

and the corresponding gene flow¹⁰, the nuclear DNA data nonetheless suggest that a possible male-mediated gene flow between colonies is insufficient to prevent a genetic substructuring even of closely neighbouring colonies.

The consequences for conservation management are obvious. The pre-programmed female reproductive behaviour makes it unlikely that the loss of a breeding habitat (because of human building operations, for example) can be compensated for by emigration to other colonies; that is, the loss of nesting sites is accompanied by the loss of specific genotypes. Thus, to preserve the genetic diversity of the *Caretta* metapopulation one needs to preserve individual nesting sites (attempts to transfer freshly deposited eggs between sites — based on the hope that hatchlings would accept a new hatching site as natal — are controversial). In addition, the described haplotypes may serve as genetic tags to identify the colony membership of individuals offshore or illegally commercialized, and also to complete our knowledge of sea turtle migratory behaviour offshore¹¹.

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1. Carr, A. F. *BioScience* **14**, 49–52 (1964).
2. Lohmann, K. J. & Lohmann, C. M. F. *Nature* **380**, 59–61 (1996).
3. Bowen, B. W. & Avise, J. C. in *Conservation Genetics: Case Histories from Nature* (eds Avise, J. C. & Hamrick, J.) 190–237 (Chapman and Hall, New York, 1995).
4. Allard, M. W. *et al. Copeia* **1994**, 34–41 (1994).
5. Bowen, B. W. *BioScience* **45**, 528–534 (1995).
6. Williams, J. G. K. *et al. Nucleic Acids Res.* **18**, 6531–6535 (1990).
7. Hadrys, H., Ballick, M. & Schierwater, B. *Mol. Ecol.* **1**, 55–63 (1992).
8. Laurent, L. *WWF Report* (International Mediterranean Program RAC/SPA, Greenpeace Mediterranean Sea Project, 1994).
9. Bowen, B. W. *et al. Conserv. Biol.* **7**, 834–844 (1993).
10. Karl, S. A., Bowen, B. W. & Avise, J. C. *Genetics* **131**, 163–173 (1992).
11. Schierwater, B. & Ender, A. *Nucleic Acids Res.* **21**, 4647–4648 (1993).
12. Excoffier, L., Smouse, P. E. & Quattro, J. M. *Genetics* **131**, 479–491 (1992).

Erratum

In the Scientific Correspondence “Competition for royalty in bees” by R. F. A. Moritz, Per Kryger and Mike H. Allsopp (*Nature* **384**, 31; 1996), the y-axis labels of Fig. 2, showing the mean offspring frequencies of various patrines of a honeybee colony, were printed incorrectly. In the top panel the scale should have read ‘0–20%’, not ‘0–2%’ as printed, and in the middle and bottom panels ‘0–80%’ rather than ‘0–8%’. □

Human effect on global climate?

SIR — The recent pattern-correlation analysis of Santer *et al.*¹ has drawn considerable attention: Nicholls², for example, stated that it represents the “clearest evidence yet that humans may have affected global climate”. This conclusion is based on the increasing similarity of the vertical temperature pattern in free-atmosphere and global-climate models incorporating combinations of carbon dioxide, sulphate aerosols and stratospheric ozone concentrations as measured by a ‘centred’ correlation statistic, $R(t)$.

Santer *et al.*¹ found significant increases in $R(t)$ in the 850–50-hPa region of the atmosphere over the period of their study (1963–87) for each of the climate model/atmospheric constituent combinations they analysed (with the exception of a sulphate-only model). We agree with Santer *et al.*¹ that this result stems largely from the pattern of stratospheric and upper tropospheric cooling and a hemispheric asymmetry in the lower and mid-troposphere. But we believe that the reported increases in $R(t)$ are almost totally caused by the higher-altitude temperature changes, and that in the lower levels the results are only a reflection of the time period chosen.

For example, using the carbon dioxide + sulphate model results from Taylor and Penner³ (Fig. 2 of ref. 1), we can compare the explained variance (obtained by squaring $R(t)$) at the end of the record (1987) between the 850–50-hPa layer (troposphere + stratosphere) to that for the 850–500-hPa layer (lower troposphere only). The values are 64% and 5%, respectively. Thus 92% of the explained variance results from the addition of data above 500 hPa. This result is typical for the models used in this study (with the exception of the sulphate-only model). Further, the calculation of $R(t)$, using four levels above 500 hPa and only two beneath, assures heavy dependence upon the upper troposphere and stratosphere.

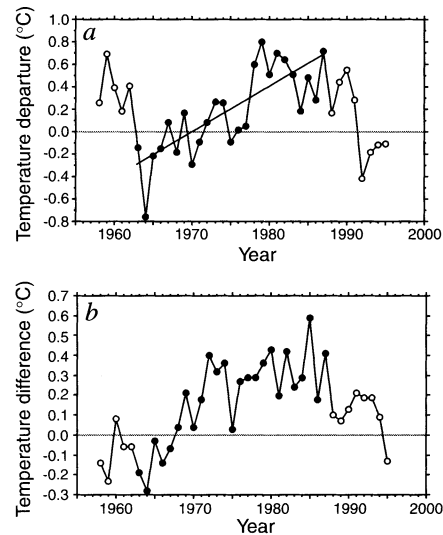
The factor that largely constitutes the hemispheric asymmetry is the warming of the mid-troposphere in the Southern Hemisphere, which appears both in the CO₂ + sulphate models (Fig. 1c, e, g–i of ref. 1) and in the observations, where it is concentrated between 850 and 300 hPa, 30 to 60° S (Fig. 1j of ref. 1). The observed data in ref. 1 are from the record of Oort and Liu⁴. This record extends to 1989 (although annual data to only 1987 are used by Santer *et al.*¹). The radiosonde record of Angell⁵ is longer (1958–95). Over the period of concurrency between refs 1 and 5, the correlation coefficient between annual anomaly values is 0.94 in the Northern Hemisphere and 0.92 in the Southern Hemisphere.

When we examine the period of record

used by Santer *et al.* in the context of the longer period available from ref. 5, we find that in the region with the most significant warming (30–60° S, 850–300 hPa) the increase is largely an artefact of the time period chosen (*a* in our figure). Although there is a statistically significant warming in the period from 1963–87, there is no significant change in the entire (1958–95) record. This has considerable bearing on the portion of $R(t)$ in ref. 1 that emanates from Southern Hemisphere mid-tropospheric warming. This result cannot be an artefact of the data, as precisely the same set of stations was used by Angell during the entire record.

Additionally, as $R(t)$ is a pattern-matching statistic, the increases in the tropospheric contribution to it are strongly dependent on the hemispheric temperature differences. The climate models predict that the Southern (sulphate-free) Hemisphere should warm relative to the Northern (sulphate-laden) Hemisphere, and for the period of record used in this study (1963–87) the observations agree. However, a close examination of upper-air temperature records again exposes this agreement as being fortuitous.

Using data from ref. 5, we present the 850–300-hPa hemispheric temperature differences (*b* in our figure). These data are highly correlated with the tropospheric $R(t)$ values calculated in ref. 1 from the CO₂ + sulphate model in ref. 3 ($r=0.80$).



a, The Angell⁵ temperature departure (°C) history for 1958–95 for the 850–300-hPa layer between 30 and 60° S latitudes. Solid circles, years studied in ref. 1. There is a statistically significant ($P<0.0001$) increase during 1963–87, whereas the overall record exhibits no significant trend. *b*, Annual temperature difference (°C) between Southern and Northern hemispheres in the 850–300-hPa layer, from ref. 5. There is a statistically significant ($P<0.05$) downward trend in this data since 1972. Solid circles, period studied by Santer *et al.*

On the basis of this relationship, estimates of $R(t)$ for the period 1988–95, coupled with the observed values in ref. 1, yield a significant downward trend since the early 1970s. This is typical of the models tested.

Even though global temperatures were dramatically reduced by the eruption of Mount Pinatubo in the latter part of this period (1991), the hemispheric temperature difference was relatively unaffected. Therefore, we believe that the eruption would have little effect on the tropospheric values of $R(t)$.

In conclusion, we suggest that the increasing pattern correlation between climate models and observations found by Santer *et al.* is primarily governed by a signal in the upper troposphere and lower stratosphere, and that in the lower and mid-troposphere, the models and observations have been drifting further apart since the early 1970s. Such a result in the troposphere cannot be considered to be a 'fingerprint' of greenhouse-gas-induced climate change. It is therefore apparent that statements about the strength of the evidence for human alteration of lower tropospheric climate must be tempered in the light of more complete data than were analysed by Santer *et al.*

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SIR — Santer *et al.*¹ present a statistical analysis in which they compare the model-generated zonal mean vertical thermal structure of the atmosphere in response to the concomitant increase of the concentration of greenhouse gases, sulphur emissions and the observed decrease in stratospheric ozone to the observed thermal structure of the atmosphere between 1963 and 1987. They find that the pattern correlation between the predicted and observed changes in zonal mean latitude height profiles of atmospheric temperature increases with time.

The authors attributed the largest amplitude signals of those trends primarily to two factors: first, the pattern comparison over 50 to 850 hPa (the signal of modelled tropospheric warming and stratospheric cooling); and second, in the troposphere, between 500 and 850 hPa, the disparity between Southern and Northern hemispheric warming due to modelled sulphate aerosol effects.

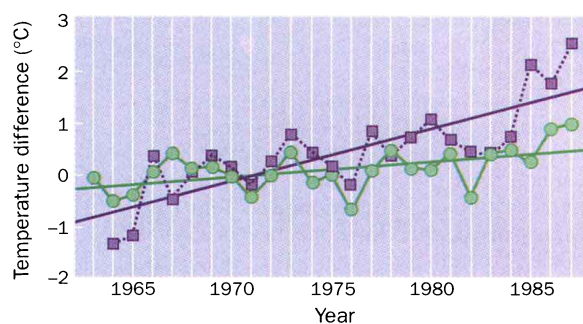
A well-known tropospheric and stratospheric temperature history for various latitudinal bands between 1958 and 1992 has already been published by Angell⁵; coverage for the Southern Hemisphere stratosphere begins in 1964. Angell's data of the layer 50–100 hPa closely correspond to the stratospheric data used by Santer *et al.*; similarly, Angell's 300–850-hPa data capture the bulk of the troposphere. A signal

of tropospheric warming and stratospheric cooling should be visible as an increasing difference between tropospheric warming and stratospheric cooling.

Taking the first point, a trend analysis of Angell's lower stratospheric data for the period 1963–87 (or 1964–87) shows a much larger cooling rate in the Southern than in the Northern Hemisphere (see figure). It appears, therefore, that most (stratospheric cooling in the Southern Hemisphere is about three times as large as in the Northern Hemisphere) of the signal pattern strength in 1963–87 relating to stratospheric/tropospheric trend differences originates in the Southern Hemisphere. This result seems to hold for the 1958–92 period as well, and for trends originating in the mid-1970s and later, which do not show any tropospheric warming in the mid-latitudes of the Southern Hemisphere, but sharply increasing stratospheric cooling and therefore increasing trend differences and strengthening of a pattern of stratospheric/tropospheric differences.

As Santer *et al.* point out, and as dynamical modelling results seem to suggest^{6–9}, the cooling of the Southern Hemisphere stratosphere, which increased sharply after about 1983 (ref. 5), may be related to stratospheric ozone depletion, which is most pronounced in the higher latitudes, but also occurs in the mid-latitudes of the Southern Hemisphere¹⁰. Therefore, the increasing signal pattern strength reported by Santer *et al.*¹ may primarily be related to Southern Hemisphere stratospheric cooling linked to ozone depletion due to CFCs (chlorofluorocarbons)^{6–9}. The possible human-induced climate effect alluded to by Santer *et al.* could then largely be attributed to stratospheric cooling by CFCs but not to the warming effect of anthropogenic greenhouse gases.

Turning to the second point, Santer *et al.* report a hemispheric-scale asymmetrical warming signal in 1963–87 only in those



Tropospheric and stratospheric temperature differences in the Northern (green symbols) and Southern (purple symbols) Hemispheres between 1963 and 1987, adapted from ref. 5. The least-squares linear trends (solid lines) are increases of 0.97 °C per decade in the Southern Hemisphere, and 0.27 °C per decade in the Northern Hemisphere.

modelling experiments that include cooling sulphate aerosol effects. The implication of this result is that the lack of sulphate emissions in the Southern Hemisphere has led to a larger tropospheric warming there.

This interhemispheric difference should be largest in the mid-latitudes, where most of the sulphur emissions occur (Fig. 1 of ref. 1). Although there is a larger 1963–87 warming in Angell's 300–850 hPa mid-latitude data (30–60 °C in both hemispheres) as well, it does not seem to be a permanent feature of the climate system in recent decades. A comparison between Northern and Southern Hemisphere trends ending in 1992 shows that trends beginning after 1964 imply increasingly greater warming in the Northern Hemisphere than in the Southern Hemisphere. This would be incompatible with increasing Northern Hemispheric cooling due to rising sulphur emissions there.

Regarding the role of natural factors, the early years of the period 1963–87 were substantially influenced by tropospheric cooling (and stratospheric warming) following the eruption of Mount Agung¹¹, whereas the end of that period was influenced by several strong El Niño events¹², which have led to some tropospheric warming and stratospheric cooling, particularly in the southern subtropics of the lower latitudes⁵. Therefore, the general tropospheric warming and stratospheric cooling trend between 1963 and 1987 has

- Santer, B. D. *et al.* *Nature* **382**, 39–46 (1996).
- Nicholls, N. *Nature* **382**, 27–28 (1996).
- Taylor, K. E. & Penner, J. E. *Nature* **369**, 734–737 (1994).
- Oort, A. H. & Liu, H. J. *J. Climatol.* **6**, 292–307 (1993).
- Angell, J. K. in *Trends '93: A Compendium of Data on Global Change* (eds Boden, T. A. *et al.*) 636–672 (ORNL/CDIAC, Oak Ridge, TN, 1994).
- Volk, C. M. *et al.* *Science* **272**, 1763–1768 (1966).
- Kiehl, J. T. *et al.* *Nature* **332**, 501–504 (1988).
- Atkinson, R. J. *et al.* *Nature* **340**, 290–294 (1989).
- Schwarzkopf, M. D. & Ramaswamy, V. *Geophys. Res. Lett.* **20**, 205–208 (1993).
- Stolarski, R. *et al.* *Science* **256**, 342–349 (1992).
- Angell, J. K. *J. Geophys. Res.* **93**, 3697–3704 (1988).
- Climate Diagnostic Bulletin August 1989* 6 (US Dept Commerce, Washington DC, 1989).
- Angell, J. K. *J. Climatol.* **1**, 1296–1313 (1988).
- Angell, J. K. *Geophys. Res. Lett.* **17**, 1093–1096 (1990).
- Gaffen, D. J. *J. Geophys. Res.* **99**, 3667–3676 (1994).
- Trenberth, K. E. & Olson, J. G. in *Greenhouse-Gas-Induced Climatic Change: A Critical Appraisal of Simulations and Observations* (ed. Schlesinger, M. E.) 249–259 (Elsevier, Amsterdam, 1991).
- Parker, D. E. & Cox, D. I. *Int. J. Climatol.* **15**, 473–496 (1995).
- Christy, J. R. *Clim. Change* **31**, 455–474 (1995).
- Kalnay, E. *et al.* *Bull. Am. Meteorol. Soc.* **77**, 437–471 (1996).
- Tett, S. F. B., Mitchell, J. F. B., Parker, D. E. & Allen, M. R. *Science* **274**, 1170–1173 (1996).
- Wigley, T. M. L. *Nature* **339**, 365–367 (1989).
- Wigley, T. M. L., Jaumann, P. J., Santer, B. D. & Taylor, K. E. *Clim. Dynam.* (in the press).
- Ramaswamy, V., Schwarzkopf, M. D. & Randel, W. J. *Nature* **382**, 616–618 (1996).
- Wigley, T. M. L. & Raper, S. C. B. *Nature* **357**, 293–300 (1992).
- Kattenberg, A. *et al.* in *Climate Change 1995: The Science of Climate Change* (eds Houghton, J. T. *et al.*) 285–357 (Cambridge Univ. Press, Cambridge, 1996).