

## Supplementary Information

*Subtropical Arctic Ocean temperatures during the Palaeocene-Eocene thermal maximum. Appy Sluijs, Stefan Schouten, Mark Pagani, Martijn Woltering, Henk Brinkhuis, Jaap S. Sinninghe Damsté, Gerald R. Dickens, Matthew Huber, Gert-Jan Reichart, Ruediger Stein, Jens Matthiessen, Lucas J. Lourens, Nikolai Pedentchouk, Jan Backman, Kathryn Moran and the Expedition Scientists*

### Relative stratigraphic position of Core 31X

The stratigraphic position of the 55 cm recovered section of Core 302-4A-31X relative to Cores 30X and 32X is uncertain (Fig. 2). It may be located anywhere in the interval between 384.54 and 388 meters composite depth (mcd). For illustration purposes, we have placed this core 1m lower than indicated in ref. 1.

### Description and calibration of $\text{TEX}_{86}'$

Originally  $\text{TEX}_{86}$  values were calculated as described in ref. 2 and converted to sea surface temperature (SST). However, the glycerol dialkyl glycerol tetraether (GDGT) lipid containing three cyclopentane rings (GDGT-3), almost always the least abundant GDGT lipid in marine sediments, was unusually high in Core 302-4A-32X (Fig. S-2A). Since terrestrial organic matter also contains isoprenoid GDGT lipids with cyclopentane rings<sup>3</sup> (and Weijers, J Schouten, S., Sinninghe Damsté, J.S., unpublished results), it is likely that the high terrestrial input in this section has disturbed the aquatic signal. The high abundance of GDGT-3 resulted in unusually high  $\text{TEX}_{86}$  values in Core 32X. To circumvent this problem,  $\text{TEX}_{86}'$  was devised which has the same definition as  $\text{TEX}_{86}$  except that isomer GDGT-3 was removed from the denominator<sup>2</sup>. The  $\text{TEX}_{86}'$  was determined for 104 marine surface sediments and found to correlate very well with annual mean SST:  $\text{TEX}_{86}' = 0.016 \times \text{SST} + 0.20$  with  $R^2 = 0.93$  (Fig S-2B). This equation was used to convert  $\text{TEX}_{86}'$  into SST. Sediments from Cores 30X and 31X show a normal marine tetraether lipid distribution, i.e., showing a very minor peak at GDGT-3, resulting in the  $\text{TEX}_{86}'$  values only slightly different from  $\text{TEX}_{86}$  values (Fig SI-2C). Interestingly, the top part of Core 30X, which shows a large terrestrial influence (Fig. 2), shows again a relatively large offset between  $\text{TEX}_{86}'$  and  $\text{TEX}_{86}$ , evidencing a larger contribution of the GDGT-3. This is consistent with a terrestrially-derived contribution to the GDGT-3 peak.

## Which temperature does TEX<sub>86</sub>' indicate in the Arctic Ocean?

TEX<sub>86</sub> and TEX<sub>86</sub>' are calibrated by core top analysis to mean annual mean SST. This empirical relation is not the same as a causal relationship. Several studies have shown that the cell number of crenarchaeota in the water column strongly depends on the season<sup>4</sup>. In most studies a negative correlation has been observed between the cell abundance of crenarchaeota and phytoplankton, likely because they compete for the nutrient ammonia. As crenarchaeota are chemoautotrophic organisms<sup>5</sup> and thus not directly depend on light it is likely that crenarchaeota predominantly thrived during times of low cell abundances of phytoplankton and low light intensities, i.e. in the Arctic winter. However, for the crenarchaeotal signal to reach the sediment floor a significant sedimentation flux is needed. In the present day ocean significant organic carbon fluxes are observed during periods of high phytoplanktonic productivity because an active food web leads to grazing and fecal pellet packaging<sup>6,7</sup>. In agreement with this we recently recorded the highest fluxes of crenarchaeotal lipids in the Arabian Sea coinciding with the periods of high productivity, despite that their relative abundance was lower in these periods (Wuchter C., Schouten S., Wakeham, S.G. and Sinninghe Damsté, J.S., unpublished results). In the Eocene Arctic Ocean the bulk of the signal will be derived from those periods with highest biological productivity, which at this high latitude must be the summer months. Hence we refined our interpretation of the TEX<sub>86</sub>' estimates by suggesting that its signal in the Eocene Arctic Ocean, although calibrated to globally-derived annual mean SSTs, may record the, on average, higher summer SSTs due to the flux-dependency of the signal.

## Preservation of the organic matter used in this study

Water column anoxia, in this paper indicated by the laminated sediments and the presence of isorenieratene derivatives, will lead to an improved preservation of certain organic compounds<sup>8</sup>. Below, we discuss the preservation of the different types of organic matter used in this paper and conclude that selective preservation did not affect our proxy records.

The presence of isorenieratene derivatives has been frequently used to detect anoxia extending up into the photic zone of the water column during the Phanerozoic<sup>9,10</sup>, including late Pliocene sapropel formation in the eastern Mediterranean<sup>11</sup>. These compounds are derived from photosynthetic green sulfur bacteria which require both light and free sulfide (i.e. euxinic conditions) to thrive. In the particular, these organisms are found in lakes and marine settings where the water column is euxinic such as the present day Black Sea<sup>12</sup>. Accordingly, the presence of isorenieratene

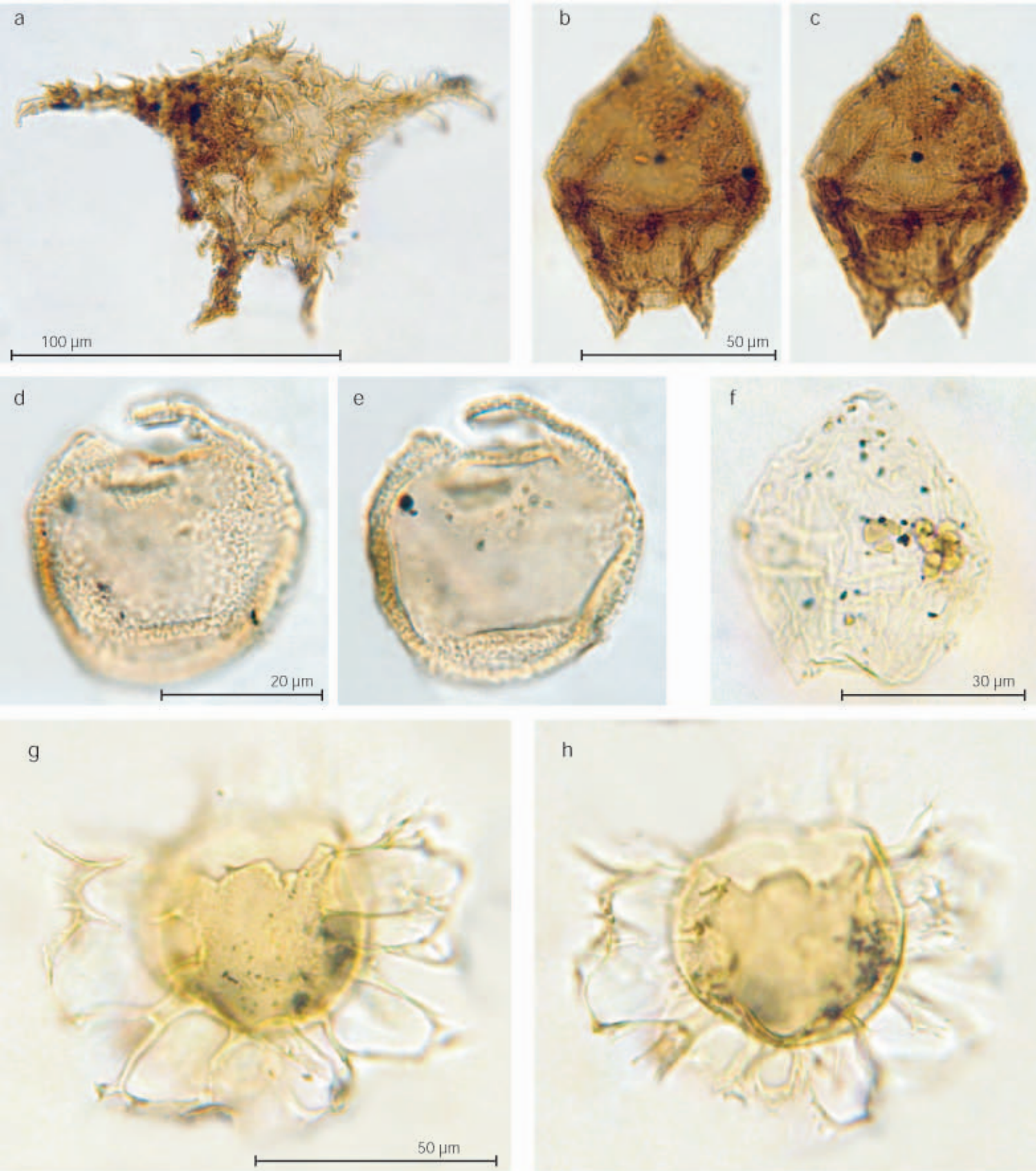
derivatives in ACEX Cores 30X and 31X indicates that euxinic conditions prevailed in the photic zone in the Arctic ocean during the PETM. These compounds preserve quite well, even when the water column is re-oxygenated and sediments are inhabited by benthic organisms<sup>13</sup>. Therefore we exclude selective preservation as possible bias on the presence/absence pattern of these compounds. Moreover, well preserved palynomorphs and substantial organic matter concentrations (~2% TOC) outside the laminated interval with isorenieratene derivatives, indicates that no severe organic matter degradation occurred. In this respect, we also exclude a bias on the terrestrial vs. marine palynomorph ratio, since it has recently been shown that preferential preservation of terrestrial palynomorphs only occurs in well-oxygenated bottom water settings<sup>14</sup>.

The terrestrial and marine compounds used in the BIT index represent structurally very similar compounds<sup>15</sup>. It has been shown that the relative distribution of the different isomers of glycerol dialkyl glycerol tetraethers (GDGTs) is independent of the oxygen concentration in the water column<sup>16</sup>. Thus, the absolute amounts of GDGTs may have increased within the laminated interval, but their relative distribution, i.e. the BIT index, has remained unaffected.

### **Bottom water anoxia at the PETM**

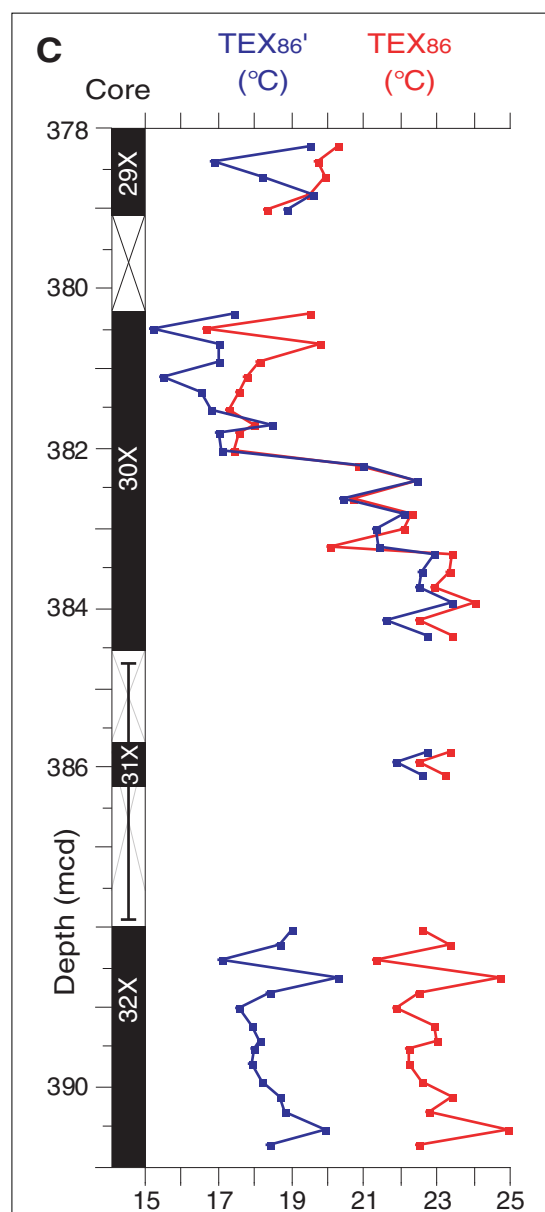
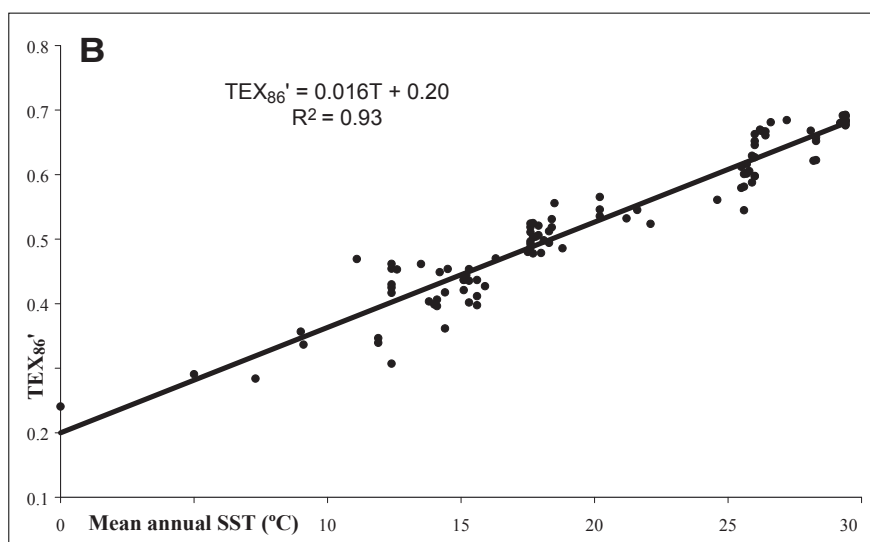
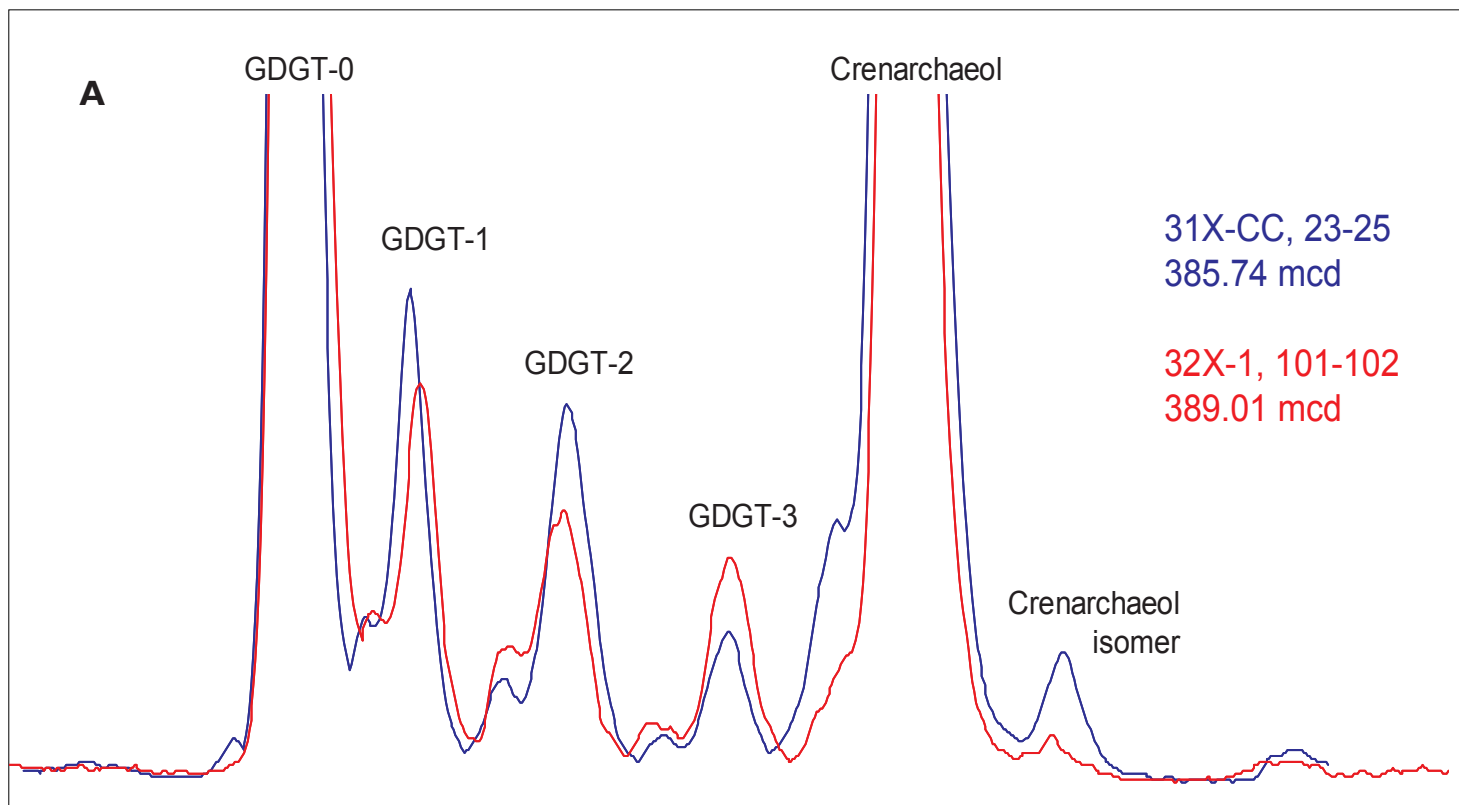
Organic linings of benthic foraminifera are common through the latest Palaeocene and earliest Eocene except during the photic zone euxinia, indicated by isorenieratane (Fig. S-3A). Sediments from the latter interval are laminated (Fig. S-3B), which implies that no bioturbation occurred after deposition of the sediments and that bottom waters were anoxic. Despite a large scatter % TOC is on average ~1.3% higher during the PETM compared to the latest Palaeocene. Unfortunately, due to the core recovery problems and potential changes in siliciclastic sediment supply related to the transgression there is relatively poor grip on sediment accumulation rates across the studied interval. However, some of the enhanced % TOC may be due to increased nutrient supply by rivers, resulting in higher phytoplankton production and, under anoxic bottom water conditions, high organic matter accumulation.

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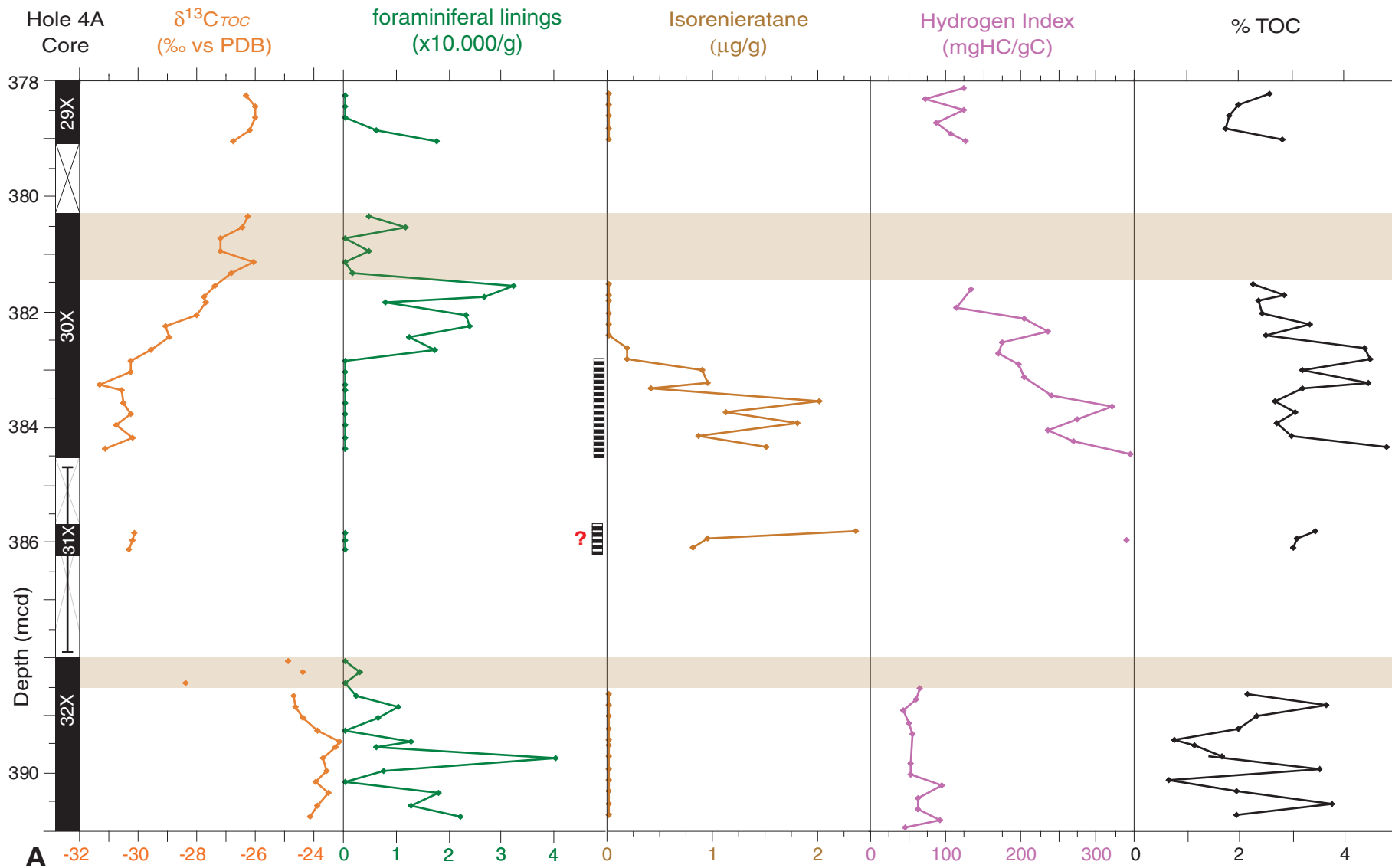


**Fig. S-1.**

Light microscope photos of dominant dinocyst taxa across the PETM of IODP Hole 302-4A. **a.** *Apectodinium augustum* (302-4A-30X, 101-103 cm; Slide 1, England Finder Coordinates S32-1); **b, c.** *Cerodinium wardenense* (302-4A-30X-3, 81-83 cm; Slide 1, F37-0); **d, e.** *Membranosphaera* spp. (302-4A-30X-1, 141-143 cm; Slide 1, L31-3); **f.** *Senegalinium* spp. (302-4A-30X-3, 101-103 cm; Slide 1, L37-3); **g, h.** *Glaphyrocysta* spp. (302-4A-30X-1, 141-143 cm; Slide 1, J43-4).



**Figure S-2. TEX<sub>86</sub>' calibration** **A)** HPLC/MS base peak chromatograms showing the chrenarchaeotal membrane lipid distribution of samples from Cores 31X and 32X. **B)** Calibration of TEX<sub>86</sub>' to mean annual sea surface temperature. **C)** Comparison between TEX<sub>86</sub>' and TEX<sub>86</sub> across the PETM of Site 302-4A.



**Figure S-3.** Additional indicators for water column euxinia and terrestrial influence. **A)** Isorenieratane and foraminifer lining distribution through the latest Palaeocene - earliest Eocene. Laminated interval indicated by stripes. Laminae are unclear in Core 31X due to drilling disturbance. The hydrogen index shows a substantial increase during the PETM, consistent with increased aquatic versus terrestrial organic matter. Despite significant scatter, average % Total Organic Carbon (TOC) increases by ~1% at the PETM. **B)** Core picture (interval 302-4A-30X-3, 123-129 cm) showing laminations.

