Planetary magnetic fields.

Observations, theory, models.

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1– Intro • Magnetic field as a signal carrying information about constitution dynamics of the interior thermal history of a planetary system. • Planetary dynamo modelling aims at retrieving information by confronting numerical models and observations investigating the theoretical difficulties of the dynamo problem.

2– Plan

- The geodynamo: where it all started
- Mars: mysteries of the lost dynamo
- Uranus/Neptune: remote is exotic
- Jupiter/Saturn: zonal flows and dynamos



4– Modelling

• Spherical shell geometry, rotation and magnetic induction are the 3 essential ingredients. One source of kinetic energy (most likely convection) is required.



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5– Solving the equations

 Navier Stokes (Boussinesq) equations + Maxwell equations = magnetohydrodynamic model





• explains the large scale space (> 1000 km) and time (>100 yrs) features of the geodynamo



Model gufm1 for 1990



Glatzmaier-Roberts simulation



• Surface measurement allows to infer inner magnetic structure and confirm dynamo mechanism.

8– Challenges for future modelling

• Are models wrong or lowpass-filtering reality? Need to have accurate ratio between viscous, thermal, compositional, magnetic time scales.





Mars Global Surveyor at $200 \ \mathrm{km}$ altitude, Purucker et al.

10– Interior of Mars

• Mars has half the Earth's radius and an iron core which is at least partially liquid.



11– Scaling of dynamo constraints: velocity amplitude

- Magnetic Reynolds number $Re_m = UD/\lambda$ has to exceed 100 for a dynamo to work.
- convective velocities U have been scaled with heat flux Q in lab experiments:

$$U = \left(\frac{\alpha g}{\rho C_p \Omega D^3}\right)^{2/5} Q^{2/5} \qquad \begin{array}{c} & & & \\ &$$

- dynamo condition implies Q > 1 GW which is very little above adiabatic, much less than typical adiabatic heat flux (~ 1 TW).
- So if there is core convection, a dynamo is extremely likely!

12– Scaling of dynamo constraints: Ohmic losses

- A working dynamo produces a lot of ohmic heat.
- Joule heating power Q_J has been scaled in numerical models:

$$au = rac{E_m}{Q_J} \propto rac{1}{Rem}$$

An asymptotic regime is reached (independant of the kinematic/magnetic diffusivity ratio): small-scale eddies dissipate a large-scale field.





• $Q_J = 1 - 10$ TW has to be extracted from the energy sources. The onset of a dynamo is coincident with a sudden "power hunger" (cf Karlsruhe experiment).

13– Possibilities?

- Why hasn't Mars a dynamo today while Earth has one?
- Death of convection? Possible if plate tectonics stopped, if inner core formation was delayed in comparison to the Earth.
- Possible later turn-on of the dynamo, stopped due to a frozen core.
- In any case the thermal history of the core needs to be clarified. Q_J is a central unknown for this history.
- Scaling will help to better estimate dissipation from surface observations.
- Present models have a high potential because they show unexpected asymptotic convergence, a hint of the closeness with the physics of real systems.

14– Remote & exotic: Uranus and Neptune



surface radial fields in gauss, Holme and Bloxham, Voyager II.

- Surface fields have a singular equatorial dipole + multipoles content.
- two instances mean that we are not "catching" a reversal.



• Dynamo action could take place in the middle conducting fluid ice layer.

• Planetary dynamos are believed to work in a

Coriolis force
$$\sim$$
 Lorentz force

equilibrium regime.

• The Elsasser number

$$\Lambda = rac{\sigma B^2}{
ho\Omega}$$

measures the ratio of the two. This should always be of order 1 and provide a convenient way of scaling planetary magnetic fields...

• ... but $\Lambda\sim 1$ in the Earth, $\Lambda\sim 0.01$ in Uranus/Neptune. How can this be? More generally, how does an equatorial dipole dynamo work?



Magnetic tension in the normal/tangential vector basis

$$T=rac{B^2}{R_c}\mathrm{e}_n+rac{\partial}{\partial s}\left(rac{B^2}{2}
ight)\mathrm{e}_s,$$

In this representation field line thickness is weighted with B^2 .

• Where thickness varies, or where thick lines are curved, work is done against magnetic tension.

18– Model predictions

• This model can be perturbed to yield either an axial, or an equatorial dipole.



• Dynamo is not efficient when dipole axis and rotation axis are orthogonal because of field line shear by vortices. Hence $\Lambda_e/\Lambda_a=0.05.$



• Exotic dynamos with equatorial dipoles and multipoles show when the dynamo shell is small. This could be a constraint for Uranus and Neptune.

20– Zonal flows in the gas giants



• The surface, and probably the deeper dynamo region of gas giants is swept by strong zonal flows.

21– The origin of zonal flows



Potential vorticity

$$q=rac{\omega+2\Omega}{H}$$

q is materially conserved and therefore develops regions with flat profiles. This means nonzero ω and therefore zonal flow.

- Zonal flows are the result of potential vorticity mixing by turbulence.
- This simply describes exchanges of angular momentum between the rotating frame and the fluid within.

22– Dynamos with zonal flows

• The magnetic field of Saturn is highly axisymmetric, but Cowling theorem prevents dynamos producing axisymmetric fields.



- The surface field looks axisymmetric on surface,
- ...but is not deeper in the shell where dynamo action takes place.
- This dynamo requires free-slip boundaries and has a different mechanism from the previous models.

23– Finally...

- Other magnetic analyses have increased the payload of space missions, for instance the discovery of liquid water oceans under the surface of jovian satellites.
- Present numerical modelling has a large potential, especially with the upcoming planetary magnetic observations.
- The knowledge of the Earth dynamo benefits from this of other natural dynamos.
- Dynamo processes are intimely connected to surface and mantle geodynamics.