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## Images of Asteroid 21 Lutetia: A Remnant Planetesimal from the Early Solar System

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Images obtained by the Optical, Spectroscopic, and Infrared Remote Imaging System (OSIRIS) cameras onboard the Rosetta spacecraft reveal that asteroid 21 Lutetia has a complex geology and one of the highest asteroid densities measured so far,  $3.4 \pm 0.3$  grams per cubic centimeter. The north pole region is covered by a thick layer of regolith, which is seen to flow in major landslides associated with albedo variation. Its geologically complex surface, ancient surface age, and high density suggest that Lutetia is most likely a primordial planetesimal. This contrasts with smaller asteroids visited by previous spacecraft, which are probably shattered bodies, fragments of larger parents, or reaccumulated rubble piles.

he European Space Agency's Rosetta mission flew by asteroid Lutetia on 10 July 2010, with a closest approach distance of 3170 km. Lutetia was chosen because of its size and puzzling surface spectrum (1, 2). The Optical, Spectroscopic, and Infrared Remote Imaging System (OSIRIS) on board Rosetta (3) took 462 images, in 21 broad- and narrowband filters extending from 240 to 1000 nm, through both its narrow-angle camera (NAC) and wideangle camera (WAC). These images covered more than 50% of the asteroid surface, mostly of the northern hemisphere (figs. S1 and S2). The resolved observations started 9 hours 30 min before the closest approach (CA) and finished 18 min after CA. At CA, the asteroid filled the field of view of the NAC with a spatial scale of ~60 m per pixel. The observations reveal a morphologically diverse surface, indicating a long and complex history.

We modeled the global shape of Lutetia, combining two techniques: stereophotoclinometry (4) using 60 NAC and WAC images, and inversion of a set of 50 photometric light curves and of contours of adaptive optics images (5, 6). The asteroid's overall dimensions are  $(121 \pm 1) \times (101 \pm 1) \times (75 \pm 13)$  km³ along the principal axes of inertia. The north pole direction is defined by a right ascension of  $51.8^{\circ} \pm 0.4^{\circ}$  and a declination of  $+10.8^{\circ} \pm 0.4^{\circ}$ , resulting in an obliquity of  $96^{\circ}$ . From the global shape model, we derived a volume of  $(5.0 \pm 0.4) \times 10^{5}$  km³. The volume error is well constrained by (i) the dynamical requirement of principal-axis rotation,

(ii) the existence of ground-based adaptive optics images from viewing directions other than that of the flyby, and (iii) the pre-flyby Knitted Occultation, Adaptive-optics and Lightcurves Approach (KOALA) model (5), which matched the shape model of the imaged part within 5%, giving us confidence that this model is accurate at this level for the southern hemisphere of Lutetia not seen during the flyby. The volume-equivalent diameter of Lutetia is 98 ± 2 km. Combining our volume estimate with the mass of  $(1.700 \pm 0.017) \times 10^{18}$  kg measured by the Radio Science Investigation (7), we obtained a density of  $3.4 \pm 0.3$  g/cm<sup>3</sup>. This value is higher than that found for most nonmetallic asteroids, whose bulk densities are in the range from 1.2 to 2.7 g/cm<sup>3</sup>, well below the average grain density of their likely meteoritic analogs. Such low densities imply large macroporosities (8) that are associated with "rubble pile" asteroids (9).

Using crater density, cross-cutting and overlapping relationships, and the presence of deformational features such as faults, fractures, and grooves, we have identified five major regions on the surface observed during CA. Two regions (Pannonia and Raetia) imaged at lower resolution were defined on the basis of sharp morphological boundaries as crater walls and ridges [Fig. 1 and see the supporting online material (SOM) for details]. The surface is covered in regolith, with slopes below the angle of repose for talus almost everywhere, but large features reveal the underlying structure. A cluster of craters close to the pole in the Baetica region is one of the most prominent features of the northern hemisphere. The most heavily cratered, and therefore oldest, regions (Noricum and Achaia) are separated by the Narbonensis region, which is defined by a crater ~55 km in diameter (Fig. 2). This crater (Massilia) contains several smaller units and is deformed by grooves and pit chains, indicating modifications that took place after its initial formation. Another large impact crater is seen close to the limb (Raetia region). A subparallel ridge formation is seen close to the terminator. A number of scarps and linear features (grooves, fractures, and faults) transecting several small craters (Fig. 2 and fig. S3) are organized along systems characterized by specific orientations for each region and with no obvious relationships with the major craters. However, in the Noricum region, a prominent scarp bounds a local topographic

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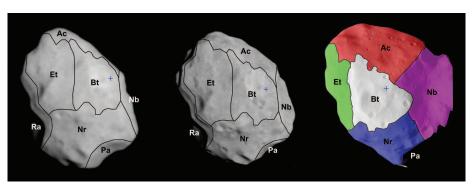
high where lineaments run almost parallel to the scarp itself and to the rims of the crater cluster in Baetica. High-resolution topography models produced by stereo image processing (10) show that one long (>10 km) groove in the Noricum region (Fig. 2C and fig. S4) is roughly 100 m deep and on a local topographic high. The linear features are similar in appearance to those on the martian moon Phobos, which are commonly interpreted as resulting from a large impact (11). On 433 Eros, the existence of similar grooves has been interpreted as evidence of competent rock below the regolith, although this asteroid is

thought to be heavily fractured (12–14). Recent work suggests that cracks can be supported in very low-strength material on a body as small as Eros (15). The pattern of grooves on Lutetia suggests strain structures or fractures within a body of considerable strength.

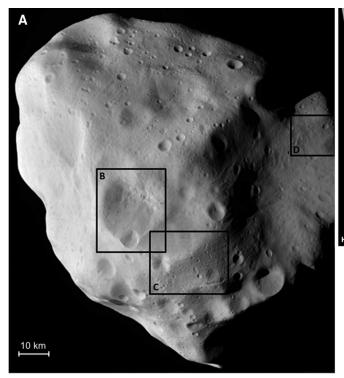
Lutetia is heavily cratered, although the crater spatial density varies considerably across the imaged hemisphere. We have identified more than 350 craters with diameters between 600 m and 55 km, which allowed us to determine Lutetia's crater retention age by measuring the crater size-frequency distribution (SFD). We chose to perform

the crater count on the Achaia region because it is a remarkably flat area imaged with uniform illumination conditions. In this region, we counted 153 craters over an area of 2800 km<sup>2</sup>. We compared Achaia's SFD with those for asteroids 253 Mathilde and 243 Ida (Fig. 3). At large crater sizes (>10 km), the crater SFD of Achaia is quite similar to that of Ida, whereas Mathilde is only slightly less cratered. There are about two or three times fewer craters at a diameter of 1 km than on Ida or Mathilde, respectively. At very small sizes (<1 km), there is a strong depletion of craters. Asteroids as large as Lutetia can be globally affected by seismic shaking; this argument has been used to explain the depletion of <200-m-diameter craters on Eros (13, 16) but cannot explain the observed paucity of craters with diameters up to 5 to 8 km (17). The apparent break in the SFD at this size range is statistically significant: According to the Kolmogorov-Smirnov test, the probability that the observed crater SFD (for diameter > 0.8 km) is consistent with a simple hard rock scaling law model (for an approximately linear crater SFD, see Fig. 3C) is only ~3%.

Small crater obliteration by Massilia crater ejecta seems unlikely given that the Achaia region does not show a systematic decrease in crater density with increasing distance to Massilia. A possible explanation for the break is a transition in the physical properties of the target. Small craters, which affect only the upper layers, form in shattered material. Larger craters, able to excavate to greater depth, form in competent rock.



**Fig. 1.** Regions on Lutetia. Three images taken at -60, -30, and -3 min before CA (left to right) showing the different regions: Bt, Baetica; Ac, Achaia; Et, Etruria; Nb, Narbonensis; Nr, Noricum; Pa, Pannonia; and Ra, Raetia. The images were taken at distances of 53, 27, and  $3.5 \times 10^6$  m and phase angles of 8°, 4°, and 52°. The resolutions of each image are approximately 1000, 500, and 60 m per pixel; Lutetia has been scaled to appear approximately the same size in each panel. The north pole is indicated by the blue cross.



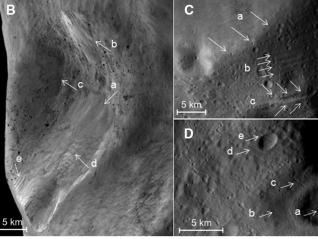


Fig. 2. Surface features. (A) CA image, with details shown under different illumination conditions in (B) to (D). (B) The central 21-km-diameter crater cluster in Baetica. Arrows a, b, and c point to landslides. Landslides a and b appear to have buried the boulders that are pervasive within the crater (with an average density of 0.4 boulders km<sup>-2</sup>). Landslide b may have exposed a rocky outcrop. A similar possible outcrop is seen opposite (e). The material at point d has a mottled appearance. (C) The boundary between Baetica (young terrain associated with the central crater cluster) and Noricum (old terrain) is extremely well defined in some places, as indicated by the arrow a. Arrows b and c highlight curvilinear features. (D) Arrows c, d, and e point to further

curvilinear features on the surface of Lutetia. In the Narbonensis region, most curvilinear features show this orientation. The curvilinear features cut the crater and its rim. Feature c cuts through the debris apron (b) of the crater (a). This implies that these linear features are younger than the craters or impact into an area with existing large-scale cracks and subsequent regolith movement.

We therefore modeled a gradual transition in the crater scaling law as strength and density increase with depth in a fractured layer (18). We determined the depth of this layer by fitting the model to the observed crater SFD (19, 20) (Fig. 3C). For typical rock properties (SOM text), the depth of the fractured layer is  $\sim$ 3 km. Based on this model, and using the lunar chronology as calibration (20), we find a crater retention age of 3.6  $\pm$  0.1 billion years for Achaia.

Scaling laws (21) and hydrocode simulations performed with the iSALE (impact Simplified Arbitrary Lagrangian Eulerian code) (22) show that the impactor that produced Massilia had a diameter ~ 8 km. According to the simulation, this impact heavily fractured but did not completely shatter Lutetia. The current main-belt impact rate suggests that such an impact occurs every ~9 billion years; therefore, the impact may have occurred relatively early in Solar System history, when the collisional environment in the asteroid belt was more intense. The early oc-

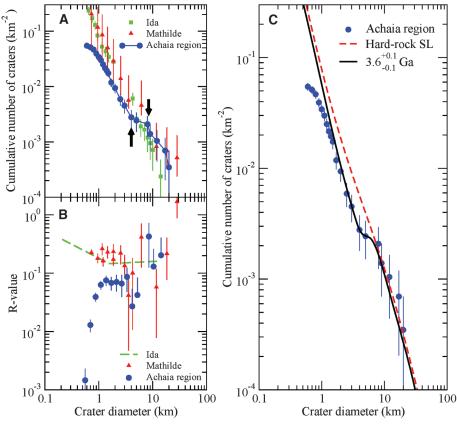
currence of such an impact is in agreement with the crater retention age for Lutetia.

The Baetica region is partially covered by smooth material that is interpreted as ejecta from the 21-km-diameter crater cluster. The images show evidence that older, smaller craters were partially buried by the ejecta. The depth of the ejecta blanket is estimated to be up to ~600 m, based on the depth-to-diameter ratios of these buried craters. The asymmetric shape of the 21-km crater cluster may be the result of internal inhomogeneity. Preexisting planes of weakness in bedrocks may control final crater shapes and facilitate the detachment of blocks and their emplacement within ejecta deposits (23). The crater interior (Fig. 2B) shows a great variety of deposits: smooth and fine deposits with boulders, gravitational taluses, and landslide accumulations. Ejecta blocks have been recorded on other asteroids (13) and Phobos (24). On Lutetia, approximately 200 blocks of up to 300 m in dimension were found around the central crater

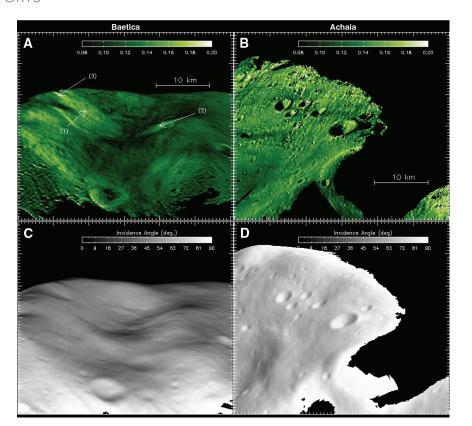
region alone. Their steep size distribution (a power law equation with an exponent of –5) is comparable to that seen on Eros (13). The presence of boulders adjacent to another impact site in the Pannonia region suggests that boulder generation is a common feature of large impacts on Lutetia, and points to excavation of shattered bedrock. The landslides appear to have been emplaced after the boulders and may have been triggered by further impacts.

To investigate the reflectance properties of the surface, OSIRIS obtained images (including several color sequences) at different asteroid rotational phases and over a range of phase angles from 0.15° to 156°. The slope of the phase curve (fig. S5) for phase angles between 5° and 30° is 0.030 mag/° for the 631-nm filter. The Lutetia disk-integrated geometric albedo was measured to be 0.194  $\pm$  0.006 at 631 nm and 0.169  $\pm$  0.009 at 375 nm, giving an average value in the V band (550 nm) of 0.19  $\pm$  0.01 and a Bond albedo of 0.073  $\pm$  0.002.

We computed disk-resolved reflectivity maps at 10° solar phase angle using the threedimensional shape model and light-scattering theory (25) in order to remove the effect of variation in illumination conditions due to the topography (Fig. 4). We detected variations of the surface reflectivity at 647 nm wavelength. The most important variegations are located inside the crater cluster in the Baetica region (Fig. 4A), where reflectivity varies up to 30% between the darkest and brightest areas. Small spatial variations in reflectivity are also present on surrounding terrain (Fig. 4B) but with a much lower contrast. In Baetica, a clear correlation is found with the local surface slope. Landslide flows or possible rock outcrops appear much brighter than the accumulation areas or surrounding cratered terrains. This suggests either a different texture of regolith or that space weathering modified the surface of the oldest areas, whereas young surfaces have been less exposed to solar radiation. Similar variations of reflectivity have been already observed on Eros, where a strong correlation between the spectral slope and the downslope movement of regolith was found (13). Diskintegrated spectrophotometry obtained 1 hour before CA reveals a flat and featureless spectrum, with a moderate spectral slope in the visible range  $(3\%/10^3 \text{ Å between 536 and 804 nm})$ , in agreement with spectra obtained from the Rosetta Visible InfraRed Thermal Imaging Spectrometer (VIRTIS) (26) and ground-based spectra taken at a similar phase angle (fig. S6). These data are consistent with both particular types of carbonaceous chondrite meteorites, namely CO3 and CV3 (1, 27), and enstatite chondrites (ECs) (28). Average bulk densities (8, 29) range from 2.96 to 3.03 g/cm<sup>3</sup> for CO and CV meteorites and 3.55 g/cm<sup>3</sup> for ECs. If Lutetia were composed purely of EC material, this would imply a bulk asteroid macroporosity of ~0 to 13% (given the uncertainty range on Lutetia's density). The low densities of COs and CVs preclude the possibility of



**Fig. 3.** Crater SFD. **(A)** Cumulative crater SFD of the Achaia region compared with those for Ida and Mathilde, the second- and third-largest asteroids imaged by spacecraft so far, respectively [data from (17)]. The arrows indicate the suggested break at 5 to 8 km in the Achaia crater SFD. **(B)** SFD shown in (A) expressed in terms of relative (R) values (cumulative crater SFD normalized to a power law with exponent -2). R values for Ida are not published, but the overall trend (dashed line) was computed from the published cumulative distribution. **(C)** Achaia crater SFD model fit. The dashed red curve represents a fit of the largest craters of the distribution (diameter > 10 km) obtained using current models for the main-belt asteroid size distribution (35) and the crater scaling law (SL) for hard rock (21). The black curve is the best fit achieved by a two-layer (fractured material over competent rock) model, which gives a crater retention age of  $3.6 \pm 0.1$  billion years.



**Fig. 4.** Slope-corrected reflectivity maps (**A** and **B**) and incidence angle maps (**C** and **D**). These are images at 647 nm of parts of the Baetica [(A) and (C)] and Achaia [(B) and (D)] regions that have been photometrically corrected with Hapke bidirectional reflectance theory (25) to remove the effect of different angles of incidence and emission for different local slopes, leaving variations in brightness due only to local albedo variations (resolution, 60 m per pixel). During the photometric correction, the Hapke model parameters describing the single scattering albedo, the coherent backscattering, the shadow hiding, the surface roughness, and the asymmetric factor were all fixed to the value that best reproduced the overall surface reflectivity. The images are corrected to a solar phase angle of 10° for both Baetica and Achaia (the original phase angles for these regions were ~70° to 95°). This phase angle was arbitrarily chosen to avoid the opposition effect that may affect the reflectivity near 0° phase angle. Large variations are visible in the younger Baetica region, whereas the older Achaia region is more uniform (aside from a dark streak associated with a crater in the left of the image). The landslide indicated by 1 and possible outcrops 2 and 3 in Baetica have a reflectivity up to 30% brighter than the accumulation area.

a pure composition of either meteorite group. If Lutetia's surface were made of these materials, this would suggest that the interior may be differentiated (30).

These macroporosities for Lutetia clearly exclude a rubble-pile structure, which typically have macroporosities >25 to 30% (9). Such a high porosity structure is also inconsistent with the extensive eiecta blankets observed around the large craters (31). If Lutetia is undifferentiated, these porosities would also exclude a completely shattered but coherent structure (total porosity in the range of 15 to 25%) (32). Partial differentiation (30) could permit much higher grain densities in the interior and therefore higher porosity and a heavily fractured body. It is therefore likely that Lutetia has survived the age of the Solar System with its primordial structure intact; i.e., it has not been disrupted by impacts. This interpretation is consistent with the current view that the collisional lifetime against catastrophic destruction of bodies with diameters ≥100 km exceeds the age of the Solar System (33). The network of curvilinear features, the crater morphology, and the crater SFD discussed above both indicate that Lutetia's interior has considerable strength and relatively low porosity as compared to that expected for primordial aggregates of fine dust. One possibility is that Lutetia is partially differentiated, with a fractured but unmelted chondritic surface overlaying a higher-density sintered or melted interior (30). In any case, Lutetia is closer to a small planetesimal than to the smaller asteroids seen by previous missions, which are thought to be shattered or rubble-pile minor bodies.

## **References and Notes**

- M. A. Barucci, M. Fulchignoni, in Rosetta: ESA's Mission to the Origin of the Solar System, R. Schulz, C. Alexander, H. Boehnhardt, K.-H. Glassmeier, Eds. (Springer, New York, 2009), pp. 55–68.
- M. Lazzarin et al., Mon. Not. R. Astron. Soc. 408, 1433 (2010).

- 3. H. U. Keller et al., Space Sci. Rev. 128, 433 (2007).
- 4. R. W. Gaskell et al., Meteorit. Planet. Sci. 43, 1049 (2008).
- 5. B. Carry et al., Astron. Astrophys. 523, A94 (2010).
- 6. M. Kaasalainen, Inverse Problem and Imaging 5, 37 (2011).
- 7. M. Pätzold et al., Science 334, 491 (2011).
- 8. G. Consolmagno, D. Britt, R. Macke, *Chemie der Erde Geochemistry* **68**, 1 (2008).
- 9. We use rubble pile to mean a "strengthless aggregate held together by gravity only," following the definition of (34)
- 10. ]. Oberst et al., Icarus 209, 230 (2010).
- P. Thomas, J. Veverka, A. Bloom, T. Duxbury, J. Geophys. Res. 84, 8457 (1979).
- 12. ]. Veverka et al., Science 289, 2088 (2000).
- R. J. Sullivan, P. C. Thomas, S. L. Murchie, M. S. Robinson, in *Asteroid III*, W. F. Bottke Jr., A. Cellino, P. Paolicchi, R. P. Binzel, Eds. (Univ. of Arizona Press, Tucson, AZ, 2002), pp. 331–350.
- A. F. Cheng, in Asteroid III, W. F. Bottke Jr., A. Cellino, P. Paolicchi, R. P. Binzel, Eds. (Univ. of Arizona Press, Tucson, AZ, 2002), pp. 351–366.
- 15. E. Asphaug, Annu. Rev. Earth Planet. Sci. 37, 413 (2009).
- J. E. Richardson Jr., H. J. Melosh, R. J. Greenberg,
  D. P. O'Brien, *Icarus* 179, 325 (2005).
- 17. D. P. O'Brien, R. Greenberg, J. E. Richardson, *Icarus* **183**, 79 (2006).
- S. Marchi, M. Massironi, E. Martellato, L. Giacomini, L. M. Prockter, *Planet. Space Sci.*, available at http://arxiv.org/abs/1105.5272.
- 19. S. Marchi et al., Planet. Space Sci. 58, 1116 (2010).
- S. Marchi, S. Mottola, G. Cremonese, M. Massironi,
  E. Martellato, Astron. J. 137, 4936 (2009).
- 21. K. A. Holsapple, K. R. Housen, Icarus 187, 345 (2007).
- K. Wünnemann, G. S. Collins, H. J. Melosh, *Icarus* 180, 514 (2006).
- 23. H. J. Melosh, in *Impact Cratering—A Geological Process* (Oxford Univ. Press, New York, 1989), p. 84.
- 24. P. C. Thomas et al., J. Geophys. Res. **105**, 15091 (2000).
- 25. B. Hapke, Icarus 157, 523 (2002).
- 26. A. Coradini et al., Science 334, 492 (2011).
- 27. M. A. Barucci et al., Astron. Astrophys. 477, 665 (2008).
- 28. P. Vernazza et al., Icarus 202, 477 (2009).
- R. J. Macke, G. J. Consolmagno, D. T. Britt, M. L. Hutson, *Meteorit. Planet. Sci.* 45, 1513 (2010).
- 30. L. T. Elkins-Tanton, B. P. Weiss, M. T. Zuber, *Earth Planet. Sci. Lett.* **305**, 1 (2011).
- 31. K. R. Housen, K. A. Holsapple, Icarus 163, 102 (2003).
- D. T. Britt, D. Yeomans, K. Housen, G. Consolmagno, in *Asteroid III*, W. F. Bottke Jr., A. Cellino, P. Paolicchi, R. P. Binzel, Eds. (Univ. of Arizona Press, Tucson, AZ, 2002), pp. 485–500.
- 33. A. Morbidelli, W. F. Bottke, D. Nesvorný, H. F. Levison, *Icarus* **204**, 558 (2009).
- 34. S. L. Wilkison et al., Icarus 155, 94 (2002).
- 35. W. F. Bottke Jr. et al., Icarus 175, 111 (2005).
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## Supporting Online Material

www.sciencemag.org/cgi/content/full/334/6055/487/DC1 SOM Text

References (36, 37)

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