period 082300 Z to 120000 Z . At point A (101730Z) the two lines intersect, as the minimum sealevel pressure at this time is plummeting downward. Based on this information, explosive deepening was forecast.

Figure 3-09-3 shows Super Typhoon Betty near maximum intensity with a welldefined eye and intense convection covering a large area around the system. Aircraft reconnaissance on the 10 th and 11th of August consistently located the maximum surface winds 10 to 15 nm ( 19 to 28 km ) from the center and radar eye diameters of 11 to 15 nm ( 20 to 28 km ). Both measurements showed the center to be very small and compact.

The threat posed by Super Typhoon Betty resulted in the evacuation of aircraft from Cubi Point Naval Air Station and Clark Air Base, as well as the movement of several ships from Subic Bay. Later, news services reported at least twenty people were killed, seven missing and more than 60,000 left homeless as a result of Betty's passage over the Philippine Islands. Damage to buildings and crops was estimated in the millions of dollars.

Betty weakened from $140 \mathrm{kt}(72 \mathrm{~m} / \mathrm{sec})$
to $110 \mathrm{kt}(57 \mathrm{~m} / \mathrm{sec})$ as it accelerated across the central Philippine Islands. The subtropical ridge continued to be the dominant synopticscale feature, extending westward into the South China Sea.

After entering the South China Sea early on the 13 th of August and still maintaining $95 \mathrm{kt}(49 \mathrm{~m} / \mathrm{sec}$ ) winds, Super Typhoon Betty began to reintensify over water as it continued on a west-northwesterly track. By 140600Z, Betty's intensity had peaked again, at $115 \mathrm{kt}(59 \mathrm{~m} / \mathrm{sec}), 390 \mathrm{~nm}(722 \mathrm{~km})$ south of Hong Kong. Betty slowly weakened as it began to interact with the mountains of Vietnam and the island of Hainan which prevented further intensification by hampering its low-level inflow. Crossing the Gulf of Tonkin in less than a day, Betty slammed into the coast of Vietnam $190 \mathrm{~nm}(352 \mathrm{~km})$ south of Hanoi. The final warning on Betty was issued at 161800 Z as the system weakened and dissipated over the mountains inland.


Figure 3-09-3. Super Typfioon Betty near maximum intensity. This exparded image shows the well-defined eye and large symmetrical area of intense convection (110057Z August DMSS visual imagery).


## TYPHOON CARY (10W)

Typhoon Cary was the second significant tropical cyclone to develop in August. It shared the western North Pacific with Super Typhoon Betty (09W) for four days; coexisted with Super Typhoon Dinah (11W) for one and a half days, and then was part of the first three-storm situation of 1987 for 12-hours with Dinah (11W) and Tropical Storm Ed (12W).

Cary was first identified on the 6th of August as an area of convection, that persisted longer than usual in the monsoon trough 200 $\mathrm{nm}(370 \mathrm{~km})$ to the southwest of the island of

Pohnpei in the eastern Caroline Islands. As a result, the cloud system was placed on the Significant Tropical Weather Advisory (ABPW PGTW) at 070600Z. The system remained broad and poorly organized over the next four days. By the 12 th, upper-level outflow had improved and was unrestricted in all quadrants. Additionally, satellite intensity analysis (Dvorak, 1984) showed winds of $25 \mathrm{kt}(13$ $\mathrm{m} / \mathrm{sec}$ ). A Tropical Cyclone Formation Alert (TCFA) followed at 120300 Z .

Aircraft reconnaissance at 122302 Z estimated the maximum surface winds at 55 kt


Figure 3-10-1. Time series from 120000Z to 171200Z October showing the natural scatter of raw intensity data and the resulting final best track intensities.


Figure 3-10-2. Plot of data from an aircoaft reconnaissance mission at 141151Z August, indicating that the circulation center at flight level ( 850 mb ) was about 18 $n \mathrm{~m}(33 \mathrm{~km})$ to the south of the aircraft fix pasition.
( $25 \mathrm{~m} / \mathrm{sec}$ ), but the minimum surface pressure reported was only 996 mb , which usually supports a maximum wind speed of 37 kt (19 $\mathrm{m} / \mathrm{sec}$ ). At 121800 Z , satellite intensity analysis determined that Cary's intensity was 35 kt ( 17 $\mathrm{m} / \mathrm{sec}$ ). Subsequent satellite intensity analysis, six hours later, indicated that Cary had winds of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$. Based on these intensity estimates the first warning on Tropical Storm Cary was issued at 130000 Z with winds of 50 kt ( $26 \mathrm{~m} / \mathrm{sec}$ ) gusting to $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$. Postanalysis revealed that Cary most probably had an intensity of $40 \mathrm{kt}(21 \mathrm{~m} / \mathrm{sec})$ at the time of the first warning, and had reached tropical storm intensity six hours earlier at 121800 Z .

The synoptic feature that dominated the low-level steering flow was the subtropical
ridge to the north. With Cary embedded in the monsoon trough east of Super Typhoon Betty (09W), the initial forecast reasoning was for Cary to track northwestward south of the ridge, closely paralleling the track of Betty (09W). The intensity was expected to increase at a normal rate, but the initial intensification and development of Cary was inhibited by Betty (09W) to the west. As Betty (09W) began to weaken as it crossed the Philippine Islands, Cary's upper-level outflow improved enough to allow development.

Satellite intensity analysis over the next 36-hours indicated that Cary developed rapidly to $90 \mathrm{kt}(46 \mathrm{~m} / \mathrm{sec})$ at 140600 Z . Postanalysis revealed that the satellite-derived intensity estimate ("T-number") was incorrect -


Figure 3-10-3. Typfroon Cary at near maximum intensity and approaching landfall on the island of Luzon (170036Z August DMSPP visual imagery).
the diameter of the cold convective cover was misinterpreted as the diameter of a central dense overcast. Aircraft reconnaissance during the same period indicated that Cary was weakening (see Figure 3-10-1). Aircraft reconnaissance at 140029 Z reported maximum winds of 50 kt ( 26 $\mathrm{m} / \mathrm{sec}$ ), however, a minimum sea-level pressure of only 1004 mb was reported, which normally supports only $21 \mathrm{kt}(11 \mathrm{~m} / \mathrm{sec})$. Aircraft reconnaissance at 141151 Z found 850 mb winds of $36 \mathrm{kt}(19 \mathrm{~m} / \mathrm{sec})$ and an 850 mb height of only 1425 meters, which extrapolated to about

1000 mb surface pressure and surface winds of 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ). The accuracy of the latter fix was especially questionable since the flightlevel winds did not support the position in the vortex data message as being the low-level center. Additionally, the Aerial Reconnaissance Weather Officer indicated there was frequent lightning in all quadrants, possible multiple centers and that a penetration of the center was not feasible on this mission. Possibly the 850 mb fix (as indicated on Figure 3-10-2) should have been made about $18 \mathrm{~nm}(33 \mathrm{~km})$ to the south as shown by the streamline analysis. Also, the excessive scatter (see Figure 3-10-1) of the intensity data acquired by different platforms during this phase of Cary's life is not often observed.

The last sçheduled western North Pacific aircraft reconnaissance mission was flown on the 15th of August. At 151405Z, the maximum 700 mb winds reported were 61 kt $(31 \mathrm{~m} / \mathrm{sec})$, and the 700 mb height was 3007 meters. This corresponds to about a 990 mb surface pressure and $46 \mathrm{kt}(24 \mathrm{~m} / \mathrm{sec})$ winds. These values represented the strongest winds and lowest pressures found by aircraft reconnaissance on this system. Earlier Dvorak intensity estimates at 150600 Z showed winds of 90 kt ( $46 \mathrm{~m} / \mathrm{sec}$ ). Post-analysis setrled on a maximum wind of about $70 \mathrm{kt}(41 \mathrm{~m} / \mathrm{sec})$ at 151800Z (see Figure 3-10-1).

Cary reached its maximum intensity of $85 \mathrm{kt}(44 \mathrm{~m} / \mathrm{sec})$ at 170600 Z , shortly before making landfall on eastern Luzon (Figure 3-103). The intensity dropped from $85 \mathrm{kt}(44 \mathrm{~m} / \mathrm{sec})$ to $50 \mathrm{kt}(26 \mathrm{~m} / \mathrm{sec})$ as Cary crossed the Philippine Islands. Extensive flooding was reported in the northern Philippine Islands. There were no reports of casualties.

Cary continued onward across the South China Sea and reintensified to 70 kt ( 36 $\mathrm{m} / \mathrm{sec}$ ) just southeast of the island of Hainan. The closest point of approach was 15 nm ( 28 km ) to the south of Hainan at 211800 Z . Cary then tracked toward the west through the Gulf of Tonkin and swept into northern Vietnam at 221200 Z . The final warning was issued at that time. The dissipating system with its residual vorticity and moisture tracked northwestward over land into Burma before finally losing its identity on satellite imagery.


## SUPER TYPHOON DINAH (11W)

Super Typhoon Dinah (11W), the most destructive typhoon to strike Okinawa and the southern islands of Japan in the past 20 years, caused extensive damage to both Japanese civilian properties and U.S. military bases and assets.

Dinah was first observed on satellite imagery as a disorganized cluster of weak convection in the near-equatorial trough on 18 August. By the 19th, convection became better organized and the disturbance was noted on the Significant Tropical Weather Advisory (ABPW PGTW) issued at 190600Z. During the next
eighteen hours, Dinah developed a low-level circulation as it passed northwest of the island of Pohnpei and moved beneath moderate directional and speed divergence at the 200 mb level. The 200000 Z satellite imagery indicated weak convective curvature and, as a result, a Tropical Cyclone Formation Alert was issued at 200427 Z . During the next eighteen hours, satellite imagery indicated a considerable increase in convection which had become more centralized (see Figure 3-11-1). The system was assigned a Dvorak intensity number ("Tnumber") of 2.0 which corresponded to maximum sustained surface winds of 30 kt ( 15


Figure 3-11-1. The initial development of Super Typfioon Dinaf was first noted as a considerafle increase in the amount of convection (202102Z August $\mathfrak{N O A A}$ visual imagery).

$\mathcal{F}$ igure 3-11-2. Super Typfioon Dinaf in the Philippine Sea near maximum intensity (260054Z August DMSP visual imagery).
$\mathrm{m} / \mathrm{sec}$ ) and an estimated MSLP of 1000 mb . At 210000 Z , the first warning was issued on Tropical Depression 11W when it was located $300 \mathrm{~nm}(556 \mathrm{~km})$ east-southeast of Guam.

Between 210000Z and 211800Z, Tropical Depression 11W assumed a more westward track in response to the strengthening subtropical ridge to the north and moved beneath an upper-level anticyclone which had associated strong speed divergence southwest of the system. The increased outflow signature on satellite imagery allowed for a Dvorak intensity estimate of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$. Based on this estimation, Tropical Depression 11 W was upgraded to Tropical Storm Dinah (11W) at 211800Z.

Over the next forty-eight hours, Dinah moved westward passing 120 nm ( 222 km ) south of Guam at 220300 Z with maximum sustained surface winds estimated at 40 kt (21 $\mathrm{m} / \mathrm{sec}$ ). Dinah did not intensify at the normal rate of one "T-number" per day. This was apparently due to $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec}) 200 \mathrm{mb}$ winds over the cyclone which created an undesirable shearing environment. However, by 240000 Z , Dinah had moved away from this unfavorable shearing environment and developed a good anticyclonic outflow pattern which was visible on satellite imagery. The 241200 Z 200 mb streamline analysis confirmed this and indicated a good cyclonic outdraft directly over Dinah's center which became anticyclonic as it moved radially outward from
the center. During the first half of this period, Dinah tracked westward and then gradually turned more toward the west-northwest. JTWC forecasts correctly predicted the system's motion which was supported by the dynamic One-Way Interactive Tropical Cyclone Model (OTCM).

A Dvorak intensity analysis of satellite imagery at 240300 Z estimated maximum sustained surface winds of $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ and an estimated MSLP of 980 mb . On the 240600Z warning, Tropical Storm Dinah was upgraded to typhoon status. At that time it was located $500 \mathrm{~nm}(926 \mathrm{~km})$ west of Guam. Between 240600Z and 250600Z, Typhoon Dinah's outflow continued to increase with some restriction northwest through northeast of the cyclone which was associated with weak
short-wave troughs passing to the north. However, those minor restrictions did not inhibit Dinah from continuing to intensify at the normal Dvorak rate.

During the next twenty-four hours, Typhoon Dinah's intensity increased at a rate much faster than the normal one "T-number" per day and by 260000 Z it reached super typhoon intensity ( 130 kt or $67 \mathrm{~m} / \mathrm{sec}$ ) at a location $500 \mathrm{~nm}(926 \mathrm{~km}$ ) east of northern Luzon (see Figure 3-11-2). Dinah remained at super typhoon intensity for only a few hours but maintained maximum sustained surface winds of $110 \mathrm{kt}(57 \mathrm{~m} / \mathrm{sec})$ or greater until 280600 Z .

From 240600 Z until 281200 Z, Dinah basically tracked toward the northwest at an average forward speed of $11 \mathrm{kt}(20 \mathrm{~km} / \mathrm{hr})$


Figure 3-11-3. Dinaf during its dissipating stage passing to the west of Okinawa, Japan (290605Z August NOAA visual imagery).
during the first twenty-four hour period, slowing to an average of $6 \mathrm{kt}(11 \mathrm{~km} / \mathrm{hr})$ by 251800 Z . The slower forward speed is typical of a very well-developed and very intense tropical cyclone as it approaches the axis of the subtropical ridge prior to recurvature.

After 281200Z, Typhoon Dinah made a turn toward a more northerly track as it moved around the western periphery of the subtropical high. During the next thirty-six hours, Dinah moved into unfavorable upper-level conditions in the form of impinging mid-level short-wave troughs moving northeastward across eastern China and Japan. As each short-wave trough passed north of Dinah, upper-level wind shear increased and the system's outflow became restricted. As a result, Dinah steadily weakened.

Although Dinah began to weaken after 281200 Z (Figure 3-11-3), it still had maximum sustained surface winds of $85 \mathrm{kt}(44 \mathrm{~m} / \mathrm{sec})$ as it passed $90 \mathrm{~nm}(167 \mathrm{~km})$ west of Kadena Air Base (Figure 3-11-4) at 291500Z. It caused considerable damage to U.S. military facilities
on Okinawa. One person was killed and six people were injured. Trees were uprooted or broken off (Figure 3-11-5), utility poles and lines were blown down, and roofs and suffered structural damage. Total damage estimates to U.S. military facilities on Okinawa were in excess of $\$ 1.3$ million. Maximum sustained surface winds on Okinawa were 63 kt ( 32 $\mathrm{m} / \mathrm{sec}$ ) with gusts from 98 to $106 \mathrm{kt}(50$ to 55 $\mathrm{m} / \mathrm{sec}$ ). Minimum sea-level pressure observed was 983 mb at 291755 Z . By 300000Z, Dinah was located $120 \mathrm{~nm}(222 \mathrm{~km})$ northwest of Okinawa with maximum sustained surface winds estimated to be $85 \mathrm{kt}(44 \mathrm{~m} / \mathrm{sec})$. A ship passing 30 nm ( 56 km ) northeast of Dinah's center at that time reported sustained winds of $75 \mathrm{kt}(39 \mathrm{~m} / \mathrm{sec})$ from the southeast and a sealevel pressure of 938.7 mb .

Dinah began to recurve by 300000 Z , assumed a north-northeasterly track and accelerated while still maintaining maximum sustained surface winds of $85 \mathrm{kt}(44 \mathrm{~m} / \mathrm{sec})$. At 301700Z, Typhoon Dinah passed 60 nm (111 km ) northwest of Sasebo Naval Base in western


Figure 3-11-4. Radarscope pfioto of Dinafi at 291224Z August (Photo courtesy of Detachment 8, 20 Weather Squadron, Kadena $\mathfrak{A B}$, Japan).


Figure 3-11-5. Trees on OKinawa were damaged and uprooted by the kigh winds associated with Dinaf's passage (Photo courtesy of Detacfiment 8, 20 Weather Squadron, Kadena AB, Japan).

Japan where maximum sustained surface winds of 60 to 65 kt ( 31 to $33 \mathrm{~m} / \mathrm{sec}$ ) with gusts to 90 $\mathrm{kt}(46 \mathrm{~m} / \mathrm{sec}$ ) were observed. Extensive damage was caused by the storm surge and tidal action on seawalls and piers. A landing craft from the USS San Bernardino was destroyed when the seawall eroded and the pier collapsed. Damage also occurred to trees, utility lines and poles, and some building structures. Damage costs to the Japanese Sasebo Navy complex were in excess of $\$ 6.7$ million, making Dinah the worst tropical cyclone to strike southwest Japan in recent history.

By 310000 Z , Dinah was becoming extratropical as it began to merge with a midlatitude frontal system that extended southwestward across the Sea of Japan. It was
beneath the polar jet stream which had winds in excess of 90 kt ( $46 \mathrm{~m} / \mathrm{sec}$ ). Dinah was downgraded to a tropical storm as its convection sheared off to the northeast. The final warning was issued at 310600 Z as the cyclone continued to accelerate toward the northeast at 33 kt ( 61 $\mathrm{km} / \mathrm{hr}$ ).

Throughout Dinah's life, JTWC consistently forecast recurvature and acceleration toward the northeast through the Sea of Japan. Forecast track errors were smaller than average. The dynamic aid OTCM performed extremely well during recurvature, while the objective aids Half Persistence and Climatology (HPAC) and climatology were used extensively as Dinah passed beneath the subtropical ridge.


Figure 3-11-5. Trees on Okinawa were damaged and uprooted by the figh winds associated with Dinah's passage (Photo courtesy of Detachment 8, 20 Weather Squadron, Kadena AB, Iapan).

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## TROPICAL STORM ED (12W)

Tropical Storm Ed (12W) was the third of four significant tropical cyclones that occurred during the month of August. Ed was a difficult system for JTWC to locate and forecast because of its fluctuations in intensity, speed and track direction, and its poorly defined cloud signature.

Ed formed during the third week of August in the western North Pacific monsoon trough about $90 \mathrm{~nm}(167 \mathrm{~km})$ east of the island of Majuro in the Marshalls. It was first detected as an area of persistent convection with a
coincident weak low-level cyclonic circulation. This suspect area appeared on the Significant Tropical Weather Advisory (ABPW PGTW) at 200600Z. For the next 24 -hours, the disturbance moved rapidly at a speed of 17 to 23 kt ( 32 to 43 $\mathrm{km} / \mathrm{hr}$ ) toward the west-northwest. Improved upper-level outflow and increased central convection prompted the first Tropical Cyclone Formation Alert (TCFA) at 210600 Z .

At 212130Z, a second TCFA was issued to supersede the first TCFA, since the disturbance was moving out of the original alert


Figure 3-12-1. Formative stage of Tropical Storm Ed (212253Z $\mathcal{A u g u s t} \mathcal{D M S S}$ visual imagery).


Figure 3-12-2. The regenerated Tropical $^{2}$ Depression $12 \mathcal{W}$ shortly before it was upgraded once more to Tropical Storm Ed (262252Z August DMSP visual imagery).
area. The disturbance continued on a westnorthwestward track at slightly lower speeds of 14 to 17 kt ( 26 to $32 \mathrm{~km} / \mathrm{hr}$ ).

Visual satellite imagery (Figure 3-12-1) showed tighter curvature of the convective cloud lines and increased cirrus outflow to the north. Also, drifting buoys in the area indicated surface wind speeds of 25 to 30 kt ( 13 to 15 $\mathrm{m} / \mathrm{sec}$ ). As a result, at 220000 Z the second TCFA was upgraded to Tropical Depression 12W. Unexpectedly, thirty-hours later Tropical Depression 12 W showed significantly decreased convection and system organization on satellite imagery. Consequently, a final warning was issued. The tropical disturbance was then placed on the ABPW PGTW and monitored for signs of future regeneration.

Ed did maintain its low-level identity even as Typhoon Dinah (11W), which was further to the west, was increasing the vertical shear aloft over it. Finally, a ragged central dense overcast persisted and the system's upperlevel outflow redeveloped. The third TCFA followed at 240800 Z . However, by 250600 Z , the TCFA was cancelled when the upper-level outflow from Super Typhoon Dinah (11W), located to the west, increased its shearing effect on Ed which caused the convection to significantly decrease.

At 262030Z, a fourth TCFA was issued when cloudiness associated with the disturbance flared-up again. Satellite intensity analysis (Dvorak, 1984) estimated the intensity of the system at $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$. This TCFA was
almost immediately upgraded as Tropical Depression 12W, with a valid time of 261800 Z , based on the receipt of a new satellite picture, which indicated that the disturbance had been developing more rapidly than previously expected (see Figure 3-12-2).

The regenerated Tropical Depression 12W was further upgraded to tropical storm intensity at 270000 Z . This upgrade was based on the Dvorak satellite intensity analysis at 261800 Z , that indicated $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ sustained surface winds. In addition, at

270600Z, Tropical Storm Ed's position was relocated on the warning due to the formation of a $60 \mathrm{~nm}(111 \mathrm{~km})$ diameter central dense overcast from a central cold cover. As a result, Ed's center location was moved $45 \mathrm{~nm}(83 \mathrm{~km})$ farther north. Later, at 271200Z, Tropical Storm Ed (12W) was relocated a second time when satellite fixes revealed that the system had moved $75 \mathrm{~nm}(139 \mathrm{~km}$ ) further north than previously forecast. When the central convection was finally stripped away from the lowlevel circulation, the last warning was issued at 280000Z (see Figure 3-12-3).


Figure 3-12-3. Tropical Storm Ed (12W) after the system fad shed its central dense overcast (270445Z $\mathfrak{A u g u s t} \mathfrak{N O}(\mathcal{A R}$ visual imagery).


TYPHOON FREDA (13W)

Freda was the first of seven significant tropical cyclones to develop during the month of September and the middle tropical cyclone (geographically) of three systems that developed at nearly the same time; namely Freda, Typhoon Gerald (14W) and Super Typhoon Holly (15W). During this three tropical cyclone outbreak, individual development and movement trends were very similar even though the systems were never closer together than 900 $\mathrm{nm}(1667 \mathrm{~km})$. Freda was unusual because although it traversed less than 10 degrees of longitude while in warning status, it moved northward for almost twenty-five degrees of latitude. Freda's thirteen day life span and fifty warnings were WESTPAC records for 1987.


With the tropical disturbance just southeast of Guam, there was heightened concern about intensification as the system moved into an area of decreased vertical shear. During the night, infrared satellite images showed a flaring of convection, rapidly expanding cirrus outflow and a speedy displacement of the cloud system toward the west. Satellite analysis at 041745 Z estimated maximum sustained surface winds of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ and supported the issuance of the first warning on Tropical Depression 13W at 041800 Z . (This was also the time JTWC went to warning on Tropical Depression 14W.) Within six hours, after the first visual satellite imagery provided a better look, Tropical Depression 13W was relocated 215 nm (398
Freda developed in $\mathcal{F}$ igure 3-13-1. First appearance of $\mathcal{F r e d a ' s ~ e y e ~ k m ) ~ e a s t - s o u t h e a s t ~ o f ~ t h e ~}$ an active monsoon trough. ( $061723 Z$ September $\mathcal{N}$ ( $\mathcal{A R}$ infrared imagery). earlier expected position. The disturbance first appeared as a persistent cluster of convection in the eastern Caroline Islands on the 1st of September. Due to the persistent convective activity it was mentioned as a new suspect area on the 030600 Z Significant Tropical Weather Advisory (ABPW PGTW). A low-level cyclonic circulation was apparent in the synoptic surface/gradient-level data beginning at 031200 Z . By 040000 Z , synoptic data indicated winds of 20 to 30 kt ( 10 to $15 \mathrm{~m} / \mathrm{sec}$ ). Satellite intensity analysis (Dvorak, 1984) estimated maximum sustained surface winds of $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$. These data, plus a distinct gradient-level circulation and a 3 mb pressure fall over the past 24 -hours (to a minimum of 1003 mb ) supported a Tropical Cyclone Formation Alert issued at 040357 Z .

Warning number two included the amplifying remarks:

Satellite imagery over the past six hours for Tropical Depression 13 W indicate that the feature previously tracked on infrared imagery, has weakened, hence the system has been relocated. The latest visual imagery shows low-level cloud lines placing the lowlevel circulation center substantially further to the east than previously expected. This also indicates a slower forward speed.


Figure 3-13-2. Typhoon $\mathcal{F r e d a}$ near maximum intensity. Note the efongation of the cloud system from east-northeast to west-soutfrwest (110030Z September DMSSP visual imagery).

Freda passed approximately 30 nm ( 56 km ) southwest of Guam while moving northwestward at $14 \mathrm{kt}(26 \mathrm{~km} / \mathrm{hr})$ with an estimated intensity of 25 to 30 kt ( 13 to $15 \mathrm{~m} / \mathrm{sec}$ ). Once past Guam, Freda developed rapidly and was upgraded to tropical storm intensity at 050600Z. (It was at this time that JTWC also began
warning on Tropical Depression 15W, thus creating the second three-storm warning situation of the year.)

Suddenly, twelve-hours later, Freda appeared to become quasi-stationary at a position approximately 250 nm ( 463 km ) to the


Figure 3-13-3. Plot of the daily amounts of precipitation at two recording stations on Guam as spiral convective arms from $\mathcal{F}$ reda passed over the island.
west-northwest of Guam. This was also the same time that Tropical Storm Gerald (14W) became quasi-stationary. The two systems were approximately $900 \mathrm{~nm}(1667 \mathrm{~km})$ apart at that time.

With the appearance of a small ragged eye on satellite imagery at $061723 Z$ (see Figure 3-13-1), Freda was upgraded to typhoon intensity. After executing a tight cyclonic loop, Freda began to move slowly westward on the 8 th. Then, on the 10 th, Freda slowed and started a tight turn toward the northeast. Concurrently with the track change, Freda reached an estimated peak intensity of 125 kt ( $64 \mathrm{~m} / \mathrm{sec}$ ), based on Dvorak satellite intensity analysis. Figure 3-13-2 shows Freda early on
the 11th as it rounds the western periphery of the subtropical ridge. Note the elongation of the cloud system into an east-northeast/westsouthwest orientation. This asymmetry is a consequence of adjustments between the tropical cyclone and the ambient flow. (One day prior to Freda's change in track toward the north, Super Typhoon Holly (15W) also moved northward. At 091200Z, the two systems were approximately $1080 \mathrm{~nm}(2000 \mathrm{~km})$ apart. Super Typhoon Holly (15W) had been steadily moving closer to Freda from the east prior to the northward bends in their tracks.)

During this prolonged northward trek, a consequence of the intense monsoonal trough and the absence of a strong subtropical ridge,

Freda started to slowly accelerate and weaken. At 1800 Z on the 13th, Freda was downgraded to tropical storm intensity.

On the 16th, Freda began to interact with an eastward-moving, mid-level trough passing to the north of the system. This interaction resulted in a curved track toward the northeast. As a result, Freda missed the southeastern tip of Honshu by approximately

180 nm ( 333 km ). Shortly thereafter, Freda began extratropical transition as vertical wind shear increased and the system entrained dry, cool, mid-latitude air. The last warning was issued by JTWC at 170000 Z as the system accelerated toward the northeast.

Guam received two distinct heavy periods of rain over five days when Freda stalled to the west (Figure 3-13-3). Specifically


Figure 3-13-4. Gradient-Level wind speeds and sea-Level pressure on Guam as Freda stalled soutfiwest of the isfand. The period of ligfter winds from the 10th through the 11th was associated with the proximity of the zone of low-level speed convergence (convergent asymptote) between Freda and Super Typfoon $\mathcal{H}$ olly (15W).
on the 4th of September, when Freda was close by NAVOCEANCOMCEN/STWC located on Nimitz Hill, Guam received 3.35 in ( 8.51 cm ) of precipitation, and the Naval Oceanography Command Detachment at the Naval Air Station Agana, a few miles further north, received 3.75 in ( 10.93 cm ). On the 8th, over 2.5 in ( 6.35 cm ) of rain fell on Guam as convection associated with a spiral band passed overhead. Due to the proximity of Freda, and later Super Typhoon Holly (15W), Guam experienced periods of
gales from the south-southwest to westsouthwest for nearly 10 days (from the 5th through the 14th) (Figure 3-13-4). The strongest observed winds reported during this period were the $40 \mathrm{kt}(21 \mathrm{~m} / \mathrm{sec})$ southwesterly gradient-level winds at 121200 Z . The resulting high seas and hazardous surf through the Marianas disrupted shipping, destroyed seawalls, damaged reefs, eroded beaches and stranded islanders; but fortunately no lives were lost.


## TYPHOON GERALD (14W)

Typhoon Gerald developed in early September in an active monsoon trough at the same time that Typhoons Freda (13W) and Holly (15W) were intensifying further to the east. Gerald was unique in that it matured within the monsoon trough and did not detach from it. The most distinctive feature of Gerald was an unusually large eye.

After Typhoon Dinah (11W) moved northward through the East China Sea and became extratropical in the Sea of Japan, the minimum sea-level pressures (MSLPs) east of the Philippine Islands remained slightly lower ( 1005 mb ) than the seasonal mean of 1007 mb . This below normal low-pressure area was not
mentioned as a suspect area on the Significant Tropical Weather Advisory (ABPW PGTW) until 020600Z September, when persistent convection appeared.

A Tropical Cyclone Formation Alert (TCFA) at 020830Z upgraded the suspect area in the Philippine Sea after a sudden flare-up of convection within the cloud system. Almost immediately the central convection fell apart as the poleward edge of the cirrus outflow flattened, restricted by the amplification of a mid-latitude trough to the north. Cancellation of the first TCFA on the monsoon depression area followed at 030800 Z (Figure 3-14-1). The arrested development of the monsoon


Figure 3-14-1. A broad band of cloudiness associated with the southwest monsoon extends eastward across the central Thilippine Isfands (030653Z. September $\mathfrak{N}$ (OAA visual imagery).


Figure 3-14-2. The 200 mb analysis at 040000Z September revealed a wind speed maximum across southwestern Japan.


Figure 3-14-3. Due to northerly flow aloft, convection associated with the low-level circulation center was confined to the southern semicircle ( $040642 Z$ September $\mathfrak{N O A A}$ visual imagery).
depression appears to be related to the movement of an upper-level wind maximum across the island of Kyushu, Japan. This resulted in an increase in northerly flow over the northern Philippine Sea (Figure 3-14-2). This increased upper-level wind shear was responsible for delaying Gerald's development beyond the monsoon depression stage.

At 040600 Z , synoptic data obtained from drifting buoy and ship reports indicated the MSLP had dropped to 1003 mb with 25 to 30 kt ( 13 to $15 \mathrm{~m} / \mathrm{sec}$ ) winds near the circulation center. Satellite imagery also showed an exposed low-level circulation was displaced slightly to the north of a single major convective band (Figure 3-14-3). These data
prompted the issuance of a second TCFA at 041000 Z . The first warning on Tropical Depression 14W followed at 041800 Z , supported by a Dvorak intensity estimate of 30 $\mathrm{kt}(15 \mathrm{~m} / \mathrm{sec}$ ) and a drifting buoy report of a 1001 mb that revealed falling surface pressures.

Since Gerald was a shallow low-level circulation in an active monsoonal trough, its movement was erratic and difficult to forecast. During the period 040000 Z to 071800 Z , the primary numerical aid, the One-Way Interactive Tropical Cyclone Model (OTCM), was used by JTWC to forecast movement.

JTWC forecast Typhoon Gerald would slowly recurve to the east of Taiwan, however,


Figure 3-14-4. Typfioon Gerald at maximum intensity. The large eye is approximately 60 nm ( 111 km ) in diameter. Prior to September, the 1987 season was characterized by an unusual number of 'midget' tropical cyclones. Typfioon $\mathcal{F r e d a}(13 W)$, which also has an eye, is located approximately $1000 \mathrm{~nm}(1852 \mathrm{~km})$ east-southeast of Gerald ( 082111 Z September DMSPP enfanced infrated imagery).
northwestward movement up the monsoonal trough began on the 7 th, as did acceleration and intensification. The 081800 Z warning signalled a major change in the expected movement of Typhoon Gerald. The forecast indicated Gerald would pass through the Luzon Strait and make landfall on the southeast coast of mainland China.

Typhoon Gerald, with a large classic eye 60 nm ( 111 km ) in diameter, reached its maximum intensity of $105 \mathrm{kt}(54 \mathrm{~m} / \mathrm{sec})$ at 081800Z (Figure 3-14-4). Later, Gerald skirted the southwest coast of Taiwan (Figure 3-14-5). The mountainous terrain reduced low-level inflow and Gerald began to weaken (Figure 3-14-6). Gerald continued to weaken over the Formosa Straits and made landfall on the China
coast 50 nm ( 93 km ) east-northeast of Amoy, a city about 245 nm ( 454 km ) east-northeast of Hong Kong. The remnants of Gerald dissipated over land and were no longer apparent on either satellite imagery or synoptic data after 110000 Z .

Typhoon Gerald caused extensive damage to Taiwan and China. In Taiwan, five people died and over $\$ 10$ million in damage was caused by heavy rain and flooding. Up to 16 inches ( 41 cm ) of rain was reported in parts of the Zhejiang Province, China (south of Beijing). Flooding inundated more than 1,950 square miles ( 505,440 hectares) of farmland, causing widespread damage to crops valued at $\$ 121$ million. The Chinese death toll from Typhoon Gerald was 122.


Figure 3-14-5. Radar presentation of the concentric rainbands of Typfroon Gerald as seen fron Hualien, Taiwan (W)MO 46699) at 090200Z September (Photograph courtesy of Central Weather Bureau, Taipei, Taiwan).


Figure 3-14-6. The effect of Land interaction on Gerald's cloud pattern. This enfranced infrared image sfows the distinct break in the central cloud mass to the north of the eye, which is related to Lee-side subsidence over western Taiwan. This cloud-minimum area paraffels the ridge line of mountainous central Taiwan (090956Z September DMSP infrared imagery).


## SUPER TYPHOON HOLLY (15W)

In early September the active monsoonal trough spawned a three tropical cyclone outbreak. Super Typhoon Holly was one of the three. Typhoons Freda (13W) and Gerald (14W) were first warned on at 041800Z September, with Holly following 12 -hours later. As these three systems matured, the monsoon trough became displaced well to the north of its "normal" location (see Figures 3-15-1 and 3-152). In fact, by the 11 th of September, Holly, Typhoon Freda (13W) and the remains of Typhoon Gerald (14W) were all north of 15 degrees North Latitude as an anticyclonic circulation developed in low-latitudes just north of the island of Pohnpei in the eastern Caroline Islands. This anomalous low-latitude high pressure suppressed additional cyclogenesis for
the next four days (see Figure 3-15-3). Monsoonal troughing began to reappear on the 171200 Z surface/gradient-level streamline analysis and was firmly re-established a day and a half later.

Holly began as a westward-moving area of persistent, but weakly organized, convection at the eastern end of the monsoon trough 560 nm ( 1037 km ) east-northeast of Kwajalein and was first mentioned on the Significant Tropical Weather Advisory (ABPW PGTW) at 010600 Z . As Holly developed over the next three days, satellite reconnaissance intensity estimates (Dvorak, 1984) of maximum sustained surface winds indicated an increase from 25 kt (13 $\mathrm{m} / \mathrm{sec})$ to $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$. Vertical wind shear


Figure 3-15-1. Gradient-fevel wind climatology for September (Sadfer, et al, 1987).


Figure 3-15-2. Low-level troughing is shown on this sutface/gradient-level streamline analysis from 110000 Z September well north of its climatological position (see Figure 3-15-1) due to the combined influences of Typhoons Freda (13W) and $\mathcal{H o l l y}$, and the remains of Typfoon Gerald (14W).


Figure 3-15-3. Ridging, shown as a band of anticyclones on this surface/gradient-level analysis for $161200 Z$ September developed in latitudes normally expected to show monsoonal troughing. This appears to have been a key element in the suppression of further low-fatitude tropical cyclone genesis through the 19th.


Figure 3-15-4. $\mathcal{H o l l y}$ near the time of its first warning (050721Z September DMSS visual imagery).


FIgure 3-15-5. Super Typhoon Holly at peakintensity. Typfoon Freda (13W) appears to the left of $\mathcal{H}$ olly (091210z September DMSSP visual imagery).
over the system remained low (not more than 10 $\mathrm{kt} \mathrm{( } 5 \mathrm{~m} / \mathrm{sec}$ )), favoring development. Surface/gradient-level streamline analysis at 040000 Z showed moderate low-level crossequatorial flow from the south into the disturbance. This was apparent from the southwesterly gradient-level winds at Truk (WMO 91334) and Pohnpei (WMO 91348) at 040000 Z . Minimum sea-level pressures were 1006 mb in Holly with the mean environmental pressures near 1009 mb . This combination, together with indications that the deepest convection was consolidating about the lowlevel circulation center, supported the issuance of a Tropical Cyclone Formation Alert at 041930Z.

The first warning on Tropical Depression 15W followed at 050600Z. At that time, the maximum sustained surface winds were $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec}$ ), with a forecast increase to $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ the next day. Satellite imagery on the 5th showed favorable upperlevel outflow and a ragged central convective mass about $21 / 2$ degrees in diameter (Figure 3-15-4). Associated convective bands southwest and east of the center implied a large-scale circulation and little competition for energy from Typhoon Freda (13W) to the west. As a consequence, Holly developed from 30 kt ( 15 $\mathrm{m} / \mathrm{sec}$ ) at the time of the first warning to 90 kt ( $46 \mathrm{~m} / \mathrm{sec}$ ) at the time of the ninth warning at 070600Z.


Figure 3-15-6. The subtropical temains of Holfy (162227Z September $\mathcal{D M S P}$ visual imagery).

Holly's track abruptly changed from northwestward to northward at a position approximately 720 nm ( 1333 km ) northeast of Guam. A maximum intensity of 140 kt ( 72 $\mathrm{m} / \mathrm{sec}$ ) was reached at 091200 Z (Figure 3-15-5). Sparse upper-air and synoptic data did not clearly show a specific weakness in the subtropical ridge to the north of Holly. As a result, the early forecast tracks called for westnorthwestward or westward movement. However, the relative movement and displacement of the monsoonal trough and the weakness of the subtropical ridge appear to have caused Holly's northward movement. (By the 10th, Typhoon Freda (13W) was about 950 $\mathrm{nm}(1759 \mathrm{~km})$ to the west-southwest and
drifting slowly west-northwestward. No binary interaction was apparent between Holly and Typhoon Freda (13W).) With no strong midlatitude systems approaching to provide stronger westerly or southwesterly steering flow, Holly (along with Typhoon Freda (13W)) drifted slowly northward in the active monsoon trough and weakened. Holly acquired subtropical characteristics after 140300 Z , and retained $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$ maximum sustained surface winds. Its remnants could still be located on satellite imagery through the 17th (Figure 3-16-6), with the final satellite fix obtained at 170600 Z . No reports of damage or loss of life were attributed to Holly during its lifetime.


## TYPHOON IAN (16W)

Typhoon Ian was the fourth of seven tropical cyclones to occur in the western North Pacific during September. Ian developed into a significant tropical cyclone six days after the second three-storm warning situation of the year involving Typhoons Freda (13W), Gerald (14W) and Super Typhoon Holly (15W) had ended on September 17th. Thirty-six hours after the first warning on Ian, it was joined by Tropical Depression 17W, which brought to seven the number of periods during 1987 that JTWC was warning on at least two systems at the same time. Even though, Tropical Depression 17W was a very short-lived system, Hurricane Peke ( 02 C ), which crossed the dateline (becoming Typhoon Peke (02C)), and Tropical Storm June (18W) soon took its place. This gave rise to the third three-storm warning situation of the year and the second to occur during September.

Forecasts verified extremely well on Typhoon Ian. The forecast track error statistics for all three verification times (i.e., 24-, 48 - and 72 -hours) were significantly less than the fiveyear average (see Chapter V, Tables 5-1A through $5-2 \mathrm{~B}$ ), though the 72 -hour forecast error of $344 \mathrm{~nm}(637 \mathrm{~km}$ ) exceeded the 1987
average. The reason for the poor 72-hour forecast errors was the unexpected slower movement of Ian between 270600 Z and 290000 Z when the system became nearly quasistationary while tracking generally toward the northwest. If this abnormal behavior had not occurred, JTWC's statistics on Ian would have been outstanding.

Ian began as a broad, poorly organized tropical disturbance 330 nm ( 611 km ) to the east-northeast of Guam. Satellite analysts from Detachment 1, 1st Weather Wing (Det 1, 1WW) alerted the Typhoon Duty Officer to the presence of a persistent area of convection showing improved upper-level outflow. This was in the region where the monsoonal trough was attempting to become re-established after being disrupted by the previous three-storm situation. On 21 September at 0600Z, JTWC added the disturbance to its Significant Tropical Weather Advisory (ABPW PGTW) and listed its potential for development as poor due to the relatively high minimum sea-level pressures (MSLPs) evident in the trough at that time. Within 24-hours, the MSLPs decreased by 2 mb and the wind speeds increased another 5 kt ( 3 $\mathrm{m} / \mathrm{sec}$ ) to $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec}$ ) (see Figure 3-16-1).


Figure 3-16-1. Ian, as a tropical disturbance witf $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$ maximum sustained winds at the surface (220007Z September DMSSP visual imagery).


Figure 3-16-2. Curvature is evident in the convective cloud lines just prior to ITWC issuing a Tropical Cyclone Formation Alert at 230130Z (222346Z September DMSS visual imagery).

Ian continued its slow-paced development. Early on the 23rd, satellite imagery (see Figure 3-16-2) showed further intensification had taken place. Curvature became evident in the low-level cloud lines. Satellite intensity analysis (Dvorak, 1984) of imagery at 222346 Z estimated $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ winds at the surface associated with this disturbance. JTWC promptly issued a Tropical Cyclone Formation Alert at 230130 Z for the Mariana Islands north of Guam.

JTWC issued the first warning on Ian (as Tropical Depression 16W) at 230600Z, with an intensity of $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$ and gusts to 35 kt


Figure 3-16-3. Typfioon Ian approximately 12 hours before reaching its maximum intensity of $110 \mathrm{Kt}(57 \mathrm{~m} / \mathrm{sec})$. Note the small circular eye and compact central dense overcast ( 251146 Z September $\mathcal{D M S P}$ enfanced infrared imagery).
( $18 \mathrm{~m} / \mathrm{sec}$ ), based on spiral bands of convection which became visible on visual and infrared satellite imagery. The system was upgraded to Tropical Storm Ian (16W) on the third warning (231800Z) as it progressed slowly westward into an area of low vertical wind shear.

At about that time, Ian began to develop at slightly greater than the normal Dvorak rate of one "T-number" per day. Wind speeds increased from $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ at 231800 Z to $60 \mathrm{kt}(31 \mathrm{~m} / \mathrm{sec})$ at 241800 Z . Midway through this period, lan turned from its westward course and began to move toward the northwest. Five radar fixes were obtained from Andersen Air


Figure 3-16-4. Plot of the slight binary interaction between Ian and June (18W) sfowing their individual tracks and the path of the midpoints.

Force Base on Guam during this same time period. The center positions were based on the convective banding features. No eye feature was apparent on radar for any of the fixes.

Between 241800Z and 250000 Z , Ian reached typhoon intensity as it moved steadily toward the northwest at $7 \mathrm{kt}(13 \mathrm{~km} / \mathrm{hr})$. It intensified at a rate of $10 \mathrm{kt}(5 \mathrm{~m} / \mathrm{sec})$ per six-hour interval (i.e., between warning times) from 241800 Z through 260000 Z (Figure 3-16-3). Note the small circular eye and the compact nature of the deepest convection. Ian reached its maximum intensity of $110 \mathrm{kt}(57 \mathrm{~m} / \mathrm{sec})$ at 260000 Z . It
was during this time of steady intensification that Tropical Depression 17W developed and then dissipated to the east of Ian. No binary interaction was apparent between them. A steady, slow decline followed. Twenty-four hours after Tropical Depression 17W had dissipated Ian slowed dramatically in forward speed as it approached a mid-latitude front lying just to the east of the Ryukyu Islands. Ian inched slowly northward between the times of 270600 Z and 290000 Z at a rate of less than 2 kt ( $4 \mathrm{~km} / \mathrm{hr}$ ). Its deep central convection decreased significantly. The movement of a mid-latitude shortwave north of Ian appeared to


Figure 3-16-5. PLot of the center-relative movement about the midpoint centroid.
have suppressed it. Once this shortwave moved off toward the east on the 29th, Ian's upper-level outflow became aligned with the jet stream (which was above the lower level front) and the system began recurving south of Japan.

Meanwhile Tropical Storm June (18W), which began its development on the 28 th, was moving rapidly northwestward at 18 to 20 kt ( 33 to $37 \mathrm{~km} / \mathrm{hr}$ ). Ian and June ( 18 W ) were close to one another at this stage (within 400 nm ( 741 km )) and eventually underwent a slight binary interaction between 300000 Z September and 020000 Z October. In Figure 3-16-4, the midway point is plotted for the times the two systems coexisted. Figure 3-16-5 shows a plot of the relative movement of each system with respect to the centroid position. As Ian and June (18W) moved northeastward and dissipated, the separation between their tracks decreased.

Ian continued to slowly weaken as this interaction took place, however JTWC forecasters and the Det 1, 1WW satellite analysts misread the changes to Ian on satellite imagery. Dvorak analysis at 010600 Z October estimated Ian's intensity at $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$, which supported a final warning and a downgrade to tropical depression intensity. Post-analysis indicates that Ian most probably transitioned to a subtropical system (rather than extratropical since the subtropical ridge was located to the north of Ian) and still had 55 kt ( $28 \mathrm{~m} / \mathrm{sec}$ ) winds at the time of the final warning.

The remnants of Ian continued to move northeastward after it transitioned to subtropical and finally dissipated $1200 \mathrm{~nm}(2222 \mathrm{~km})$ to the east of Japan on the 4th of October. No damage or deaths were attributed to Ian.


## TROPICAL DEPRESSION 17W



Figure 3-17-1. Tropical Depression 17W slowly developed from a tropical disturbance $250 \mathrm{~nm}(463 \mathrm{~km})$ nortfr of the Marshall Islands during the same period ITWC was waming on Typhoon Ian (16W). It was first detected on satelfite imagery on the 23rd of September and mentioned on the Significant Tropical Weather Advisory (ABPW) PGTW) as a new suspect area at 0600Z. ITWC issued a Tropical Cyclone Formation Mlert nearly 24-fours Later at 240330Z when this system displayed increased convective organization. Maximum sustained surface winds, at that time, were estimated at 15 to 25 kf ( 8 to $13 \mathrm{~m} / \mathrm{sec}$ ). The first warning on Tropical Depression 17 W was issued at $241800 Z$ based on satelfite intensity estimate (Dvorak, 1984) of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec}$ ) winds at the surface. $\mathcal{A}$ welf. established mid-Level ridge was located to the northeast of Tropical Depression 17W. At the same time, an eastwardprogressing, mid-latitude trough to the north was beginning to influence the system. This trough continued to suppress Tropical Depression 17W's development even after it had passed to the northeast of the disturbance. Concurrently, Typhoon Ian's (16W) upper-Level outflow (at the left of the image) restricted Tropical Depression 17W's outflow in the nortfwest quadrant. This combination of factors appears to have stopped further intensification and induced dissipation over water. The last warning was issued on the 26 th of September at 0600Z. The satelfite image above shows Tropical Depression 17W shortly before the second warning was issued, whife it was at its maximum intensity of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ and $960 \mathrm{~nm}(1778 \mathrm{~km})$ east-southeast of Typhioon Ian ( 16 W ). $\mathcal{A}$ partially exposed lowLevel circulation is visible slightly west of the heaviest convection (242305Z September DMSS visual imagery).


## TYPHOON PEKE (02C)

Typhoon Peke (02C) was the first hurricane during the past twenty years (since Typhoon Sara (28) in September 1967) to form in the central North Pacific and cross the dateline. Peke was the only significant tropical cyclone to cross the dateline north of the equator this year. The first twenty-five advisories were issued by the Central Pacific Hurricane Center (CPHC) in Honolulu, Hawaii (the Naval Western Oceanography Center at Pearl Harbor, Hawaii issued the corresponding warnings for the Department of Defense customers). The final twenty-three warnings were issued by JTWC.

Peke began as a broad area of convection about $480 \mathrm{~nm}(889 \mathrm{~km})$ to the southsoutheast of Johnston Island in the west central North Pacific on the 20th of September. The system tracked toward the west and increased in convection and organization over the next 24hours. The upper-level outflow was initially restricted by an upper-level trough to the north of the system. The first advisory on Tropical Depression 02C was issued by CPHC at 211800 Z . Satellite imagery indicated a lowlevel cyclonic circulation with spiral banding. This developed after the tropical cyclone had moved toward the west past the restricting influence of the upper-level trough to the north.

Over the next 18 -hours, the amount of convection continued to increase. Upper-level outflow was unrestricted to the south and was becoming less restricted to the north, prompting the upgrade to Tropical Storm Peke (02C). During this time, Peke changed its track from westward to northwestward in response to a mid-level weakness in the subtropical ridge. It continued to intensify at a normal rate (Dvorak, 1984) and began to track more toward the north. CPHC upgraded the system to Hurricane Peke ( 02 C ) at 231800 Z based on the formation of a banding-type eye, but Dvorak intensity postanalysis indicated that the system most probably did not reach hurricane intensity for another 6to 12 -hours. Peke continued to intensify and reached a first peak, of $85 \mathrm{kt}(44 \mathrm{~m} / \mathrm{sec})$, at 250000 Z .

Peke continued moving northward until 270600 Z . After which time, it tracked toward the west-northwest in response to the strong mid-level flow around the subtropical ridge lying to the north of the system. CPHC issued their last advisory on Hurricane Peke (02C) at 271800Z September. It was approximately 30 $\mathrm{nm}(56 \mathrm{~km})$ to the east of the dateline when JTWC issued its first warning (warning number 26) and redesignated the system as Typhoon Peke (02C) at 280000Z. Peke crossed the dateline at 280600Z. After having maintained a steady $75 \mathrm{kt}(39 \mathrm{~m} / \mathrm{sec})$ intensity for over 36 hours, it began to re-intensify. Peke reached its second peak intensity (of 100 kt ( $51 \mathrm{~m} / \mathrm{sec}$ )) between 290600 Z and 291200 Z , as upper-level outflow to the north improved.

Shortly afterward, JTWC was issuing warnings on three western North Pacific systems. Typhoon Ian (16W) was over 2000 $\mathrm{nm}(3704 \mathrm{~km})$ to the west, and having little effect directly on Peke. About 1000 nm (1852 km ) to the west of Peke, Tropical Storm June (18W) was beginning to organize (Figure 3-02C-1). Peke, together with Typhoon Ian (16W) and Tropical Storm June (18W), modified the environment and forced the subtropical ridge axis even further north to beyond 35 degrees North Latitude.

Earlier, as Peke crossed the dateline, it accelerated over a 48 -hour period from near 7 kt ( $13 \mathrm{~km} / \mathrm{hr}$ ) forward speed to about 16 kt ( 30 $\mathrm{km} / \mathrm{hr}$ ). At that time, Peke began to entrain drier air into its central region. Satellite imagery at 291800 Z indicated that a banding eye was present instead of an eye within a central dense overcast. Peke maintained its intensity and was still at $100 \mathrm{kt}(51 \mathrm{~m} / \mathrm{sec})$ at 301200 Z , when the system began to decelerate. At that time, recurvature was forecast along with rapid weakening due to strong westerly flow aloft. Peke became nearly quasi-stationary at 010600 Z October, prior to recurvature. Within six hours, Peke was recurving toward the north-northeast and had steadily weakened from $100 \mathrm{kt}(51 \mathrm{~m} / \mathrm{sec}$ ) at 301200 Z September to $75 \mathrm{kt}(39 \mathrm{~m} / \mathrm{sec})$ at 011200 Z October. Over


Figure 3-02C-1. Typhoon Peke (02C) with Tropical Storm June (18W) to the west (302243Z September DMSP infrared imagery).
the next day, Peke weakened even more, to 55 $\mathrm{kt}(28 \mathrm{~m} / \mathrm{sec})$, but instead of tracking toward the north-northeast, the low-level drifted toward the southeast, while the upper-level tracked toward the south-southeast in response to weak steering flow and increasing vertical wind shear. The last warning on Tropical Depression 02C was
issued by JTWC at 031200 Z . The remnants of Peke moved erratically over the next three and one half days in response to weak steering flow, first tracking toward the east, then toward the northwest and finally back toward the southeast until it could no longer be identified on satellite imagery.


## TROPICAL STORM JUNE (18W)



Figure 3-18-1. Tropical Storm Iune (18W) was the seventh and final tropical cyclone to occur in the western $\mathcal{N}$ (orth Pacific during the month of September. It followed in the wake of Tropical Depression 17W. June was of significance due to the fact that, as it developed, two other tropical cyclones were also being wamed on by gTWC. Iune struggled to develop for four days and was first noted as a suspect area on the Significant Tropical Weather $\mathcal{A d v i s o r y}(\mathcal{A B P W} \mathcal{P} G T W$ ) on September 27tf at 0600Z. $\mathcal{A} t 280600 \mathrm{Z}$, ITWC issued a Tropical Cyclone Formation Alert Gased on increased convective organization and satelfite intensity estimates (Dvorak, 1984) of $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec}$ ). During the next 18 -hours, the convection associated with 9 une increased significantly north of the partially exposed Low-Level circulation center. This, plus a satelfite intensity analysis of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$, prompted the first warning, at 290000Z. Throughout its existence, June's upper-Cevel outflow was restricted in the west quadrant due to the strength of Typhoon Ian's (16W) outflow. The final warning on Iune was issued on the 1st of October at 0000Z when the increased vertical wind shiear finally stripped away the central cloudiness. The Low-Level remnants continued to track toward the northeast and were detected on satellite imagery until the 2nd of October. The above image shows Iune approximately 12 -fiours after reacfing its peak intensity of $40 \mathrm{~K} \ddagger(21 \mathrm{~m} / \mathrm{sec})(292016 \mathrm{Z}$ September DMSSP visual imagery).


## TYPHOON KELLY (19W)

Typhoon Kelly was the first of two significant tropical cyclones to develop during the month of October. It moved steadily on a northward track, reaching its maximum intensity at the point of recurvature near 28 degrees North Latitude. Kelly made landfall on the Japanese island of Shikoku about 100 nm ( 185 km ) southwest of Osaka with typhoonforce winds, then crossed west central Honshu, the main Japanese island, before moving into the Sea of Japan.

After an outbreak of three tropical cyclones in early September, an aclimatological surface ridge developed in the low latitudes which proved to be unfavorable for tropical cyclones genesis for about six days. A similar occurrence took place during the first week of October with Typhoon Ian (16W), Tropical Storm June (18W) and Typhoon Peke (02C). The strong surface ridge was the primary synoptic feature in an area normally dominated by the monsoon trough. The existence of a low-latitude ridge during the height of the tropical cyclone season appeared to be a readjustment mechanism for the unusual northward displacement of the active monsoon trough to 25 degrees North Latitude associated with both multiple-storm outbreaks.

On 6 October crossequatorial flow returned to the low latitudes from the southern Philippine Islands to the area south of the island of Pohnpei in the eastern Caroline Islands, allowing the monsoonal trough to reestablish itself along 5 degrees North Latitude. Moonlight visual satellite imagery on 7 October indicated a circulation was developing $190 \mathrm{~nm}(352 \mathrm{~km})$ south of the island of Yap (Figure 3-19-1). The 071400 Z Significant Tropical Weather Advisory (ABPW PGTW) mentioned the area and classified its potential for development into a significant tropical cyclone as fair, Figure 3-19-1. Moonfight visual imagery sfowing Typfoon Kelly at an based on an initial satellite intensity early stage of development (071243Z October $\mathcal{D M S S}$ visual imagery).


Figure 3-19-2. $\mathcal{A}$ plot of the surface pressures for Yap and Koror for the time period 050000Z to 100000 Z October (missing values are extrapolated for continuity purposes). Although pressures were dropping at both stations, the lower surface pressures and more rapid faffs at Yap indicated the low pressure center was passing close by.


Figure 3-19-3. The Low-Level circulation center located west of the convective cloud mass is partially obscured by cirrus 6 low-off. $A$ convective band south of Kelly identifies an area of intense convergence which extended as far west as the isfand of Mindanao (100043Z October DMSPP visual imagery).
persistent convection around the exposed lowlevel circulation center. Satellite imagery over the next 24 -hours showed a steady increase in the amount of convection within the cloud system. The first warning on Tropical Depression 19W was issued at 100000 Z , supported by a satellite intensity estimate of 30 $\mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$. At the time of the first warning, the low-level center was still located about 60 nm ( 111 km ) west of the dense convection (see Figure 3-19-3).

A $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ ship report 30 nm ( 56 km ) north-northwest of the circulation center at 101200 Z supported the earlier upgrade to tropical storm and previous satellite estimates that Kelly had attained tropical storm intensity. The low-level center position remained 95 nm $(176 \mathrm{~km})$ southwest of the upper-level circulation center at that time. This separation between the low- and upper-level positions


Figure 3-19-4. Typfioon Kelly approximately twelve fours prior to its reaching its maximum intensity of 95 kt ( $49 \mathrm{~m} / \mathrm{sec}$ ) ( $150042 Z$ October DMMSP visual imagery). resulted in initial position relocations on the 101200 Z and 101800 Z warnings.

Once the low- and upper-level centers became aligned on the 11th, Kelly slowly intensified. Minimal typhoon intensity of 65 kt ( 33 $\mathrm{m} / \mathrm{sec}$ ) was reached at 120000 Z . Kelly's intensity peaked at $95 \mathrm{kt}(49 \mathrm{~m} / \mathrm{sec})$ near the point of
recurvature at 151200 Z (Figure 3-19-4).
By that time, Typhoon Kelly posed a serious threat to Japan. As it began to slow its forward speed and recurve near the 28th parallel, the forecast question was whether the system would continue to track northward across Japan or recurve south of Japan. Synoptic data at 150000 Z indicated the upperlevel westerly winds south of Japan were nearly zonal, which would tend to steer Kelly toward the east-northeast, favoring the south of Honshu scenario. This reasoning prevailed and recurvature south of Japan was forecast. By 160000 Z however, a mid-level long-wave trough, anchored near the Yellow Sea, deepened as an intense short wave came in phase with the trough axis. Consequently, the steering flow ahead of the trough shifted from westerly to south-southwesterly and Kelly continued its course across Japan instead of recurving sharply northeastward.

Typhoon Kelly weakened only slightly as it began to assume extratropical characteristics on the 16 th. Synoptic reports indicated Kelly did not dissipate as rapidly as implied from satellite imagery. Upper-air reports at Shiono, Japan (WMO 47778) revealed that Kelly still packed winds of 95 kt $(49 \mathrm{~m} / \mathrm{sec})$ at the 850 mb level at 161200 Z .

Eventually, Typhoon Kelly made landfall on the southern coast of the island of Shikoku. At least 13,000 homes were flooded and another 30 were badly damaged by mudslides triggered by as much as 20 inches ( 51 cm ) of rain. Wind gusts were reported as high as $120 \mathrm{kt}(62 \mathrm{~m} / \mathrm{sec}$ ) as typhoon-force winds battered southern Japan. At least eight people were killed.

After crossing the islands of Shikoku and Honshu, Kelly moved offshore and became extratropical over the Sea of Japan. Later, Misawa Air Base (WMO 47580), located near the northern tip of Honshu, reported maximum surface winds of $32 \mathrm{kt}(16 \mathrm{~m} / \mathrm{sec}$ ) and a surface pressure of 985 mb as the extratropical low passed between the islands of Honshu and Hokkaido at 171200Z. The residual circulation of Typhoon Kelly was no longer visible on satellite imagery after 180300 Z .


## SUPER TYPHOON LYNN (20W)

Super Typhoon Lynn was the third tropical cyclone of 1987 to produce maximum sustained surface winds of $140 \mathrm{kt}(72 \mathrm{~m} / \mathrm{sec})$ with gusts to $170 \mathrm{kt}(87 \mathrm{~m} / \mathrm{sec})$ and the second to attain an estimated minimum sea-level pressure (MSLP) of 898 mb (only Super Typhoons Betty (09W) and Nina (22W) were lower with a MSLP of 891 mb ). It was also the fifth super typhoon of the year. Lynn, during its latter stages, also had a devastating impact on Taiwan and caused some concern in the Hong Kong area, as well.

Lynn began as a broad, poorly organized area of convection in the monsoon trough about 200 nm ( 370 km ) north-northeast of Kwajalein Atoll in the Marshall Islands. After the convection had persisted for 24 -hours, it was added as a new suspect area to the Significant Tropical Weather Advisory (ABPW

PGTW) at 150600 Z . Maximum sustained surface winds were estimated at 15 to 20 kt ( 8 to $10 \mathrm{~m} / \mathrm{sec}$ ); the MSLP was estimated to be 1008 mb . Over the next 18 -hours, upper-level outflow and the amount of convection increased significantly as the MSLP decreased to 1001 mb . For these reasons, a Tropical Cyclone Formation Alert was issued at 160030 Z , when the system was located about $360 \mathrm{~nm}(667 \mathrm{~km})$ north-northwest of the island of Pohnpei in the eastern Caroline Islands. Six hours later at 160600Z, the first warning on Tropical Storm Lynn (20W) was issued, based on the satellite intensity estimate (Dvorak, 1984) of 35 kt ( 18 $\mathrm{m} / \mathrm{sec}$ ). Until 171800Z, Lynn had been moving toward the west along the southern periphery of the subtropical ridge. Before reaching tropical storm intensity, Lynn had been moving at speeds in excess of $20 \mathrm{kt}(37 \mathrm{~km} / \mathrm{hr})$. But as it began to intensify, it decelerated. By 161200 Z ,


Figure 3-20-1. Tropical Storm Lymn (20W), shortly before being upgraded to typfioon intensity (180528Z October NOAA visual imagery).


Figure 3-20-2. $\mathcal{H A N} \mathcal{D A R}$ observations for the islands of Saipan and Rota. These observations ilfustrate the closest point of approack of Typhoon Lynn to Saipan and Rota was between 181500Z and 181600Z.

Lynn was moving at a speed of only $6 \mathrm{kt} \mathrm{(11}$ $\mathrm{km} / \mathrm{hr}$ ). At 180600 Z , it was upgraded to typhoon status when visual and infrared satellite imagery, plus radar observations from Andersen Air Force Base on Guam, indicated Lynn had formed an eye 20 nm ( 37 km ) in diameter (see Figure 3-20-1). Satellite analysis at that time estimated Lynn's intensity at $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$. (Post-analysis on the system indicated that Lynn was most probably a typhoon at 180000 Z .)

As it intensified, Typhoon Lynn was starting to track toward the west-northwest, away from Guam towards the island of Saipan. Consequently, JTWC amended its 180600 Z warning which had forecast a more westward track. At 181200Z, Typhoon Lynn made its closest point of approach (CPA) to Guam when it tracked 75 nm ( 139 km ) to the northeast of the island. Maximum sustained surface winds recorded on the island were $36 \mathrm{kt}(19 \mathrm{~m} / \mathrm{sec}$ ) with a peak gust of $57 \mathrm{kt}(29 \mathrm{~m} / \mathrm{sec})$ at Agana (WMO 91212). A maximum rainfall accumulation of 6.08 inches ( 154.4 mm ) was recorded at Andersen Air Force Base (WMO 91218).

Lynn's approach had a profound effect on the island of Guam. Apra Harbor on the west side of Guam was closed after four U.S.

Navy ships sortied to open waters. Military airfields evacuated aircraft and secured some aircraft in hangars. All commercial flights to and from Guam were cancelled on 18 October. Most villages on Guam reported flooding in low-lying areas, broken windows, and power and water outages. The power outages were caused mainly by the high winds which knocked vegetation onto power lines, and required several days for Guam Power Authority to repair. Perhaps the most serious damage from Typhoon Lynn was to local agriculture.

At 181500Z, Lynn made its CPA to the island of Tinian - $15 \mathrm{~nm}(28 \mathrm{~km})$ southwest of the island. The automatic weather station observations at Rota, 53 nm ( 98 km ) southsouthwest of Tinian, and at Saipan, $5 \mathrm{~nm}(9 \mathrm{~km})$ to the northeast of Tinian, for 181500 Z and 181600 Z recorded Lynn's passage(see Figure 3-20-2). Maximum sustained surface winds of 45 $\mathrm{kt}(23 \mathrm{~m} / \mathrm{sec}$ ), with a peak gust of 65 kt ( 33 $\mathrm{m} / \mathrm{sec}$ ) were recorded on Saipan. All commercial airline flights to and from Saipan were cancelled. Schools and government offices on Saipan were closed on 19 and 20 October. The islands of Saipan and Rota both experienced island-wide power outages on the evening of 18 October.


Figure 3-20-3. Super Typfroon Lymn at its maximum intensity of $140 \mathrm{Kt}(72 \mathrm{~m} / \mathrm{sec})(192240 Z$ October N(OAA visual imagery).

Once past the Marianas, Lynn intensified rapidly from $80 \mathrm{kt}(41 \mathrm{~m} / \mathrm{sec})$ to its peak intensity of $140 \mathrm{kt}(72 \mathrm{~m} / \mathrm{sec}$ ), reaching super typhoon intensity ( 130 kt or $67 \mathrm{~m} / \mathrm{sec}$ ) shortly after 191800Z. Super Typhoon Lynn maintained its 140 kt ( $72 \mathrm{~m} / \mathrm{sec}$ ) intensity (Figure 3-20-3) until 210000Z.

As Lynn began weakening after 210000Z, its track became westerly. Prior to that time, the forecast track had been westnorthwesterly to northwesterly. Numerical guidance provided by the One-Way Interactive Tropical Cyclone Model (OTCM) appeared to be accurate. The 210600 Z warning echoed this guidance, however Lynn persisted on its westward track. A closer look at the lowertropospheric and deep-layer mean flow fields
north of the typhoon provided a clue as to why Lynn was not behaving as expected. Because of Lynn's synoptic size cyclonic circulation, the integrated effect on the low- and mid-level steering flow was to eliminate the narrow subtropical ridge. Perhaps, OTCM interpreted the large-scale storm-induced circulation as being the synoptic steering flow and therefore, did not detect the narrow subtropical ridge. In contrast, Lynn was large enough to be resolved by the Navy Operational Global Atmospheric Prediction System (NOGAPS) and European Center for Medium-Range Weather Forecasting (ECMWF) numerical models, which in turn provided more accurate forecast guidance. The next warning (number 22 at 211200Z) put the forecast back "on track" and OTCM became the less-favored alternate scenario.

A close encounter by a merchant vessel on 22 October provided testimony to the fury of the typhoon. Excerpts from the ship's log include:
"Sea and swell were of height and steepness that we couldn't turn around anymore ... Seas are approximately 2 $1 / 2$ times the bridge height and breaking all around us. At 1000 we recorded the lowest barometric pressure of 969 HPA", (approximately $75 \mathrm{~nm}(139 \mathrm{~km})$ from the center of Typhoon Lynn, at that time). "During passage of "Lynn" visibility was reduced to 000.0 mtr . Wind above comprehension ... our ears on the bridge were popping due to pressure change with pitching of vessel."

At 240000Z, Lynn was tracking through the Luzon Strait, moving toward the northwest. From 24 through 26 October, it devastated portions of Taiwan (Figure 3-20-4). The island received high winds because of the strong pressure gradient between Lynn's low central pressure and the large high pressure area over mainland China. These high winds, caused rapid orographic lifting along the steep mountains of Taiwan, producing torrential precipitation. Although the center of Lynn passed $110 \mathrm{~nm}(204 \mathrm{~km})$ to the southwest of Taiwan, it produced heavy weather over the northernmost parts of the island. News services reported 68 inches ( 173 cm ) of rainfall on the capital city of Taipei from 24 to 26 October! In Taipei, torrential rainshowers caused landslides that smashed houses and killed 14 people. Over 2,200 people were stranded by floodwaters from


Figure 3-20-4. Tropical Storm Lynn during its rapid weakening phase (250013Z October $\mathcal{T O R A}$ visual imagery).
the Keelung River in Taipei making travel impossible. The Central Weather Bureau in Taipei reported $84 \mathrm{kt}(43 \mathrm{~m} / \mathrm{sec}$ ) maximum sustained surface winds on 24 October and 61 kt ( $31 \mathrm{~m} / \mathrm{sec}$ ) maximum sustained surface winds on 25 October. The port city of Keelung reported over five million dollars worth of damage from Lynn. Lynn created 20 ft ( 6.1 meter) high waves at Hengchun on the extreme southern tip of Taiwan and nine children were swept away. The result of Lynn's passage was Taiwan's worst flooding in 40 years; 42 people perished and 18 were reported missing.

Visual and infrared satellite imagery at 260300 Z , indicated that Lynn was being sheared apart. Subsequent satellite imagery showed the low-level circulation center moving
toward the west-southwest away from the convective mass. At 270000 Z , the final warning was issued on Tropical Depression 20W (Lynn). Although the tropical cyclone had lost its central convection, the remaining lowlevel circulation center still had an impact on the Hong Kong area. The wind speeds and precipitation amounts the Hong Kong area received were higher than expected. The strong pressure gradient between the residual low offshore and the high over mainland China fostered gales along the south coast. The Hong Kong (WMO 45005) 280000Z upper-air sounding revealed maximum winds of 55 kt ( 28 $\mathrm{m} / \mathrm{sec}$ ) from the east-southeast, at 900 and 850 mb . The low tracked south of Hong Kong, into the Pearl River Estuary and eventually dissipated over land.


## TROPICAL STORM MAURY (21W)



Figure 3-21-1. Tropical Storm Maury was a relatively weak, but persistent, tropical cyclone that tracked westward across the Philippine and South China Seas. It was the first of three significant tropical cyclones to form in November and the second system of the year to regenerate over water. Tropical Storm Maury formed in early November in the westem Nortf Pacific near-equatorial trough about $300 \mathrm{~nm}(556 \mathrm{~km})$ to the southeast of Guam. It was first detected as an area of deep convection with a well-defined, Low-Level circulation center but a poor upperLevel outflow and was first mentioned on the Significant Tropical Weather $\mathcal{A d}$ visory ( $\mathcal{A B P W}$ PGTW) at 070600Z. Due to unfavorable vertical shear, the system appeared to fiave a poor cfiance for further development and showed little sign of intensification over the next three days. However, the central convection did increase. At 110130Z, a Tropical Cyclone Formation $\mathfrak{A l e r t}$ (TCFA) was issued. The first warning on Tropical Depression 21W followed at 110600Z, based upon the presence of a central dense overcast and satellite internsity estimates (Dvorak, 1984) of 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ). Tropical Depression 21W maintained its organization and convection over the next 6 - to 9 -fours, 6ut began to slow and weaken, as an eastward-moving, Low- to mid-Level trough passed to the north. At 120000Z, IT'WC issued a final warning, but continued to monitor the remnants for possible regeneration. With the trough in the polar westerlies displaced to the east, the convection in the remnants of Tropical Depression 21 W flared-up, prompting the issuance of a second TCFA at 130330Z. Afmost immediately, an abbreviated warning (the first of the year) was issued (at 130530Z) as the cloud system gained in organization. Over the next 12 -hours, the regenerated Tropical Depression 21W maintained its organization, but remained below tropical storm intensity (see above image). By 131800Z, as it approached the Philippine IsLands, Tropical Depression 21W showed signs of becoming less organized, as the central convection diminished. Prior to Tropical Depression 21W crossing the Philippine Islands, yTWC altered its forecast philosophy from 'dissipating over Land', to 'regeneration' in the South China Sea. As Tropical Depression 21W entered the Soutf China Sea, the deep convection increased significantly. Dvorak satellite intensity analysis at $161200 Z$ estimated maximum surface winds of $35 \mathrm{Kt}(18 \mathrm{~m} / \mathrm{sec})$, prompting $\operatorname{ITWC}$ to upgrade $T_{\text {ropical }}$ Depression 21W to Tropical Storm Maury (21W). Maury tracked westward across the South China Sea and reached the maximum intensity of $45 \mathrm{k} \neq(23 \mathrm{~m} / \mathrm{sec})$ at 170600 Z . Later, it made landfalf on the southeast coast of Vietnam 25 nm ( 46 km ) north of Cam Rank Bay at 190400Z. The final waming was issued at 190600 Z as Maury dissipated over Land. N(o reports of severe damage or loss of life were received ( $131009 Z$ N (ovember DOMSP infrared imagery).


Super Typhoon Nina was the most intense and most destructive tropical cyclone to develop in the western North Pacific in 1987. During its track toward the west, it devastated the Truk Atoll in the eastern Caroline Islands, decimated the north central Philippine Islands and then executed a final dramatic loop in the South China Sea south of Hong Kong. Nina was the second of three significant tropical cyclones to develop during November.

Nina developed in low latitudes just west of the dateline. At 150000 Z , satellite intensity analysis (Dvorak, 1984) estimated a cloud system center had maximum sustained surface winds of $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec}$ ). For two days this disturbance showed marked diurnal fluctuations in convection. It was first mentioned on the 170600 Z Significant Tropical Weather Advisory (ABPW PGTW) as a system with fair potential to develop into a significant tropical cyclone. The system displayed good upper-level outflow and increasing convection over a broad area.

A Tropical Cyclone Formation Alert was issued at 190100 Z as deep convection consolidated in the center of the tropical disturbance. Synoptically, the system appeared to be well-established in the low levels up to 400 mb (Figure 3-22-1) with 25 to 30 kt ( 13 to $15 \mathrm{~m} / \mathrm{sec}$ ) easterlies aloft. With speed and directional divergence aloft, Nina continued its rapid organization. At 191200 Z , the first warning was issued on Tropical Depression 22W. By that time, Nina had formed a curved band of convection. Satellite imagery (Figure 3-22-2) suggested unrestricted upper-level outflow over the system; however, the upperlevel rawinsonde reports showed that the anticyclonic circulation (at 200 mb ) was displaced to the east of the center of cirrus outflow (Figure 3-22-3).

At the time of the first warning, working plots of satellite fix positions indicated Nina was slowing down its west-northwestward movement. (To the contrary, post-analysis revealed that Nina did not slow down while intensifying but actually accelerated slightly.


Figure 3-22-1. Heights of the standard pressure levels and surface pressure at Pofinpei (WMO 91348), Truk (WMO 91334) and Guam (WMO 91217) at 190000Z. At this time, Nina was 230 $\mathrm{nm}(426 \mathrm{~km})$ southeast of the island of Pohnpei. In comparison with Truk and Guam, the Lower heights at Pohnpei (W)MO 91348) at the 400 mb Cevel and below are due to the approaching tropical cyclone.


Figure 3-22-2. Satelfite imagery indicating improving organization in the central convection and good upper-Cevel outflow associated with the tropical disturbance that was to become Nina (191457Z November DMSSP visual imagery).

This acceleration along the working best track would greatly affect the forecast movement. For example, if a cyclone is moving at 2 kt ( 4 $\mathrm{km} / \mathrm{hr}$ ) faster than forecast, it will travel in 72hours an additional $144 \mathrm{~nm}(267 \mathrm{~km}$ ).)

Nina continued to intensify and accelerate. By the time of the second warning at $191800 Z$, Nina was upgraded to tropical storm intensity. More rapid westward movement was supported by upper-level data at Guam (WMO 91217), Truk (WMO 91334) and Pohnpei (WMO 91348), which indicated 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) easterly mid-level flow during this time (Figure 3-22-4). At 201600Z, Nina passed $40 \mathrm{~nm}(74 \mathrm{~km})$ south of Moen Island in the Truk Atoll while moving west-northwestward at 18 kt ( $33 \mathrm{~km} / \mathrm{hr}$ ). Satellite intensity analysis estimated winds between 45 and 50 kt ( 23 to 26 $\mathrm{m} / \mathrm{sec}$ ). Maximum winds reported at Moen Island were $60 \mathrm{kt}(31 \mathrm{~m} / \mathrm{sec})$ with gusts up to 80 $\mathrm{kt}(41 \mathrm{~m} / \mathrm{sec}$ ). (Note: The difference in intensity may be due to the fact that winds in
the right front quadrant of a tropical cyclone are a combination of its kinetic energy and the vector addition of its forward movement.) The lowest pressure recorded was 987 mb , which correlates (Atkinson and Holliday, (1977)) to 50 $\mathrm{kt}(26 \mathrm{~m} / \mathrm{sec})$.

Nina passed the Truk Atoll during the early morning hours on the 21st of November. Civil Action Teams reported that five people were killed, 38 seriously injured, and most of the more than 40,000 residents were homeless and without electrical power. The Truk Atoll was declared a federal disaster area in order to compensate for the $\$ 30$ to $\$ 40$ million in damage to housing, businesses and agriculture. In addition, U. S. Armed Forces airlifted supplies into the ravaged islands.

After Nina passed the Truk Atoll, it slowly decelerated. The rate of intensification also slowed. Nevertheless, Nina was upgraded to typhoon intensity at 211200 Z . Nina passed


Figure 3-22-3. Symoptic data at 1912002, showing the center of the upper-Cevel anticyclonic circulation at 200 mb Cocated considerably east of the center of the cirrus outflow associated witf Trovical Devression 22W.




Figure 3-22-4. Upper-air winds and heigfits at standard pressure fevels for Guam (WMO 91217), Truk(WMO 91334) and Pofinpei (WMO 91348) for the period 180000Z to 200000Z. Notice the 30 $K_{t}(15 \mathrm{~m} / \mathrm{sec})$ mid-level easterfies which would support the rapid movement of $\mathcal{N}$ (ina toward the west.


Figure 3-22.5. Satellite imagery showing the well-defined eye of Super Typfioon Nina as it approached the Philippine Islands (250701Z November NO凡Я visual imagery).


Figure 3-22.6. Matching infrared image for visual in Figure 3-22-5 (250701Z November $\mathcal{N}$ (OAA infrared imagery).

| Prognostic Reasoning |  |  |  |
| :---: | :---: | :---: | :---: |
| DTG | Direction | Guidance | Speed |
| 2600002 | WNW for $36-\mathrm{hrs}$ then NW | OTCM \& COSMOS | Decelerating |
| 2612007 | WNW 12- to 18hrs then NW | Break in 400 mb ridges | ----- |
| Alternate scenario: | $\begin{aligned} & \text { Recurve in } \\ & 48 \text { - to } 72 \text {-hrs } \end{aligned}$ | OTCM | ----- |
| 2nd altern scenario: | ate <br> Move west | CSUM | ----- |
| 270000 z | WNW for 48-hrs | NE surge interaction, ECMWF \& NORAPS <br> \& HPAC | Slowly decelerating |
| Alternate scenario: | West | --------- | ------ |
| 271200 Z | NW then west | NE surge, NORAPS \& NOGAPS \& ECMWE | ------ |
| 2800002 | North for 27-hrs then West | NE surge, $700 \& 400 \mathrm{mb}$ progs \& HPAC |  |
| 2812007 | ENE | Strong mid- to upper-level westerly flow \& CSUM | Accelerating |
| 2900002 | ENE | Strong mid- to upperlevel southwesterly flow | Accelerating |

Figure 3-22-7. A66reviated Prognostic Reasoning for the 260000Z through 290000Z November time period.

60 nm ( 111 km ) north of the island of Ulithi and $95 \mathrm{~nm}(176 \mathrm{~km})$ north of Yap at 221000 Z and 221600Z, respectively. Later, on November 25th, an overflying Navy aircraft observed moderate flood damage to the Ulithi's agricultural areas. Twenty percent of the buildings had received structural damage. No damage was reported on Yap.

Nina began to slowly accelerate and rapidly intensify (Holliday and Thompson,
1979), dropping approximately 4 mb per six hours, as it approached the Philippine Islands. Beginning at 241200 Z , Nina began to explosively deepen (Holliday and Thompson), dropping approximately 8 mb per six hours. Nina displayed a symmetrical eye that was 18 $\mathrm{nm}(33 \mathrm{~km})$ in diameter (Figures 3-22-5 and 3-22-6). Nina slammed into the southern tip of Luzon at 251500 Z with maximum winds estimated at $145 \mathrm{kt}(75 \mathrm{~m} / \mathrm{sec})$ with gusts to 175


Figure 3-22-8. Difference between the working and final best tracks as $\mathfrak{N i}$ (ina was sheared apart by the strong surge of the $\mathfrak{N}$ (ortheast Monsoon.
kt ( $90 \mathrm{~m} / \mathrm{sec}$ ). At least 687 people perished in the north central Philippine Islands. As with the Truk Atoll, Nina struck at night. Philippine authorities declared a state of emergency for 18 provinces that were battered by Nina. Overall more than 500,000 people were either rendered homeless, evacuated, or lost their sources of income. Croplands were heavily damaged. News sources reported that Nina was the most destructive typhoon to hit the Philippine Islands in nearly 20 years.

Nina traversed between the islands of Luzon and Mindoro and entered the South China Sea with 95 kt ( $49 \mathrm{~m} / \mathrm{sec}$ ) winds. Although satellite imagery could not detect an eye, land-based radar continued to track the cloud covered eye. Shortly thereafter, Nina was packing $100 \mathrm{kt}(51 \mathrm{~m} / \mathrm{sec})$ winds.

Once Nina was in the South China Sea, the forecast philosophy attempted to keep up with the changing synoptic situation. Figure 3-22-7 provides an abbreviated look at the specifics of each prognostic reasoning message for the 260000 Z through 290000 Z November time period. Basically what initially appeared to be straight-forward, wasn't! The decoupling of Nina's lower- and upper-level circulations developed into a complex event; culminating in the 270000 Z prognostic reasoning message,
which became a classic example of being wrong for all the right reasons. The net result was a very tense situation for Hong Kong and the southern China coast.

As the system began to move northward, an eye became visible at 280300 Z . Within 6- to 12 -hours, Nina was sheared apart by the shallow, but strong, low-level surge in the northeast monsoon flow and strong westerly winds at the mid- and upper-levels. During the shearing, the deep convection, which was poorly defined and being positioned as an upper-level circulation by satellite, accelerated east-northeastward along the quasi-stationary front. As a consequence, the forecast philosophy embraced a cloud system moving rapidly through the Luzon Strait and becoming extratropical. (Post-analysis found that the lowlevel most probably separated from the upperlevel circulation center at 280600 Z . This resulted in a $340 \mathrm{~nm}(630 \mathrm{~km}$ ) difference at $24-$ hours between the working best track and the final best track points as seen in Figure 3-22-8.) The upper-level cloudiness did move eastnortheastward; however, the low-level circulation center executed an anticyclonic loop and headed slowly southward with the monsoonal flow. The residual low-level vorticity and cloudiness rapidly dissipated over the South China Sea.


## TROPICAL STORM OGDEN (23W)



Figure 3-23-1. The third significant tropical cyclone of November, Tropical Storm Ogden, developed into a tropical storm in the South China Sea and quickly made landfall over southern Vietnam. Ogden was first detected late on $\mathcal{N}$ (ovember 20th as a poorly organized area of convection just east of the Philippines. Once in the South China Sea, the development of spiralling low-level cloud lines led to the system's first mention on the Significant Tropical Weather $\mathfrak{A d v i s o r y}$ (ABPW PGTW) at 230600Z. At 240400Z, a Tropical Cyclone Formation Alert (TCFA) was issued based on the improved low-Level cloud organization and symoptic reports of a closed surface circulation with maximum winds of 15 to $25 \mathrm{kt}(8$ to $13 \mathrm{~m} / \mathrm{sec}$ ). Shortly thereafter, Dvorak intensity analysis of satelfite imagery estimated 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) winds which prompted the first warning on Tropical Depression $23 \mathcal{W}$ at 240600 Z (see image above). At 241800Z, Ogden reached a maximum intensity of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$ just prior to making landfalf. Ogden made landfalf on the east coast of Vietnam $18 \mathrm{~nm}(33 \mathrm{~km})$ south of Tuy Hoa at 241900Z. The final warning was issued at 250000 Z as the system moved infand and dissipated (240712Z November $\mathfrak{N O M A}$ visual imagery).


## TROPICAL STORM PHYLLIS (24W)

Typhoon Phyllis was the only significant tropical cyclone to develop in the western North Pacific in December and the third to regenerate over water in 1987 (reference Tropical Storms Ed (12W) and Maury (21W)). It struck the central Philippine Islands three weeks after Super Typhoon Nina (22W) and added further misery to that ravaged nation.

Phyllis began as an area of weakly organized convection in the eastern Caroline Islands $150 \mathrm{~nm}(278 \mathrm{~km})$ southeast of the Truk Atoll. It was mentioned for the first time on the 091030Z December Significant Tropical Weather Advisory (ABPW PGTW) after
exhibiting a rapid increase in the amount and organization of convection. The potential development was listed as fair due to the preexistence of a low-level circulation and unrestricted upper-level outflow.

A Tropical Cyclone Formation Alert (TCFA) was issued the next day at 100230 Z when a satellite intensity estimate (Dvorak, 1984) indicated $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$ winds at the surface. The first warning, on Tropical Depression 24W, came at 101800 Z as the estimate of the surface winds increased to 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) and the associated deep convection became more centralized. At that time, Tropical


Figure 3-24-1. The well-defined, Low-Level circulation center of Tropical $\mathcal{D e p r e s s i o n ~} 24 \mathcal{W}$ is reveafed by the spiral convective bands of cloudiness ( 1105472 December $\mathfrak{N}$ (OAA visual imagery).

Depression 24W was located 370 nm ( 685 km ) south-southeast of Guam and was moving toward the northwest. Twenty-four hours later, it made its closest point of approach to Guam ( 210 nm ( 389 km ) to the southwest) and was upgraded to a tropical storm based on the development of a large cloud system and improved upper-level outflow in the southwest quadrant (see Figure 3-24-1). Early dissipation was forecast (beginning with the third warning at 110600 Z ). The approach of an eastwardmoving, mid-latitude trough would increase the vertical wind shear. As the short wave moved eastward from mainland China, Phyllis slowed
its forward motion until 130600Z, then abruptly changed course and accelerated toward the west-southwest. After downgrading the tropical cyclone to a tropical depression at 130000 Z , the final warning followed at 140000 Z . The displacement of central convection to the northeast of the low-level circulation center and the entrainment of cooler, drier air appeared to have started an irreversible weakening process.

However, within 18 -hours (once the vertical wind shear decreased), Phyllis began to reestablish its central convection under a favorable upper-level outflow pattern. This


Figure 3-24-2. Tropical Storm Phylfis sfortly after regeneration (142136Z December DMSSP infrared imagery).


Figure 3-24-3. Tightly wrapped, small eye of Typhoon Phyllis during landfall on the isfand of Samar ( $152123 Z$ December DMSSP infrared imagery).
resulted in the issuance of a second TCFA at 141630 Z . The (first regenerated) warning followed at 141800 Z (as warning number 15 on the system) based on the satellite intensity estimate of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec}$ ) (see Figure 3-24-2). Intensification continued until 150000 Z when Phyllis peaked at $100 \mathrm{kt}(51 \mathrm{~m} / \mathrm{sec})$ while making landfall on the island of Samar in the central Philippine Islands (Figure 3-24-3). Phyllis left ten people dead and thirteen more were listed as missing when a ferry boat sank off of northern Samar.

After peaking, Phyllis weakened slowly for 24-hours while traversing the central Philippine Islands. Weakened by the frictional effects of the surrounding mountainous island terrain, Phyllis entered the South China Sea and was downgraded to a tropical depression at 180000 Z . The forecast to dissipate within 48hours over water was basically correct, however, the tropical cyclone did briefly reintensify to $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ on the 19 th . No other reports of deaths or serious damage were received.

## 3. NORTH INDIAN OCEAN TROPICAL CYCLONES

Eight significant tropical cyclones developed in the North Indian Ocean during 1987. That set a new all-time record and surpassed the previous high of seven systems in 1979. This was in sharp contrast with 1986 when only three significant tropical cyclones were observed. The long-term mean is approximately four per year. These eight systems (all of tropical storm intensity) developed during the Spring and Fall transition periods (i. e., the intervals of weak opposing wind flow between the Northeast and Southwest Monsoons). Tables 3-5 and 3-6 provide a summary of information for 1987 and comparison with earlier years.

| TABLE 3.5 NORTH INDIAN OCEAN |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 1987 SIGNIFICANT TROPICAL CYCLONES |  |  |  |  |  |
| 2RCPICAL CYCLOME | PERID Of of harking | CALENOAR DAYS OF HARNIMG | MuHER O marnings ISSUED |  | $\begin{aligned} & \text { ESTIMATED } \\ & \text { HSLP - KB } \end{aligned}$ |
| TC 01B | 01 FER-03 FEB | 3 | 11 | 55 (28) | 384 |
| 760 ${ }^{283}$ | $02 \mathrm{JT-} 05 \mathrm{JW}$ | 4 | 12 | 55 (28) | 983 |
| ${ }^{\text {Tc }}$ c 038 |  | 5 | 18 | 50 (26) | 987 |
|  | 15 © ${ }^{\text {c }}$ - 16 OT | ? |  | 45 (23) | 991 |
| TC OSB | 31 CT - 03 MV | 1 | 14 | 40 (21) | 994 |
| Tc 068 | 11 MOV - 13 MOV | 3 | 6 | 50 (26) | 987 |
| Tr 07a | O8 DEC - 111 DEC | 4 | 14 | 45 (23) | 991 |
| $\pi{ }^{088}$ | $\underline{18 D E C-190 E C}$ | 2 | s | 35 (18) | 997 |
|  | 1987 TOIALS: | $26^{*}$ | 83 |  |  |

*Overlapping days are counted only once in sum.

| TABLE 3-6. |  |  | FREQUENCY OF NORTH INDIAN OCEAN TROPICAL CYCLONES |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| YEAR | JAN | FEB | MAR | APR | MAY | JUN | JUL | AUG | SEP | OCT | nov | DEC | TOTAL |
| 1971* | - | - | - | - | - | 0 | 0 | 0 | 0 | 1 | 1 | 0 | 2 |
| 1972* | 0 | 0 | 0 | 1 | 0 | 0 | 0 | 0 | 2 | 0 | 1 | 0 | 4 |
| 1973* | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 1 | 2 | 1 | 4 |
| 1974* | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 1 | 0 | 1 |
| 1975 | 1 | 0 | 0 | 0 | 2 | 0 | 0 | 0 | 0 | 1 | 2 | 0 | 6 |
| 1976 | 0 | 0 | 0 | 1 | 0 | 1 | 0 | 0 | 1 | 1 | 0 | 1 | 5 |
| 1977 | 0 | 0 | 0 | 0 | 1 | 1 | 0 | 0 | 0 | 1 | 2 | 0 | 5 |
| 1978 | 0 | 0 | 0 | 0 | 1 | 0 | 0 | 0 | 0 | 1 | 2 | 0 | 4 |
| 1979 | 0 | 0 | 0 | 0 | 1 | 1 | 0 | 0 | 2 | 1 | 2 | 0 | 7 |
| 1980 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 1 | 1 | 2 |
| 1981 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 1 | 1 | 1 | 3 |
| 1982 | 0 | 0 | 0 | 0 | 1 | 1 | 0 | 0 | 0 | 2 | 1 | 0 | 5 |
| 1983 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 1 | 0 | 1 | 1 | 0 | 3 |
| 1984 | 0 | 0 | 0 | 0 | 1 | 0 | 0 | 0 | 0 | 1 | 2 | 0 | 4 |
| 1985 | 0 | 0 | 0 | 0 | 2 | 0 | 0 | 0 | 0 | 2 | 1 | 1 | 6 |
| 1986 | 1 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 2 | 0 | 3 |
| 1987 | 0 | 1 | 0 | 0 | 0 | 2 | 0 | 0 | 0 | 1 | 2 | 2 | 8 |
| (1975-1987) |  |  |  |  |  |  |  |  |  |  |  |  |  |
| AVERAGE | 0.2 | 0.1 | 0.0 | 0.1 | 0.7 | 0.5 | 0.0 | 0.1 | 0.2 | 1.0 | 1.5 | 0.5 | 4.7 |
| CASES | 2 | 1 | 0 | 1 | 9 | 6 | 0 | 1 | 3 | 13 | 19 | 6 | 61 |

[^1]


## TROPICAL CYCLONE 01B



Figure 3.01B-1. Tropical Cyclone $01 B$ was the first significant tropical cyclone to form in the Bay of Bengal during 1987. It was detected as an amorphous area of convection about $500 \mathrm{~nm}(926 \mathrm{~km})$ east of Sri Lanka on the 29th of January and was noted on the 301800Z Significant Tropical Weather Advisory (ABIO PGTW). Satelfite imagery at that time showed upper-Level anticyclonically curved outflow over a weak, low-level circulation. Witfin the next 24 . fours, the organization and amount of convection steadily increased. A Tropical Cyclone Formation Alert was issued at 311900Z. Satelfite imagery showed convective banding had continued to increase, Gut sparse synoptic data showed no Low surface pressures. At 0000Z on February 1st, the first warning was issued with the appearance of a central dense overcast and unrestricted outflow in all quadrants. The system then tracked steadily northeastward. The intensity peaked at $55 \mathrm{kt}(28 \mathrm{~m} / \mathrm{sec})$ at 020600 Z as the system began interaction with upper-Cevel southwesterfies, which sent a long plume of cirrus northeastward across Burma. A partially exposed Cow-Level circulation center became apparent at 030000Z, as increased vertical wind shear from the southwesterfies afoft stripped away the central cloudiness. Six fiours later the low-level vortex was fully exposed (see above imagery). At 0312002, ITWC issued the final warning on the $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ weakening tropical cyclone. The remnants of Tropical Cyclone $01 \mathcal{B}$ continued to track toward the northeast and dissipation occurred after it made landfalf on $\mathcal{F e b r u a n y}$ 4th over the northwest coast of Burma (030805Z February $\mathcal{N O} \mathcal{A R}$ visual imagery).


## TROPICAL CYCLONE 02B



Figure 3-02B-1. Tropical Cyclone 02B was the second significant tropical cyclone to form in the Bay of Bengal. It was detected on satelfite imagery as an area of organizing convection about $220 \mathrm{~mm}(407 \mathrm{~km})$ soutfwest of Rangoon, Burma and was first mentioned as a new suspect area on the 3018002 May Significant Tropical Weather $\mathcal{A d}$ visory ( $\mathcal{A B I O}$ PGT'W). The development of strong central convection prompted a Tropical Cyclone Formation ALert on Iune 1st at 0600Z. The first tropical cyclone warning followed a day later at $020600 Z$ as a result of continued development. The forecast track toward the nortfiwest, which agreed closely with the Half Persistence and Climatology ( $\mathcal{H}$ (PAC) guidance, changed during the subsequent 24-hours, as mid-Level ridging caused Tropical Cyclone $02 \mathcal{B}$ to assume a recurvature track toward the northeast (see above imagery). The One-Way Interactive Tropical Cyclone Model (OTCSM), correctly predicted this recurvature toward the northeast; however, the guidance was discounted due to the previous poor performance of the model in this region. $\mathcal{A t} 040600 \mathrm{Z}$, Tropical Cyclone 02B reached its maximum intensity of $55 \mathrm{kt}(28 \mathrm{~m} / \mathrm{sec})$ and developed a magged eye. This intensity was maintained until the system made Landfall over Bangladesh at 041200Z. Rapid dissipation folfowed. No reports of damage or loss of Life were received (030421Z June DMSS visual imagery).'


## TROPICAL CYCLONE 03A



Figure 3-03A-1. Tropical Cyclone 03A began on Iune 4th as a monsoon depression with the supporting convection displaced from the Low-Level circulation center. Throughout the life of Tropical Cyclone 03Я, brisk southwesterly flow dominated the Low-levels witf an overlying tropical easterly jet in the upper-Levels. The Low pressure center developed $250 \mathrm{~nm}(463 \mathrm{Km})$ southeast of central Oman and moved slowly eastward along the edge of the low-level southwesterlies. $\mathcal{A n}$ expanded nadius of over $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ winds in the south semicircle resulted from interaction between the tropical cyclone and the already brisk monsoonal flow. YTWC issued its first warning at 050600Z without a preceding Tropical Cyclone Formation Alert, due to the rapid consolidation of convection around an exposed Low-level circulation center. Post-analysis indicated the system frad actually reached tropical storm intensity 18-fiours earfier. Tropical Cyclone $03 \mathcal{A}$ reached a maximum intensity of $50 \mathrm{Kt}(26 \mathrm{~m} / \mathrm{sec})$ at 060000 Z shortly before it abruptly changed its track toward the nortf. This intensity was maintained until 080000Z when the track became westerly. The final warning was issued at 091200Z as the system rapidfy weakened due to increased vertical wind shear. Cloudiness associated with the remains of the Cow-Levef circulation flared up for a short time between 110400 Z (see above imagery) and 111200Z; fowever, complete dissipation over water occurred within the next 24-fiours ( 1105012 gune $\mathcal{D M S P}$ visuaf imagery).


## TROPICAL CYCLONE 04B



Figure 3-04B-1. Tropical Cyclone 04B began as a monsoon depression in the Andaman Sea on the 12th of October and tracked toward the west-northwest. By 13 October at 0600Z, the cloud system fiad separated from the general monsoonal cloudiness. At that time, ST'WC added the system to its Significant Tropical Weather Pdvisory (ABIO PGTW), and indicated its potential for continued development was fair. At 142030Z, gTWC issued a Tropical Cyclone Formation Alert based on the appearance of a central dense overcast and an associated higher Dvorak intensity estimate of $30 \mathrm{Kt}(15 \mathrm{~m} / \mathrm{sec})$. The first warning followed at 150000 Z as a result of a Dvorak estimate of 35 Kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) surface winds (see above imagery). This first waming, which indicated lardfall witfin the next 12 - to 24-hours, was also desigrated as the final warning. However, this forecast proved to be overly optimistic. At 151800Z, Tropical Cyclone 04B peaked at $45 \mathrm{Kt}(23 \mathrm{~m} / \mathrm{sec})$ and a second warning. was issued. This was necessary Gecause Tropical Cyclone 04B was on track, Gut stifl over water. At 160000Z, Tropical Cyclone 04B was finalfed for a second time as it moved infand and weakened (150408Z October DMSP visuaf imagery).


## TROPICAL CYCLONE 05B



Figure 3-05B-1. Tropicaf Cyclone 05B spawned from the monsoon trough over the southern Bay of Bengal midway betzueen Sri Lanka and nortfiern Sumatra in late October. Its detection on satellite imagery resufted in the reissuance of the Significant Tropical Weatfier Rdvisory (ABIO PGITW) at 300300 Z and assignment of a fair potential for development. Intensification continued and a Tropical Cyclone Formation Alert was issued at 3015572. The first wanning was issued at 310000Z. Tropical Cyclone 05B moved toward tfie nortf and was initially forecast to continue moving northward, then tumn northeastward, crossing soutfiem Bangladesfi and northern Burma. Instead, it assumed a northwestwurd track artd showed in forward speed. Once the system began to take a definite track toward the northwest, the forecast philosopfy was changed and the One-Way Tropical Cyclone Model (OTCM) guidance was followed with excelfert results. The peak intensity of $55 \mathrm{kt}(28 \mathrm{~m} / \mathrm{sec})$ was reached at 1200 Z on the 1 st of November and maintained until the system was close insfore as evidenced by the well-defined convective cloud band on the satellite image above. Issuance of the final waming occurred at 030600Z as Tropical Cyclone 05B was dissipating over land ( 0202212 November $\mathfrak{K}$ (OMA visual imagery).


## TROPICAL CYCLONE 06B



Figure 3-06B-1. Tropical Cyclone 06B became evident on satellite imagery on 8 November as a weakly organized area of convection in the nortfiern Andaman Sea. Initially, it was associated with a broad band of monsoonal choudiness which extended from the southern India eastward to the central Mndaman Sea. First mention of the suspect area occurred on the Significant Tropical Weather Advisory ( $\mathfrak{A B I O}$ PGTW) at 101800Z. Sparse synoptic data implied a closed, Low-Level cyclonic circufation and associated upper-Level divergent flow. As the tropical cyclone's organization increased, a Tropical Cyclone Formation Alert was issued at 1020272 (see above imagery), followed by the first warning at 111800Z. Tropical Cyclone $06 B$ reached a peak intensity of $50 \mathrm{Kf}(26 \mathrm{~m} / \mathrm{sec})$ at 121800 Z (see above imagery) as it turned northwestward. Four fours later it made fandfalf and rapidly weakened whife moving into the Eastern Ghats mountains along the coast. The final warning was issued at 130000Z (120929Z November NOAA visual imagery).


## TROPICAL CYCLONE 07A



Figure 3-07A-1. Tropical Cyclone 07凡 was the first significant tropical cyclone in the Arabian Sea during the month of December since 1980. It afso marked the first time since 1979 that seven significant tropical cyclones have occurred in the North Indian Ocean. Tropical Cyclone $07 \mathcal{A}$ initialfy developed as an exposed Low-Fevel cinculation on December $2 n d$. It slowly intensified, reaching an intensity of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec}$ ) sfiortly before making fandfall on the southeast coast of India at $041900 \mathrm{Z}, 150 \mathrm{~nm}(278 \mathrm{~km})$ south of the city of Madras. No warnings were issued on this tropical depression in the Bay of Bengal, however it was mentioned on the 041800Z Significant Tropical Weather Advisory ( $\mathcal{A B I O} P G T W$ ) as faving poor potential to develop into a significant tropical cyclone due to its proximity to Cand. Symoptic data indicated the disturbance maintained its identity as it tracked acrass the southern tip of India. Once the system moved out over water it reintensified in the Arabian Sea, $97^{\prime W} \mathrm{CC}$ issued a Tropical Cyclone Formation ALert at 080930Z. The first warming followed a few hours Later at 081200Z, with winds of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ based on a satellite intensity aralysis (Doorak, 1984). $\mathcal{A}$ maximum intensity of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$ was reached at 101200 Z prior to Tropical Cyclone 07凡 recurving northward through a break in the subtropical ridge. It then headed toward the western coast of India where increasing vertical wind shear on the 11th weakened Tropical Cyclone 07 $\mathfrak{A}$ before it made landfaff at $120000 \mathrm{Z}, 90 \mathrm{~mm}(167 \mathrm{~km})$ south of Bombay. Nop reports of extensive damage or lass of life were received. The above stored data mosaic shows the system just prior to reaching maximum intersity ( $100202 Z$ December DMKSP visual imagery).


## TROPICAL CYCLONE 08B



Figure 3-08B-1. Tropical Cyclone 08B was the record-setting eighth significant tropical cyclone to develop in the North Indian Ocean in 1987 and the second system to occur in December. It began as a rapidfy organizing tropical disturbance 375 nm ( 695 km ) east of Sri Lanka on the 16th. Tropical Cyclone $08 \mathcal{B}$ was mentioned for the first time on the $161800 Z$ Significant Tropical Weather Advisory (ABIO PGTW) and was classified as faving fair potential to develop into a significant tropical cyclone based on a Low-Level cyclonic circulation evident in the synoptic data and an improving upper-Level outflow pattern. A Tropical Cyclone Formation $\mathfrak{A l e r t}$ (TCFA) was issued the following day at 170800 Z when it became apparent the system was increasing in the amount of convection and in organization. Sateflite intensity analysis (Dvorak, 1984) estimated $25 \mathrm{kt}(13 \mathrm{~m} / \mathrm{sec})$ winds at the surface, at that time. The first waming on Tropical Cyclone 08B came six hours later when it developed a central dense overcast and the sateffite intensity estimate reached $30 \mathrm{~K} \pm(15 \mathrm{~m} / \mathrm{sec})$. Tropical Cyclone $08 B$ was initially forecast to move infand near Madras, India and dissipate within 48 -hours; however, the system slowed dramatically on the 18tf. It changed course at 190000 Z and headed toward the northeast. ITWC issued a final warning at 1900002 when it appeared the upperand Lower-Level centers frad become displaced by strong vertical shear. The exposed Low-Level circulation center maintained its identity during the subsequent 24 -fiour period and re-developed its central convection. $\mathcal{A}$ second TCFA was issued at $200630 Z$ as the remnants of Tropical Cyclone $08 \mathcal{B}$ tracked northeastward and improved in organization. No further warnings were issued, however, despite several satelfite intensity estimates of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ on the 20 th because ITWC believed the intensity was at, or below, warning criteria and not expected to develop further. On the 21st, the tropical cyclone began to weaken and Looped unexpectedly back toward the Indian subcontinent (see above imagery). It made landfall on the 23 rd on the Indian coast, $165 \mathrm{~nm}(306 \mathrm{~km})$ south of Madras, India. $\mathcal{N}(\mathrm{o}$ reports of major damage or loss of life were received (211229Z December DMSSP visual imagery).

## CHAPTER IV - SUMMARY OF SOUTH PACIFIC AND SOUTH INDIAN OCEAN TROPICAL CYCLONES

## 1. GENERAL

The JTWC area of responsibility (AOR) was expanded on 1 October 1980 -- to include the southern hemisphere from 180 degrees Longitude westward to the east coast of Africa. Details on tropical cyclones in this region for July 1980 to June 1982 are contained in Diercks et al, (1982). For the July 1982 through June 1984 period, reference the NOCC/JTWC TECH NOTE 86-1. As in earlier reports, data on tropical cyclones forming in, or moving into, the South Pacific Ocean east of 180 degrees Longitude, which is the Naval Western Oceanography Center's (NAVWESTOCEANCEN) AOR, are included for completeness. JTWC provides the sequential
numbering for all South Pacific and South Indian Ocean significant tropical cyclones. The current convention (as stated in USCINCPACINST 3140.1 (series)) for labelling tropical cyclones that develop in the South Indian Ocean (west of 135 degrees East Longitude) is to add the suffix "S" to the assigned tropical cyclone number, while those originating in the South Pacific Ocean (east of 135 degrees East Longitude) receive a "P" suffix. The "P" suffix also applies to significant tropical cyclones which form east of 180 degrees Longitude in the South Pacific Ocean. Also, it should be noted that to encompass the southern hemisphere tropical cyclone season, which


| 015 -------- | 01 | AUG | - 03 | AUG | 3 | 4 | 40 | (21) | 994 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 02P OSEA | 22 | NOV | - 25 | NOV | 4 | 7 | 55 | (28) | 984 |
| 03P PATSY | 14 | DEC | - 18 | DEC | 5 | 8 | 55 | (28) | 984 |
| 04P RAJA | 23 | DEC | - 01 | JAN | 10 | 18 | 90 | (46) | 953 |
| 05P SALLY | 28 | DEC | - 04 | JAN | 8 | 16 | 65 | (33) | 976 |
| 06 S -------- | 07 | JAN | - 09 | JAN | 3 | 5 | 45 | (23) | 991 |
| 07 S -------- | 10 | JAN | - 12 | JAN | 3 | 5 | 55 | (28) | 984 |
| 08P TUSI | 16 | JAN | - 20 | JAN | 5 | 10 | 100 | (51) | 943 |
| 095 ALININA | 16 | JAN | - 20 | JAN | 5 | 8 | 75 | (39) | 967 |
| 09 S ALININA* | 22 | JAN | - 23 | JAN | 2 | 4 | 65 | (33) | 976 |
| 10 S CONNIE | 17 | JAN | - 20 | JAN | 4 | 6 | 55 | (28) | 984 |
| 11P IRMA | 19 | JAN | - 20 | JAN | 2 | 3 | 30 | (15) | 1000 |
| 12 S DAMIEN | 01 | FEB | - 05 | FEB | 5 | 9 | 50 | (26) | 987 |
| 13 P | 04 | FEB | - 05 | FEB | 2 | 4 | 40 | (21) | 994 |
| 14P UMA | 05 | FEB | - 09 | FEB | 5 | 9 | 80 | (41) | 963 |
| 15P JASON | 07 | FEB | - 13 | FEB | 7 | 13 | 65 | (33) | 976 |
| 16P VELI | 08 | FEB | - 09 | FEB | 2 | 3 | 30 | (15) | 1000 |
| 17 S CLOTILDA | 11 | FEB | - 16 | EEB | 6 | 11 | 50 | (26) | 987 |
| 18 S ELSIE | 22 | FEB | - 25 | FEB | 4 | 7 | 60 | (31) | 980 |
| 19P -------- | 28 | FEB | - 01 | MAR | 2 | 3 | 40 | (21) | 994 |
| 20P WINI | 01 | MAR | - 06 | MAR | 6 | 9 | 65 | (33) | 976 |
| 215 DAODO | 03 | MAR | - 15 | MAR | 13 | 25 | 75 | (39) | 967 |
| 22P YALI | 08 | MAR | - 12 | MAR | 5 | 8 | 65 | (33) | 976 |
| 23P KAY | 08 | APR | - 16 | APR | 9 | 17 | 65 | (33) | 976 |
| 24 S -------- | 23 | APR | - 26 | APR | 4 | 8 | 75 | (39) | 967 |
| 25 P ZUMAN | 23 | APR | - 26 | APR | 4 | 8 | 55 | (28) | 984 |
| 26 S | 24 | APR | - 26 | APR | 3 | 5 | 45 | (23) | 991 |
| 27P BLANCHE | 22 | MAY | - 25 | MAY | 4 | 7 | 55 | (28) | 984 |
| 28 S -------- | 25 | JUN | - 27 | JUN | 3 | 5 | 35 | (18) | 997 |
| 1987 TOTALS: |  |  |  |  | 94 | 245 |  |  |  |

[^2]NOTE: NAMES OF CYCONES GIVEN BY REGIONAL WARNING CENTERS (NANDI, BRISBANE, DARWIN, PERTH AND MAURITIUS) AND ARE APPENDED TO JTWC WARNINGS, WHEN AVAILABLE.

## TABLE 4-2.

FREQUENCY OF CYCLONES BY MONTH AND YEAR

| YEAR | JUL | AUG | SEP | OCT | Nov | DEC | JAN | FEB | MAR | APR | MAY | JUN | TOTAL |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| (1959-1978) |  |  |  |  |  |  |  |  |  |  |  |  |  |
| AVERAGE* | - | - | - | 0.4 | 1.5 | 3.6 | 6.1 | 5.8 | 4.7 | 2.1 | 0.5 | - | 24.7 |
| 1981 | 0 | 0 | 0 | 1 | 3 | 2 | 6 | 5 | 3 | 3 | 1 | 0 | 24 |
| 1982 | 1 | 0 | 0 | 1 | 1 | 3 | 9 | 4 | 2 | 3 | 1 | 0 | 25 |
| 1983 | 1 | 0 | 0 | 1 | 1 | 3 | 5 | 6 | 3 | 5 | 0 | 0 | 25 |
| 1984 | 1 | 0 | 0 | 1 | 2 | 5 | 5 | 10 | 4 | 2 | 0 | 0 | 30 |
| 1985 | 0. | 0 | 0 | 0 | 1 | 7 | 9 | 9 | 6 | 3 | 0 | 0 | 35 |
| 1986 | 0 | 0 | 1 | 0 | 1 | 1 | 9 | 9 | 6 | 4 | 2 | 0 | 33 |
| 1987 | 0 | 1 | 0 | 0 | 1 | 3 | 6 | 8 | 3 | 4 | 1 | 1 | 28 |
| (1981-1987) |  |  |  |  |  |  |  |  |  |  |  |  |  |
| AVERAGE | 0.4 | 0.1 | 0.1 | 0.6 | 1.4 | 3.4 | 7.0 | 7.3 | 3.9 | 3.4 | 0.7 | 0.1 | 28.6 |
| CASES | 3 | 1 | 1 | 4 | 10 | 24 | 49 | 51 | 27 | 24 | 5 | 1 | 200 |

occurs from January through April, the limits of each tropical cyclone year are defined as 1 July to 30 June. Thus, the 1987 southern hemisphere tropical cyclone year is from 1 July 1986 to 30 June 1987. (This is in contrast to the
convention of labelling northern hemisphere tropical cyclones which is based on the calendar year - 1 January to 31 December - to include the seasonal activity from May through December.)

TABLE 4-3. YEARLY VARIATION OF TROPICAL CYCLONES BY OCEAN BASIN

| YEAR | SOUTH INDIAN (105 E WESTWARD) | AUSTRALIAN $(105 E-165 E)$ | SOUTH PACIFIC (165 E EASTWARD) | TOTAL |
| :---: | :---: | :---: | :---: | :---: |
| (1959-1978) |  |  |  |  |
| AVERAGE* | 8.4 | 10.3 | 5.9 | 24.7 |
| 1981 | 13 | 8 | 3 | 24 |
| 1982 | 12 | 11 | 2 | 25 |
| 1983 | 7 | 6 | 12 | 25 |
| 1984 | 14 | 14 | 2 | 30 |
| 1985 | 14 | 15 | 6 | 35 |
| 1986 | 14 | 16 | 3 | 33 |
| 1987 | 9 | 8 | 11 | 28 |
| (1981-1987) | 11.9 | 11.1 | 5.6 | 28.6 |
| AVERAGE |  |  |  |  |
|  | 83 | 78 | 39 | 200 |
| CASES |  |  |  |  |

* (GRAY, 1979)
TABLE 4-4. MAXIMUM SUSTAINED SURFACE WINDS VERSUS MINIMUM SEA-LEVEIPRESSURE (ATKINSON AND HOLLIDAY, 1977)
MAXIMUM SUSTAINED
SURFACE WIND (KT)
EQUIVALENT MINIMUM
SEA-LEVEL PRESSURE (MB
30 ..... 1000
35 ..... 997
40 ..... 994
45 ..... 991
50 ..... 987
55 ..... 984
60 ..... 980
65 ..... 976
70 ..... 972
75 ..... 967
80 ..... 963
85 ..... 958
90 ..... 954
95 ..... 948
100 ..... 943
105 ..... 938
110 ..... 933
115 ..... 927
120 ..... 922
125 ..... 916
130 ..... 910
135 ..... 904
140 ..... 898
145 ..... 892
150 ..... 885
155 ..... 879
160 ..... 872
165 ..... 865
170 ..... 858


## 2. SOUTH PACIFIC AND SOUTH INDIAN OCEAN TROPICAL CYCLONES

The 1987 year (1 July 1986 through 30 June 1987) was active, with 28 significant tropical cyclones (see Table 4-1) reaching warning status. This did not exceed the total of 33 tropical cyclones for 1986 (1 July 1985-30 June 1986). Eleven tropical cyclones occurred in the South Pacific, east of 165 degrees East Longitude, which is about twice the long-term mean. The Australian area ( 105 to 165 degrees East Longitude) accounted for eight tropical cyclones compared to the climatological mean of 10.3 - two less than normal. Nine tropical cyclones developed in the South Indian Ocean, which is about one more than the long-term mean of 8.4 cyclones (see Tables 4-2 and 4-3). Meteorological satellite surveillance of tropical cyclones has been updating climatologies since the early 1960s. (This meteorological watch from space detects tropical cyclones that might have previously gone undetected over the conventional data sparse oceanic areas.) Thus, tropical cyclone climatologies should continue to benefit from increased surveillance from space in some areas, for example, the South Indian Ocean.

Caveat: Intensity estimates for southern hemisphere tropical cyclones are derived primarily from satellite imagery evaluation (Dvorak, 1984) and from intensity estimates reported by other regional centers. Only, in extremely rare instances are the intensity estimates based on surface observational data. Estimates of the minimum sea-level pressure are usually derived from the Atkinson and Holliday (1977) relationship between the maximum sustained one-minute surface wind and the minimum sea-level pressure (Table 4-4). This relationship has been shown to be representative for tropical cyclones in the western North Pacific and is also used by the Australian regional warning centers to provide intensity estimates. However, since these pressure estimates are usually based on wind intensities that were derived from interpretation of satellite imagery, considerable caution should be exercised when using these resultant pressure values in future tropical cyclone work.



## CHAPTER V - SUMMARY OF FORECAST VERIFICATION

## 1. ANNUAL FORECAST VERIFICATION

## a. Western North Pacific Ocean

The positions given for warning times and those at the 24-, 48- and 72-hour forecast times were verified against the final best track positions at the same valid times. The (scalar) forecast, along-track and cross-track errors (illustrated in Figure 5-1) were then calculated for each tropical cyclone and are presented in Tables 5-1A, 5-1B, 5-1C and 5-1D. Figures 52A through 5-2C provide, respectively, the frequency distributions of forecast errors in 30 nm increments for 24-, 48 -, and 72 -hour forecasts of all 1987 tropical cyclones in the western North Pacific. A summation of the mean forecast errors, as calculated for all tropical cyclones in each year, is shown in Table 5-2A. Table 5-2B includes along-track and cross-track errors for 1987. A comparison of the annual mean forecast errors for all tropical cyclones as compared to those tropical cyclones that reached typhoon intensity can be seen in Table 5-3. The mean forecast errors for 1987 as compared to the ten previous years are graphed in Figure 5-3.


Figure 5-1. Definition of cross-track error (XTIE), along-track error (ATE), and forecast track error ( $\mathcal{F T E}$ ). In tfis example, the XTE is positive (to the right of the Best Track) and the $\mathcal{A T E}$ is negative (6efind or sfower than the Best Track).


| FORFCAST FRRRORS (NM) |  |
| :--- | :---: |
| MFAN : | 107 |
| MAGOUR |  |
| MFDIAN: | 106 |
| STANDARD |  |
| DEVIATION: | 60 |
| CASES : | 563 |

$\mathcal{F i g u r e}$ 5-2A. Frequency distribution of the 24 -hour forecast errors in 30 nm ( 56 km ) increments for all significant tropical cyclones in the western North Pacific during 1987.




TABLE 5-1D.
72-HOUR FORECAST ERROR SUMMARY FOR THE WESTERN NORTH PACIFIC OCEAN SIGNIFICANT TROPICAI CYCLONES OF 1987 (ERRORS IN NM)

| TROPICAI |  | CYCLONE | FORECASI ERROR | ALONG-TRACK ERROR |  |  | CROSS-TRACK ERROR |  |  | NUMBER OF WARNINGS |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | ABS |  | MAG | BIAS | ABS | MAG | BIAS |  |
| (01w) | TY |  | ORCHID | 234 |  |  | -100 |  | 3 | 79 | 11 |
| (02W) | TS | PERCY | 206 |  |  | * |  | 3 | * | 4 |
| (03W) | TS | RUTH |  |  |  |  |  |  | --- | 0 |
| (04W) | TY | SPERRY | 421 |  |  | * |  | 8 | * | 9 |
| (05w) | STY | THELMA | 479 |  |  | -270 |  | 4 | 90 | 24 |
| (06W) | TY | VERNON | 224 |  |  | * |  | 5 | * | 10 |
| (07w) | TY | WYNNE | 332 |  |  | -207 | 1 |  | -86 | 30 |
| (08W) | TY | ALEX | 330 |  |  | -245 |  | 2 | -97 | 12 |
| (09W) | STY | BETTY | 257 |  |  | -178 | 1 | 55 | 136 | 17 |
| (10w) | TY | CARY | 248 |  | 5 | 2 | 2 | 2 | 218 | 28 |
| (11W) | STY | DINAH | 222 |  |  | -6 |  | 3 | 3 | 30 |
| (12W) | TS | ED | 278 |  |  | * |  | 5 | * | 5 |
| (13W) | TY | FREDA | 327 |  |  | 6 | 2 | 2 | -190 | 34 |
| (14W) | TY | GERALD | 193 |  |  | -31 | 1 | 9 | 57 | 14 |
| (15W) | STY | HOLIY | 455 |  |  | -35 | 3 | 2 | -299 | 31 |
| (16W) | TY | IAN | 344 | 25 |  | 122 | 1 | 1 | -52 | 21 |
| (17w) | TD | 17 W |  |  |  | 122 |  |  | , | 0 |
| (02C) | TY | PEKE | 247 |  |  | * |  |  | * | 5 |
| (18W) | TS | JUNE | 264 |  |  | * |  |  | * | 1 |
| (19\%) | TY | KEL工Y | 289 | 21 |  | -159 | 1 | 4 | -39 | 17 |
| (20W) | STY | LYNN | 298 |  |  | 27 |  | 2 | -18 | 31 |
| $(21 w)$ | TS | MAURY | 183 |  |  | 101 |  | 0 | 47 | 12 |
| (22W) | STY | NINA | 279 |  |  | -191 | 15 |  | 93 | 30 |
| (23W) | TS | OGDEN |  |  |  | 191 |  |  | 9 | 0 |
| (24W) | TY | PHYLIIS | 196 |  |  | 24 |  | 6 | 64 | 13 |
| TOTALS |  |  | 303 |  |  | -78 | 1 |  | -13 | 389 |

TABLE 5-2A. ANNUAL MEAN FORECAST ERRORS FOR THE WESTERN NORTH PACIFIC

| YEAR | 24-HOUR |  | 48-HOUR |  | FORECAST 72 -HOUR ${ }^{\text {72 }}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| YEAR |  | RIGHT ANGLE | FORECAST | RIGHT ANGLE | FORECAST | RIGHT ANGLE |
| 1971 | 111 | 64 | 212 | 118 | 317 | 117 |
| 1972 | 117 | 72 | 245 | 146 | 381 | 210 |
| 1973 | 108 | 74 | 197 | 134 | 253 | 162 |
| 1974 | 120 | 78 | 226 | 157 | 348 | 245 |
| 1975 | 138 | 84 | 288 | 181 | 450 | 290 |
| 1976 | 117 | 71 | 230 | 132 | 338 | 202 |
| 1977 | 148 | 83 | 283 | 157 | 407 | 228 |
| 1978 | 127 | 75 | 271 | 179 | 410 | 297 |
| 1979 | 124 | 77 | 226 | 151 | 316 | 223 |
| 1980 | 126 | 79 | 243 | 164 | 389 | 287 |
| 1981* | 123 | 75 | 220 | 119 | 334 | 168 |
| 1982* | 113 | 67 | 237 | 139 | 341 | 206 |
| 1983* | 117 | 72 | 259 | 152 | 405 | 237 |
| 1984* | 117 | 66 | 233 | 137 | 363 | 231 |
| 1985* | 117 | 66 | 231 | 134 | 367 | 214 |
| 1986 | 121 | ** | 261 | ** | 394 | ** |
| 1987 | 107 | ** | 204 | ** | 303 | ** |

* THE TECHNIQUE FOR CALCULATING RIGHT-ANGLE ERROR WAS REVISED IN 1981 ; THEREFORE, A DIRECT CORRELATION IN RIGHT-ANGLE STATISTICS CANNOT BE MADE FOR THE ERRORS COMPUTED BEFORE 1981 AND THE ERRORS COMPUTED SINCE 1981.
** IN 1986, THE RIGHT-ANGLE ERROR WAS REPLACED BY CROSS-TRACK ERROR (SEE FIGURE 5-1 FOR THE DEFINITION OF CROSS-TRACK ERROR).

TABLE 5-2B. 1987 MEAN FORECAST, ALONG-TRACK AND CROSS-TRACK ERRORS FOR THE WESTERN NORTH PACIFIC OCEAN (ERRORS IN NM).

| TIMES: | FORECAST ERROR | $\begin{gathered} \text { ALONG-TRACK } \\ \text { ERROR } \end{gathered}$ |  |  | CROSS-TRACK |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | ABS | MAG | BIAS | ABS | MAG | BIAS |
| 24 -HOUR | 107 | 71 |  | -30 | 64 |  | -8 |
| 48-HOUR | 204 | 134 |  | -58 | 127 |  | -12 |
| 72-HOUR | 303 | 198 |  | -78 | 186 |  | -13 |

TABLE 5-3. ANNUAL MEAN FORECAST ERRORS FOR THE WESTERN NORTH PACIFIC

| YEAR | 24-HOUR |  | 48-HOUR |  | 72-HOUR |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1950-1958 |  | 170 |  |  |  |  |
| 1959 |  | $117 * *$ |  | 267** |  |  |
| 1960 |  | 177** |  | $354 * *$ |  |  |
| 1961 |  | 136 |  | 274 |  |  |
| 1962 |  | 144 |  | 287 |  | 476 |
| 1963 |  | 127 |  | 246 |  | 374 |
| 1964 |  | 133 |  | 284 |  | 429 |
| 1965 |  | 251 |  | 303 |  | 41 C |
| 1966 |  | 136 |  | 280 |  | 432 |
| 1967 |  | 125 |  | 276 |  | 414 |
| 1968 |  | 105 |  | 229 |  | 337 |
| 1969 |  | 111 |  | 237 |  | 349 |
| 1970 | 104 | 98 | 190 | 281 | 279 | 272 |
| 1971 | 111 | 99 | 212 | 203 | 317 | 308 |
| 1972 | 117 | 116 | 245 | 245 | 381 | 382 |
| 1973 | 108 | 102 | 197 | 193 | 253 | 245 |
| 1974 | 120 | 114 | 226 | 218 | 348 | 357 |
| 1975 | 138 | 129 | 288 | 279 | 450 | 442 |
| 1976 | 117 | 117 | 230 | 232 | 338 | 336 |
| 1977 | 148 | 140 | 283 | 266 | 407 | 390 |
| 1978 | 127 | 120 | 271 | 241 | 410 | 459 |
| 1979 | 124 | 113 | 226 | 219 | 316 | 319 |
| 1980 | 126 | 116 | 243 | 221 | 389 | 362 |
| 1981 | 123 | 117 | 220 | 215 | 334 | 342 |
| 1982 | 113 | 114 | 237 | 229 | 341 | 337 |
| 1983 | 117 | 110 | 259 | 247 | 405 | 384 |
| 1984 | 117 | 110 | 233 | 228 | 363 | 361 |
| 1985 | 117 | 112 | 231 | 228 | 367 | 355 |
| 1986 | 121 | 117 | 261 | 261 | 394 | 403 |
| 1987 | 107 | 201 | 204 | 211 | 303 | 318 |

* FORECASTS WERE VERIFIED WHEN THE TROPICAL CXCLONE INTENSITIES WERE OVER 35 KT (18 M/SEC).
** FORECAST POSITIONS NORTH OF. 35 DEGREES NORTH LATITUDE WERE NOT

WESTERN NORTH PACIFIC


Figure 5-3. Annual mean forecast errors (in nm) for aff significant tropical cyclones in the western North Pacific.

## b. North Indian Ocean

The positions given for warning times and those at the $24-, 48$-, and 72 -hour valid times were verified for tropical cyclones in the North Indian Ocean by the same methods used for the western North Pacific. It should be noted that despite the record-setting eight North Indian Ocean tropical cyclones, these error statistics should not be taken as representative of any trend due to the small sample number.

Table 5-4 is the forecast along-track and crosstrack error summary for the North Indian Ocean. Table 5-5A contains a summary of the annual mean forecast errors for each year. Table 5-5B includes along-track and cross-track errors for 1987. Forecast errors are plotted in Figure 5-4 (Seventy-two hour forecast errors were evaluated for the first time in 1979). There were no verifying 72 -hour forecast in 1983 and 1985.




Figure 5-4. Annual mean forecast errors (in nm) for all tropical cyclones in the N(ortfi Indian Ocean.
c. South Pacific and South Indian Oceans

The positions given for warning times and those at the 24-, 48 -, and 72 -hour valid times were verified for tropical cyclones in the Southern Hemisphere by the same methods used for the western North Pacific. It should be noted that due to the lack of verifying ground-
truth data, these error statistics should not be taken as representative of any trend. Table 5-6 is the forecast along-track and cross-track error summary for the Southern Hemisphere. Tables 5-7A and B contains a summary of the annual mean forecast errors for each year. Table 5-8 includes along- and cross-track errors for 1987. Forecast errors are plotted in Figure 5-5.

| TABLE 5-6. |  |  | 1987 FORECAS |  | ERROR SUMMARY FOR THE SOUTH PACIFIC AND SOUTH INDIAN OCEAN SIGNIFICANT TROPICAL CYCLONES (ERRORS IN NM) |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  | 24-HOURS |  | CROSS-TRACK |  | NO OF WRNGS | FCSTERROR | 48-HOURS ALONG-TRACK |  | CROSS-TRACK |  |
| TROPICAL | NO OF | INITTAL | NO Or | FCST | ALONG-T | RACK |  |  |  |  |  |  |  |  |
| CYCLONE | WRNGS | POS ER | WRNGS | ERROR | ABS MAG | BIAS | ABS MAG | BIAS |  |  | ABS MBG | BIAS | ABS MAG | BIAS |
| TC 015 | 4 | 25 | 3 | 90 | 56 | ** | 57 | ** | 1 | 196 | 49 | ** | 190 | ** |
| TC 02P OSEA | 6 | 33 | 4 | 141 | 125 | ** | 61 | ** | 2. | 292 | 248 | ** | 145 | ** |
| TC 03P PATSY | 6 | 42 | 4 | 160 | 67 | ** | 135 | ** | 2 | 340 | 141 | ** | 296 | ** |
| TC 04P RAJA | 18 | 35 | 16 | 171 | 117 | -99 | 114 | 0 | 14 | 333 | 259 | -177 | 176 | 21 |
| TC 05P SALLY | 16 | 41 | 14 | 139 | 62 | -39 | 106 | 53 | 11 | 281 | 248 | -129 | 97 | -24 |
| TC 06S | 5 | 38 | 4 | 203 | 119 | ** | 155 | ** | 4 | 365 | 105 | ** | 339 | ** |
| TC 07s | 5 | 24 | 3 | 119 | 73 | ** | 80 | ** | 2 | 194 | 133 | ** | 138 | ** |
| TC OAP TUSI | 9 | 25 | 7 | 104 | 45 | ** | 86 | ** | 5 | 254 | 75 | ** | 236 | ** |
| TC 09S ALININA | 12 | 37 | 8 | 188 | 124 | ** | 111 | ** | 5 | 422 | 220 | ** | 331 | ** |
| TC 10S CONNIE | 6 | 15 | 5 | 92 | 34 | ** | B0 | ** | 4 | 201 | 58 | ** | 183 | ** |
| TC 11P IRMA | 3 | 13 | 1 | 17 | 12 | ** | 12 | ** | 0 | -- | - | - | -- | - |
| TC 12S DAMIEN | 9 | 35 | 7 | 138 | 62 | ** | 111 | ** | 5 | 135 | 118 | ** | 49 | ** |
| TC 13P | 3 | 36 | 2 | 139 | 53 | ** | 125 | ** | 0 | -- | -- | -- | -- | -- |
| TC 14P UMA | 8 | 347 | 6 | 101 | 77 | ** | 55 | ** | 4 | 256 | 199 | ** | 115 | ** |
| TC 15P JASON | 12 | 27 | 8 | 120 | 109 | ** | 35 | ** | 2 | 369 | 344 | ** | 129 | ** |
| TC 16P VELI | 3 | 67 | 1 | 424 | 422 | ** | 44 | ** | 0 | -- | --- | -- |  | -- |
| TC 17S CLOTILDA | 10 | 51 | 7 | 109 | 61 | ** | 86 | ** | 6 | 227 | 151 | ** | 154 | ** |
| TC 18S | 7 | 11 | 6 | 147 | 76 | ** | 120 | ** | 4 | 369 | 140 | ** | 304 | ** |
| TC 19P | 3 | 5 | 1 | 332 | 325 | ** | 69 | ** | 0 | -- | --- | --- | - | - |
| TC 20p WINI | 9 | 47 | 6 | 146 | 122 | ** | 66 | ** | 4 | 286 | 220 | ** | 146 | ** |
| TC 21s DAODO | 25 | 47 | 20 | 188 | 110 | -86 | 136 | -14 | 18 | 307 | 182 | -38 | 204 | 0 |
| TC 22P Yali | 7 | 50 | 5 | 96 | 81 | ** | 42 | ** | 3 | 212 | 89 | ** | 192 | ** |
| TC 23P KAY | 17 | 25 | 12 | 90 | 50 | -6 | 63 | -24 | 12 | 166 | 105 | 1 | 113 | -82 |
| TC 24S | 8 | 35 | 6 | 209 | 150 | ** | 112 | ** | 4 | 519 | 446 | ** | 174 | ** |
| TC 25P 2UMAN | 7 | 49 | 5 | 136 | 86 | ** | 84 | ** | 4 | 273 | 200 | ** | 163 | ** |
| TC 26S | 5 | 47 | 3 | 161 | 147 | ** | 54 | ** | 1 | 344 | 341 | ** | 46 | ** |
| TC 27P BLANCHE | 7 | 22 | 5 | 93 | 70 | ** | 48 | ** | 5 | 138 | 106 | ** | 77 | ** |
| TC 28S | 5 | 56 | 3 | 194 | 106 | ** | 132 | ** | 1 | 450 | 374 | ** | 250 | ** |
| TOTALS | 235 | 46 | 172 | 145 | 94 | -57 | 90 | 13 | 123 | 280 | 195 | -102 | 161 | 6 |
| ** the statistical parameter bias does not compute for instances of ten cases, or less.* |  |  |  |  |  |  |  |  |  |  |  |  |  |  |

TABLE 5-7A. ANNUAL MEAN FORECAST ERRORS FOR THE SOUTHERN HEMISPHERE

> 24-HOUR 48-HOUR

|  | $24-$ HOUR |  | 48 -HOUR |  |
| :--- | :---: | :---: | :---: | :---: |
| YEAR | FORECAST | RIGHT ANGLE | FORECAST | RIGHT ANGLE |
|  |  |  |  |  |
| 1981 | 165 | 119 | 315 | 216 |
| 1982 | 144 | 91 | 274 | 174 |
| 1983 | 154 | 84 | 288 | 150 |
| 1984 | 133 | 73 | 231 | 124 |
| 1985 | 138 | 78 | 242 | 133 |
| 1986 | 133 | $\star * *$ | 268 | $\star * *$ |
| 1987 | 145 | $* * *$ | 280 | $* * *$ |


| TABLE 5-7B. | 1987 | MEAN FORECAST, ALONG-TRACK AND CROSS-TRACK ERRORS |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| : |  | ALONG | ACK | CROSS | аск |
|  | FORECAST | ERR |  | ERR |  |
| TIMES : | ERROR | ABS MAG | BIAS | ABS MAG | BIAS |
| 24-HOUR | 145 | 94 | -57 | 90 | 13 |
| 48-HOUR | 280 | 195 | -102 | 161 | 6 |



Figure 5-5. Annual mean forecast errors (in nm ) for all tropical cycfones in the South Pacific and South Indian Oceans.

## 2. COMPARISON OF OBJECTIVE TECHNIQUES

## a. General

Objective techniques used by JTWC are divided into five main categories:
(1) extrapolation;
(2) climatological and analog techniques;
(3) model output statistics;
(4) dynamic models; and
(5) empirical and analytical techniques;

In September 1981, JTWC began to initialize its array of objective forecast techniques (described below) on the six-hour old preliminary best track position (an interpolative process) rather than the forecast (partially extrapolated) warning position, e.g. the 0600 Z warning is now supported by objective techniques developed from the 0000 Z preliminary best track position. This operational change has yielded several advantages:
*Techniques can now be requested much earlier in the warning development time line, i.e. as soon as the track can be approximated by one or more fix positions after the valid time of the previous warning;
*Receipt of these techniques is virtually assured prior to the development of the next warning; and
*Improved (mean) forecast accuracy. This latter aspect arises because JTWC now has more reliable approximation of the short-term tropical cyclone movement. Further, since most of the objective techniques are biased towards persistence, this new procedure optimizes their performance and provides more consistent guidance on short-term movement, indirectly yielding a more accurate initial position estimate as well as lowering 24-hour forecast errors.
b. Description of Objective Techniques
(1). XTRP -- Forecast positions for 24- and 48-hours are derived from the extension of a straight line which connects the most recent and 12 -hour old preliminary best track positions.
(2). CLIM -- A climatological aid providing 24 -, 48 -, and 72 -hour tropical cyclone forecast positions (and intensity changes in the western North Pacific) based upon the position of the tropical cyclone. The output is based upon data records from 1945 to 1981 for the western North Pacific Ocean and 1900 to 1981 for the North Indian Ocean.
(3). HPAC -- Forecast positions are generated from a blend of climatology and persistence. The 24 -, 48 - and 72 -hour positions are equally weighted between climatology and persistence. Persistence is a straight line extension of a line connecting the current and 12-hour old positions. Climatology is based on data from 1945 to 1981 for the western North Pacific Ocean and 1900 to 1981 for the North Indian Ocean.
(4). TYAN -- An updated analog program which combines the earlier versions TYFN 75 and INJAN 74. The program scans a 30-year climatology with a similar history (within a specified acceptance envelope) to the current tropical cyclone. For the western North Pacific Ocean, three forecasts of position and intensity are provided for $24-$, 48 -, and $72-$ hours: RECR - a weighted mean of all tropical cyclones which were categorized as "recurving" during their best track period; STRA - a weighted mean of all accepted tropical cyclones which were categorized as moving "straight" (westward) during their best track period: TOTL - a weighted mean of all accepted tropical cyclones, including those used in the RECR and STRA forecast. For the North Indian Ocean, a single (total) forecast track is provided for the 12 -hour intervals to 72 -hours.
(5). COSMOS -- A model output statistics (MOS) routine based on the geostrophic steering at the $850-, 700-$, and $500-$ mb levels. The steering is derived from the HATTRACK point advection model run on Global prognostic fields from the FLENUMOCEANCEN's NOGAPS prediction system. The MOS forecast is then blended with the 6hour past movement to generate the forecast track.
(6). Colorado State University Model (CSUM) -- A statistical method developed Matsumoto (1984) utilizes synoptic and persistence predictors by discretizing the forecast timeframe into three 24 -hour time steps. Climatology is incorporated into the forecast via a stratification scheme based on the position of the tropical cyclone relative to the 500 mb subtropical ridge. Three sets of regression equations are used to determine the north-south and east-west displacements depending on whether the tropical cyclone is south, on or north relative to the ridge.
(7). One-way Interactive Tropical Cyclone Model (OTCM) -- A coarse-mesh, three-layer in the vertical, primitive equation model with a 205 km grid spacing over a 6400 x 4700 km domain. The model's fields are computed around a bogused, digitized cyclone vortex using FLENUMOCEANCEN's' Numerical Variational Analysis (NVA) or NOGAPS prognostic fields for the specified valid time. The past motion of the tropical cyclone is compared to initial steering fields and a bias correction is computed and applied to the model. FLENUMOCEANCEN's NOGAPS global prognostic fields are used at 12-hour intervals to update the model's boundaries. The resultant forecast positions are derived by locating the 850 mb vortex at six-hour intervals to 72 -hours.
(8). TAPT -- An empirical technique which utilizes upper-tropospheric wind fields to estimate acceleration associated with the tropical cyclone's interaction with the mid-latitude westerlies. It includes guidelines for the duration of acceleration, upper-limits, and probable path of the cyclone.

## (9). CLIPER -- A statistical

regression technique based on climatology, current intensity, position and past movement. This technique is used as a crude measure of real forecast skill when verifying forecast accuracy.
(10). THETA-E -- An empirically derived relationship between a tropical cyclone's minimum sea-level pressure (MSLP) and 700 mb equivalent potential temperature (Theta-E) was developed by Sikora (1976) and Dunnavan (1981). By monitoring MSLP and trends, the forecaster can evaluate the potential for sudden, rapid deepening of a tropical cyclone.
(11). WIND RADIUS -- Following an analytical model of the radial profiles of sealevel pressures and winds in mature tropical cyclones (Holland, 1980), a set of radii for $30-$, $50-$, and $100-\mathrm{knot}$ winds based on the tropical cyclone's maximum winds have been produced to aid the forecaster in determining forecast wind radii.
(12). DVORAK -- An estimation of tropical cyclone's current and 24 -hour forecast intensity is made from interpolation of satellite imagery (DVORAK, 1984) and provided to the forecaster. These intensity estimates are used in conjunction with other intensity-related data and trends to forecast tropical cyclone intensity.

JTWC uses HPAC, TAPT, TYAN78, COSMOS, OTCM and CSUM operationally to develop track forecasts.

## c. Testing and Results

A comparison of selected techniques is included in Table 5-8 for all western North Pacific tropical cyclones, Table 5-9 for all North Indian Ocean tropical cyclones. In these tables, " x-axis " refers to techniques listed vertically. For example (Table 5-8) in the 507 cases available for a (homogeneous) comparison, the average forecast error at 24 -hours was 118 nm ( 219 km ) for TOTL and 120 nm ( 222 km ) for RECR. The difference of $2 \mathrm{~nm}(4$ km ) is shown in the lower right. (Differences are not always exact, due to computational round-off which occurs for each of the cases available for comparison).



TABLE 5-10. 1987 ERROR STATISTICS FOR SELECTED OBJECTIVE TECHNIQUES IN THE SOUTHERN HEMISPHERE 24-HOUR MEAN FORECAST ERROR (NM)

|  | JTWC |  | OTCM |  | toti |  | hPAC |  | CLIM |  | XTRP |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| JTWC | 172 | 145 |  |  |  |  |  |  |  |  |  |  |
|  | 145 | 0 |  |  |  |  |  |  |  |  |  |  |
| OTCM | 100 | 141 | 111 | 136 |  |  |  |  |  |  |  |  |
|  | 138 | -3 | 136 | 0 |  |  |  |  |  |  |  |  |
| toti | 78 | 135 | 80 | 129 | 88 | 194 |  |  |  |  |  |  |
|  | 202 | 67 | 166 | 37 | 194 | 0 |  |  |  |  |  |  |
| hPac | 71 | 130 | 73 | 126 | 70 | 126 | 81 | 126 |  |  |  |  |
|  | 124 | -6 | 119 | -7 | 127 | 1 | 126 | 0 |  |  |  |  |
| CLIM | 71 | 130 | 73 | 126 | 70 | 126 | 81 | 126 | 81 | 271 |  |  |
|  | 168 | 38 | 164 | 38 | 165 | 39 | 171 | 45 | 171 | 0 |  |  |
| XTRP | 99 | 137 | 99 | 137 | 78 | 203 | 81 | 126 | 81 | 171 | 112 | 146 |
|  | 148 | 11 | 133 | -4 | 136 | -67 | 118 | -8 | 118 | -53 | 146 | 0 |

48-HOUR MEAN FORECAST ERROR (NM)

\begin{tabular}{|c|c|c|c|c|c|c|c|c|c|c|c|c|}
\hline \& \multicolumn{2}{|l|}{JTWC} \& \multicolumn{2}{|l|}{OTCM} \& \multicolumn{2}{|l|}{TOTL} \& \multicolumn{2}{|l|}{HPAC} \& CLIM \& \multicolumn{3}{|l|}{XTRP} <br>
\hline JTWC
OTCM \& 123
280

71

298 \& $$
\begin{array}{r}
280 \\
0 \\
283 \\
15
\end{array}
$$ \& \[

$$
\begin{array}{r}
82 \\
287
\end{array}
$$

\] \& \[

$$
\begin{array}{r}
287 \\
0
\end{array}
$$

\] \& \& \& \& \& \& | JTWC |
| :--- |
| OTCM |
| TOTL |
| HPAC |
| CIM |
| XTRP | \& - ORF

- ONE
- TOT
- HAL
- CKI
- $12-1$ \& | IAL JTwC FORECAST |
| :--- |
| ay tropical cyclone model analog (tyan 18) |
| Clithatology aid persistance blend tolocy |
| Extrapolation | <br>

\hline \multirow[t]{2}{*}{TOTL} \& 58 \& 271 \& 60 \& 274 \& 65 \& 319 \& \& \& \& \& \& <br>
\hline \& 330 \& 59 \& 283 \& 9 \& 319 \& 0 \& \& \& \& \& \& <br>
\hline \multirow[t]{2}{*}{HPAC} \& 51 \& 264 \& 52 \& 270 \& 50 \& 240 \& 60 \& 215 \& \& \& \& <br>
\hline \& 218 \& -46 \& 218 \& -52 \& 215 \& -25 \& 215 \& 0 \& \& \& \& <br>
\hline \multirow[t]{2}{*}{CLIM} \& 51 \& 264 \& 52 \& 270 \& 50 \& 240 \& 60 \& 215 \& 60 \& 287 \& \& <br>
\hline \& 291 \& 27 \& 288 \& 18 \& 262 \& 22 \& 287 \& 72 \& 287 \& 0 \& \& <br>

\hline XTRP \& $$
\begin{array}{r}
69 \\
300
\end{array}
$$ \& \[

$$
\begin{array}{r}
279 \\
21
\end{array}
$$

\] \& \[

$$
\begin{array}{r}
71 \\
278
\end{array}
$$

\] \& \[

279

\] \& \[

$$
\begin{array}{r}
56 \\
275
\end{array}
$$

\] \& \[

329

\] \& \[

$$
\begin{array}{r}
60 \\
247
\end{array}
$$

\] \& \[

$$
\begin{array}{r}
215 \\
32
\end{array}
$$

\] \& \[

$$
\begin{array}{r}
60 \\
247
\end{array}
$$

\] \& \[

$$
\begin{array}{r}
287 \\
-40
\end{array}
$$

\] \& \[

$$
\begin{array}{r}
84 \\
304
\end{array}
$$

\] \& \[

$$
\begin{array}{r}
304 \\
0
\end{array}
$$
\] <br>

\hline
\end{tabular}

72-HOUR MEAN FORECAST•ERROR (NM)


# CHAPTER VI - TROPICAL CYCLONE SUPPORT SUMMARY 

## 1. NAVAL ENVIRONMENTAL PREDICTION RESEARCH FACILITY

The Pocket Tropical Cyclone Model (PTCM)

(Evans, J.L., Monash University, Australia and J.H. Chu, NAVENVPREDRSCHFAC)

PTCM is a linear tropical cyclone motion prediction scheme incorporating the effects of a large-scale environmental flow and the beta-effect. The model is based on the equations developed by Holland (1983) and has been operational in a modified form in the Australian region for a number of years. The current version of the model has been developed by Evans and Holland to be a purely objective forecasting tool, and is presently undergoing operational testing in the Australian region.

PTCM is being incorporated in the NEPRF ATCF system and a series of case studies are planned to test its effectiveness in the Northwest Pacific region. In addition, the model will be expanded to include additional terms for diagnosis of tropical cyclone motions.

## THE ADVANCED TROPICAL CYCLONE MODEL (ATCM)

(Hodur, R.M., NAVENVPREDRSCHFAC)

The Advanced Tropical Cyclone Model (ATCM) was installed at the Fleet Numerical Oceanography Center in 1987 for evaluation by JTWC forecasters. Although testing in 1986 indicated that the ATCM could perform better than the OTCM, these results were not obtained during real-time runs during the 1987 season. In particular, the ATCM demonstrated a large right bias and nearly always weakened the storms with time. These effects were particularly noticeable in the ATCM forecasts of Typhoon Lynn in October, 1987.

Experiments are being performed to isolate the reason(s) for these ATCM forecast errors. A new version of the ATCM has been developed which incorporates some of the features of NOGAPS 3.0. These include a 15 -
level optimum interpolation analysis and an increase in the number of model levels from 12 to 21 in order to include a high resolution planetary boundary layer. In addition, the cumulus parameterization has been modified so that the ATCM can maintain the tropical storm circulation during the forecast. Also, sensitivity experiments are being conducted to find the best structure of the initial bogus and to study the effect of increasing the horizontal resolution.

## Navy Tactical Applications Guide (NTAG), Vol. 6

(Fett, R.W., NAVENVPREDRSCHFAC)

An effort is now underway to develop a series of examples demonstrating the use of high quality satellite data for analysis and forecasting in the tropics. Both polar orbital and geostationary satellite data are used to study the evolution of certain weather effects or of a particular weather phenomenon at a given time. These examples are intended for publishing in the NTAG Volume 6, Part I, Tropical Weather Analysis and Forecast Applications, and Volume 6, Part II, Tropical Cyclone Weather Analysis and Forecast Applications. NTAG Volume 6, Part I was distributed in June 1986. Part II is still in the research process. Publication is anticipated in 1988/89.

## Tropical Cyclone Condition Setting Aid for Sasebo and Iwakuni, Japan

## (Jarrell, J.D., Sci. Appl. International Corporation)

A forecast aid has been developed for predicting tropical cyclone associated winds at Sasebo and Iwakuni, Japan. The aid consists of two parts. The first part is a collection of charts which relate winds observed at the two stations to the maximum sustained winds at the center of a tropical cyclone as a function of cyclone locations. The second part of the aid is a collection of diagrams which estimate the worst case arrival time of $50-\mathrm{kt}$ winds.

## Improvements to Combined Confidence Rating System

Harry Hamilton (ST Systems, Monterey, CA)
The Combined Confidence Rating System (CCRS) has been improved via a redesign of its weighting function. The new weighting function is derived from the following: the inverse of a covariance matrix which is a combination of the historical cross-track and along-track covariance matrices, and the objective aid forecasts. The weights are generated as follows:
a. Let $\mathrm{Q}_{\mathrm{X}}$ and $\mathrm{Q}_{\mathrm{y}}$ be the cross-track and along-track covariance matrices, respectively. The desired combination of these two, $Q$, is equal to $Q_{x}+a Q_{y}$, where a has been determined empirically to be 0.25 .
b. The weight for each available objective forecast technique is the sum of all terms of the relevant technique divided by the sum of all terms of $Q^{-1}$. The sum of the weights for all available objective techniques must equal 1.0.

The Combined Confidence Weighted Forecast (CCWF) is generated for JTWC by summing the selected objective forecasts used in the calculations.

## Automated Tropical Cyclone Forecasting System

(Tsui, T.L., Miller, R.J., and A.J. Schrader, NAVENVPREDRSCHFAC)

The Automated Tropical Cyclone Forecasting (ATCF) system is an IBM PC compatible software package currently being developed for the Joint Typhoon Warning Center (JTWC). ATCF is designed to allow JTWC forecasters to display graphically tropical cyclone forecast information, merge and analyze synoptic wind fields, provide objective fix guidance, select optimum objective forecast
aid, and expedite the issuance of tropical cyclone warnings. One great advantage of using ATCF is the standardization of the tropical cyclone forecasting procedure, so that during the course of the tropical cyclone warning preparation, forecasters will not neglect consideration of any decisional steps or available options. ATCF automatically saves all tropical cyclone data, computes the real-time and post-storm statistics, and allows forecasters to randomly access any past storm data. A communication package included in ATCF simplifies the data transfer procedure between JTWC and Fleet Numerical Oceanography Center in Monterey, CA.

The ATCF will be installed at JTWC in January 1988 for test and evaluation. Modifications on the system will be followed to make the system be compatible with the design of the JTWC Automation Project.

## North Pacific Tropical Cyclone Climatology

(Miller, R.J. and T.L. Tsui, NAVENVPREDRSCHFAC)

A tropical cyclone climatology for the North Pacific has been developed and now is being reviewed by EGPACOM. Data used for the western basin were taken from the JTWC Tropical Cyclone Data Base and covered a period of 40 years, 1945-84. Eastern basin data spanned the 34 -year period from 1949 to 1982 and were obtained from the consolidated worldwide tropical cyclone data base at National Climatic Data Center, Asheville, N.C. Storms for both basins were sorted according to month/day of the year into twenty four 31-day overlapping periods. For each period, four charts are supplied: 1) actual storm paths; 2) mean storm paths; 3) average storm speed; and 4) storm constancy and frequency.

JTWC has evaluated and offered suggestions for modifications of the climatology. The final version of the compilation should be completed in March 1988.

## EOF Post-Processing Forecast Technique

(Tsui, T.L. and J.H. Chu, NAVENVPREDRSCHFAC)

NEPRF is adapting the Empirical Orthogonal Function (EOF) post-processing tropical cyclone forecast scheme developed by Naval Postgraduate School (NPS) on the Fleet Numerical Oceanography Center computer system. The NPS EOF technique objectively
recognizes the salient patterns of large-scale horizontal wind fields with respect to the center of a tropical cyclone. This information, in terms of the EOF coefficients, will be used to modify the tropical cyclone track forecasts produced by the numerical models. The skill of this method is derived from the regression equations between the EOF coefficients and the forecast tracks of the One-way Tropical Cyclone Model (OTCM) in the western North Pacific during the period from 1979 to 1983.

## 2. JOINT TYPHOON WARNING CENTER

Joint Typhoon Warning Center Automation Project (JTWC-AP)

LT Brian J. Williams, USN, Typoon Duty Officer, JTWC Automation Officer.

A comprehensive effort is currently underway to provide JTWC with state-of-theart, automated tools to aid the Typhoon Duty Officer (TDO) in the collection, presentation, and analysis of data. These tools will also streamline the production of the warning messages and provide decision-making aids for the TDO. Automation of JTWC will take place in two phases. The first phase is the implementation in January 1988 of the Automated Tropical Cyclone Forecasting system (ATCF). The ATCF consists of a "suite" of program modules designed to run on IBM-AT compatible microcomputers. The concept and design of the ATCF (described above by Dr. Tsui and Mr. Miller) is a cooperative effort between NEPRF and JTWC. The second phase of automation will be the implementation of the more comprehensive JTWC-AP in FY 89. The JTWC-AP will integrate features of the ATCF with a more complete advanced data base archival and retrieval system, satellite imagery looping, overlay, and increased emphasis on expert systems that make the TDO's watch routine more efficient and effective.

The hardware suite that will run the ATCF programs (described above by Dr. Tsui and Mr. Ron Miller) has five workstations connected by a file server network to share common data files (see Figure 6-1). A dedicated
terminal will provide the send/receive interface with the Automated Weather Network (AWN). Numerical forecast aids, FNOC analyses and prognostic fields, as well as near-real time synoptic data (as a back-up to the AWN) will be received via remote requests over the TYMNET public data network. The TYMNET connects the JTWC microcomputers to FNOC mainframes. Outgoing messages to customers without access to AWN are inserted into the AUTODIN system via paper tape sent to the local Navy Telecommunications Command Center (NTCC). The ATCF software and hardware implementation represents the first step toward automation of JTWC.

A major feature of the future JTWC-AP will be the reference roster data base. This data base will contain critical data about customers in JTWC's AOR. It will include storm haven information, telephone points-ofcontact, notification criteria for threatened customers, geographical information, local area forecasting rules of thumb, weather reporting station locations, etc. Whenever a customer is threatened, the reference roster will automatically prompt the TDO with customerspecific information. JTWC is currently working to compile the data reference roster for the JTWC-AP project manager. The reference roster will be easily edited to add or delete information as conditions change. This feature should significantly improve the level of support to JTWC's customers.

Another important feature of the JTWCAP is a training or playback mode which will call up archived data to realistically recreate
previous forecast scenarios. This will be possible due to the integration of satellite imagery, numerical analyses, prognostic fields, and "raw" data into one data base "tagged" by time, geography, or event ( $e . g .$, a tropical cyclone). This feature will provide the ability to display, analyze and recreate the timing of receipt of all data that was available for a past storm. This will allow a controlled training environment, especially in the off-season, as well as an outstanding tool for forecast "bust" reviews. In addition, this function will provide a complete and rich data base for post-analysis, case studies, and other research.

Future plans for the JTWC-AP include the implementation of decision-making aids (such as decision trees developed at the Naval Postgraduate School) and expert systems to aid in forecasting genesis, motion, intensity and dissipation. The JTWC-AP will provide a comprehensive real-time and archived tropical cyclone data base as well as the tools to manipulate data. This system is expected to significantly improve JTWC's operational support, while providing an excellent means for studying, improving, and "fine tuning" tropical cyclone forecasting methods and operational procedures.


Figure 6-1. Physical layout of the ATCF in IT'WC's working spaces.

## ANNEX A

## 1. GENERAL

Due to the rapid growth of the use of microcomputers in the meteorological community and to save publishing costs, tropical cyclone track data (with best track, initial warning, 24-, 48- and 72-hour JTWC forecasts) and fix data (satellite, aircraft, radar and synoptic) are now available separately upon request. The data will be available in ASCII format on 5.25 inch "floppy" diskettes. The data sets are available on four diskettes, which include the western North Pacific and North Indian Ocean (1 January - 31 December 1987)

## 2. WARNING VERIFICATION STATISTICS

## a. WESTERN NORTH PACIFIC

This section includes verification statistics for each warning in the western North
on two and the South Pacific and South Indian Oceans (1 July 1986-30 June 1987) on the other two. Agencies or individuals desiring these data sets should send the appropriate number of "floppy" diskettes (four if both data sets are desired, two if only one desired) to NAVOCEANCOMCEN/JTWC Guam with their request. When the request is received, the data will be copied onto your diskettes and returned with an explanation of the data formats. The use of floppy diskettes should facilitate the transfer of these rather large data files into your computer.

Pacific Ocean during 1987. Pre- and postwarning best track positions are not printed, but are available on floppy diskettes by request.

Typhoon Orchid (01w)

|  | 00 h | 24 h | 48 h | 72 h |
| :--- | ---: | ---: | ---: | ---: |
| Average | 17 | 91 | 158 | 234 |
| (Cases | 23 | 19 | 15 | 11 |


| DTG | 埧 | BTIAT | BT LON | POS ER | 24_ER | 48 ER | $72 . E R$ | BT WN | WW ER | 24. WE | 48.WE | 72. WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87010812 | 1 | 7.2N | 150.6E | 37.7 | 37.6 | 126.7 | 204.5 | 35 | -10 | -5 | -25 | -20 |
| 87010818 | 2 | 7.2 N | 150.0 E | 65.7 | 53.6 | 130.7 | 168.2 | 35 | -5 | -10 | -35 | -10 |
| 87010900 | 3 | 7.3N | 149.2E | 6 | 152.2 | 252.5 | 242.1 | 40 | -5 | -10 | -40 | 5 |
| 87010906 | 4 | 7.5N | 148.0E | 18.8 | 191.9 | 266.9 | 206.2 | 40 | -5 | -20 | -35 | 5 |
| 87010912 | 5 | 7.9 N | 146.5E | 13.4 | 119.5 | 139.2 | 144 | 45 | -5 | -25 | -20 | 5 |
| 87010918 | 6 | 8.4N | 144.9E | 29.8 | 101.6 | 89 | 106.7 | 50 | -10 | -35 | -10 | 10 |
| 87011000 | 7 | 8.8N | 143.4 E | 13.3 | 49.4 | 60.9 | 147.2 | 55 | -5 | -25 | 20 | 40 |
| 87011006 | 8 | 9.2N | 141.9 E | 29.7 | 26.5 | 47.4 | 207 | 65 | -5 | -10 | 25 | 40 |
| 87011012 | 9 | 9.7N | 140.4 E | 5.9 | 50.4 | 151.1 | 330.5 | 75 | -10 | 5 | 20 | 25 |
| 87011018 | 10 | 10.0N | 139.1E | 13.4 | 75.6 | 187 | 430 | 85 | -15 | 20 | 25 | 25 |
| 87011100 | 11 | 10.2N | 137.9E | 5.9 | 11.8 | 93.5 | 390.8 | 95 | -5 | 45 | 40 | 40 |
| 87011106 | 12 | 10.5N | 136.8 E | 8.4 | 12 | 79.9 | N/A | 90 | 10 | 50 | 55 | N/A |
| 87011112 | 13 | 10.8N | 135.8 E | 6 | 35.8 | 159.7 | N/A | 80 | 20 | 35 | 45 | N/A |
| 87011118 | 14 | 10.9N | 134.9E | 5.9 | 70.7 | 260.1 | N/A | 70 | 25 | 20 | 35 | N/A |
| 87011200 | 15 | 11.1 N | 134.0E | 6 | 34.5 | 331.8 | N/A | 60 | 10 | 0 | 5 | N/A |
| 87011206 | 16 | 11.4N | 133.3E | 6 | 91.1 | N/A | N/A | 60 | 10 | 5 | N/A | N/A |
| 87011212 | 17 | 11.7N | 132.7E | 0 | 165.8 | N/A | N/A | 65 | 0 | 5 | N/A | N/A |
| 87011218 | 18 | 11.7N | 132.2E | 21.3 | 267.5 | N/A | N/A | 60 | 0 | 5 | N/A | N/A |
| 87011300 | 19 | 11.3N | 131.6E | 0 | 192.8 | N/A | N/A | 50 | 5 | 5 | N/A | N/A |
| 87011306 | 20 | 10.8N | 131.1 E | 13.4 | N/A | N/A | N/A | 45 | 5 | N/A | N/A | N/A |
| 87011312 | 21 | 10.1N | 130.7E | 12 | N/A | N/A | N/A | 40 | 5 | N/A | N/A | N/A |
| 87011318 | 22 | 8.7N | 130.4E | 78.9 | N/A | N/A | N/A | 35 | 5 | N/A | N/A | N/A |
| 87011400 | 23 | 6.8N | 129.9E | 13.4 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |


| Tropical Storm Percy (02W) |  | 00 h | 24 h | 48 h | 72 h |
| ---: | ---: | ---: | ---: | ---: | ---: |
|  | Average | 19 | 69 | 116 | 206 |
|  | U Cases | 9 | 8 | 4 | 4 |


| DTG | W1 | BT LAT | BT LON | POS ER | 24.ER | 48 ER | 72 ER | BT WN | HW ER | 24.WE | 48. WE | 72.WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87041100 | 1 | 9.3N | 142.9E | 8.4 | 43 | 92.5 | 163 | 40 | 0 | 20 | 25 | 30 |
| 87041106 | 2 | 9.8 N | 142.4 E | 24 | 81.5 | 30.6 | 100.6 | 40 | 0 | 20 | 25 | 30 |
| 87041112 | 3 | 10.3 N | 140.7E | 59.3 | 85.8 | 152.1 | 279.8 | 35 | 5 | 15 | 15 | 5 |
| 87041118 | 4 | 10.6 N | 141.0E | 18.9 | 70.6 | 191.8 | 280.2 | 30 | 10 | 10 | 10 | 5 |
| 87041200 | 5 | 10.8 N | 140.2E | 18 | 51.2 | N/A | N/A | 30 | 0 | 0 | N/A | N/A |
| 87041206 | 6 | 11.0 N | 139.3E | 13.4 | 50.2 | N/A | N/A | 25 | 5 | 0 | N/A | N/A |
| 87041212 | 7 | 11.2 N | 138.4E | 11.8 | 67 | N/A | N/A | 25 | 5 | 5 | N/A | N/A |
| 87041218 | 8 | $11.4 N$ | 137.7E | 6 | 102.6 | N/A | N/A | 25 | 5 | 0 | N/A | N/A |
| 87041300 | 9 | 11.8 N | 137.3E | 16.8 | N/A | N/A | N/A | 25 | 5 | N/A | N/A | N/A |




| Surer Typhoon Thelma (05w) |  |  | Average * Cases | $\begin{array}{r} 00 \mathrm{~h} \\ \hline 18 \\ 34 \end{array}$ | $\begin{array}{r} 24 \mathrm{~h} \\ 146 \\ 31 \end{array}$ | $\begin{array}{r} 48 \mathrm{~h} \\ 311 \\ 28 \end{array}$ | $\begin{array}{r} 72 \mathrm{~h} \\ 479 \\ 24 \end{array}$ | BT. WN | WW ER | 24 WE | 48.WE | 72.WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |  |  |
| DTG | W\# | BT LAT | BT LON | POS ER | $24.5 B$ | 48 EB | 72 ER |  |  |  |  |  |
| 87070718 | 1 | 13.0 N | 149.5E | 84.8 | 238.2 | 352.7 | 402.6 | 30 | 0 | 0 | -30 | -25 |
| 87070800 | 2 | 13.9 N | 148.9 E | 43.6 | 263.3 | 401 | 491.1 | 30 | 0 | -5 | -30 | -45 |
| 87070806 | 3 | 14.5 N | 147.4 E | 58.4 | 280.1 | 510 | 995.5 | 30 | 0 | 0 | -30 | -70 |
| 87070812 | 4 | 14.6 N | 145.5 E | 37.6 | 202.8 | 462.2 | 864.3 | 35 | -5 | -20 | -40 | -80 |
| 87070818 | 5 | 15.0N | 143.8E | 34.1 | 8.3 | 128.1 | 350.8 | 40 | -10 | -25 | -45 | -60 |
| 87070900 | 6 | 15.6 N | 142.3E | 18.9 | 139.9 | 318.5 | 636.9 | 45 | 5 | -20 | -45 | -55 |
| 87070906 | 7 | 16.2 N | 141.0E | 8.3 | 162.4 | 357.8 | 615.1 | 50 | 0 | -25 | -55 | -60 |
| 87070912 | 8 | 16.5N | 139.7E | 13.3 | 109.6 | 336.4 | 529.5 | 60 | -5 | -25 | -55 | -45 |
| 87070918 | 9 | 16.7 N | 138.3E | 11.5 | 127.8 | 250.1 | 187.5 | 70 | -15 | -35 | -55 | -25 |
| 87071000 | 10 | 17.0 N | 136.9 E | 6 | 116.8 | 271.7 | 234.3 | 80 | 0 | -10 | -5 | 30 |
| 87071006 | 11 | 17.3N | 135.4E | 18.9 | 165.2 | 292.1 | 258.2 | 85 | 0 | -15 | -5 | 35 |
| 87071012 | 12 | 17.7N | 134.0E | 8.3 | 181.4 | 324 | 288.5 | 90 | 0 | -25 | 0 | 35 |
| 87071018 | 13 | 17.8 N | 132.5E | 6 | 158.6 | 269.6 | 195.4 | 100 | -5 | -5 | 30 | 40 |
| 87071100 | 14 | 17.9N | 131.2E | 0 | 192.1 | 276.8 | 147.6 | 110 | -5 | 10 | 30 | 25 |
| 87071106 | 15 | 17.9N | 129.9E | 18 | 122.7 | 168.2 | 108.5 | 120 | -10 | 10 | 35 | 30 |
| 87071112 | 16 | 17.8 N | 128.8E | 8.3 | 123.3 | 189 | 212.4 | 130 | -10 | 25 | 45 | 25 |
| 87071118 | 17 | 17.6 N | 128.0E | 16.6 | 104.9 | 188.2 | 307 | 125 | -5 | 15 | 20 | 10 |
| 87071200 | 18 | 17.6 N | 127.3E | 6 | 163.5 | 421.1 | 675.7 | 120 | 0 | 15 | 15 | 20 |
| 87071206 | 19 | 17.9N | 126.8E | 18.9 | 222.7 | 485.7 | 820.2 | 120 | -5 | 15 | 10 | 15 |
| 87071212 | 20 | 18.5N | 126.6E | 8.3 | 141.3 | 382.1 | 762.7 | 115 | -10 | 5 | -5 | 5 |
| 87071218 | 21 | 19.2N | 126.4E | 29.4 | 190 | 418.3 | 857.9 | 105 | -5 | 0 | -5 | 20 |
| 87071300 | 22 | 19.9N | 126.2E | 8.2 | 92.3 | 192.6 | 540.2 | 95 | -5 | -10 | -10 | 20 |
| 87071306 | 23 | 20.9N | 126.0E | 18.9 | 67.1 | 196.9 | 497.6 | 90 | -5 | -15 | -15 | 15 |
| 87071312 | 24 | 22.1N | 125.6 E | 12.6 | 47.1 | 217.8 | 524.5 | 90 | -15 | -25 | -15 | 10 |
| 87071318 | 25 | 23.4N | 125.3E | 17.6 | 74 | 308 | N/A | 90 | -15 | -25 | -5 | N/A |
| 87071400 | 26 | 24.7N | 125.2E | 6 | 62.7 | 354.2 | N/A | 90 | -5 | -10 | 15 | N/A |
| 87071406 | 27 | 25.9N | 124.9E | 0 | 54.9 | 312 | N/A | 90 | -5 | -10 | 20 | N/A |
| 87071412 | 28 | 27.4N | 124.8E | 8 | 102.1 | 346.4 | N/A | 90 | -5 | -5 | 20 | N/A |
| 87071418 | 29 | 28.6N | 125.0E | 8 | 186.1 | N/A | N/A | 85 | -5 | 5 | N/A | N/A |
| 87071500 | 30 | 29.9N | 125.5E | 12 | 228.2 | N/A | N/A | 80 | 0 | 30 | N/A | N/A |
| 87071506 | 31 | 31.9N | 126.3E | 13 | 219.7 | N/A | N/A | 80 | -5 | 20 | N/A | N/A |
| 87071512 | 32 | 34.4N | 127.2E | 24 | N/A | N/A | N/A | 70 | 0 | N/A | N/A | N/A |
| 87071518 | 33 | 36.8N | 128.2E | 30 | N/A | N/A | N/A | 55 | 5 | N/A | N/A | N/A |
| 87071600 | 34 | 39.0 N | 129.4E | 101.3 | N/A | N/A | N/A | 40 | -5 | N/A | N/A | N/A |


| Typhoon Vernon (06) |  | 00h | 24h | 48h | 72h |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | Average | 33 | 119 | 180 | 224 |
|  | \% Cases | 21 | 18 | 14 | 10 |


| DTG | Win | ET IAT | BT IOX | POS ER | 24.EB | 48 mP | 12.58 | BT WN | MTE ER | 24. WF | 48 WE | 72. WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87071618 | 1 | 12.1N | 137.5E | 11.7 | 82.2 | 125.9 | 282.5 | 30 | 0 | 10 | 10 | -15 |
| 87071700 | 2 | 12.2N | 135.3E | 54.1 | 42.9 | 151 | 209.5 | 30 | 0 | 10 | 5 | -15 |
| 87071706 | 3 | 12.2N | 133.1E | 18.6 | 66.6 | 160.9 | 207.6 | 35 | -5 | 5 | 0 | -10 |
| 87071712 | 4 | 12.3N | 131.0E | 42.6 | 153.1 | 188.8 | 152.8 | 35 | -5 | 5 | -5 | 10 |
| 87071718 | 5 | 12.7N | 129.7E | 16.8 | 181.1 | 216.2 | 162.6 | 35 | -5 | 0 | -5 | 10 |
| 87071800 | 6 | 13.1N | 128.9E | 94.2 | 279.7 | 280.9 | 196.1 | 35 | -5 | -10 | -10 | 15 |
| 87071806 | 7 | 13.6 N | 128.5E | 112.3 | 212 | 207.6 | 219.4 | 40 | -10 | -20 | -15 | 5 |
| 87071812 | 8 | 14.1N | 128.1E | 34.2 | 139.2 | 151.2 | 112.7 | 45 | -5 | -10 | 10 | 25 |
| 87071818 | 9 | 15.0N | 127.5E | 58 | 140.9 | 159.8 | 260.9 | 50 | -10 | -10 | 10 | 30 |
| 87071900 | 10 | 15.9N | 126.7E | 23.9 | 125 | 234.8 | 442.9 | 55 | 0 | 0 | 25 | 60 |
| 87071906 | 11 | 16.9N | 125.9E | 42 | 103.4 | 214.5 | N/A | 60 | 0 | 10 | 35 | N/A |
| 87071912 | 12 | 18.0N | 125.3E | 12 | 84.7 | 81.4 | N/A | 65 | 0 | 20 | 50 | N/A |


| 87071918 | 13 | 18.7N | 124.7E | 24.8 | 75.7 | 128.8 | N/A | 65 | 0 | 20 | 55 | N/A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87072000 | 14 | 19.3N | 124.2E | 5.7 | 73.9 | 217.1 | N/A | 65 | 0 | 15 | 35 | N/A |
| 87072006 | 15 | 20.1N | 123.5E | 12.8 | 101.5 | N/A | N/A | 60 | 5 | 15 | N/A | N/A |
| 87072012 | 16 | 21.1N | 123.0E | 16.4 | 116.1 | N/A | N/A | 55 | 0 | 20 | N/A | N/A |
| 87072018 | 17 | 22.0 N | 122.5E | 20.6 | 110.1 | N/A | N/A | 55 | 0 | 25 | N/A | N/A |
| 87072100 | 18 | 23.1N | 122.2E | 8.2 | 66.8 | N/A | N/A | 50 | 5 | 30 | N/A | N/A |
| 87072106 | 19 | 24.6N | 121.8E | 32.4 | N/A | N/A | N/R | 45 | 5 | N/A | N/A | N/A |
| 87072112 | 20 | 26.0 N | 121.6 E | 36.9 | N/A | N/A | N/A | 35 | 5 | N/A | N/A | N/A |
| 87072118 | 21 | 27.3N | 121.6E | 22.1 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |




| 87081500 | 25 | 16.3 N | 110.8 E | 13 | 64 | 281.9 | N/A | 115 | -5 | 5 | 40 | N/A |
| :--- | :--- | :--- | :--- | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 87081506 | 26 | 16.7 N | 110.1 E | 0 | 74.5 | 343.9 | N/A | 105 | -5 | 10 | 30 | N/A |
| 87081512 | 27 | 17.0 N | 109.2 E | 5.7 | 104.3 | N/A | N/A | 95 | -5 | 15 | N/A | N/A |
| 87081518 | 28 | 17.3 N | 108.3 E | 0 | 116 | N/A | N/A | 90 | -10 | 20 | N/A | N/A |
| 87081600 | 29 | 17.5 N | 107.5 E | 0 | 189.3 | N/A | N/A | 80 | -10 | 15 | N/A | N/A |
| 87081606 | 30 | $17.7 N$ | 106.8 E | 0 | 278.3 | N/A | N/A | 65 | -5 | 15 | N/A | N/A |
| 87081612 | 31 | 18.0 N | 105.6 E | 23.6 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 87081618 | 32 | $18.1 N$ | 104.3 E | 62.4 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |




| 87082300 | 5 | $12.2 N$ | 157.1 E | 35.2 | 215.7 | 282.2 | 411.5 | 30 | -5 | 15 | 20 | 30 |
| ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 87082306 | 6 | 12.9 N | 156.7 E | 50.7 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 25 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87082618 | 7 | 16.1 N | 149.1 E | 37.8 | 178.6 | 138.5 | $\mathrm{~N} / \mathrm{A}$ | 35 | -5 | 15 | 35 | $\mathrm{~N} / \mathrm{A}$ |
| 87082700 | 8 | 16.4 N | 148.4 E | 54.3 | 165.6 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | 25 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87082706 | 9 | 17.2 N | 147.7 E | 61.1 | 152.9 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 30 | 5 | 20 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87082712 | 10 | 17.8 N | 146.8 E | 26.6 | 36.2 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 30 | 0 | 15 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87082718 | 11 | 18.2 N | 145.9 E | 18.9 | 92.4 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 30 | 0 | 20 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87082800 | 12 | 18.4 N | 145.2 E | 50.7 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 25 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |



| 87091512 | 44 | 28.6N | 139.9 E | 12.1 | 46.6 | N/A | N/A | 50 | 0 | -5 | N/A | N/A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87091518 | 45 | 29.6N | 140.0E | 8 | 43.2 | N/A | N/A | 50 | 0 | -5 | N/A | N/A |
| 87091600 | 46 | $30.4 N$ | 140.3 E | 5.2 | 26 | N/A | N/A | 45 | 0 | 0 | N/A | N/A |
| 87091606 | 47 | 31.1N | 140.7E | 16.5 | 75.1 | N/A | N/A | 45 | 0 | 10 | N/A | N/A |
| 87091612 | 48 | 31.9N | 141.4 E | 6 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 87091618 | 49 | 32.6 N | 142.3 E | 30 | N/A | N/A | N/A | 40 | 0 | N/A | N/A | N/A |
| 87091700 | 50 | 33.4 N | 143.0E | 18.7 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |


| Typhoon | ld | (14W) |  | 00 h | 24h | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 20 | 97 | 163 | 193 |  |  |  |  |  |
|  |  |  | * Cases | 24 | 20 | 18 | 14 |  |  |  |  |  |
| DTG | W | BT LAT | BT_ION | POS ER | 24 ER | 48 ER | 72 ER | BT WN | HW_ER | $24 . \mathrm{WE}$ | 48 WE | 72.WE |
| 87090418 | 1 | 16.3 N | 126.3E | 21.4 | 167.8 | 263.2 | 294.2 | 35 | -10 | -10 | -30 | -45 |
| 87090500 | 2 | 16.6 N | 126.6E | 71.3 | 184.2 | 262.6 | 265 | 35 | -5 | -15 | -30 | -45 |
| 87090506 | 3 | 16.9 N | 126.5E | 53.2 | 143.9 | 230.9 | 240.7 | 40 | -5 | 0 | -10 | -25 |
| 87090512 | 4 | 17.1N | 126.5 E | 51.8 | 122.3 | 155.9 | 94.1 | 45 | -5 | -10 | -15 | -30 |
| 87090518 | 5 | 17.3N | 126.5 E | 17.2 | 70.7 | 63.9 | 82.9 | 45 | -5 | -10 | -20 | -35 |
| 87090600 | 6 | 17.4N | 126.5E | 30.5 | 79.8 | 54.3 | 138.1 | 50 | -5 | -5 | -15 | -25 |
| 87090606 | 7 | 17.5 N | 126.5 E | 24.7 | 48.3 | 66.7 | 156.8 | 55 | -5 | -5 | -15 | -15 |
| 87090612 | 8 | 17.6 N | 126.4 E | 13.3 | 12 | 102.7 | 170 | 60 | -5 | -10 | -30 | -10 |
| 87090618 | 9 | 17.8 N | 126.4 E | 6 | 66.4 | 105.1 | 134.2 | 60 | -5 | -15 | -35 | -5 |
| 87090700 | 10 | 18.0 N | 126.3 E | 12.9 | 115.7 | 198.6 | 206.8 | 65 | -5 | -15 | -25 | 20 |
| 87090706 | 11 | 18.4N | 126.0E | 11.4 | 105.3 | 200.6 | 245.5 | 70 | -5 | -10 | -15 | 40 |
| 87090712 | 12 | 19.0N | 125.6 E | 12.8 | 100.3 | 186.7 | 232.6 | 75 | -5 | -20 | -5 | 70 |
| 87090718 | 13 | 19.5N | 125.1E | 8.2 | 87.5 | 138.7 | 208 | 80 | 5 | 0 | 35 | 70 |
| 87090800 | 14 | 20.0N | 124.7E | 13.3 | 78.7 | 155 | 241.8 | 85 | 5 | 0 | 45 | 65 |
| 87090806 | 15 | 20.2N | 124.0E | 13.3 | 106.4 | 178.7 | N/A | 90 | 0 | 5 | 40 | N/A |
| 87090812 | 16 | 20.4N | 123.2E | 12.7 | 74.7 | 212.7 | N/A | 100 | 0 | 20 | 70 | N/A |
| 87090818 | 17 | 20.6 N | 122.4 E | 0 | 78.2 | 156.3 | N/A | 105 | -5 | 0 | 30 | N/A |
| 87090900 | 18 | 20.9N | 121.9 E | 16.8 | 98 | 199.3 | N/A | 105 | 0 | 10 | 20 | N/A |
| 87090906 | 19 | 21.3N | 121.2 E | 8.2 | 100.2 | N/A | N/A | 100 | 0 | 20 | N/A | N/A |
| 87090912 | 20 | 21.7N | 120.5E | 8.2 | 111 | N/A | N/A | 90 | 0 | 40 | N/A | N/A |
| 87090918 | 21 | 22.5N | 120.0E | 16.6 | N/A | N/A | N/A | 80 | 0 | N/A | N/A | N/R |
| 87091000 | 22 | 23.2N | 119.5 E | 12.6 | N/A | N/A | N/A | 70 | -5 | N/A | N/A | N/A |
| 87091006 | 23 | 24.1N | 119.0E | 12.5 | N/A | N/A | N/A | 55 | 0 | N/A | N/A | N/A |
| 87091012 | 24 | 25.0N | 118.5E | 12.4 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |


| Super Ty | n | (15w) |  | 00 h | 24h | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 23 | 122 | 275 | 455 |  |  |  |  |  |
|  |  |  | * Cases | 43 | 34 | 33 | 31 |  |  |  |  |  |
| DTG | W1 | BT LAT | BT LON | POS ER | 24E8 | 48 ER | 72.EB | BT WN | HW ER | 24-WE | 48 WE | 72. WE |
| 87090506 | 1 | 12.2N | 168.3 E | 11.7 | 60.6 | 65.1 | 78.9 | 30 | 0 | -20 | -45 | -55 |
| 87090512 | 2 | 12.2N | 167.8E | 34.8 | 113.3 | 95.8 | 106.1 | 35 | -5 | -25 | -50 | -55 |
| 87090518 | 3 | 12.3 N | 167.1E | 56.2 | 130.7 | 108.3 | 131.9 | 45 | -10 | -25 | -50 | -55 |
| 87090600 | 4 | 12.4N | 166.2 E | 12 | 72.2 | 183.6 | 306.1 | 45 | 0 | -15 | -30 | -25 |
| 87090606 | 5 | 12.6 N | 165.5E | 37.1 | 120.2 | 258.9 | 421 | 55 | -5 | -20 | -30 | -30 |
| 87090612 | 6 | 12.9N | 164.6 E | 6 | 49.4 | 150.1 | 333.4 | 65 | -5 | -30 | -35 | -50 |
| 87090618 | 7 | 13.3N | 163.7E | 26.3 | 110.5 | 221.6 | 428.9 | 70 | -5 | -20 | -25 | -35 |
| 87090700 | 8 | 13.8N | 163.0E | 26.7 | 106.6 | 211.9 | 423.2 | 80 | -10 | -25 | -25 | -20 |
| 87090706 | 9 | 14.3N | 162.2E | 21.4 | 112.1 | 225.6 | 466.1 | 90 | -5 | -10 | -15 | 0 |
| 87090712 | 10 | 14.8N | 161.4E | 32.2 | 122.4 | 269.6 | 569.3 | 100 | -5 | 0 | -10 | 10 |
| 87090718 | 11 | 15.3N | 160.7E | 25 | 124.7 | 313.9 | 643.6 | 105 | 0 | 0 | -5 | 15 |
| 87090800 | 12 | 15.8N | 159.9E | 8.3 | 67.6 | 310.9 | 675.3 | 110 | -5 | 0 | 0 | 10 |
| 87090806 | 13 | 16.3N | 159.2E | 13 | 72.5 | 344.4 | 737.5 | 115 | -5 | -5 | 5 | 15 |


| 87090812 | 14 | 16.8N | 158.4E | 5.7 | 109.6 | 413.6 | 810.4 | 115 | 0 | -10 | 10 | 20 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87090818 | 15 | 17.3N | 157.6E | 5.7 | 173.2 | 439.5 | 755.5 | 120 | -5 | -10 | 10 | 30 |
| 87090900 | 16 | 17.8 N | 156.8E | 13.3 | 191.8 | 490.7 | 757 | 125 | 0 | 0 | 15 | 40 |
| 87090906 | 17 | 18.3N | 156.2E | 23.6 | 192.3 | 505.2 | 743.1 | 135 | -5 | 10 | 20 | 45 |
| 87090912 | 18 | 18.6N | 155.9E | 21.3 | 135.6 | 392.1 | 568.8 | 140 | -5 | 20 | 30 | 65 |
| 87090918 | 19 | 18.9N | 155.8 E | 26.6 | 144 | 350 | 506.4 | 140 | -5 | 15 | 35 | 65 |
| 87091000 | 20 | 19.4N | 155.6E | 18 | 132.9 | 292.1 | 395.5 | 135 | 5 | 15 | 35 | 35 |
| 87091006 | 21 | 19.9 N | 155.5 E | 12.8 | 134.8 | 242.3 | 376.7 | 130 | 5 | 10 | 30 | 25 |
| 87091012 | 22 | 20.5N | 155.5 E | 12 | 121.9 | 233.6 | 378.7 | 125 | 5 | 15 | 45 | 45 |
| 87091018 | 23 | 21.2N | 155.6E | 8.2 | 118.7 | 146.4 | 288.6 | 125 | 5 | 15 | 40 | 45 |
| 87091100 | 24 | 22.0 N | 155.8 E | 8.2 | 76 | 168.1 | 340.3 | 120 | 10 | 10 | 10 | 15 |
| 87091106 | 25 | 23.1N | 156.0E | 6 | 113.8 | 442.1 | 445.8 | 115 | 10 | 15 | 5 | 10 |
| 87091112 | 26 | 24.0N | 156.0E | 12 | 161.6 | 334.8 | 495.1 | 110 | 5 | 20 | 10 | 15 |
| 87091118 | 27 | 24.8N | 155.7E | 13.2 | 159.6 | 314.9 | 482 | 100 | 5 | 15 | 15 | 20 |
| 87091200 | 28 | 25.6N | 155.3E | 18 | 170.4 | 334.3 | 448.3 | 90 | 5 | 10 | 20 | 20 |
| 87091206 | 29 | 26.0N | 155.2E | 10.8 | 117.4 | 264.7 | 330.9 | 80 | 10 | 5 | 15 | 20 |
| 87091212 | 30 | 26.4 N | 155.2E | 12.3 | 131.5 | 263.2 | 334.1 | 70 | 15 | 15 | 20 | 15 |
| 87091218 | 31 | 26.7N | 155.4 E | 24.1 | 138.7 | 234.3 | 349.8 | 65 | 10 | 15 | 20 | 15 |
| 87091300 | 32 | 26.9N | 155.6E | 16.1 | 102.1 | 157.1 | N/A | 65 | 0 | 10 | 15 | N/A |
| 87091306 | 33 | 27.2N | 156.0E | 144 | 135.5 | 297.9 | N/A | 65 | 0 | 5 | 10 | N/A |
| 87091312 | 34 | 27.4N | 156.5E | 55.7 | 125.4 | N/A | N/A | 55 | 0 | 5 | N/A | N/A |
| 87091318 | 35 | 27.7N | 157.1E | 27.2 | N/A | N/A | N/A | 50 | 0 | N/A | N/A | N/A |
| 87091400 | 36 | 28.0N | 157.5E | 49.1 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 87091406 | 37 | 28.3N | 157.9E | 51.7 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 87091412 | 38 | 28.6 N | 158.2E | 31.9 | N/A | N/A | N/A | 35 | 10 | N/A | N/A | N/A |
| 87091418 | 39 | 28.9N | 158.3E | 12.1 | N/A | N/A | N/A | 30 | 10 | N/A | N/A | N/A |
| 87091500 | 40 | 29.2N | 158.5E | 16.8 | N/A | N/A | N/A | 30 | 5 | N/A | N/A | N/A |
| 87091506 | 41 | 29.5N | 158.5E | 20.8 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |
| 87091512 | 42 | 29.8 N | 158.6E | 5.2 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |
| 87091518 | 43 | 30.0 N | 158.7E | 15.9 | N/A | N/A | N/A | 25 | 0 | N/A | N/A | N/A |



| 87092812 | 22 | 23.6N | 138.0E | 5.5 | 30.5 | 88.3 | N/A | 70 | -5 | -10 | -5 | N/A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87092818 | 23 | 23.7 N | 138.0 E | 12 | 24.6 | 72.7 | N/A | 65 | -5 | -10 | -10 | N/A |
| 87092900 | 24 | 23.8N | 137.8 E | 21.1 | 8.1 | N/A | N/A | 65 | -5 | -15 | N/A | N/A |
| 87092906 | 25 | 23.9N | 137.5E | 11 | 31.9 | N/A | N/A | 65 | -10 | -20 | N/A | N/A |
| 87092912 | 26 | 23.9 N | 137.3E | 8.1 | 24.9 | N/A | N/A | 65 | -10 | -15 | N/A | N/A |
| 87092918 | 27 | 24.0N | 137.1E | 17.5 | N/A | N/A | N/A | 60 | -10 | N/A | N/A | N/A |
| 87093000 | 28 | 24.2N | 136.9 E | 8.1 | 149.4 | N/A | N/A | 60 | -15 | -20 | N/A | N/A |
| 87093006 | 29 | 24.5N | 136.8 E | 12 | N/A | N/A | N/A | 60 | -20 | N/A | N/A | N/A |
| 87093012 | 30 | 24.8N | 136.7E | 6 | N/A | N/A | N/A | 55 | -15 | N/A | N/A | N/A |
| 87093018 | 31 | 25.3 N | 137.1E | 12.4 | N/A | N/A | N/A | 55 | -20 | N/A | N/A | N/A |
| 87100100 | 32 | 26.1N | 137.9E | 12 | N/A | N/A | N/A | 55 | -20 | N/A | N/A | N/A |
| 87100106 | 33 | 26.9N | 138.3E | 18.8 | N/A | N/A | N/A | 55 | -25 | N/A | N/A | N/A |

Tropical Dopression 17W

|  | 00 h | 24 h | 48 h | 72 h |
| :--- | ---: | ---: | ---: | ---: |
| Average | 21 | 81 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| Cases | 7 | 3 | 0 | 0 |


| DTG | W迷 | BT LAT | BT ION | Pos EB | $24 . E R$ | 48 ER | 72.ER |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87092418 | 1 | 17.0 N | 161.3 E | 68.9 | 155.5 | N/A | N/A |
| 87092500 | 2 | 17.0N | 160.7E | 17.2 | 45.7 | N/A | N/A |
| 87092506 | 3 | 16.9N | 160.1E | 8.3 | 43.5 | N/A | N/A |
| 87092512 | 4 | 16.8 N | 159.4E | 13 | N/A | N/A | N/A |
| 87092518 | 5 | 16.8N | 158.7E | 8.3 | N/A | N/A | N/A |
| 87092600 | 6 | 16.7 N | 157.9E | 23.8 | N/A | N/A | N/A |
| 87092606 | 7 | 16.4N | 157.3E | 8.3 | N/A | N/A | N/A |
| Tropical | mim | (18w) |  | 00h | 24h | 48h | 72h |
|  |  |  | Average | 33 | 165 | 66 | N/A |
|  |  |  | * Cases | 9 | 2 | 2 | $0$ |


| DTG | W1 | BT LAT | BT LON | POS ER | 24 ER | 48 ER | 72.ER | BT_WN | WW ER | 24.WE | 48.WE | 72.WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87092900 | 1 | 23.7N | 155.8E | 18.8 | 169.8 | 84.7 | N/A | 35 | 0 | 5 | 0 | N/A |
| 87092906 | 2 | 25.2N | 155.0E | 54 | 161.4 | 47.6 | N/A | 40 | -5 | 5 | 5 | N/A |
| 87092912 | 3 | 26.8N | 154.1E | 49.2 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |
| 87092918 | 4 | 28.0N | 152.3E | 40.1 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |
| 87093000 | 5 | 28.2N | 149.7E | 29.1 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/R |
| 87093006 | 6 | 28.1N | 148.7E | 63.5 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |
| 87093012 | 7 | 28.4N | 148.6E | 5.3 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 87093018 | 8 | 29.2N | 148.7E | 31.8 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/R |
| 87100100 | 9 | 29.9N | 149.1E | 7.9 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |



| 87101206 | 10 | 16.8N | 137.2E | 23 | 41.5 | 94.6 | 92.2 | 65 | 0 | 10 | 20 | 25 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87101212 | 11 | 17.3N | 137.2E | 8.3 | 24.7 | 117.3 | 148.2 | 70 | 0 | 20 | 20 | 20 |
| 87101218 | 12 | 17.8 N | 137.1E | 6 | 47.7 | 147.8 | 169.1 | 70 | 5 | 20 | 25 | 25 |
| 87101300 | 13 | 18.3N | 136.8E | 8.3 | 76.8 | 174 | 182.7 | 75 | 5 | 10 | 20 | 0 |
| 87101306 | 14 | 18.9N | 136.4 E | 8.3 | 44.9 | 92.1 | 177.9 | 75 | 0 | -15 | -35 | -45 |
| 87101312 | 15 | 19.5 N | 136.1 E | 8.2 | 88.5 | 103.3 | 276.8 | 75 | 0 | -20 | -40 | -35 |
| 87101318 | 16 | 20.2N | 135.7E | 12 | 102.8 | 92.5 | 335.6 | 80 | -5 | -20 | -40 | -25 |
| 87101400 | 17 | 21.3 N | 135.2E | 8.2 | 97.9 | 88 | 336.9 | 80 | -5 | -20 | -30 | 0 |
| 87101406 | 18 | 22.1N | 134.5E | 16.4 | 80.1 | 88.2 | N/A | 80 | -5 | -20 | -30 | N/A |
| 87101412 | 19 | 23.0 N | 133.7E | 13.2 | 90.7 | 193.1 | N/A | 85 | -10 | -25 | -25 | N/A |
| 87101418 | 20 | 24.0N | 133.0E | 12.5 | 53.1 | 216.7 | N/A | 85 | -10 | -30 | -20 | N/A |
| 87101500 | 21 | 25. 0 N | 132.2E | 5.4 | 38.7 | 294.4 | N/A | 90 | -5 | -15 | 10 | N/A |
| 87101506 | 22 | 25.8N | 131.9E | 6 | 151.6 | N/A | N/A | 90 | -5 | -15 | N/A | N/A |
| 87101512 | 23 | 26.7N | 131.8 E | 6 | 126.1 | N/A | N/A | 95 | -15 | -10 | N/A | N/A |
| 87101518 | 24 | 27.7N | 131.9 E | 10.6 | 211.5 | N/A | N/A | 95 | -15 | -5 | N/A | N/A |
| 87101600 | 25 | 29.0 N | 132.4 E | 28.8 | 279.7 | N/A | N/A | 90 | -20 | 5 | N/A | N/A |
| 87101606 | 26 | 30.4N | 133.0 E | 16.6 | N/A | N/R | N/A | 90 | -30 | N/A | N/A | N/A |
| 87101612 | 27 | 32.2N | 133.4 E | 19.4 | N/A | N/A | N/A | 80 | -25 | N/A | N/A | N/A |
| 87101618 | 28 | 34.2N | 134.4E | N/A | N/A | N/A | N/A | 70 | N/ A | N/A | N/A | N/A |


| Super Typhoon Lymn (20W) |  | 00 h | 24 h | 48h | 72h |
| :--- | :--- | ---: | ---: | ---: | ---: | ---: |
|  | Average | 17 | 89 | 184 | 298 |
|  | \# Cases | 44 | 41 | 39 | 31 |


| DTG | WH | BT Lat | BTION | POS ER | 24.ER | 48 ER | $72 . E 8$ | BT WN | WW ER | 24. WE | 48 WE | 72 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87101606 | 1 | 13.0N | 155.3E | 35.1 | 238.8 | 341.6 | 412.8 | 35 | 0 | -10 | -15 | -30 |
| 87101612 | 2 | 13.2N | 153.9E | 24.1 | 147.2 | 170.7 | 248.3 | 35 | 0 | -5 | -5 | -20 |
| 87101618 | 3 | 13.3N | 153.2E | 13.1 | 72.6 | 68.1 | 164.5 | 40 | 0 | -5 | -5 | -25 |
| 87101700 | 4 | 13.3N | 152.3E | 16.7 | 36.5 | 114.1 | 126.1 | 45 | 0 | 0 | -5 | -30 |
| 87101706 | 5 | 13.3 N | 151.3E | 8.4 | 76.7 | 144.5 | 133.1 | 55 | -10 | -5 | -15 | -30 |
| 87101712 | 6 | 13.2N | 150.1E | 23.4 | 100.9 | 157 | 145.8 | 55 | -5 | -10 | -30 | -30 |
| 87101718 | 7 | 13.3N | 148.9 E | 8.4 | 99 | 166.6 | 190.9 | 60 | -5 | -10 | -35 | -25 |
| 87101800 | 8 | 13.6N | 147.6E | 12 | 146.8 | 223.4 | 251.2 | 65 | -5 | -10 | -45 | -25 |
| 87101806 | 9 | 14.0 N | 146.6 E | 5.8 | 26 | 12.9 | 53.7 | 70 | -5 | -15 | -40 | -5 |
| 87101812 | 10 | 14.6N | 145.7E | 13.1 | 34.5 | 37.3 | 89 | 75 | -5 | -20 | -35 | -15 |
| 87101818 | 11 | 15.1N | 145.0E | 11.6 | 6 | 13.3 | 131.3 | 80 | -10 | -35 | -40 | -10 |
| 87101900 | 12 | 15.7N | 144.3 E | 16.7 | 40.1 | 98 | 245.8 | 90 | -5 | -40 | -30 | 20 |
| 87101906 | 13 | 16.2 N | 143.5E | 8.3 | 26.6 | 149.4 | 353.7 | 100 | -10 | -35 | -10 | 25 |
| 87101912 | 14 | 16.6 N | 142.7E | 6 | 39.9 | 145.7 | 256.6 | 115 | -20 | -25 | -5 | 20 |
| 87101918 | 15 | 16.9 N | 141.8 E | 8.3 | 49.6 | 173.1 | 280.2 | 125 | -10 | -5 | 20 | 25 |
| 87102000 | 16 | 17.3N | 141.1E | 13.3 | 107.5 | 290.9 | 472.6 | 140 | 0 | 5 | 40 | 40 |
| 87102006 | 17 | 17.6N | 140.2E | 18.9 | 148.7 | 332.3 | 530.4 | 140 | 0 | 15 | 25 | 25 |
| 87102012 | 18 | 17.8N | 139.2E | 11.4 | 146.5 | 378.9 | 559 | 140 | 0 | 15 | 25 | 20 |
| 87102018 | 19 | 18.0N | 138.0E | 17.1 | 168.1 | 392.4 | 558.7 | 140 | -5 | 10 | 20 | 10 |
| 87102100 | 20 | 18.0 N | 136.7E | 30.9 | 145.9 | 342.6 | 466.6 | 140 | -10 | 20 | 20 | 5 |
| 87102106 | 21 | 18.0N | 135.3 E | 24.8 | 140.4 | 308.5 | 395.7 | 125 | 0 | 15 | 15 | 5 |
| 87102112 | 22 | 18.0N | 134.0E | 11.4 | 81.8 | 117.9 | 118.9 | 125 | -5 | 10 | 5 | 5 |
| 87102118 | 23 | 18.0N | 132.5E | 26.6 | 74.4 | 34.5 | 146.7 | 115 | 0 | 10 | 0 | 5 |
| 87102200 | 24 | 18.0N | 130.9E | 8.3 | 36.4 | 99.6 | 270.2 | 100 | 5 | 5 | -15 | -5 |
| 87102206 | 25 | 18.0 N | 129.3E | 5.7 | 62.4 | 162.2 | 347.7 | 100 | 0 | -10 | -25 | -20 |
| 87102212 | 26 | 18.1N | 127.8E | 6 | 45.9 | 198 | 393.6 | 100 | -5 | -20 | -15 | -15 |
| 87102218 | 27 | 18.1N | 126.3E | 12.9 | 97.8 | 297.3 | 530.8 | 95 | -5 | -25 | -10 | -10 |
| 87102300 | 28 | 18.3N | 124.9E | 12.9 | 113.6 | 315.8 | 514 | 90 | 0 | -25 | -10 | -10 |
| 87102306 | 29 | 18.5N | 123.6E | 6 | 107.5 | 260 | 391 | 90 | 0 | -20 | -5 | 5 |
| 87102312 | 30 | 18.8N | 122.5 E | 11.4 | 118.5 | 235.5 | N/A | 90 | -5 | -10 | -5 | N/A |
| 87102318 | 31 | 19.1N | 121.6E | 17 | 118.2 | 257.5 | N/A | 90 | -5 | 0 | 5 | N/A |
| 87102400 | 32 | 19.4N | 120.9E | 12.8 | 75.3 | 174.2 | 229.4 | 90 | -5 | 0 | 10 | 5 |


| 87102406 | 33 | 19.7 N | 120.3 E | 16.5 | 98.4 | 178.9 | 230.8 | 85 | -5 | 0 | 15 | 5 |
| ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 87102412 | 34 | 20.0 N | 119.9 E | 0 | 78.6 | 149.1 | $\mathrm{~N} / \mathrm{A}$ | 75 | 0 | 5 | 15 | $\mathrm{~N} / \mathrm{A}$ |
| 87102418 | 35 | 20.3 N | 119.6 E | 30.6 | 126.4 | 196.6 | $\mathrm{~N} / \mathrm{A}$ | 65 | 5 | 10 | 0 | $\mathrm{~N} / \mathrm{A}$ |
| 87102500 | 36 | 20.6 N | 119.3 E | 28.7 | 95.6 | 147.4 | $\mathrm{~N} / \mathrm{A}$ | 65 | 0 | 5 | 0 | $\mathrm{~N} / \mathrm{A}$ |
| 87102506 | 37 | 20.9 N | 119.1 E | 24.6 | 17.8 | 68.1 | $\mathrm{~N} / \mathrm{A}$ | 60 | 5 | 10 | 5 | $\mathrm{~N} / \mathrm{A}$ |
| 87102512 | 38 | 21.1 N | 119.0 E | 12.7 | 69.7 | 179.4 | $\mathrm{~N} / \mathrm{A}$ | 55 | 5 | 10 | 5 | $\mathrm{~N} / \mathrm{A}$ |
| 87102518 | 39 | 21.3 N | 118.9 E | 12.7 | 91 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 45 | 10 | 5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87102600 | 40 | 21.5 N | 118.7 E | 12.7 | 68.7 | 49.3 | $\mathrm{~N} / \mathrm{A}$ | 45 | 0 | -5 | -5 | $\mathrm{~N} / \mathrm{A}$ |
| 87102606 | 41 | 21.7 N | 118.2 E | 23.1 | 73.4 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 40 | -5 | -10 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87102612 | 42 | 21.6 N | 117.8 E | 66.9 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | -5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87102618 | 43 | 21.5 N | 117.6 E | 73.5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | -5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87102700 | 44 | 21.3 N | 117.3 E | 11.2 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | -10 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |


| Tropical | Storm Maury (21w) |  | 00h |  | 24h | 48h | 72 h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 27 | 107 | 162 | 183 |  |  |  |  |  |
|  |  |  | \# Cases | 29 | 19 | 13 | 12 |  |  |  |  |  |
| DTG |  | bT IAt | BT ION | POS ER | $24 . E R$ | 48 EB | 72.ER | BT WN | WWLEE | 24.WE | 48.WE | 72 WE |
| 87111106 | 1 | 14.7N | 133.5E | 8.3 | 133.4 | 217.2 | 339.6 | 30 | 0 | 25 | 35 | 30 |
| 87111112 | 2 | 15.0 N | 132.9 E | 23.2 | 141.9 | 223.6 | 279.7 | 30 | 0 | 30 | 35 | 45 |
| 87111118 | 3 | 15.0N | 132.4E | 13.3 | 159.7 | 288.9 | 324.9 | 25 | 5 | 15 | 20 | 35 |
| 87111200 | 4 | 14.5 N | 132.3E | 42 | N/A | N/A | N/A | 20 | 5 | N/A | N/A | N/A |
| 87111306 | 5 | 13.7 N | 127.6E | 18.9 | 92.6 | 151.9 | 171.5 | 30 | 0 | 15 | 15 | 20 |
| 87111312 | 6 | 13.6 N | 126.9E | 66.7 | 221.7 | 349.7 | 395.1 | 30 | 0 | 10 | 20 | 15 |
| 87111318 | 7 | 13.5 N | 126.2E | 88.4 | 221.8 | 304.3 | 317.1 | 30 | 0 | 0 | 10 | 5 |
| 87111400 | 8 | 13.5 N | 125.6E | 5.8 | 25.1 | 88.5 | 141.8 | 30 | 0 | 5 | 0 | 0 |
| 87111406 | 9 | 13.5 N | 125.0E | 16.7 | 83.7 | N/A | N/A | 25 | 0 | 0 | N/A | N/A |
| 87111412 | 10 | 13.5 N | 124.3E | 8.4 | N/A | N/A | N/A | 25 | 0 | N/A | N/A | N/A |
| 87111418 | 11 | 13.4 N | 123.6E | 26.2 | N/A | N/A | N/A | 25 | 0 | N/A | N/A | N/A |
| 87111500 | 12 | 13.2N | 122.7E | 13.1 | N/A | N/A | N/A | 25 | 0 | N/A | N/A | N/A |
| 87111506 | 13 | 13.0N | 122.0E | 21.3 | 46.8 | 111.6 | 50.1 | 25 | 5 | 10 | 0 | 20 |
| 87111512 | 14 | 12.8 N | 121.0E | 13.1 | 76.9 | 102.2 | 50.7 | 25 | 5 | 5 | 10 | 25 |
| 87111518 | 15 | 12.8 N | 119.9 E | 12 | 76.8 | 47.4 | 42.4 | 25 | 5 | 0 | 10 | 30 |
| 87111600 | 16 | 12.8 N | 118.6 E | 21.3 | 87.1 | 52.9 | 40 | 30 | 0 | 0 | 10 | 20 |
| 87111606 | 17 | 12.9N | 117.3E | 12 | 72.6 | 62.9 | 47.3 | 30 | 0 | -5 | 15 | 10 |
| 87111612 | 18 | 13.3 N | 116.1E | 30.6 | 42.5 | 113.8 | N/A | 35 | 0 | 5 | 0 | N/A |
| 87111618 | 19 | 13.5N | 115.1E | 12 | 78.8 | N/A | N/A | 40 | -5 | 0 | N/A | N/A |
| 87111700 | 20 | 13.5 N | 114.2E | 29.5 | 144.3 | N/A | N/A | 40 | -5 | 5 | N/A | N/A |
| 87111706 | 21 | 13.4N | 113.3E | 72.5 | 200.5 | N/A | N/A | 45 | -10 | -5 | N/A | N/A |
| 87111712 | 22 | 13.0 N | 112.8 E | 110.5 | N/A | N/A | N/A | 40 | -5 | N/A | N/A | N/A |
| 87111718 | 23 | 13.0N | 112.3 E | 30.6 | N/A | N/A | N/A | 40 | -10 | N/A | N/A | N/A |
| 87111800 | 24 | 13.0 N | 111.9E | 39.4 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |
| 87111806 | 25 | 13.0 N | 111.5E | 8.4 | 42.6 | N/A | N/A | 35 | 5 | 5 | N/A | N/A |
| 87111812 | 26 | 12.9 N | 110.9E | 5.8 | 88.7 | N/A | N/A | 35 | 5 | 10 | N/A | N/A |
| 87111818 | 27 | 12.8 N | 110.3 E | 8.4 | N/A | N/A | N/A | 30 | 5 | N/A | N/A | N/A |
| 87111900 | 28 | 12.7N | 109.6E | 21.3 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |
| 87111906 | 29 | 12.3N | 108.8E | 18.9 | N/A | N/A | N/A | 25 | 0 | N/A | N/A | N/A |



| 87112006 | 4 | 5.8 N | 154.6E | 8.5 | 182.6 | 353 | 466.7 | 45 | 0 | 5 | 25 | 25 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87112012 | 5 | 6.5 N | 152.9E | 13.3 | 114 | 224.7 | 281.4 | 50 | -5 | 0 | 15 | 25 |
| 87112018 | 6 | 7.2N | 151.3E | 6 | 148.8 | 222.5 | 236.7 | 60 | -15 | 0 | 10 | 20 |
| 87112100 | 7 | 8.0 N | 149.6E | 18.8 | 148.6 | 257.9 | 240.5 | 60 | -5 | 5 | 5 | 25 |
| 87112106 | 8 | 8.7N | 147.8E | 30.3 | 129.7 | 206.1 | 155.3 | 60 | -5 | 0 | 0 | 15 |
| 87112112 | 9 | 9.5N | 146.0E | 13.4 | 81 | 180.3 | 194.8 | 65 | -10 | -5 | -5 | 0 |
| 87112118 | 10 | 10.0 N | 144.3 E | 29.5 | 34.8 | 78.9 | 156.1 | 65 | -5 | 0 | 0 | -5 |
| 87112200 | 11 | 10.3N | 142.7E | 8.4 | 64.5 | 114.2 | 162.1 | 65 | 0 | -5 | 5 | -15 |
| 87112206 | 12 | 10.7N | 141.0E | 0 | 84.2 | 148.2 | 206.3 | 70 | 0 | -5 | 0 | -20 |
| 87112212 | 13 | 11.0 N | 139.3 E | 12 | 69.5 | 150.5 | 235.1 | 75 | 0 | 5 | 5 | -30 |
| 87112218 | 14 | 11.1N | 137.8E | 24.7 | 56.8 | 146.9 | 213.7 | 80 | 5 | 10 | 0 | -15 |
| 87112300 | 15 | 11.2N | 136.1E | 6 | 100.5 | 165 | 115.3 | 85 | 5 | 20 | -10 | 5 |
| 87112306 | 16 | 11.5 N | 134.6E | 13.4 | 137 | 180.2 | 77.6 | 90 | 0 | 0 | -35 | -15 |
| 87112312 | 17 | 11.8 N | 133.4E | 13.2 | 76.9 | 63 | 110.9 | 90 | 0 | -20 | -70 | -30 |
| 87112318 | 18 | 11.9 N | 132.5E | 6 | 42 | 97.4 | 179.6 | 95 | -5 | -30 | -55 | -35 |
| 87112400 | 19 | 11.9 N | 131.6E | 5.9 | 36.5 | 122.2 | 219.5 | 90 | 0 | -30 | -30 | -35 |
| 87112406 | 20 | 12.0 N | 130.6E | 6 | 46.4 | 187 | 276.7 | 100 | -10 | -55 | -30 | -35 |
| 87112412 | 21 | 12.1 N | 129.6E | 8.4 | 63.2 | 205 | 304.7 | 105 | -20 | -70 | -30 | -30 |
| 87112418 | 22 | 12.3 N | 128.5E | 5.9 | 111.7 | 241.6 | 290.7 | 115 | -20 | -30 | -20 | -10 |
| 87112500 | 23 | 12.6N | 127.4E | 18.6 | 181.3 | 329.1 | 309.1 | 125 | -30 | -10 | -15 | 0 |
| 87112506 | 24 | 12.8 N | 126.1E | 5.9 | 188 | 333.5 | 264.5 | 135 | -20 | 10 | 5 | 10 |
| 87112512 | 25 | 13.0 N | 124.5E | 13.1 | 108.8 | 199.4 | 281.2 | 145 | -10 | 5 | 15 | 30 |
| 87112518 | 26 | 13.2 N | 122.8E | 13.1 | 92.8 | 228.5 | 356.1 | 130 | 0 | 5 | 10 | 30 |
| 87112600 | 27 | 13.5 N | 121.0E | 8.4 | 106.9 | 162.6 | 205.1 | 110 | 0 | -5 | 5 | 45 |
| 87112606 | 28 | 14.0 N | 119.3E | 13.1 | 99.8 | 156.9 | 240.9 | 95 | 0 | 0 | -10 | 45 |
| 87112612 | 29 | 14.6 N | 117.8E | 6 | 72.2 | 137.3 | 361.9 | 95 | -10 | -5 | -15 | 10 |
| 87112618 | 30 | 15.2N | 116.4E | 5.8 | 96.7 | 215.9 | 400 | 100 | 0 | 5 | 10 | 10 |
| 87112700 | 31 | 15.9N | 115.1E | 6 | 195.8 | 539.8 | N/A | 100 | 0 | -5 | 5 | N/A |
| 87112706 | 32 | 16.7N | 113.8E | 8.3 | 249.3 | N/A | N/A | 95 | 5 | -15 | N/A | N/A |
| 87112712 | 33 | 17.5 N | 113.0E | 0 | 311.1 | N/A | N/A | 95 | -5 | -20 | N/A | N/A |
| 87112718 | 34 | 18.6 N | 112.7E | 5.7 | 238.9 | N/A | N/A | 95 | 0 | 10 | N/A | N/A |
| 87112800 | 35 | 19.3 N | 112.8E | 5.7 | 226.2 | 448.2 | N/R | 95 | 5 | 30 | 20 | N/A |
| 87112806 | 36 | 19.8N | 113.1E | 8.2 | 301.8 | N/A | N/A | 100 | 0 | 35 | N/A | N/A |
| 87112812 | 37 | 20.3N | 114.0E | 11.3 | N/A | N/A | N/A | 80 | -5 | N/A | N/A | N/A |
| 87112818 | 38 | 20.0 N | 115.0E | 37 | N/A | N/A | N/A | 60 | 10 | N/R | N/A | N/A |
| 87112900 | 39 | 19.1N | 115.4 E | 169.8 | N/A | N/A | N/A | 50 | 10 | N/A | N/A | N/A |
| 87112906 | 40 | 18.2N | 115.4E | N/A | N/A | N/A | N/A | 40 | N/A | N/A | N/A | N/A |


| Tropicel Stom Ogden (23F) |  | 00 h | 24 h | 48 h | 72 h |
| :--- | :--- | ---: | ---: | ---: | ---: |
|  | Average | 26 | 99 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
|  | \# Cases | 4 | 2 | 0 | 0 |


| DTG | Hif | BT IAT | BT.ION | ROS AR | 24_ER | 48 ER | 72.ER | BT.WN | HW ER | 24-6E | 48 HE | 72. WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87112406 | 1 | 12.0N | 111.3 E | 13.2 | 94 | N/A | N/A | 35 | -5 | 5 | N/A | N/A |
| 87112412 | 2 | 12.3N | 110.4E | 18.6 | 104.1 | N/A | N/A | 40 | -10 | 10 | N/A | N/A |
| 87112418 | 3 | 12.7N | 109.4E | 47.5 | N/A | N/A | N/R | 45 | -10 | N/A | N/A | N/A |
| 87112500 | 4 | 13.5 N | 108.5E | 26.7 | N/A | N/A | N/A | 35 | -10 | N/A | N/A | N/A |


| Typhoon Pkyllis |  | (24W) |  | 00 h | 24b | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | Average \# Cases | 16 | 129 | 171 | 196 |  |  |  |  |  |
|  |  | 34 | 25 | 13 | 13 |  |  |  |  |  |
| DTE | W14 |  | BT LAT | BTILON | RQS_ER | 24.EB | 48 ER | 72 ER | BT WN | WW ER | 24.WE | 48 WE | 72 hE |
| 87121018 | 1 | 7.3N | 146.1E | 21.6 | 122.3 | 169.3 | 232.9 | 25 | 5 | 10 | 25 | 50 |
| 87121100 | 2 | 8.1 N | 144.6 E | 8.4 | 131.6 | 246.1 | 300.4 | 30 | 0 | 5 | 25 | 35 |
| 87121106 | 3 | 8.9 N | 143.5E | 13.3 | 152.9 | N/A | N/A | 30 | 0 | 5 | N/A | N/A |
| 87121112 | 4 | 10.0N | 142.9E | 8.4 | 172.2 | N/A | N/A | 30 | 0 | 5 | N/A | N/A |



## b. NORTH INDIAN OCEAN

This section includes verification statistics for each warning in the North Indian

Ocean during 1987. Pre- and post- warning best track positions are not printed, but are available on floppy diskettes by request.

JHWC FORECAST TRACK AND IANENSITY ERRORS BY WARNING

TROPICAL CYCLOND 018

| DTG | 组 | BT IAT | BT ION | POS_EB | 24.ER | 48 ER | 72.8 B | BT WN | Fim ER | 24.ER | 48 EB | 72.ER |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87020100 | 1 | 8.7N | 85.2E | 34.7 | 181.6 | 254.1 | 311.1 | 75 | 0 | 0 | 20 | 40 |
| 87020106 | 2 | 9.6 N | 85.4 E | 38.1 | 133.1 | 129.5 | 197.1 | 35 | 0 | -5 | 20 | 30 |
| 87020112 | 3 | 10.6 N | 85.8E | 26.5 | 73.3 | 74.5 | N/A | 40 | 0 | 0 | 30 | N/A |
| 87020118 | 4 | 11.6 N | 86.3E | 18.6 | 17.4 | 29.1 | N/A | 45 | -5 | 5 | 30 | N/A |
| 87020200 | 5 | 12.5 N | 86.8 E | 17.6 | 28.8 | 21.3 | N/R | 50 | 0 | 20 | 40 | N/A |
| 87020206 | 6 | 13.5 N | 87.4 E | 17.5 | 74.9 | 34.4 | N/A | 55 | 0 | 15 | 20 | N/A |
| 87020212 | 7 | 14.4N | 87.8 E | 21.2 | 66.3 | $N / \mathrm{A}$ | N/A | 50 | 5 | 15 | N/A | N/A |
| 87020218 | 8 | 15.3N | 88.1E | 25 | 54.3 | N/A | N/A | 45 | 5 | 15 | N/A | N/A |
| 87020300 | 9 | 16.1N | 88.4 E | 34 | 34.4 | N/A | N/A | 40 | 10 | 20 | N/A | N/A |
| 87020306 | 10 | 16.9N | 89.1E | 11.5 | 111.7 | N/A | N/A | 35 | 10 | 20 | N/A | N/A |
| 87020312 | 11 | 17.7N | 89.7E | 12.9 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |


| TROPICAL | CYCLONE | 028 | Average | O0h | 24h | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | 33 | 166 | N/A | N/A |  |  |  |  |  |
|  |  |  | \# Cases | 12 | 7 | 0 | 0 |  |  |  |  |  |
| DTG | H娄 | BT IAT | BT ION | ROS ER | 24_ER | 48 EB | 72 EB | BT WN | HW ER | 24.ER | 48 EB | 72.ER |
| 87060206 | 1 | 17.0N | 89.3E | 34.9 | 95.7 | N/A | N/A | 35 | -5 | 15 | N/A | N/A |
| 87060212 | 2 | 17.2N | 88.7E | 34.4 | 159.9 | N/A | $N / A$ | 35 | 0 | 15 | N/A | N/A |
| 87060218 | 3 | 17.5 N | 88.5E | 28.6 | 203.9 | N/A | N/A | 35 | 0 | 10 | N/A | N/A |
| 87060300 | 4 | 18.0N | 88.5E | 54.5 | 247.9 | N/A | N/A | 35 | 5 | -5 | N/A | N/A |
| 87060306 | 5 | 18.5 N | 88.6E | 37.7 | 261.7 | N/A | N/A | 40 | 0 | -25 | N/A | N/A |
| 87060312 | 6 | 19.1 N | 89.15 | 46.4 | 147.4 | N/A | N/A | 40 | -5 | -20 | N/A | N/A |
| 87060318 | 7 | 19.7 N | 89.6E | 34.4 | 49 | N/A | N/A | 45 | -5 | -5 | N/A | $\mathrm{N} / \mathrm{A}$ |
| 87060400 | 8 | 20.3N | 90.1 E | 12 | N/A | N/A | N/A | 50 | -5 | N/A | N/A | N/A |
| 87060406 | 9 | 21.1 N | 90.6 E | 8.2 | N/A | N/A | N/A | 55 | 0 | N/A | N/A | N/A |
| 87060412 | 10 | 21.9N | 90.8E | 33.2 | N/A | N/A | N/A | 50 | 5 | N/A | N/A | N/A |
| 87060418 | 11 | 22.8 N | 90.8 E | 40.8 | N/A | N/A | N/A | 40 | 5 | N/A | N/A | N/A |
| 87060500 | 12 | 23.7 N | 90.9E | 32 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | $\mathrm{N} / \mathrm{A}$ |


| TRORICAT CYCIONE 03A |  | $00 h$ | $24 h$ | $48 h$ | $72 h$ |
| ---: | ---: | ---: | ---: | ---: | ---: |
|  | Average | 62 | 165 | 183 | 208 |
|  | \# Cases | 18 | 12 | 12 | 12 |


| DTG | What | BT TAT | BT_ION | POS EB | 24.EB | 48_EB | 72 EBP | BT WN | MW ER | 24.ER | 48.ER | 72.EB |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87060506 | 1 | 16.1 N | 62.2 E | 28.8 | 42.1 | 107.2 | 150 | 45 | 0 | 10 | 20 | 35 |
| 87060512 | 2 | 16.1 N | 62.6 E | 13 | 98.9 | 134.8 | 123.3 | 45 | 0 | 5 | 15 | 35 |
| 87060518 | 3 | 16.1 N | 63. OE | 62.4 | 126 | 90.6 | 337.1 | 45 | 5 | 10 | 20 | 45 |


| 87060600 | 4 | 16.1 N | 63.5 E | 44.2 | 40.6 | 169.6 | 291.5 | 50 | 0 | 10 | 20 | 45 |
| :--- | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 87060606 | 5 | 16.1 N | 64.2 E | 51.9 | 127.3 | 209.6 | 139 | 50 | 5 | 20 | 25 | 25 |
| 87060612 | 6 | 16.1 N | 64.9 E | 80.9 | 175.6 | 198 | 95.4 | 50 | 5 | 15 | 25 | 25 |
| 87060618 | 7 | 16.1 N | 65.7 E | 126.8 | 223.5 | 234.9 | 109.3 | 50 | 5 | 10 | 20 | 20 |
| 87060700 | 8 | 17.0 N | 66.2 E | 116.8 | 234 | 206.3 | 99.7 | 50 | 0 | 0 | 15 | 20 |
| 87060706 | 9 | 17.9 N | 66.1 E | 90.7 | 247.1 | 245.1 | 220.5 | 50 | 0 | 5 | 15 | 25 |
| 87060712 | 10 | 18.9 N | 66.0 E | 102.2 | 217.9 | 208.8 | 255.4 | 50 | 0 | 15 | 30 | 45 |
| 87060718 | 11 | 19.8 N | 65.9 E | 92.8 | 211.1 | 204 | 313.2 | 50 | 0 | 20 | 25 | 25 |
| 87060800 | 12 | 20.6 N | 65.5 E | 132.6 | 236.3 | 195.4 | 363.7 | 50 | 0 | 15 | 25 | 20 |
| 87060806 | 13 | 21.0 N | 64.9 E | 45 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 45 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87060812 | 14 | 21.0 N | 64.3 E | 32 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 40 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87060818 | 15 | 20.9 N | 63.7 E | 28.7 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87060900 | 16 | 20.4 N | 63.4 E | 32 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87060906 | 17 | 19.9 N | 63.7 E | 37.5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87060912 | 18 | 19.9 N | 64.3 E | 13.3 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 30 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |


| TROPICAL CYCLONE 04B |  | 00 h | 24 h | 48 h | 72 h |
| ---: | ---: | ---: | ---: | ---: | ---: |
|  | Average | 12 | N/A | N/A | N/A |
|  | \# Cases | 3 | 0 | 0 | 0 |


| DTG | 鯺 | BT LAT | BT ION | PQS ER | 24ER | 48 ER | 72 ER | BT WN | WW ER | 24ER | 48 ER | $72 . E R$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87101500 | 1 | 14.5 N | 83.7E | 13.1 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 87101518 | 2 | 16.0 N | 81.5E | 17.3 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 87101600 | 3 | 16.6 N | 80.5E | 8.3 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |



| TROPICAL CYCLORE O6B |  |  |  | 00h | 24h | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 16 | 113 | N/A | N/A |  |  |  |  |  |
|  |  |  | * Cases | 6 | 2 | 0 | 0 |  |  |  |  |  |
| DTG | W\# | BT LAT | BT LON | ROS ER | 24 ER | $48 . \mathrm{ER}$ | 72 ER | BT WN | WW ER | 24 EB | 48 ER | 72 ER |
| 87111118 | 1 | 14.7 N | 87.3E | 8.3 | 185.3 | N/A | N/A | 35 | -5 | 0 | N/A | N/A |
| 87111200 | 2 | 14.6N | 85.5E | 13.3 | 42 | N/A | N/A | 35 | 0 | -10 | N/A | N/A |
| 87111206 | 3 | 14.2N | 84.0E | 13.3 | N/A | N/A | N/A | 40 | -5 | N/A | N/A | N/A |
| 87111212 | 4 | 14.4N | 82.5 E | 12 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 87111218 | 5 | 15.3N | 81.2 E | 24.7 | N/A | N/A | N/A | 50 | 0 | N/A | N/A | N/A |
| 87111300 | 6 | $16.3 N$ | 80.0 E | 29.4 | N/A | N/A | N/A | 40 | 0 | N/A | N/A | N/A |
| TROPICAL | criclosis | 07A |  | 00h | 24h | 48 h | 72h |  |  |  |  |  |
|  |  |  | Average | 38 | 119 | 307 | 421 |  |  |  |  |  |
|  |  |  | \# Cases | 14 | 11 | 6 | 6 |  |  |  | . |  |
| DTG | W | BT LAT | BTILON | POS ER | 24ER | 48ER | 72.EB | BT WN | WW ER | 24ER | 48 ER | 72 ER |
| 87120812 | 1 | 10.9N | 70.7E | 0 | 129.1 | 326 | 483.2 | 35 | 0 | 15 | 0 | 5 |
| 87120818 | 2 | 11.2N | 70.3 E | 16.8 | 176.3 | 356.3 | 485.1 | 35 | 0 | 10 | 0 | 10 |
| 87120900 | 3 | 11.6 N | 70.0 E | 37.3 | 220.7 | 402.2 | 517.3 | 35 | 0 | 10 | 5 | 10 |
| 87120906 | 4 | 12.0 N | 69.9 E | 23.5 | 160.8 | 286.3 | 381.2 | 35 | 0 | 5 | 10 | 20 |
| 87120912 | 5 | 12.6 N | 70.0 E | 84.3 | 254.8 | 336.3 | 407.9 | 35 | 0 | 0 | 10 | 20 |
| 87120918 | 6 | 13.3N | 70.1E | 60.3 | 96 | 139.7 | 256.5 | 40 | -5 | 0 | 15 | 25 |
| 87121000 | 7 | 13.9N | 70.3 E | 84.8 | 109.3 | N/A | N/A | 40 | -5 | -10 | N/A | N/A |
| 87121006 | 8 | 14.3N | 70.6 E | 8.4 | 5.8 | N/A | N/A | 45 | -10 | -5 | N/A | N/A |
| 87121012 | 9 | 14.7 N | 71.0E | 21.4 | 37.5 | N/A | N/A | 45 | -10 | -5 | N/A | N/A |
| 87121018 | 10 | 15.0N | 71.38 | 24 | 54.3 | N/A | N/A | 45 | 0 | 5 | N/A | N/A |
| 87121100 | 11 | 15.4N | 71.6 E | 36.5 | 74 | N/A | N/A | 40 | 0 | -5 | N/A | N/A |
| 87121106 | 12 | 15.9N | 71.9 E | 37.8 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 87121112 | 13 | 16.4 N | 72.3 E | 45.4 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |
| 87121118 | 14 | 16.9 N | 72.7E | 54.3 | N/A | N/A | N/A | 30 | -5 | N/A | N/A | N/A |
| tropical | cyclonis | 08B |  | 00\% | 24. | 48h | 72h |  |  |  |  |  |
|  |  |  | Average | 123 | 192 | N/A | N/A |  |  |  |  |  |
|  |  |  | * Cases | 5 | 3 | 0 | 0 |  |  |  |  |  |
| DTG | W1 | BT LAT | BT ION | POS ER | 24ER | 48.ER | 72 ER | BT WN | WW ER | 24EB | 48 ER | 72 ER |
| 87121800 | 1 | 10.7 N | 82.6 E | 24.7 | 159 | N/A | N/A | 30 | 0 | 10 | N/A | N/A |
| 87121806 | 2 | 10.5N | 82.2E | 76 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |
| 87121812 | 3 | 10.4N | 81.8 E | 128.2 | 176.1 | N/A | N/A | 30 | 0 | 0 | N/A | N/A |
| 87121818 | 4 | 10.3 N | 81.4E | 181.9 | 240.6 | N/A | N/A | 30 | 0 | 0 | N/A | N/A |
| 87121900 | 5 | 10.5N | 81.1E | 206.6 | N/A | N/A | N/A | 25 | 0 | N/A | N/A | N/A |

## c．SOUTHERN HEMISPHERE

This section includes verification statistics for each warning in the South Indian and western South Pacific Oceans from 1 July

1986 to 30 June 1987．Pre－and post－warning best track positions are not printed，but are available on floppy diskettes by request．

JHWC FORECAST TRACK AND INTIENSITY ERRORS BY FARNING

| Tropical | Cyclona | 01S |  | 00\％ | 24h | 48 h |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 25 | 90 | 196 |  |  |  |  |
|  |  |  | ＊Cases | 4 | 3 | 1 |  |  |  |  |
| DTG | H星 | BT＿IST | BT ION | POS ER | 24．EB | 48 EB | BT WN | WW＿ER | 24．7E | 48 WE |
| 86080112 | 1 | 7.6 N | 78.3 E | 24.5 | 91 | 196.4 | 40 | 0 | 30 | 40 |
| 86080200 | 2 | 7．6N | 76.9 E | 8.4 | 81.5 | N／A | 35 | 5 | 5 | N／A |
| 86080212 | 3 | 7．4N | 75.5 E | 32.3 | 99.2 | N／A | 30 | 5 | 0 | N／A |
| 86080300 | 4 | 6．9N | 74．0E | 36.5 | N／A | N／A | 30 | －5 | N／A | N／A |
| Tropical | Cyclono | $02 P$ |  | 00h | 24h | 48 h |  |  |  |  |
|  |  |  | Average | 33 | 141 | 292 |  |  |  |  |
|  |  |  | ＊Cases | 6 | 4 | 2 |  |  |  |  |
| DTG | W暑 | BT 工发T | BT ION | POS ER | 24＿EB | 48．ER | BI WN | WW＿EB | 24WE | 48．WE |
| 86112218 | 1 | 12.7 N | 168．8E | 21.5 | 29 | 196.9 | 35 | －5 | －15 | －10 |
| 86112306 | 2 | 13．6N | 168．7E | 8.4 | 75.3 | 388.2 | 45 | －10 | －5 | 5 |
| 86112318 | 3 | 14.7 N | 169．2E | 31.4 | 212.7 | N／A | 50 | 0 | 20 | N／A |
| 86112406 | 4 | 15．7N | 170．1E | 40.9 | 250.3 | N／A | 55 | 5 | 25 | N／A |
| 86112418 | 5 | 16.7 N | 172．3E | 53.1 | N／A | N／A | 55 | 5 | N／A | N／A |
| 86112506 | 6 | 17.4 N | 175．3E | 43.9 | N／A | N／A | 55 | －5 | N／A | N／A |
| Tropical | Cyclone | 035 |  | 00h | 24h | 48h |  |  |  |  |
|  |  |  | Average | 42 | 160 | 340 |  |  |  |  |
|  |  |  | ＊Cases | 6 | 4 | 2 |  |  |  |  |
| DTG | 昰 | BT IAT | BT－ION | POS EB | 24．EB | 48 ER | BT＿MN | MNLES | 24．W5 | 48．N15 |
| 86121418 | 1 | 13．0N | 168．0E | 59.7 | 169.5 | 343.3 | 35 | 0 | 0 | 30 |
| 86121506 | 2 | 14.2 N | 167．4E | 38 | 232.2 | 338 | 50 | 5 | 25 | 50 |
| 86121518 | 3 | 16.0 N | 166．3E | 55 | 148.7 | N／A | 55 | 10 | 40 | N／A |
| 86121606 | 4 | 18.1 N | 166．0E | 39.9 | 91.6 | N／A | 50 | 25 | 55 | N／A |
| 86121618 | 5 | 19.7 N | 166.4 E | 34.5 | N／A | N／A | 45 | 15 | N／A | N／A |
| 86121706 | 6 | 22.0 N | 167．5E | 28.4 | N／A | N／A | 35 | 10 | N／A | N／A |
| Tropical | Cyclona | 042 |  | 00 h | 24h | 48h |  |  |  |  |
|  |  |  | Average | 35 | 171 | 333 |  |  |  |  |
|  |  |  | ＊Cases | 18 | 16 | 14 |  |  |  |  |
| DTG | 표ํ | BT＿IAT | BT ION | POS ER | 24．ER | 48 EB | BT HN | FWTEB | 24 FEE | 48 FIE |
| 86122318 | 1 | 12.1 N | 176.35 | 8.4 | 274.1 | 606.8 | 35 | 0 | 5 | 5 |
| 86122406 | 2 | 12.7 N | 176．15 | 66.8 | 366 | 640.1 | 35 | 5 | －5 | 5 |
| 86122418 | 3 | 13.6 N | 177．6E | 17.5 | 66.7 | 204.7 | 45 | 0 | 0 | 5 |
| 86122506 | 4 | 14.4 N | 179．0E | 46.9 | 117.2 | 312 | 55 | 0 | 5 | －5 |
| 86122518 | 5 | 14.9 N | 180．6E | 41 | 51 | 216.6 | 55 | 0 | 10 | 5 |
| 86122606 | 6 | 14.9 N | 182.2 E | 21.4 | 134.7 | 353.3 | 55 | 5 | 0 | －10 |
| 86122618 | 7 | 14．9N | 182．6E | 12 | 268.2 | 552.1 | 55 | 10 | 5 | －20 |


| 86122706 | 8 | 14.9 N | 182.3 E | 18.4 | 159.6 | 404.6 | 65 | 5 | 0 | -15 |
| ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 86122718 | 9 | 15.0 N | 181.6 E | 11.6 | 78.2 | 82.8 | 65 | 10 | 0 | -10 |
| 86122806 | 10 | 15.0 N | 180.8 E | 35.3 | 151.9 | 292.1 | 75 | 0 | 0 | -20 |
| 86122818 | 11 | 15.2 N | 180.2 E | 8.3 | 158.8 | 362.5 | 80 | 0 | -5 | -15 |
| 86122906 | 12 | 16.1 N | 180.1 E | 5.8 | 75.3 | 196.1 | 80 | 5 | -15 | 0 |
| 86122918 | 13 | 17.4 N | 180.4 E | 0 | 102.5 | 256.1 | 90 | 0 | -5 | 35 |
| 86123006 | 14 | 19.0 N | 181.2 E | 8.3 | 127.1 | 186 | 90 | -5 | 5 | 20 |
| 86123918 | 15 | 21.2 N | 182.0 E | 22.4 | 175.3 | $\mathrm{~N} / \mathrm{A}$ | 90 | -5 | 20 | $\mathrm{~N} / \mathrm{A}$ |
| 86123106 | 16 | 23.6 N | 183.1 E | 36 | 434.2 | $\mathrm{~N} / \mathrm{A}$ | 65 | 0 | 10 | $\mathrm{~N} / \mathrm{A}$ |
| 86123118 | 17 | 25.0 N | 183.4 E | 44.3 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 45 | 0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 87010106 | 18 | 25.0 N | 182.0 E | 238 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 30 | 5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |


| Tropical Cyclone 05P |  | 00 h | 24 h | $\frac{48 \mathrm{~h}}{}$ |
| :---: | :---: | ---: | ---: | ---: |
|  | Average | 41 | 139 | 281 |
|  | \# Cases | 16 | 14 | 11 |


| DTG. | Wis | Bt lat | BT ION | POS ER | 24_EB | 48 EB | BT WN | HW ER | 24 WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 86122800 | 1 | 14.7N | 196.8E | 41.8 | 98.3 | 161 | 30 | 10 | 15 | 15 |
| 86122812 | 2 | 15.3N | 196.8E | 17.4 | 168.3 | 418.5 | 35 | 15 | 20 | 10 |
| 86122900 | 3 | 15.9N | 196.8E | 26.6 | 170.8 | 432.7 | 45 | 20 | 15 | -5 |
| 86122912 | 4 | 16.6N | 196.9E | 13.3 | 82.4 | 203.2 | 50 | 15 | 15 | 0 |
| 86123000 | 5 | 17.0 N | 197.4E | 40.2 | 157.2 | 193.3 | 55 | 25 | 10 | 5 |
| 86123012 | 6 | 17.1N | 198.2E | 46.7 | 151 | 154.8 | 60 | 20 | 15 | 10 |
| 86123100 | 7 | 17.0 N | 199.0E | 18.9 | 144.5 | 311.9 | 65 | 10 | 0 | -5 |
| 86123112 | 8 | 16.4 N | 198.8 E | 11.5 | 72 | 111.3 | 60 | 15 | 25 | 20 |
| 87010100 | 9 | 17.2N | 198.5E | 45.4 | 42 | 92.3 | 60 | 15 | 15 | 15 |
| 87010112 | 10 | 17.8 N | 198.5E | 39.9 | 118.3 | 376.1 | 55 | 15 | 20 | 10 |
| 87010200 | 11 | 18.8 N | 198.8E | 45.4 | 272.2 | 638.3 | 55 | 10 | 5 | 0 |
| 87010212 | 12 | 19.8N | 199.7E | 43.5 | 225 | N/A | 50 | 20 | 5 | N/A |
| 87010300 | 13 | 20.7N | 201.3E | 65.6 | 117.9 | N/A | 50 | 15 | 5 | N/A |
| 87010312 | 14 | 21.8 N | 203.6E | 83.4 | 129.1 | N/A | 50 | 15 | 15 | N/A |
| 87010400 | 15 | 23.2N | 205.7E | 97.8 | N/A | N/A | 45 | 5 | N/A | N/A |
| 87010412 | 16 | 24.8N | 207.8E | 29 | N/A | N/A | 40 | 0 | N/A | N/A |


| Tropical Cyclone 06S |  | 00 h | 24 h | 48 h |
| :--- | ---: | ---: | ---: | ---: |
|  | Average | 38 | 203 | 365 |
|  | \# Cases | 5 | 4 | 4 |


| DTG | W年 | BTIAT | BT ION | POS ER | 24-ER | 48 ER | BT WN | WW ER | 24 WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87010712 | 1 | 7.1 N | 76.6 E | 21.6 | 172.7 | 312.1 | 35 | -5 | 10 | -5 |
| 87010800 | 2 | 8.0 N | 76.5E | 24.5 | 287.2 | 460 | 35 | 0 | 15 | 25 |
| 87010812 | 3 | 9.4 N | 77.9E | 37.5 | 162.8 | 313.5 | 35 | 0 | 10 | 10 |
| 87010900 | 4 | 10.9N | 79.15 | 29.7 | 189.4 | 376.3 | 35 | 10 | 10 | -5 |
| 87010912 | 5 | 11.2 N | 79.7E | 78.9 | N/A | N/A | 35 | 0 | N/A | N/A |


| Tropical Cyclone 07s |  | 00 h | 24 h | 48 h |
| :--- | :--- | ---: | ---: | ---: |
|  | Average | 24 | 119 | 194 |
|  | * Cases | 5 | 3 | 2 |


| DTE | Wi | BT Lat | BT LON | ROS ER | 24 ER | 48 ER | BT WN | Him ER | 24.WE | 48 Wr |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87011012 | 1 | 9.1N | 97.5E | 11.8 | 64.5 | 63.2 | 50 | 0 | 25 | 35 |
| 87011100 | 2 | 10.7 N | 98.2E | 43 | 189.6 | 325.9 | 50 | -5 | -5 | 0 |
| 87011112 | 3 | 11.7N | 99.0 E | 8.4 | 103.5 | N/A | 45 | -5 | -10 | N/A |
| 87011200 | 4 | 12.5N | 99.3 E | 21.5 | N/A | N/A | 45 | -10 | N/A | N/A |
| 87011212 | 5 | 12.6 N | 99.4 E | 36.5 | N/A | N/A | 45 | -15 | N/A | N/ |


| Tropical Cyclone |  | 089 |  | 00h | 24h | 48h |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | Average | 25 | 119 | 194 |  |  |  |  |
|  |  | * Cases | 9 | 7 | 5 |  |  |  |  |
| DTG | W\% |  | BT Lat | BT LON | POS EB | 24.58 | 48 ER | BT WN | HM ER | 24 WE | 48 WE |
| 87011600 | 1 |  | 10.7 N | 189.0E | 26.7 | 163 | 229.6 | 35 | 5 | 10 | -25 |
| 87011612 | 2 | 10.9 N | 189.7E | 41.7 | 42.6 | 91.4 | 45 | 0 | -10 | -35 |
| 87011700 | 3 | 10.7 N | 189.7E | 21.4 | 115.6 | 291.3 | 55 | 0 | -20 | -15 |
| 87011712 | 4 | 11.8 N | 190.0E | 24.3 | 120.4 | 265.5 | 70 | -5 | -20 | 15 |
| 87011800 | 5 | 13.6 N | 190.4E | 13.1 | 202.1 | 392.2 | 90 | -10 | 5 | 50 |
| 87011812 | 6 | 15.3N | 191.4E | 47.8 | 5.7 | N/A | 100 | -10 | 15 | N/A |
| 87011900 | 7 | 17.4N | 192.9E | 8.3 | 79.8 | N/A | 90 | -5 | 25 | N/A |
| 87011912 | 8 | 19.5 N | 194.1E | 18.9 | N/A | N/A | 70 | -5 | N/A | N/A |
| 87012000 | 9 | 21.5N | 196.4E | 28.7 | N/A | N/A | 40 | 5 | N/A | N/A |
| Tropical | Cyclone | 098 |  | 00h | 24h | 48h |  |  |  |  |
|  |  |  | Average | 37 | 188 | 422 |  |  |  |  |
|  |  |  | * Cases | 12 | 8 | 5 |  |  |  |  |
| DTG | W1 | BTIAT | BT LON | ROS ER | 24.ER | 48 ER | BT WN | HMER | 24.WE | 48 WE |
| 87011618 | 1 | 13.9N | 66.9E | 35.5 | 303.9 | 630.7 | 45 | -10 | -10 | 5 |
| 87011706 | 2 | 14.8N | 67.3 E | 13.1 | 249.5 | 619.7 | 55 | -15 | -15 | 15 |
| 87011718 | 3 | 15.6 N | 68.2E | 64.7 | 265.7 | 514.3 | 65 | -15 | 5 | 40 |
| 87011806 | 4 | 17.2N | 68.9E | 12 | 103.3 | 208.8 | 75 | -15 | 25 | 35 |
| 87011818 | 5 | 19.1 N | 69.9E | 46.4 | 128.3 | 138.8 | 65 | 0 | 40 | 20 |
| 87011906 | 6 | 20.4N | 70.7E | 39.8 | 140.5 | N/A | 55 | -5 | 0 | N/A |
| 87011918 | 7 | 21.7 N | 71.1E | 17.8 | N/A | N/A | 40 | 0 | N/A | N/A |
| 87012006 | 8 | 23.3N | 71.2E | 84.3 | N/A | N/A | 35 | 0 | N/A | N/A |
| 87012206 | 9 | 26.5N | 63.2E | 21 | 103.3 | N/A | 65 | -15 | 15 | N/A |
| 87012218 | 10 | 29.2N | 63.5 E | 5.2 | 217.4 | N/A | 55 | 5 | -10 | N/A |
| 87012306 | 11 | 33.0 N | 66.5E | 70.5 | N/A | N/A | 55 | -5 | N/A | N/A |
| 87012318 | 12 | 36.8N | 72.0E | 36 | N/A | N/A | 55 | -15 | N/A | N/A |
| Tropical | Cyclone | 10 s |  | 00h | 24h | 48h |  |  |  |  |
|  |  |  | Average | 15 | 92 | 201 |  |  |  |  |
|  |  |  | * Cases | 6 | 5 | 4 |  |  |  |  |
| DTG | W违 | BT LAT | BT LON | POS AB | 24.ER | 48 ER | BT WN | HW ER | 24 WE | 48. WE |
| 87011712 | 1 | 17.2N | 120.9E | 23.7 | 54.3 | 106.2 | 30 | 5 | 0 | 15 |
| 87011800 | 2 | 17.8 N | 120.5E | 8.3 | 43.5 | 102.6 | 45 | -5 | 0 | 0 |
| 87011812 | 3 | 18.7N | 119.7E | 5.7 | 47.6 | 196.8 | 55 | -5 | 15 | 15 |
| 87011900 | 4 | 19.6 N | 118.7E | 13.3 | 100.9 | 397.6 | 55 | -15 | 0 | 20 |
| 87011912 | 5 | 20.7N | 117.9E | 30.5 | 215.8 | N/A | 50 | 15 | 15 | N/A |
| 87012000 | 6 | 22.1N | 117.6E | 11.1 | N/A | N/A | 50 | -10 | N/A | N/A |
| Tropical | Cyclane | 118 |  | 00h | 24h | 48h |  |  |  |  |
|  |  |  | Average | 13 | 17 | N/A |  |  |  |  |
|  |  |  | * Cases | 3 | 1 | 0 |  |  |  |  |
| DTG | H2010 | BT IAT | BT ION | POS ER | 24.EB | 48 ER | BT MN | HW ER | 24.WE | 48 WE |
| 87011918 | 1 | 13.0N | 137.7E | 13.3 | 17.4 | N/A | 30 | 5 | 10 | N/A |
| 87012006 | 2 | 13.4 N | 136.5E | 8.4 | N/A | N/A | 30 | 10 | N/A | N/A |
| 87012018 | 3 | 14.2N | 134.8E | 18.5 | N/A | N/A | 30 | 10 | N/A | N/A |


| Tropical Cyclone 12S |  | 00 h | 24 h | 48 h |
| :--- | ---: | ---: | ---: | ---: |
|  | Average | 35 | 138 | 135 |
|  | \# Cases | 9 | 7 | 5 |


| DTG | Wı | BT LAT | BT LON | POS 58 | 24-ER | 48 ER | BT WN | HW ER | 24.6E | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87020100 | 1 | 15.6 N | 123.4 E | 121.5 | 236.3 | 298.6 | 35 | 0 | 15 | 20 |
| 87020112 | 2 | 16.0N | 122.7E | 18.3 | 121.3 | 97.1 | 40 | 5 | 20 | 35 |
| 87020200 | 3 | 16.3N | 122.8E | 34 | 77.3 | 42.7 | 40 | 0 | 0 | 0 |
| 87020212 | 4 | 16.5 N | 122.5E | 13.3 | 69.1 | 46.1 | 40 | 0 | 0 | 10 |
| 87020300 | 5 | 17.5N | 121.1E | 5.7 | 66.2 | 193.6 | 45 | 0 | 0 | 50 |
| 87020312 | 6 | 17.8N | 119.9 E | 72.3 | 275.9 | N/A | 45 | 5 | 5 | N/R |
| 87020400 | 7 | 17.9N | 118.9E | 45.7 | 125.4 | N/A | 50 | 0 | 15 | N/A |
| 87020412 | 8 | 18.0N | 118.0E | 0 | N/A | N/A | 40 | 5 | N/A | N/A |
| 8702050 | 9 | 18.2N | 117.3 E | 6 | N/A | N/A | 30 | 10 | N/A |  |


| Tropical Cyclone 13P |  | 00h | 24h | 48h |
| :--- | :--- | ---: | ---: | ---: |
|  | Average | 36 | 139 | N/A |
|  | $\#$ Cases | 3 | 2 | 0 |


| DTG | 䦃 | BT IAT | BT LON | POS ER | 24.E8 | 48 ER | BT WN | Wh Es | 24. WE | 48 We |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87020400 | 1 | 16.0 N | 189.7E | 46.1 | 91.6 | N/A | 40 | - | 15 | N/A |
| 87020412 | 2 | 16.0 N | 191.1E | 18.3 | 186.9 | N/A | 40 | 0 | 20 | N/A |
| 87020500 | 3 | 15.8 N | 192.7E | 45.4 | N/A | N/A | 35 | 0 | N/R | N/A |


|  |  | 00 h | 24 h | 48 h |
| :---: | ---: | ---: | ---: | ---: |
|  | Average | 347 | 101 | 256 |
|  | H Cases | 8 | 6 | 4 |


| DTG | Hit | BT LAT | BT LON | POS ER | 24.ER | 48 ER | BT HN | Hit Er | 24-WE | 48.WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87020500 | 1 | 13.2 N | 163.3 E | 24.1 | 102 | 190.9 | 35 | 5 | -10 | -10 |
| 87020512 | 2 | 13.7N | 164.3E | 12 | 96.2 | 180.8 | 45 | 0 | -5 | 35 |
| 87020600 | 3 | 14.7 N | 165.2 E | 24 | 92.9 | 243.8 | 65 | 0 | 0 | 40 |
| 87020612 | 4 | 15.9 N | 166.1 E | 18.3 | 64.5 | 409.9 | 70 | 0 | 25 | 45 |
| 87020700 | 5 | 17.0 N | 167.0E | 13.3 | 68.7 | N/A | 75 | 5 | 45 | N/A |
| 87020712 | 6 | 18.3 N | 168.2E | 23.6 | 182.3 | N/A | 55 | 25 | 45 | N/A |
| 87020800 | 7 | 20.2N | 170.4E | 43.4 | N/A | N/A | 35 | 25 | N/A | N/A |
| 87020812 | 8 | 22.0N | 175.0E | 101.8 | N/A | N/A | 25 | 25 | N/A | N/A |


| Tropical Cyclone 15P |  | 00 h | 24 h | 48 h |
| ---: | ---: | ---: | ---: | ---: |
|  | Average | 27 | 120 | 369 |
|  | \# Cases | 12 | 8 | 2 |


| DTG | W4 | BT Lat | BT LON | POS ER | 24ER | 48 ER | BT WN | WW ER | 24.WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87020706 | 1 | 13.2N | 140.3E | 51.2 | 102.6 | 223.2 | 30 | 5 | 15 | -30 |
| 87020718 | 2 | 13.2N | 139.5E | 41.9 | 146.6 | N/A | 35 | 5 | -10 | N/A |
| 87020806 | 3 | 13.2N | 138.6E | 52.6 | 18.9 | N/A | 40 | 10 | -15 | N/A |
| 87020818 | 4 | 13.2N | 137.5E | 37.8 | N/A | N/A | 50 | -10 | N/A | N/A |
| 87020906 | 5 | 13.1N | 136.8E | 5.8 | 78.9 | N/A | 55 | -10 | -30 | N/A |
| 87020918 | 6 | 12.9N | 136.3E | 25.1 | 17.6 | N/A | 55 | -15 | -35 | N/A |
| 87021006 | 7 | 12.6N | 136.0E | 11.7 | N/A | N/A | 60 | -20 | N/A | N/A |
| 87021018 | 8 | 12.3N | 135.5E | 5.9 | 205.9 | 514.8 | 60 | -20 | -15 | 5 |
| 87021106 | 9 | 13.0N | 135.4E | 18 | 282.8 | N/A | 65 | -25 | -5 | N/A |
| 87021118 | 10 | 13.9N | 137.0E | 24.1 | 106.8 | N/A | 60 | -20 | -5 | N/A |
| 87021206 | 11 | 14.9N | 138.4E | 18.4 | N/A | N/A | 50 | 15 | N/A | N/A |
| 87021218 | 12 | 15.7N | 140.0E | 34 | N/A | N/A | 40 | 30 | N/A | N/A |


| Tropical Cyclone 169 |  |  |  | 00h | 24b | 48h |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 67 | 424 | N／A |  |  |  |  |
|  |  |  | ＊Cases | 3 | 1 | 0 |  |  |  |  |
| DTG | W＊ | BT Lat | BTILON | POS EB | 24．ER | 48 ER | BT HN | HW ER | 24 WE | 48 WE |
| 87020818 | 1 | 18.5 N | 173．6E | 45.5 | 424.1 | N／A | 25 | 20 | 30 | N／A |
| 87020906 | 2 | 21．9N | 177．7E | 13.2 | N／A | N／A | 25 | 5 | N／A | N／A |
| 87020918 | 3 | 26.1 N | 180．8E | 144.3 | N／A | N／A | 25 | 5 | N／A | N／A |
| Tropical Cyclone |  | 175 |  | 00h | 24h | 48 h |  |  |  |  |
|  |  |  | Average | 51 | 109 | 227 |  |  |  |  |
|  |  |  | ＊Cases | 10 | 7 | 6 |  |  |  |  |
| DTG | W考 | bT LAT | BT LON | POS EB | 24＿EB | 48 EB | BT WN | WW ER | 24＿WE | 48 WE |
| 87021106 | 1 | 21．1N | 53．8E | 84.2 | 79.7 | 81.1 | 30 | 5 | 10 | 5 |
| 87021118 | 2 | 21.0 N | 54．0E | 83.8 | 125.1 | 147.9 | 35 | 0 | 0 | 5 |
| 87021206 | 3 | 20．9N | 54．3E | 21.2 | 45 | 131.7 | 35 | 5 | 5 | 15 |
| 87021218 | 4 | 20．8N | 54．8E | 18.9 | 95.1 | 285 | 45 | 10 | 20 | 15 |
| 87021306 | 5 | 20．9N | 55．6E | 28.7 | 16.4 | 229.5 | 45 | 10 | 25 | 25 |
| 87021318 | 6 | 20．9N | 56．4E | 12.7 | 171.4 | 486.9 | 45 | 5 | 15 | 15 |
| 87021406 | 7 | 20．7N | 56.4 E | 55.1 | 233.7 | N／A | 40 | 10 | 5 | N／A |
| 87021418 | 8 | 21．1N | 55．7E | 119.3 | N／A | N／A | 35 | 0 | N／A | N／A |
| 87021506 | 9 | 22．1N | 54.6 E | 23 | N／A | N／A | 30 | 5 | N／A | N／A |
| 87021518 | 10 | 23.0 N | 53．0E | 68.7 | N／A | N／A | 30 | 5 | N／A | N／A |
| Tropical Cyalan 183 |  |  |  | 00h | 246 | 48h |  |  |  |  |
|  |  |  | Average | 11 | 147 | 369 |  |  |  |  |
|  |  |  | ＊Cases | 7 | 6 | 4 |  |  |  |  |
| DTG | H14 | bTiAt | BT ION | POS ER | 24＿E8 | $48 . \mathrm{ER}$ | BT HiN | MWE ER | 24 WE | 48－6E |
| 87022212 | 1 | 15.6 N | 123.0 E | 6 | 106.6 | 389.7 | 30 | 0 | 10 | 15 |
| 87022300 | 2 | 15.8 N | 121.4 E | 0 | 175.2 | 6812.6 | 35 | 0 | 5 | 5 |
| 87022312 | 3 | 15．8N | 120．7E | 8.3 | 178.9 | 317.4 | 40 | 0 | 5 | 5 |
| 87022400 | 4 | 16.2 N | 121．2E | 29.3 | 198.9 | 427.8 | 45 | 0 | 0 | 20 |
| 87022412 | 5 | 17.2 N | 121．6E | 16.6 | 29.4 | N／A | 45 | 0 | －15 | N／A |
| 87022500 | 6 | 18.6 N | 121.2 E | 5.7 | 193.3 | N／A | 55 | 0 | 0 | N／A |
| 87022512 | 7 | 19.9 N | 121．0E | 12.8 | N／A | N／A | 55 | 0 | N／A | N／A |
| Tropical Cyclone 19P |  |  |  | 00h | 24 h | 48h |  |  |  |  |
|  |  |  | Average | 5 | 332 | N／A |  |  |  |  |
|  |  |  | ＊Cases | 3 | 1 | 0 |  |  |  |  |
| DTG | W年 | BT IAT | BT LON | POS EB | 24 ER | 48 ER | BT WN | HW ER | 24－WE | 48 WE |
| 87022818 | 1 | 14.9 N | 197.9 E | 0 | 332.1 | N／A | 35 | 0 | 5 | N／A |
| 87030106 | 2 | 18．5N | 200．0E | 0 | N／A | N／A | 35 | 0 | N／A | N／A |
| 87030118 | 3 |  | 202．8E | 16.3 | N／A | N／A | 40 | －5 | N／A | N／A |
| Tropical Cyclone 20P |  |  |  | －00h | 24 h | 48h |  |  |  |  |
|  |  |  | Average | 47 | 146 | 286 |  |  |  |  |
|  |  |  | －Cases | 9 | 6 | 4 |  |  |  |  |
| DTG | 迷 | BT IAT | BT Low | ROS ER | 24EER | 48＿ER | BT WN | WW ER | 24 WE | 48．WE |
| 87030112 | 1 | 13．5N | 187．4E | 55.5 | 183.9 | 244.3 | 35 | 0 | －10 | －15 |
| 87030200 | 2 | 13．9N | 190．5E | 83.3 | 170.8 | 261.5 | 40 | 0 | －5 | －15 |


| 87030212 | 3 | 15.3N | 193.3 E | 34.7 | 157 | 306.3 | 45 | -5 | -20 | -25 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87030312 | 4 | 20.5N | 194.5E | 89.6 | 185.4 | 333.6 | 60 | -5 | -15 | -20 |
| 87030400 | 5 | 22.2N | 195.3E | 51.5 | 108 | N/A | 65 | -10 | -15 | N/A |
| 87030412 | 6 | 23.5N | 196.5E | 24.6 | 76.3 | N/A | 65 | 0 | -20 | N/A |
| 87030500 | 7 | 24.9N | 198.2E | 16.2 | N/A | N/A | 55 | 0 | N/A | N/A |
| 87030512 | 8 | 26.2N | 200.3E | 36 | N/A | N/A | 65 | 0 | N/A | N/A |
| 87030600 | 9 | 28.1N | 203.6E | 36.5 | N/A | N/A | 65 | 0 | N/A | N/A |


| Tropical Cyclone 21s |  | 00 h | 24 h | 48 h |
| :--- | :--- | ---: | ---: | ---: |
|  | Average | 47 | 188 | 307 |
|  | * Cases | 25 | 20 | 18 |


| DTG | What | BT LAT | BI LON | POS EB | 24ER | 48 ER | BT WN | HW ER | 24.WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87030300 | 1 | 13.7 N | 71.5E | 34.7 | 242.5 | 449.6 | 30 | 5 | 15 | 25 |
| 87030312 | 2 | 13.2 N | 71.4E | 8.4 | 108.9 | 277.2 | 35 | 10 | 25 | 25 |
| 87030400 | 3 | 12.5 N | 71.5E | 44.8 | 307.4 | 566.7 | 40 | 10 | 20 | 25 |
| 87030412 | 4 | 13.0 N | 72.1E | 72.6 | 273.2 | 468.5 | 40 | 15 | 25 | 30 |
| 87030500 | 5 | 13.0N | 72.7E | 123.4 | 327.6 | 449 | 40 | 25 | 30 | 40 |
| 87030512 | 6 | 13.5 N | 74.3E | 62.7 | 282.9 | 493 | 40 | 20 | 10 | 0 |
| 87030600 | 7 | 13.7N | 75.5E | 42 | 288.4 | 431.8 | 40 | 10 | 10 | -5 |
| 87030612 | 8 | 13.0N | 75.6E | 104.3 | 247.1 | 286.8 | 35 | 15 | 5 | -20 |
| 87030700 | 9 | 12.1N | 76.78 | 21.5 | 136.6 | 157 | 35 | 0 | -5 | -40 |
| 87030712 | 10 | 12.0 N | 76.6 E | 119 | 358.7 | N/A | 35 | 0 | -25 | N/A |
| 87030800 | 11 | 12.1N | 75.9E | 59.2 | 242.1 | 345.7 | 40 | -5 | -35 | -35 |
| 87030812 | 12 | 12.7 N | 75.1E | 12 | 24.7 | 54.9 | 55 | -10 | -10 | 0 |
| 87030900 | 13 | 13.4N | 74.0 E | 16.7 | 94.1 | 197.1 | 70 | -10 | 10 | 0 |
| 87030912 | 14 | 14.5N | 73.1 E | 13.3 | 13 | 136.3 | 70 | 0 | 0 | -5 |
| 87031000 | 15 | 15.3N | 72.4 E | 5.8 | 62.9 | 232.9 | 65 | 0 | -10 | -20 |
| 87031012 | 16 | 16.4 N | 71.9 E | 23.8 | 154.2 | 332.5 | 65 | -5 | -15 | -15 |
| 87031100 | 17 | 17.0N | 71.7E | 18 | 115.6 | 279.7 | 65 | -10 | -20 | -10 |
| 87039112 | 18 | 17.7N | 72.0 E | 41.8 | 191.6 | 189.7 | 65 | -10 | -15 | -15 |
| 87031200 | 19 | 18.8 N | 72.9E | 61.7 | 202.9 | 188.2 | 65 | -15 | -20 | -15 |
| 87031212 | 20 | 20.1N | 74.4 E | 51.1 | 84.7 | N/A | 55 | -5 | -20 | N/A |
| 87031300 | 21 | 22.1N | 75.3 E | 45.7 | N/A | N/A | 50 | -5 | N/A | N/A |
| 87031312 | 22 | 23.6 N | 74.9 E | 89.8 | N/A | N/A | 50 | -10 | N/A | N/A |
| 87031400 | 23 | 24.6N | 73.9 E | 18 | N/A | N/A | 45 | -5 | N/A | N/A |
| 87031412 | 24 | 25.5N | 73.4 E | 63.8 | N/A | N/A | 40 | -5 | N/A | N/A |
| 87031500 | 25 | 27.2N | 73.5E | 42.3 | N/A | N/A | 40 | -10 | N/A | N/A |


| Tropical Cyclone 22P |  | 00 h | 24 h | 48 h |
| ---: | ---: | ---: | ---: | ---: |
|  | Average | 50 | 96 | 212 |
|  | \# Cases | 7 | 5 | 3 |


| DTG | [1* | BT.LAT | BT ION | POS ER | 24.E8 | 48_ER | BT WN | HW ER | 24-NE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 87030818 | 1 | 16.6N | 163.9E | 29.2 | 109.6 | 243.7 | 40 | -5 | -20 | 0 |
| 87030906 | 2 | 17.3N | 163.7E | 18.9 | 74 | 138 | 65 | -10 | 5 | 35 |
| 87030918 | 3 | 17.9N | 163.6E | 39.9 | 132.5 | 256.1 | 65 | 0 | -5 | 5 |
| 87031006 | 4 | 18.2N | 163.7E | 11.4 | 41.3 | N/A | 65 | -10 | 0 | N/A |
| 87031018 | 5 | 18.8N | 164.3E | 13.3 | 124 | N/A | 55 | 0 | 10 | N/A |
| 87031106 | 6 | 19.5 N | 165.0E | 22.6 | N/A | N/A | 40 | 5 | N/A | N/A |
| 87031118 | 7 | 20.2N | 167.2E | 216.2 | N/A | N/A | 35 | 0 | N/A | N/A |



| Tropical Cyclone 26S |  |  |  | 00h | 24h | 48 h |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 47 | 161 | 344 |  |  |  |  |
|  |  |  | * Cases | 5 | 3 | 1 |  |  |  |  |
| DTG | W\% | BT LAT | BT LON | POS ER | $24 . E R$ | 48 ER | BT WN | WW ER | 24.WE | 48 WE |
| 87042412 | 1 | 15.0N | 74.2E | 11.6 | 144.9 | 344.2 | 40 | -5 | 5 | 35 |
| 87042500 | 2 | 14.6 N | 74.9 E | 62.6 | 156.1 | N/A | 40 | -5 | 0 | N/A |
| 87042512 | 3 | 13.9N | 74.8 E | 54 | 182.9 | N/A | 45 | -5 | 25 | N/A |
| 87042600 | 4 | 13.3N | 74.4 E | 101.3 | N/A | N/A | 35 | 10 | N/A | N/A |
| 87042612 | 5 | 12.9N | 73.9E | 5.8 | N/A | N/A | 25 | 5 | N/A | N/A |
| Tropical | Cyclone | 278 |  | 00h | 24h | 48 h |  |  |  |  |
|  |  |  | Average | 22 | 93 | 138 |  |  |  |  |
|  |  |  | * Cases | 7 | 5 | 5 |  |  |  |  |
| DTG | W\# | BT LAT | BT LON | POS EB | 24ER | 48 ER | BT WN | WW ER | 24 WE | 48 WE |
| 87052218 | 1 | 12.3 N | 159.4 E | 13.4 | 87.7 | 67.7 | 35 | -5 | 10 | 20 |
| 87052306 | 2 | 13.7N | 158.6 E | 13.1 | 45.8 | 155.7 | 40 | -5 | 0 | 10 |
| 87052318 | 3 | 15.0 N | 158.1E | 23.9 | 99.6 | 255.6 | 45 | 0 | 5 | 10 |
| 87052406 | 4 | 15.9 N | 157.5E | 8.3 | 102.6 | 108.6 | 45 | 0 | 10 | 15 |
| 87052418 | 5 | 16.6 N | 157.1E | 37.8 | 132.1 | 101.7 | 45 | 10 | 25 | 20 |
| 87052506 | 6 | 16.6 N | 156.8E | 24.7 | N/A | N/A | 35 | 5 | N/A | N/A |
| 87052518 | 7 | 17.5N | 156.3E | 33.2 | N/A | N/A | 30 | 0 | N/A | N/A |
| Tropical | Cyclona | 285 |  | 00h | 24h | 48 h |  |  |  |  |
|  |  |  | Average | 56 | 194 | 450 |  |  |  |  |
|  |  |  | * Cases | 5 | 3 | 1 |  |  |  |  |
| DTG | W\# | bT Lat | BTILN | POS ER | 24.EB | 48 EB | BT WN | WW EB | 24-WE | 48.45 |
| 87062500 | 1 | 8.0 N | 87.4E | 165.7 | 321.3 | 449.6 | 30 | 5 | 30 | N/A |
| 87062512 | 2 | 7.4 N | 87.3 E | 16.9 | 55.1 | N/A | 30 | 5 | 15 | N/A |
| 87062600 | 3 | 7.6 N | 86.7E | 45.4 | 206.9 | N/A | 25 | 10 | 0 | N/A |
| 87062612 | 4 | 8.5 N | 85.8 E | 18.8 | N/A | N/A | 30 | 5 | N/A | N/R |
| 87062700 | 5 | 9.9N | 86.6E | 36.5 | N/A | N/A | 30 | 0 | N/A | N/A |

## APPENDIX I

## DEFINITIONS

BEST TRACK - A subjectively smoothed path, versus a precise and very erratic fix-to-fix path, used to represent tropical cyclone movement.

CENTER - The vertical axis or cone of a tropical cyclone. Usually determined by wind, temperature, and/or pressure distribution.

CYCLONE - A closed atmospheric circulation rotating about an area of low pressure (counterclockwise in the northern hemisphere).

EPHEMERIS - Position of a body (satellite) in space as a function of time; used for gridding satellite imagery. Since ephemeris gridding is based solely on the predicted position of the satellite, it is susceptible to errors from vehicle wobble, orbital eccentricity and the oblateness of the earth.

EXPLOSIVE DEEPENING - A decrease in the minimum sea-level pressure of a tropical cyclone of 2.5 $\mathrm{mb} / \mathrm{hr}$ for 12 hours or $5.0 \mathrm{mb} / \mathrm{hr}$ for six hours (Holliday and Thompson, 1979).

EXTRATROPICAL - A term used in warnings and tropical summaries to indicate that a cyclone has lost its "tropical" characteristics. The term implies both poleward displacement froom the tropics and the conversion of the cyclone's primary energy sources from release of latent heat of condensation to baroclinic processes. The term carries no implications as to strength, size or intensity.

EYE - A term used to describe the central area of a tropical cyclone when it is more than half surrounded by wall cloud.

FUJIWHARA EFFECT - An interaction in which tropical cyclones within $700 \mathrm{~nm}(1296 \mathrm{~km})$ of each other begin to rotate about one another. When intense tropical cyclones are within about $400 \mathrm{~nm}(741 \mathrm{~km})$ of each other, they may also begin to be drawn closer to each other (Brand, 1970) (Dong and Neumann, 1983).

INTENSITY - The maximum wind speed, typically within one degree of the center of the tropical cyclone.

MAXIMUM SUSTAINED WIND - Highest surface wind speed averaged over a one-minute period of time. Peak gusts over water average 20 to 25 percent higher than sustained winds.

RAPID DEEPENING - A decrease in the minimum sea-level pressure of a tropical cyclone of $1.25 \mathrm{mb} / \mathrm{hr}$ for 24 hours (Holliday and Thompson, 1979).

RECURVATURE - The turning of a tropical cyclone from an initial path toward the west or northwest to a path toward the northeast.

SIGNIFICANT TROPICAL CYCLONE - A tropical cyclone becomes "significant" with the issuance of the first numbered warning by the responsible warning agency.

SIZE - The areal extent of the tropical cyclone measured radially outward from the center (e.g., radius of the outer closed isobar ).

STRENGTH - The average wind speed of the inner low-level circulation, usually measured within one to three degrees of the center of the tropical cyclone.

SUPER TYPHOON (HURRICANE) - A typhoon or hurricane in which the maximum sustained surface wind (one-minute mean) is $130 \mathrm{kt}(67 \mathrm{~m} / \mathrm{sec}$ ) or greater.

TROPICAL CYCLONE - A non-frontal, migratory low-pressure system of usually synoptic scale developing over tropical or subtropical waters and having a definite organized circulation.

TROPICAL DEPRESSION - A tropical cyclone in which the maximum sustained surface wind (one-minute mean) is $33 \mathrm{kt}(17 \mathrm{~m} / \mathrm{sec})$ or less.

TROPICAL DISTURBANCE - A discrete system of apparently organized convection - generally 100 to 300 nm ( 185 to 556 km ) in diameter - originating in the tropics or subtropics, having a non-frontal migratory character and having maintained its identity for 12- to 24hours. It may or may not be associated with a detectable perturbation of the wind field. As such, it is the basic generic designation which, in successive stages of identification, may be classified as a tropical depression, tropical storm or typhoon (hurricane).

TROPICAL STORM - A tropical cyclone with maximum sustained surface winds (one-minute mean) in the range of 34 to 63 kt ( 17 to $32 \mathrm{~m} / \mathrm{sec}$ ) inclusive.

TROPICAL UPPER-TROPOSPHERIC TROUGH (TUTT) - A dominant climatological system (upper-level trough) and a daily synoptic feature, of the summer season, over the tropical North Atlantic, North Pacific and South Pacific Oceans (Sadler, 1979).

TYPHOON / HURRICANE - A tropical cyclone in which the maximum sustained surface wind (one-minute mean) is $64 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ or greater. West of 180 longitude degrees they are called typhoons and east of 180 degrees longiude they are called hurricanes. Foreign governments use these and other terms for tropical cyclones and may apply different intensity criteria.

WALL CLOUD - An organized band of cumuliform clouds immediately surrounding the central area of a tropical cyclone. The wall cloud may entirely enclose or partially surround the center.

## APPENDIX II

NAMES FOR TROPICAL CYCLONES

| Column 1 | Column 2 | Column 3 | Column 4 |
| :--- | :--- | :--- | :--- |
| ANDY |  |  |  |
| BRENDA | ABBY | ALEX | AGNES |
| CECIL | CARMEN | BETTY | BILL |
| DOT | DOM | CARY | CLARA |
| ELLIS | ELLEN | ED | DOYLE |
| FAYE | FORREST | FREDA | ELSIE |
| GORDON | GEORGIA | GERALD | FABIAN |
| HOPE | HERBERT | HOLLY | HAL |
| IRVING | IDA | IAN | IRMA |
| JUDY | JOE | JUNE | JEFF |
| KEN | KIM | KELLY | KIT |
| LOLA | LEX | LYNN | LEE |
| MAC | MARGE | MAURY | MAMIE |
| NANCY | NORRIS | NINA | NELSON |
| OWEN | ORCHID | OGDEN | ODESSA |
| PEGGY | PERCY | PHYLLIS | PAT |
| ROGER | RUTH | ROY | RUBY |
| SARAH | SPERRY | SUSAN | SKIP |
| TIP | THELMA | THAD | TESS |
| VERA | VERNON | VANESSA | VAL |
| WAYNE | WYNNE | WARREN | WINONA |

NOTE:
Names are assigned in rotation, alphabetically. When the last name (WINONA) has been used, the sequence will begin again with "ANDY".

SOURCE: CINCPACINST 3140.1 (series)

## APPENDIX III

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## APPENDIX IV

## PAST ANNUAL TROPICAL CYCLONE REPORTS

Copies of the past<br>Annual Tropical Cyclone<br>Reports<br>can be obtained through:<br>National Technical Information Service<br>5285 Port Royal Road<br>Springfield, Virginia 22161

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BLOCK 18 (CONTINUED)<br>TROPICAL CYCLONE BEST TRACK DATA<br>TROPICAL CYCLONE FORECASTING<br>AIRCRAFT RECONNAISSANCE<br>DYNAMIC TROPICAL CYCLONE MODELS<br>TYPHOON ANALOG MODEL<br>TROPICAL CYCLONE STEERING MODEL<br>CLIMATOLOGY/PERSISTENCE TECHNIQUES<br>TROPICAL CYCLONE FIX DATA

The Black Swan insignia of the 54 WRS Typfioon Chasers is included in dedication to the squadron's forty-thiree year's of support to the tropical cyclone forecast and warning mission. The nature and extent of their involvement with the USPACOM tropical cyclone warning system is best described by their last Commander, Lt. Col. Don $\mathcal{H}$. Owen, during fis 1987 Tropical Cyclonc Conference briefing, "As we deactivate on 1 October, we will not be flying with you, literally, but our fiearts and wishes for successful typfioon forecasting remain with you in the future."


Looking across the starboard wing and outboard engine of the 54th Weather Reconnaissance Squadrom (54 WhS) WC-130 aircraft, the camera captures a section of the spectacular inner wall of Typhoon Wynne's (07W) eye. The tight curvature of the eye wall cloud shows across the top of the image. (The pentagons near the center of the photograph are due to sunlight glare off the Lens diaphragm of the camera.) At the time the picture was taken, 260010Z $3 u l y$ 1987, the aircraft reconnaissance mission ( $\mathcal{A F} 8610507$ WYON()(E) was fixing the eye over the Philippine Sea, $250 \mathrm{~nm}(463 \mathrm{~km}$ ) north of Guam and $48 \mathrm{~nm}(89 \mathrm{~km}$ ) west of the isfand of Pagan in the northern Marianas. The circular eye diameter was $16 \mathrm{~nm}(30 \mathrm{~km}$ ) and the minimum sea-Level pressure at that time was 922 mb (Photo courtesy of Detackiment 3, 1st Weather Wing and Kenneth W. Rease, Captain, USMF).


The Air Force Air Weather Service celebrated its 50th anniversary on 1 Iufy 1987. The insignia that was designed for the occasion is displayed here in recognition of Detackment 1, 1st Weather Wing's enormous contribution to the $\mathcal{N}$ (AVOCEAXCOMCEN(/ITTWC tropical cyclone forecast and warning support to the U. S. Pacific Command. Twenty-eight years of cooperation between Navy and Air Force personnel working at $\mathcal{N A V O C E A X C O M C E N} /$ ITWC Gave resulted in one of the finest tropical cyclone forecasting centers in the world.


[^0]:    $1_{\text {Perronal }}$ commonication with Mi. Paol M. Hattori of the U.S. Geological Surwey (Depertment of Intorior) revealed the following event of inverest. Pior to $251145 Z$ July thore was no increase in micromeismic activity on the World Wide Stmonardizod Seimograph Network (WWSSN)
     E), Gum. From $251145 Z$ to $251800 Z$, the shart-period microscirms increased nopiceabiy and
     Atwor 2518002 , the background renumed to epproximately pro-251145Z levola. Also, the lonk-
    period microseisms epproximately doubled in amplitude from 2510002 to $251700 Z$. The period microseisms epproximetaly doubled in amplitude from 2510002 to 2517002 . The microsim being in the anst-weat conrponent manoe, the postibility exints that this mique transit of the Merina Tronch. Similar activity has been obelarved with other tropical cycloneas but with less clarity. With an absence of other organized weather systems in the arca, the typhocris compact sise mend distmes from the WWSSN station on Gram reselved in en unpasully cloar recording of Wymid's approch to the northern Merianal.

[^1]:    * JTWC WARNING RESPONSIBILITY BEGAN ON 4 JUN 71 FOR THE BAY OF BENGAL, EAST OF 90 DEGREES EAST LONGITUDE. AS DIRECTED BY
    CINCPAC, JTWC ISSUED NARNINGS ONLY FOR THOSE TROPICAL CYCLONES THAT DEVELOPED OR TRACKED THROUGH THAT PORTION OF THE BAY OF CINCPAC, JTWC ISSUED WARNINGS ONLY FOR THOSE TROPICAL CYCLONES THAT DEVELOPED OR TRACKED THROUGH THAT PORTION OF THE BAY OF THE WRSTERN PORTION OF THE BAY OF BENGAL AND THE ENTIRE ARABIAN SEA,

    FORMATION ALERTS: 7 OE 8 FORMATION ALERTS DEVELOPED INTO SIGNIFICANT TROPICAL CYCLONES. TROPICAL CYCLONE FORMATION ALERTS WERE ISSUED FOR ALL OF THE SIGAIFICANT TROPICAL CYCLONES THAT DEVELOPED IN 1987, EXCEPT TROPICAL CYCLONE O3A.

    MARNINGS:

    | NUMBER OF CALENDAR WARNING DAYS: | 26 |
    | :--- | :--- |
    | NUMBER OF CALENDAR WARNING DAYS |  |
    | WITH TWO TROPICAL CYCLONES: | 1 |
    | NUMBER OF CALENDAR WARNING DAYS |  |
    | WITH THREE TROPICAL CYCLONES: | 0 |

[^2]:    * REGENERATED
    ** OVERLAPPING DAYS INCLUDED ONLY ONCE IN SUM.

