A peer-reviewed version of this preprint was published in PeerJ on 7 July 2016.

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Prado GMEM, Anelli LE, Petri S, Romero GR. (2016) New occurrences of fossilized feathers: systematics and taphonomy of the Santana Formation of the Araripe Basin (Cretaceous), NE, Brazil. PeerJ 4:e1916 https://doi.org/10.7717/peerj.1916

New occurrences of fossilized feathers: systematics, taphonomy, and paleoecology of the Santana Formation of the Araripe Basin (Cretaceous), NE, Brazil

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Feathers are the most complex and diversified integuments in vertebrates. Their complexity are provided by the different forms and functions, and they occur both in nonavian and avian-dinosaurs. Despite their rareness, feathers are found throughout the world, and the Santana Formation (comprised by Crato and Romualdo formations) of the Araripe Basin is responsible for the majority of these records in Brazil. Most occurrences is consisted by isolated feathers, where downy-feathers is the recurrent morphotype, two coelurosaurs and one enantiornithe bird. The sedimentary deposition of this unit is consisted by a lacustrine (Crato Fm) and lagoonal (Romualdo Fm) environments, where reducing conditions prevailed, precluding the activity of bottom dwelling organisms that favored the exquisite preservation. Despite the arid and hot conditions during the Cretaceous, life teemed in the adjacency of both paleolakes, however, feathered non-avian dinosaurs were not found yet in the Crato Member. By the great diversity of life that existed in the paleolake surroundings, is possible to recognize, through the fossil record, that a complex and diversified trophic chain was well established during the time period of sedimentation of this unit. When the remains reached the bottom of the paleolakes, the subsequent isolation from the environment allowed their preservation. In this work, three fossilized feathers, consisted of two downy and one contour feather, extracted from the laminated limestone of the Crato Member of the Santana Formation, were described and identified according to morphological and evolutionary models. We also used the terminology commonly applied to extant organisms. Relying on the fossil record of this unit and the adjacencies formations and basins (by autochthonous condition), taxonomic inferences can be made when the lowest hierarchy level is considered, and hence, is possible to propose the plausible taxa that could bear these elements. Taphonomic and paleoecological aspects, such as the preservation of these structures, and the presence of dinosaurs, were also reviewed, as well as the future perspectives about the study of these elements. Despite the virtual low significance, the pragmatical study of fossilized feathers, can help with the understanding of the evolution and paleobiology of dinosaurs, especially on the South Hemisphere.

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NEW OCCURRENCES OF FOSSILIZED FEATHERS: SYSTEMATICS, TAPHONOMY, AND PALEOECOLOGY OF THE SANTANA FORMATION OF THE ARARIPE BASIN (CRETACEOUS), NE, BRAZIL

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ABSTRACT

17 Feathers are the most complex and diversified integuments in vertebrates. Their complexity are provided by the different forms and functions, and they occur both in non-avian and 18 avian-dinosaurs. Despite their rareness, feathers are found throughout the world, and the 19 20 Santana Formation (comprised by Crato and Romualdo formations) of the Araripe Basin is 21 responsible for the majority of these records in Brazil. Most occurrences is consisted by 22 isolated feathers, where downy-feathers is the recurrent morphotype, two coelurosaurs and 23 one enantiornithe bird. The sedimentary deposition of this unit is consisted by a lacustrine 24 (Crato Fm) and lagoonal (Romualdo Fm) environments, where reducing conditions prevailed, 25 precluding the activity of bottom dwelling organisms that favored the exquisite preservation. 26 Despite the arid and hot conditions during the Cretaceous, life teemed in the adjacency of 27 both paleolakes, however, feathered non-avian dinosaurs were not found yet in the Crato 28 Member. By the great diversity of life that existed in the paleolake surroundings, is possible to recognize, through the fossil record, that a complex and diversified trophic chain was well 29 30 established during the time period of sedimentation of this unit. When the remains reached the 31 bottom of the paleolakes, the subsequent isolation from the environment allowed their 32 preservation. In this work, three fossilized feathers, consisted of two downy and one contour 33 feather, extracted from the laminated limestone of the Crato Member of the Santana 34 Formation, were described and identified according to morphological and evolutionary 35 models. We also used the terminology commonly applied to extant organisms. Relying on the 36 fossil record of this unit and the adjacencies formations and basins (by autochthonous 37 condition), taxonomic inferences can be made when the lowest hierarchy level is considered, 38 and hence, is possible to propose the plausible taxa that could bear these elements. 39 Taphonomic and paleoecological aspects, such as the preservation of these structures, and the 40 presence of dinosaurs, were also reviewed, as well as the future perspectives about the study 41 of these elements. Despite the virtual low significance, the pragmatical study of fossilized 42 feathers, can help with the understanding of the evolution and paleobiology of dinosaurs, 43 especially on the South Hemisphere.

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45 Keywords: Fossil Feathers, Santana Formation, Systematic Paleontology, Paleoecology

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48 INTRODUCTION

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Feathers are the most complex integuments of vertebrates, due to their variety of forms and roles. This structure is responsible in the thermoregulation, display, protection against radiation, toxicity, buoyancy and even to produce sound (Lucas & Stettemheim, 1972; Dumbacher *et al.*, 2004; Bostwick & Prum, 2005; Clark *et al.*, 2011; Dimond *et al.*, 2011).

Recent studies of molecular data of feathers, suggest the possible phylogenetic hierarchy in the emergence of these elements, tying with the genesis of the tetrapod integuments. During the period of land transition, which was marked by the aerial exposure and its consequences (*e.g.*, the radiation and free O₂), many amphibians shared the same toolkit for the development of this novelty. Because of this, the integumentary filaments can now be considered a symplesiomorphy, once it is also found in mammals, pterosaurs, non-avian and avian dinosaurs (Greenwold & Sawyer, 2013; Lowe *et al.*, 2014). Even though, they are distinct between these groups, where α-keratin is synthesized by mammals and β-keratin by reptiles. As evidences indicate a plausible multiple origin of these structures within Dinosauria (Clarke, 2013), is possible that the first feathers were present even in the base of the superorder (Xu, 2006; Norell, 2011). Nevertheless, the presence of these elements in basal dinosaurs remains uncertain (Barret, 2013).

Since filaments were reported in a basal ornithischian (Godefroit *et al.*, 2014), a possible occurrence of filament-feathers in dinosauromorphs, or even, in early saurischians, may fills the gap between dinosaurs and other archosaurs (*e.g.*, pterosaurs). However, remains of these animals are often fragmented or unearthed in coarse grain sediments (Langer *et al.*, 2014; Nesbitt *et al.*, 2013; Benton *et al.*, 2014) that precludes the preservation of this fragile structures. Nevertheless, both filament and true feathers were reported in Jurassic theropods (megalosauroids and coelurosaurs), suggesting that this character maybe present early in early theropods (Rauhut *et al.*, 2012; Foth *et al.*, 2014; Xu *et al.*, 2014).

Despite the broad distribution within this clade, true feathers (*i.e.* modern morphotypes) of β -keratin, are only synthesized by more derived theropod dinosaurs (Prum & Brush, 2002; Xu, 2006; Zhang *et al.*, 2006; Xu & Guo, 2009; Norell, 2011; Clarke, 2013). One remarkable finding of these elements in Mesozoic dinosaurs, is an imprintation of wing and tail feathers in a enantiornithine embryo of Liaoning province in China, suggesting that precocity is indeed, an ancient characteristic of birds (Zhou & Zhang, 2004).

In the past, the genesis of feathers were related with the advent of flight (Feduccia, 1993; Martin, 1998), but recent studies indicate that this function originated in a different way, which required the fully formation of these elements (Dial, 2003; Dial *et al.*, 2008; Heers & Dial, 2012). 82 In addition, evidences of simple morphotypes in non-avian dinosaurs, such as Psittacosaurus (Mayr 83 et al., 2002), Tvaniulong (Zheng et al., 2009) and Kulindadromeus (Godefroit et al., 2014), makes this hypothesis even more unlikely. Other interpretations, consider that feathers originated to 84 85 perform the thermoregulation functions, nevertheless, recent studies indicate that dinosaurs were 86 mesothermic, suggesting that their origin do not have any relation with the advent of homeothermy 87 (Grady et al., 2014). Furthermore, the endothermy may have only originated during or briefly after 88 the advent of the flight, since this activity require a high metabolism, with high consumption of O_2 89 and a low accumulation of $C_3H_6O_3$ (lactic acid). To have an efficient thermoregulation control, is 90 possible that this condition evolved with the help of the feathers (Ruben & Jones, 2000) since the 91 non-avian theropods already were covered with them.

A tactile function, as the possible cause of the origin of these elements, was recently proposed by Persons & Currie (2015). This hypothesis explains satisfactorily the origin of filamentous-type integuments, that could be located in the face of its owners for semi-fossorial habits. Despite the difficulty of these elements to be preserved in the fossil record, their proposal require further evidences, once they are inexistent up to the present date.

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Suggested by their distribution throughout the dinosauria clade, the possible reason for this integumentary novelty can be assigned to the ability in maintain social interactions (by visual communications or physiological features), where feathers were important elements in this activity. Despite being a theoretical concept, this model also explains the evolution of the morphotypes, as well as its wide range of color patterns that arisen in avian-dinosaurs, once the sexual selection could be the main driver for their evolution (Dimond *et al.*, 2011; Koschowitz *et al.*, 2014)

103 Because feathers are delicate, their presence in the fossil record, can be associated with soft-104 tissue preservation, and by so, they can be related to konservat-lagerstätten deposits. These 105 elements rarely survive the physicochemical processes that follows the burial, thus, they are usually 106 found as carbonized and iron traces, inclusions in ambers and coprolites, and as imprints (Wetmore, 107 1943; Martins-Neto & Kellner, 1989; Davis & Briggs, 1995; Perrichot, 2004; Perrichot et al., 2008; 108 Vinther et al., 2008; McKellar et al., 2011; Vitek et al., 2013).

109 Only a few deposits possess this kind of vestiges, not exceeding 50 around the world 110 (Kellner, 2002). Despite its rareness, there is a relatively cosmopolitan distribution of these 111 structures, that extends from the Middle Jurassic to the Neogene. Feathers were found in Mesozoic 112 and Cenozoic sedimentary rocks and ambers on the North Hemisphere, in countries such as Canada, 113 China, France, Germany, United States, Russia, Slovenia and Spain (Williston, 1896; Wetmore, 114 1943; Shufeldt, 1913; Kurochkin, 1985; Sanz et al., 1988; Laybourne, 1994; Grimaldi & Case, 1995; Alonso et al., 2000; Buffetaut et al., 2002; Grimaldi et al., 2002; Perrichot, 2004; Zhang et 115

al., 2006; 2010; Perrichot *et al.*, 2008; Schweigert *et al.*, 2009; Knight *et al.*, 2011; McKellar *et al.*,
2011; Zelenitsky *et al.*, 2012; Thomas *et al.*, 2014). On the South Hemisphere, they were found in
Australia (Talent *et al.*, 1966; Waldman 1970), Brazil (Kellner, 2002), Peru (Clarke *et al.*, 2010),
and Antarctica (Mansilla *et al.*, 2013).

The first occurrence of fossil feathers in Brazil, was reported by Shufeldt (1916), in which a primary remex was unearthed from the pelitic sediments of the Tremembé Formation, since then, specimens with different morphotypes and from three geologic units. It consists of the Early Cretaceous (Aptian-Albian) of the Santana Formation of the Araripe Basin (NE, Brazil); the Paleogene (Oligocene) of the Tremembé Formation of the Taubaté Basin (SE, Brazil); and, from the Neogene (Early Miocene) of the Pirabas Formation of the Barreirinhas Basin (N, Brazil) (Shufeldt, 1916; Santos, 1950; Ackerman, 1964; Kellner, 2002; Sayão *et al.*, 2011). While the Pirabas Formation exhibited a single occurrence of feathers since 1964, both Santana and Tremembé formations are responsible for the major records of this type of fossil (Kellner, 2002; Sayão *et al.*, 2011). The occurrences of fossilized feathers in Brazil can be seen in Table 1.

On this paper, we report three new occurrences of fossil feathers, from Cretaceous of Crato Formation of the Araripe Basin. Here, we propose the systematic approach to these fossils, according to the available data. Preliminary discussions about the taphonomy and paleoecology are presented; the presence of avian dinosaurs and their paleoecology are also discussed.

Geologic Setting

The Araripe Basin (Fig. 1) is located in the northeast of Brazil, and its extension is in the order of the approximately 5.500 to 8.000 Km², extending through three states (Ceará, Pernambuco and Piauí), and is situated between 38° 30' to 40° 50' W of longitude, and 7° 05' to 7° 50' S of latitude (Coimbra *et al.*, 2002; Vianna & Neumann, 2002). By the exceptional fossil preservation, this sedimentary unit is described as a *Konservat-Lagerstätte* (Martill *et al.*, 2007). The geology of this basin has been studied since the XIX century (Carvalho & Santos, 2005), with different interpretations that depend of each authors approaches (Maisey, 1991; Assine, 1992; Martill, 1993; Carvalho & Santos, 2005; Assine, 2007; Martill *et al.*, 2007a).

Because Assine (1992; 1994; 2007) have been working with the entire basin, in a detailed level and in accordance with Brazilian Code of Stratigraphy, we prefer to adopt his descriptions in this paper. The Santana Formation, is comprised by the Crato Member (in the bottom), and the Romualdo Member (in the atop). Both units have different lithologies that reflect their distinct deposition environment. Many of the exquisitely preserved fossils of the Araripe Basin, come from 150 the Crato Member strata, which is characterized by micritic laminated limestones intercalated with 151 shales and mudstones of varied thickness. This unit was formed in a lacustrine environment, with brackish water, and reducing conditions in the bottom strata of the controversially depth water 152 153 column (Assine, 1992; 1994; 2007; Martill, 2007; Heimhofer et al., 2010).

154 Through the occurrences of ostracods and palinomorphs, Coimbra et al. (2002) correlated the Araripe Basin with other deposits that was suggested with same age. However, to Barbalha and 155 156 Santana formations, ostracodes could not assign to any biozone. The assignment was made by the study of palynomorphs, that indicated an Albian age (~120 Ma) to Crato Member of the Santana 157 158 Formation (Coimbra et al., 2002).

160 **MATERIALS AND METHODS**

Three specimens were studied and described following the terminology of Lucas & Stettemhein (1972); Sick (1984) and Proctor & Lynch (1993). These fossils were apprehended by the Brazilian Federal Police and the IPHAN (Institute of National Historical and Artistic Heritage) **-J**65 and are deposited in the Paleontological Collection of the Laboratory of Systematic Paleontology from the Geosciences Institute, of the University of São Paulo, in the city of São Paulo. The specimens received the registered numbers: GP/2E-7853, GP/2E-7854 and GP/2E- 8771. The acronyms used in the collection assign the "GP" to Geology and Paleontology sets, and '2E', to the vertebrate set.

170 All specimens were photographed using a millimeter-scale stand with Canon EOS REBEL T3 with aperture of 100 mm and under a stereomicroscope Carl Zeiss with a capture system 171 172 AxioCam ICC3 and using the AxioVision LE software. The specimens were measured with a 173 caliper and the AxioVision LE software. We measured specifics portions of the feathers such as 174 barbs and rachis. The difference between each portion that was compared with the total size of the 175 length. These measurements were used to infer the morphology to classify them according to the 176 literature of extant feathers (Lucas & Stettemheim, 1972; Sick, 1984; Proctor & Lynch, 1994).

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178 **RESULTS**

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180 **Systematic Palaeontology**

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182	Order Saurischia Seeley 1888
183	Suborder Therapoda Marsh 1881

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- Division Coelurosauria Von Huene 1914 sensu Gauthier 1986
- Family *Incertae sedis*
 - (Figure 2, A B)

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188 Material: GP/2E-7853 (fig. 2, A).

- 189 Horizon: Crato Member, Santana Formation, Araripe Basin.
- 190 Lithology: Weathered (beige) micritic laminated limestone.
- 191 Age: Lower Cretaceous (Aptian/Albian).

192 Description: This specimen is a complete feather with reduced dimensions compared to other 193 morphotypes (i.e., contour feathers and pennaceous feathers) and orange coloration. Barbules are 194 not clearly visible and are presented only in some regions of the barbs. The rachis consists of a 195 slight line. The distal extremity presents ramifications where barbs with a diverse length originates. 196 The calamus is not present.

197 Measures: See table 2, first row. Dimensions are in mm.

198 **Taphonomy:** The color of this specimen (orange/reddish), indicate that the fossil may be preserved as an iron oxide. The matrix presents features such as a light beige coloration due slight weathering, calcified filaments, crystals of sphalerite, etc. (Martill & Briggs, 2001; Heimhofer et al., 2010).

199 200 201 **Diagnosis:** Despite having a fairly generic morphotype, this specimen has a typical morphology of 202 the plumulaceous feathers due to the presence of rachis very well delineated and barbs of varying 203 sizes, this feather is interpreted as a downy feather (Lucas & Stettemhein, 1972; Sick, 1984; Proctor 204 & Lynch, 1994). It is not possible to observe the presence of the calamus. Generally, because of 205 their fragility and their small size (in life it could represented only 1.5% of the total length of the 206 feather), this portion does not preserve widely in the fossil record (Lucas & Stettenheim, 1972; Kellner, 2002). Also, in the matrix, a nearly complete skeleton of a small fish is associated (Fig. 2, 207 208 D), that here we interpreted belonging to the Dastilbe genus (Maisey, 1991; Dietze, 2007; Martill et 209 al., 2007a).

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Family Incertae sedis

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213 Material: GP/2E-7854 (fig. 2, B).

214 Horizon: Crato Member, Santana Formation, Araripe Basin.

- 215 Lithology: Weathered (beige) micritic laminated limestone.
- 216 Age: Lower Cretaceous (Aptian/Albian).
- **Description:** The proximal portion is degraded, although, the rachis are visible. Several barbs with 217

- different length originate from them. Also, it is possible to notice the presence of vestigial barbules. 218
- 219 As the GP/2E-7853, the calamus is a slight line.
- 220 Measures: See table 2, second row. Dimensions in mm.
- 221 **Taphonomy:** Similar to GP/2E-7853, this feather possesses a small dimension and it is complete. It
- 222 is possible to notice that there is color variation between the proximal to distal portion of the vanes,
- 223 with light to darker brownish tones that represents the preservation by carbonaceous traces (Davis 224 & Briggs, 1995).
- 225 **Diagnosis:** This specimen also presents the typical morphotype of the plumulaceous feathers, 226 classified as downy feathers. On the umbilicus proximallis portion (Fig. 2, E), the slight line 227 structure is consisted of an external molt that is interpreted as the vestige of the calamus. By the 228 preservational characteristics (e.g., external mould, lack of organic remains), the evidence suggests 229 that this portion was degraded during the taphocenosis, or geochemical processes that followed the 230 burial (diagenesis).

Order Saurischia Seeley 1888					
Suborder Therapoda Marsh 1881					
Division Coelurosauria Von Huene 1914 sensu Gauthier 1986					
Subdivision Maniraptoriformes Holtz 1995					
Family Incertae sedis					
(Figure 2, C)					

239 Material: GP/2E-8771 (fig. 2, C).

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- 240 Horizon: Crato Member, Santana Formation, Araripe Basin.
- 241 Lithology: Grayish micritic laminated limestone.
- 242 Age: Lower Cretaceous (Aptian/Albian).

243 **Description:** This specimen it is a complete feather with a bigger dimension, compared with the 244 two previously described (tables 4 and 5). A diverse barbs with variable lengths originates from a 245 slight rachis. The barbules are clearly visible, and they also vary in size. In extant feathers, vanes 246 are united by the 'hooklets' (structures similar to hooks) (Lucas & Stettenheim, 1972; Sick, 1984), 247 however they are not preserved in this fossil.

- Measures: See table 2, third row. Dimensions are in mm. 248
- Taphonomy: As well as GP/2E-7853 and GP/2E-7854, this specimen is also occurs in a limestone 249
- 250 matrix. Due to the blackish color of the fossil, this feather may be preserved as carbonized trace,
- 251 since it is the common type of preservation of feathers in rocks (Davis & Briggs 1995).

252 Diagnosis: According to morphology, GP/2E-8771 is associated to the typical extant contour 253 feathers or semiplumes. Attached to the basal part (the umbilicus), an structure similar to afterfeather emerged forming an V shape. This element is larger than the vanes (Lucas & 254 255 Stettenheim, 1972). Nevertheless, it is unlikely that this element represent an afterfeather once it 256 does not show any diagnose feature, such as a "slight rachis" or umbilical origin. The barbules are 257 present and they suggest some degree of cohesion between barbs, however, there are no 'hooklets' 258 preserved on this specimen. The characteristic that distinguish this specimen to the other two 259 described previously, is the color hue of the matrix. This feature is an indicative that this sediment 260 may not be exposed to weathering processes that usually changes the rock color (Martill & Frey, 1995). In an attempt to make more attractive to the illegal sale, some portions of the feather were 261 262 degraded with a scraper tool, especially on the portion where the calamus was supposed to be 263 found. The GP/2E-8771 sample is the only Mesozoic feather described here, that is truly secure to 264 be assigned to the crown group Aves, since all of its characteristics are very similar to modern 265 morphotypes.

Systematic analysis

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268 269 Two specimens (GP/2E-7853 and GP/2E-7854) possess a simplest and generic morphotype. They resembled the 'IIIb' evolutionary-developmental stage (Prum & Brush, 2002) and the '4th' 270 271specimen of the morphological models reported in the fossil record (Xu & Guo, 2009). We have 272 tried to classify the other feather, by the same process. Based on the morphology, we compared 273 specific portions (table 3) that enabled us to classify to the plumulaceous morphotype (table 4).

274 Both feathers GP/2E-7853 and GP/2E-7854 have a plumulaceous appearance that is similar 275 to downy feathers that main character is its dimension, where the rachis is shorter than the longest barb. The difference between the rachis and the longest barbs of GP/2E-7854 is 48,21%, although, 276 277 the rachis of GP/2E-7853 is 8,27% longest than the larger barb. Despite the fact that GP/2E-7853 278 have a longest rachis than the larger barbs, by the fluffy aspect, dimension, and morphology, we 279 prefer to assign this feather (together with GP/2E-7854) as a downy feather. The occurrence of this 280 morphotype is wide in the extant class Aves, once they are present beyond the semiplumes and are located in the apterium portions of most birds. According to the fossil record, these structures could 281 282 also belong to non-avian dinosaurs, making the taxonomic assignment even hard to be inferred 283 (Lucas & Stettenheim, 1972; Prum & Brush, 2002; Zhang et al., 2006; Xu & Guo, 2009).

284 The specimen GP/2E-8771, possess a morphotype similar to semiplumes, with a apparent 285 aftershaft on the proximallis portion. However, this structure may not represent this element. The 286 rachis is 43,95% shorter than the longest barb. The morphology of this feather is similar to the type 287 of afterfeathers, that possess a long, narrow with shorter vanes. The hyporachis is almost with the 288 same length of the afterfeather. In extant cases, these feathers are related mainly to birds of the 289 order of Galliformes and Trogoniformes, but also to Tinamiformes birds (Lucas & Stettenheim, 290 1972).

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292 DISCUSSION

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294 Isolated feathers have been described formally in many works (Kellner, 2002), however, 295 differently from fish scales, mollusks shells or plant trunks and leaves, none of them received a 296 proper taxonomic treatment until the present date. The mainly reason for the lack systematic 297 procedure, may be caused by its rareness in the fossil record, that hampers the exact taxonomic 298 determination. Despite the taphonomic significance, this inappropriate treatment, can also be an 299 issue that systematists simply ignores once these elements demonstrate low taxonomic interest (*i.e.*, 300 low potential to assign a new taxon). Nevertheless, Rautian (1978) applied a different taxonomic 301 approach to these elements, once their existence represented a diagnose of new bird taxon. 302 Nowadays, this method prove to be problematic, since non-avian dinosaurs also possessed them, 303 demanding a different way to assess their taxonomic value. Here, we propose a simple and 304 parsimonious approach to describe fossilized feathers, assigning their morphotypes to the basal 305 animals that possessed them according to the fossil record of non-avian and avian dinosaurs.

306 Because the specimens studied here were product of apprehension (illegal fossil trade), the 307 exactly stratigraphic position cannot be determined. However, the laminated limestones (LL) of the 308 Crato Member, possess many characteristics that is well known by worldwide geologists and 309 paleontologists. Since the LL only occur in this unit, it is plausible to assign these fossils to this 310 specific strata.

311 Two specimens (GP/2E-7853 and GP/2E7854) deserved more attention, since they are from 312 the Mesozoic Era (period marked by "evolutionary experiments"), and possess a generic 313 morphotype that resembles early, and then, basal feathers. Despite the controversy about the 314 morphotype diversity provided by the squeeze during diagenesis (Foth, 2012), an apparent decrease 315 is suggested by the fossil record (Zhang et al., 2006; Xu & Guo, 2009; Xu et al., 2010). Nevertheless, it was possible to associate both specimens aforementioned, to the evolutionary stages 316 317 as it is proposed by the literature (Prum and Brush, 2002; Zhang et al., 2006, Xu & Guo, 2009). 318 Features such as presence of barbs that originates from a scanty rachis, absence of barbules, small 319 dimension between morphotypes, bigger length of the barb than the rachis, and fluffy aspect, are

320 present in both feathers. By these characteristics, they represent a basal morphotype and we 321 classified them belonging to the 'IIIb' stage of the evolutionary model (Prum and Brush, 2002) and to '4th morphotype' commonly found on the fossil record (Xu and Guo, 2009). 322

323 The preservation of the macro-structures, identification of morphotype, and size, allow us to 324 suggest the possible roles during life, their placement throughout their body, and proportion of the owners (Lucas & Stettemheim, 1972; Sick, 1984; Proctor & Lynch, 1993). Considering that these 325 326 two feathers (GP/2E-7853 and GP/2E-7854) were identified as auricular feathers (length between 327 2.5 to 17 mm), the animal that possessed these elements may not have a big dimension, *i.e.*, not 328 exceeding the domestic chicken size. Thus, it must had a role similar to the extant birds, where the 329 mainly function is in ear protection (Lucas & Stettenheim, 1972). The other feather (GP/2E-8771), 330 a contour feather, we suggest that it may have also taken the same protective function. However, it 331 might functioned in the thermoregulation of its owners. Nevertheless, even in basal coelurosaurs, 332 they may had other roles such as in display, shielding nests, etc. (Turner et al., 2007). Other lines of evidences suggests that dinosaurs already possessed visual acuity, with nocturnal or crepuscular behavior, and abilities to communicate visually may have been present in the Mesozoic (Varrichio, et al., 2007; Xu et al., 2009; Schmitz & Motani, 2011; Koschowitz et al., 2014).

The morphotype of GP/2E-8771, and its position throughout the body, indicate that this feather may favored the camouflage and communication between the owners, as seen in modern birds (Gluckman & Cardoso, 2010). In addition, it could have also assumed -- hypothetically -- a sexual role, similar to extant birds with iridescent and colorful feathers, such as peacocks (Zi et al., 2003) and birds-of-paradise (Irestedt et al., 2009).

The process of preservation explains the absence of hooklets in all specimens, since these 342 structures are very delicate, their presence is not expected. Because of this feature, they are not 343 common in feathers preserved in rock matrix, only present in feathers included in amber (Davis & Briggs, 1995; Laybourne, 1994; Perrichot et al., 2008; Mckellar et al., 2011; Thomas et al., 2014). 344

346 Sedimentary Deposition, Paleoenvironment, and Taphonomy

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348 Over the Aptian-Albian, the Santana Formation had two different depositional systems that 349 followed the evolution of the basin throughout the late Mesozoic. The Crato Member is represented 350 by a restricted lacustrine environment with brackish waters; and the Romualdo Member is 351 suggested to be formed by a lagoonal, with seasonally contact with marine waters, or even to a 352 moments of marine regression-transgressions (Assine, 1994; 2007; Martill et al., 2007a). The Crato and Romualdo formations, are separated by an unconformity marked at the top of the Crato 353

354 Member. A fossiliferous strata of shales and evaporites that characterizes the 'Ipubi Layers'. This layer occur with varied thickness and lateral continuum, suggesting the possible shallowing of the 355 water column (Assine, 2007; Martill et al., 2007b). The shales of this unit is responsible for the 356 357 single occurrence of an ichonofossil in the entire Santana Formation, where it was assigned to a 358 turtle (Dentzien-Dias et al., 2010). Fielding et al. (2005) mentioned a turtle remains from this layer. Despite her inaccurate taxonomic assignment (as Araripemys), Oliveira et al. (2011) described this 359 360 fossil correctly, noting the affinity with a undetermined pleurodiran.

361 According to paleontological and sedimentary evidences, such as palynomorphs and 362 evaporites, the Crato Member was formed in a clear and relatively shallow waters during an arid 363 and dry climate, where the calcium carbonate sediments were deposited in a low energetic input 364 with formation of halite and anhydrite minerals (Assine, 1994; Silva et al., 2003; Assine, 2007; 365 Martill et al., 2007). As suggested elsewhere (Martill et al., 2007a), this anoxic and hypersaline 366 environment prevented the presence of the bottom-dwelling organisms, once the salinity level might 367 have be higher to the osmotic toleration (Martill, 1993; Martill et al., 2007ab; Martill et al., 2008b).

368 Despite the controversy, the source of the sediment may be due the events of algal blooms, 369 where the carbonate, was organically precipitated by pico- and phytoplankton (Martill et al., 370 371 2007a). However, but albeit meager, evidences of microbial mats may also indicate the origin of this sediment, *i.e.*, precipitated by these microorganisms (Martill et al., 2007a; Martill et al., 372 2008b). The presence of articulated, undisturbed fossils, and pseudomorphs of pyrite and marcasite, 373 indicate that the reducing condition prevailed in the bottom of the paleolake, enabling the exquisite 374 preservation, allowing the presence of non-resistant tissues in the fossil record (Fielding et al., 2005; Martill et al., 2007a; Pinheiro et al., 2012; Simões et al., 2014; Barling et al., 2015). The high 375 376 degree of articulation and the exquisite preservation, suggests that a low energy environment 377 prevailed, without any or significant carcass transportation, as well as disturbance by scavenging organisms (Fielding et al., 2005; Martill et al., 2007a; Bráez et al., 2009; Figueiredo & Kellner, 378 379 2009; Pinheiro et al., 2012; Barling et al., 2015).

380 Despite this "harsh" environment, this sedimentary unit is remarkably noted by the abundant 381 biota that is preserved with a high degree of fidelity. The vertebrate fauna is composed primarily by 382 crocodiles, turtles, frogs, birds, pterosaurs, and numerous fishes (Maisey, 1991; Martill, 1993; 1997; 383 Fielding et al., 2005; Martill et al., 2007a; Martill et al., 2008a; Figueiredo & Kellner, 2009; 384 Pinheiro et al., 2012; Simões et al., 2014; Oliveira & Kellner, 2015). The invertebrate animals was 385 also abundant, with the mainly occurrences of arthropods and mollusks (Maisey, 1991; Martill, 386 1993; Grimaldi & Engel, 2005; Martill et al., 2007a; Barling et al., 2015). Despite this faunal richness, the flora was also exuberant (Martill et al., 2007; 2012; Mohr et al., 2015), with a 387

diversity characterized by the presence of macro and microfossils of pteridophytes, gymnosperms,
angiosperms, palynomorphs, pollens, seeds, etc. (cf. Maisey, 1991; Martill, 1993; Martill *et al.*,
2007a; 2012).

The fauna of the Crato Member may have been autochthonous (Naish *et al.*, 2004), however, the terrestrial vertebrates of the Santana Formation may had a different geographical origin through time, indicated by evidences in other adjacent basins (Carvalho & Gonçalves, 1994; Carvalho, 1995; Carvalho & Araújo, 1995; Carvalho *et al.*, 1995; Carvalho & Pedrão, 1998). Nevertheless, in both *lagerstätten* units (Crato and Romualdo members), animals may have be well adapted to the arid and dry climate (Naish *et al.*, 2004; Martill *et al.*, 2007a; Heimhofer *et al.*, 2010). Many of the animals may lived in the surroundings of the paleolake, that could bear a high diversity of plants, especially angiosperms. By these singular taphonomic features, we suggest that, the presence of larger vertebrates in this unit, is derived by the occasional incursions into shallow waters, in order to prey fishes, insects, and other land animals, preserving these animals *in situ*.

On the other hand, it can be also speculated that the Romualdo fauna may have been allochtonous. The presence of a very diversified animals, indicate that they lived in the nearby shorelines, since they was also well established to this environment (Naish *et al.*, 2004). Especially to the vertebrate remains, their presence can be explained by the drifting hypothesis. As proposed by Naish *et al.* (2004), the carcasses of the animals who lived in the nearby regions, may have been transported by rivers through tens of kilometers (or less), before it reached the paleolake. Although, as Assine (1994) noted, during the Romualdo Member deposition, the environment was characterized by the maximum marine transgression, where the salty waters entered this basin with N-NW direction. Therefore, the carcass of these animals may have been dragged by these moments of water incursions, not having any relationship with perennial rivers.

Nevertheless, their presence can be also explained by the attraction of the dead fish, that 411 412 could be exposed on the shores of the paleolake during the mass mortality events, caused by 413 chemoclinal alterations or by the remobilization of the anoxic layers to upper strata. Once in the 414 shoreline, these large vertebrates may have stuck on the soft and deep sediments, where they died 415 by natural means (Olson & Alvarenga, 2002, Varricchio et al., 2008). After death and despite the 416 scavenger activity that might followed, the carcass of these vertebrates could be finally transported 417 into the paleolake. Once there, they were preserved by the process of 'encapsulation', also known as 418 "The Medusa effect" (Martill, 1989), that is a characteristic mode of preservation of the Romualdo 419 Member nodules. This hypothesis could also explain the preservation of the isolated and often 420 disarticulated bones of pterosaurs, dinosaurs and other aerial and terrestrial vertebrates. 421 Nonetheless, the absence of ichnofossils in the rocks of this unit, makes this hypothesis difficult to

422 be established. However, further works might concentrate in these evidences.

423 Besides the previous 'trapping hypothesis', Duncan & Jensen (1976) presented evidence of 424 mass mortality of modern aquatic birds by toxins, consisted mainly by *Clostridium botulinum*. As 425 pointed by these authors, C. botulinum can colonize similar environments, and evidences of these 426 microorganisms in other fossil sites suggests their responsibility to the mass mortality events 427 (Varrichio, 1995; Lingham-Soliar, 2012). In the Santana Formation, the activity of these organisms 428 could also be responsible for the presence of the vertebrates in the fossil record (Duncan & Jensen, 429 1976; Varrichio, 1995). However, only the mass mortality of fishes is commonly observed, and 430 further evidences of this phenomenon is needed, such as, a high bone concentration of different 431 vertebrates in the same strata (Varrichio, 1995; Martill, 1997; Martill et al., 2008a).

From the three specimens studied, only GP/2E-7853 shows a coloration (reddish/orange) that is typical of the iron oxides-hydroxides, possibly limonite. This type of preservation was also observed in others feathers from the same provenance (Maisey, 1991; Martins-Neto & Kellner, 1988; Martill & Frey, 1995; Martill & Davis, 2001). The remaining specimens, may be preserved as incarbonization, once it is one of the most common type of preservation of organic molecules, as it presents the characteristic dark black hue (Tegelaar *et al.*, 1989; Davis & Briggs, 1990; 1995; Kellner, 2002; Briggs, 2003). Although, the explanation of their chemical composition, needs further geochemical investigation.

The main hypothesis that explains the presence of the isolated feathers in the fossil record, especially in the Santana Formation, assumes that these elements may have been blown into the paleolake by events of strong winds. Once they have reached the waters of the lake, these feathers may have sunk quickly, reaching the bottom in seconds to few minutes, where they were rapidly buried (Martill & Davis, 2001). Despite the fact that this hypothesis satisfy this question, others mechanisms (and educated guesses) may also be praised.

446 During their life, birds tend to lose feathers by ontogeny, breeding season, and under high 447 stress situations. This latter mechanism is called 'fright molt'. Generally, they tend to release 448 rectrices and semiplumes (Sick, 1984). It is possible that the animals that possessed these structures, 449 released them during similar conditions. The presence of the isolated feathers could occur by animal 450 hunting, *i.e.*, by fish, crocodiles, dinosaurs, or, pterosaurs. Some extant fishes occasionally include 451 birds on their diet by opportunism (Davenport, 1979; French, 1981; O'Brien, 2014; Perry et al., 452 2013). A good prey-predator example in the Santana Formation, is indicated by the evidence of a 453 conical tooth assigned to a spinosaurid dinosaur, reported inserted in a cervical vertebrae of the 454 ornithocheirid pterosaur (Buffetaut et al., 2004). But, despite the possibility that birds were also 455 prey, the fossil record of the established trophic chain, do not show yet, these animals as a food

456 source of the other organisms, in a way that coprolites did not provide yet, evidence of this diet 457 (Maisey, 1991; Martill, 1993; Lima *et al.*, 2007).

458 Other hypothesis may also explain satisfactorily the occurrence of isolated feathers and the 459 absence of bones associated with them. Nevertheless, the "wind hypothesis" remain the most valid 460 and satisfactory.

461 The carbonate concretions of the Romualdo Member, provided a record of at least four non-462 avian dinosaurs, although, only theropods were found in this unit. The review of the previously 463 considered an ornithischian ischium, is now interpreted as a rib of a unknown theropod dinosaur 464 (Machado & Kellner, 2007). Thus, the dinosaur fauna of the Araripe Basin consists of two 465 spinosaurids, Irritator challengeri (Martill et al., 1996) and its possible synonym, Angaturama **A66** limai (Kellner & Campos, 1996); and two coelurosaurs, Santanaraptor placidus (Kellner, 1999) and 467 *Mirischia asymmetrica* (Naish *et al.*, 2004). The latter dinosaur, possess at least at the family level, 468 two feathered species (Ji & Ji, 1996; Ji et al., 2007). It is speculated, that filament feathers were 469 present even in megalosauroid dinosaurs (Rauhut et al., 2012). But according to previous reports, 470 evidences of feathers were not detected in any taxa of non-avian dinosaurs of the Araripe Basin -471 (Kellner, 1999; Naish et al., 2004). This absence is odd, by the vast record of soft tissues in both 472 members (Crato and Romualdo members). They include insect muscle fibers (Grimaldi & Engel, 473 2005; Barling et al., 2015), dinosaurs blood vessels (Kellner, 1996a), pterosaur wing membranes, 474 muscle fibers, and headcrest (Martill & Unwin, 1989; Kellner, 1996b; Pinheiro et al., 2012), fish 475 muscle tissue and stomach contents (Martill, 1989; 1990; Wilby & Martill, 1992), skin impressions 476 of turtle (Fielding et al., 2005), fossilized microbodies related to pigmentation (Vinther et al., 2008), among others (cf. Martill, 1993; Martill et al., 2007a). 477

478 The first record of Mesozoic avian dinosaur in Brazil, was only described recently with a 479 fossil unearthed from the Crato Member of the Santana Formation (Carvalho et al., 2015). The 480 feathers of this fossil, an enantiornithine undetermined, possess interesting features, showing an 481 extremely long rectrices, secondary remiges, alular feathers, and filamentous feathers. Regarding to 482 the remex and rectrices, there is no doubt that they were pennaceous feathers. Nevertheless, the 483 filamentous aspect may be a taphonomic artifact (Foth, 2013), indicating that these structures most 484 likely were, contour feathers or downy feathers. It is possible to note that in rectrices, patches with 485 granulate spots may be associated with color patterns. However, no other evidence of its hue is 486 given by the authors. Two other possible specimens of enantiornithine, from the same provenance, 487 are briefly mentioned elsewhere (Fielding et al., 2005; Naish, 2007). These specimens are not 488 accessible, and so, they are not formally described, since these specimens seemingly were lost to 489 private collectors that they acquired through the illegal fossil trade.

490 The presence of this undetermined enantiornithine, suggests doubtless, that birds were 491 present in the Araripe Basin landscape. Even though, non-avian dinosaurs with feathers, still 492 remains unknown to this deposit. Since the absence of feathers associated with bones of non-avian 493 dinosaurs makes an unusual event in both Lagerstätten (Crato and Romualdo members), some 494 possibilities emerges: (i) the non-avian dinosaurs found in this deposit were glabrous (i.e., they did 495 not possessed feathers) or were low in coverings; (ii) a selective taphonomic or geological process 496 acted obliterating their preservation; (iii) during the time of deposition, taphonomic conditions were 497 very different between both members or even to the same unit, disallowing their preservation; (iv) 498 all possibilities may have happened simultaneously, or consecutively for the case of taphonomical 499 and diagenetical processes; (v) feathered non-avian dinosaurs (and most birds) were lost by 500 geological processes (diagenesis, erosion, and weathering); or, (vi) these animals were not 501 discovered yet. With the exception of the two latter, we consider these possibilities to be very 502 problematic because the process of fossilization preserved tissues that is more prone to degradation, 503 but did not allowed feathers that is relatively more resistant. This can wrongly suggests, that a 504 differential preservation may have occurred. But, the considerable numbers of non-resistant tissues 505 reported in both members, makes this hypothesis very unlikely. It is important to state, that 506 analogous deposits with similar lithology (limestone rocks) and depositional settings, e.g., the Las 507 Hoyas Formation in Spain (Sanz et al., 1988; 1996) and the Solnhofen Formation in Germany (Barthel et al., 1994); hold records of dinosaurs preserved similarly to the Araripe Basin. Especially 508 509 in Solnhofen, feathers are present in the Archaeopteryx specimens, but they are not in the 510 Compsognathus longipes (Barthel et al., 1994). This suggests that the third hypothesis may be true, 511 once the selective taphonomic/geological events can determinate the differential preservation of 512 carcasses in the same depositional conditions. Especially to Santana fossils, the formation of 513 concretions of the Romualdo Member, may be the responsible for obliterating these integumentary 514 tissues, differently from the Crato Member that preserved the enantiornithine with feathers.

By their localization throughout the body, it is possible that these elements were more exposed to geochemical reactions during the initial phase of decay that followed the burial, being degraded early in diagenesis or by the weathering that succeeded their exposure. However, to nodules, the geochemical alteration by weathering may not be the responsible for the absence of these elements, since the dinosaur tissues remained relatively isolated from the surrounding environment throughout the geological time. Is expected that further studies may enlighten this odd feature.

522 Until the present day, only a few records of feathers, or filaments considered as 523 'protofeathers', were found associated with ornithischians dinosaurs (Mayr *et al.*, 2002; Xu *et al.*,

2009; Zheng et al., 2009; Sevaliev & Alifanov, 2014; Godefroit et al., 2014). Evidences of sauropod 524 525 integuments, indicate that they were consisted mainly by scales and osteoderms (Czerkas, 1992). Although, the poor record of feathers in individuals of ornithischia, and the absence in sauropoda, 526 527 may be also a taphonomic artifact. Despite the occasional events of great sediment deposition, the 528 preservation of the carcass of these animals (often huge), required more time to be completely 529 buried, and so, preserved. This slow process opposes to the rapid burial of feathers as suggested 530 elsewhere (Martill & Davis, 2001), explaining their absence alongside sauropods and great 531 ornithischians bones.

Another taphonomic feature, is the type and grain size of the sediment that buried these animals. Siliciclastic coarse grains, tend to preserve only larger hard parts of the animals (*i.e.*, bones, keratinous beaks, tooths, and claws). Generally, these sediments are related to high energetic depositional systems, with unidirectional flows, such as rivers and streams (Holz & Simões, 2002). Even though, fine-grained sandstones, such as found in the Ediacara Hills in Australia (Gehling, 1999) and in Horseshoe Canyon Formation in Canada (Quinney *et al.*, 2013), preserved soft tissues such as the Ediacara organisms and feathers, respectively. Nevertheless, the lack of recognition during the bone extraction, or a careless excavation also represents a destructive potential of these elements (Zelenitsky *et al.*, 2012).

Despite their scarcity, evidences of enanthionithines and maniraptorans were found in Cretaceous deposits of Brazil, emphasizing their presence in the Brazilian landscapes (Chiappe & Calvo, 1994; Carvalho & Pedrão, 1998; Alvarenga & Nava, 2005; Novas *et al.*, 2005; Naish, 2007; Machado *et al.*, 2008; Candeiro *et al.*, 2012a; Candeiro *et al.*, 2012b; Marsola *et al.*, 2014; Tavares *et al.*, 2014; Carvalho *et al.*, 2015; Delcourt & Grillo, 2015). Even though, the poor record of Mesozoic birds, added to the absence of feathered dinosaurs, makes the taxonomic assignment even harder to infer (Naish, 2007; Sayão *et al.*, 2011). It is unlikely that these feathers belonged to nontheropod dinosaurs, such as sauropods or ornithischians, once they are also absent in the Araripe Basin (Naish, 2004). On the other hand, a large numbers of evidences indicate that only non-avian theropods possessed these structures (Unwin, 1998; Chiappe & Witmer, 2002; Prum & Brush, 2002; Norell & Xu, 2005; Xu, 2006; Zhang *et al.*, 2006; Xu & Guo, 2009; Norell, 2011; Mckellar, 2011).

Because modern-type of feathers was already present in the Middle Jurassic basal and derived animals (Foth *et al.*, 2012; 2014b), it hinders the taxa assignment of these isolated feathers at the genus level. The occurrence of these elements attached to bones of known feathered dinosaurs, is very important to determine the possible taxon and morphotype. But despite this difficulty, the taxonomic attribution can be made at least at the division and subdivision levels. As filament-feathers were found in basal theropods; *e.g.*, *Sinocalliopteryx prima* (Ji & Ji, 1996); 558 Sinocalliopteryx gigas (Ji et al., 2007); Sciurumimus albersdoerferi (Rauhut et al., 2012), and 559 Yutyrannus huali (Xu et al., 2012); we consider that the Araripe Basin non-avian dinosaurs were 560 also covered with these elements. However, their were restricted in distribution throughout the body 561 as suggested by evidences in the other specimens. We consider more parsimonious to assign both 562 feathers, GP/2E-7853 and GP/2E-7854, to the Coelurosauria clade. Because true pennaceous feathers were found in Ornithomimus edmontonicus (Zelenitsky et al., 2012), we assign 563 564 GP/2E-8771 to the Maniraptoriformes clade. As pointed by the large numbers of evidences, both 565 groups (Fig. 3) are responsible for these integuments in dinosaurs (Clarke, 2013).

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567 Future perspectives

569 In a striped contour feather from the Araripe Basin described by Martill & Frey (1995), 570 Vinther et al. (2008) have found oblate microbodies restricted only to the dark portions of the 571 specimen. The light portions, were markedly preserved as imprintation. Those structures were 572 previously interpreted as autolithified bacteria (Wuttke, 1983; Davis & Briggs, 1995), but 573 574 575 576 subsequent studies revealed them as evidence of fossilized melanosomes (Vinther et al. 2008; Zhang et al., 2010; Barden et al., 2011; Glass et al., 2012; 2013; McNamara et al., 2013; Vitek et al., 2013; Barden et al., 2014; Li et al., 2014; Egerton et al., 2015; Vinther, 2015). This interpretation, enabled reconstructions of ancient color patterns of extinct animals, such as 577 dinosaurs, birds, reptiles and fishes. (Vinther et al., 2008; 2010; Clarke et al., 2010; Carney et al., 578 2012; Field et al., 2013; Li et al., 2010, 2012, Lindgren et al., 2012, 2014). Grueling evidence of 579 melanin in fossilized feathers still remains unknown, however, works on this subject has been focus 580 of investigations (Colleary & Vinther, 2013).

581 Further investigations using the Scanning Electron Microscopy equipped with Energy Dispersive X-Ray Spectroscopy (SEM-EDS), will help on the identification of the presence of the 582 583 ultrastructures such as minerals, melanosomes, and other possible elements in this feathers. In 584 addition, other techniques, such as Raman Spectroscopy (RAMAN), X-Ray Fluorescence (XRF), 585 Gas Chromatography-Mass Spectrometry (GC-MS), among others, also can give information about 586 their chemistry (Wogelius et al., 2011; Egerton et al., 2015), indicating possible taphonomic 587 processes that occurred after the deposition (Davis & Briggs, 1995; Schweitzer et al., 2008; 588 McNamara, 2013). Besides the study with ancient pigmentation, the application of these techniques 589 are important, once it provides more information about these fossils. Theses approaches not only 590 allow a better understand of the taphonomic and diagenetic processes, but it also enables 591 paleoenvironmental and paleoecological reconstructions.

593 CONCLUSION

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595 Despite the difficulty on the systematic approach, it is possible to identify the isolated 596 feathers, basing on the lower taxonomic rank, relying on the fossil record of the unit. For the 597 Santana Formation of the Araripe Basin, the maximum taxonomic status reached, is the Division 598 (Coelurosauria) and Subdivision level (Maniraptoriformes). Based on the extinct and modern 599 morphotypes, and, on evolutionary model of feathers, the fossils were identified as, two downy 600 feathers (GP/2E-7853 and GP/2E-7854) and one semiplume (GP/2E-8771).

These feathers may be preserved as limonite (GP/2E-7853) and carbonized traces (GP/2E-7854 and GP/2E-8771); and the mechanisms which allowed the preservation of these elements was briefly discussed. As suggested by Martill & Davis (2001), we also consider that these feathers have been transported into the paleolake by strong winds. Once in the waters, they sunk and were buried rapidly in the anoxic bottom. The absence of oxygen, has an important role, once it prevented the activity of scavenging organisms, allowing its preservation. Nevertheless, we also considered other possible causes, *e.g.*, by predation (by fright molt).

608 While there are records of non-avian dinosaurs in the Araripe Basin, there is not yet formally descriptions of avian dinosaurs. Nor, presence of feathers associated directly with bones. Even 609 610 though, soft tissues were found in many animals, including dinosaurs. Although unlikely, its 611 possible that a differential taphonomic process happened, preserving these non-resistant tissues 612 instead of feathers. Further geochemical investigations may reveal this process and how this specimens were preserved. Future investigations may also focus on the identification of the 613 614 ultrastructures in addition to its chemical composition, offering the possible roles in life. Despite 615 their rareness and low taxonomic potential, fossilized feathers can offer insights about the 616 paleobiology of its owners and the paleoecology of the Araripe Basin.

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618 ACKNOWLEDGEMENTS

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We would like to thank Luis Fábio Silveira (Ornithological Collection of the Zoology Museum – University of São Paulo) for the helpful comments and aid on the identification of these specimens. We also thank Ivone C. Gonzales (Institute of Geosciences - University of São Paulo) for the support in the accessing the Paleontological Collection and these specimens. And we are deeply thankful of Gabriel L. Osés, Bruno B. Kerber (Institute of Geosciences - University of São Paulo), Cibele G. Voltani (Institute of Geosciences and Exact Sciences - São Paulo State 626 University), and Mirian L. A. F. Pacheco (Federal University of São Carlos - Sorocaba) with the 627 support and invaluable comments.

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629 Author contributions: G.M.E.M.P. and L.E.A. designed the project and performed the research. 630 G.M.E.M.P. and G.R.R. analyzed data, wrote the manuscript and designed figures.

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632 Additional information: The authors declare no competing financial interests.

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634 Financial Support: This project was supported by the Pró-Reitoria de Pesquisa of University of 635 São Paulo under the undergraduate research program granted to G.M.E.M.P. between 2013 to 2014.

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Table 1(on next page)

Brazilian feather occurrences

Table 1. The Brazilian fossil record of feathers (formally described specimens).

FEATHER	DEPOSIT	AGE	PRESERVATION	OBSERVATIONS	REFERENCE
One primary remex	Tremembé Fm	Paleogene (Oligocene)	Carbonized	First record in Brazil	Shufeldt (1916)
One contour feather	Tremembé Fm	Paleogene (Oligocene)	Carbonized	Feather assigned to a Turdidae (<i>Turdus rufiventris</i>)	Santos (1950)
Two pennaceous feathers	Pirabas Fm	Neogene (Miocene)	Carbonized	Possible semiplumes	Ackerman (1964)
One primary remex	Santana	Cretaceous (Aptian/Albian)	Limonitc/ Imprint	Asymmetrical feather assigned to birds	Martins-Neto & Kellner (1988)
Contour feathers	Tremembé Fm	Paleogene (Oligocene)	Carbonized/ Imprint	Several feathers associated with skeleton of the bird <i>Taubacrex granivora</i>	Alvarenga (1988)
Semiplume	Santana Bm	Cretaceous (Aptian/Albian)	Carbonized	Feather assigned to passerine birds	Martill & Filgueira (1994)
Down feather	Santana Fm	Cretaceous (Aptian/Albian)	Carbonized	Feather attributed to thermoregulation function of a bird	Kellner et al. (1994)
Contour feather	Santana Fm	Cretaceous (Aptian/Albian)	Carbonized (with melanosome preservation)	Feather with (banded) color pattern preserved.	Martill & Frey (1995)
One symmetrical feather	Santana Fm	Cretaceous (Aptian/Albian)	Carbonized	The biggest isolated feather associated with ectoparasite eggs. Assigned to a bird.	Martill & Davis, (1998; 2001)
Plumulaceous feathers	Santana Fm	Cretaceous (Aptian/Albian)	No data. Presumably carbonized	One plume and one semiplume	Sayão & Uejima (2009)
Down feather	Santana Fm	Cretaceous (Aptian/Albian)	Carbonized	Feather assigned to a bird	Sayão et al. (2011)
Several rectrices, remiges and filamentous feathers (possibly contour feathers)	Santana Fm	Cretaceous (Aptian/Albian)	Carbonized	Several feathers associated with a skeleton of an enantiornithe bird. First record of Mesozoic bird in Brazil.	Carvalho et al. (2015)

Table 1. The Brazilian fossil record of feathers (formally described).

Figure 1(on next page)

Araripe Basin locality and lithology

Figure 1. The Araripe Basin locality, the stratigraphic columns, units and chronology. (Adapted and modified from Coimbra et al., 2002; Vianna & Neumann, 2002; Assine, 2007)

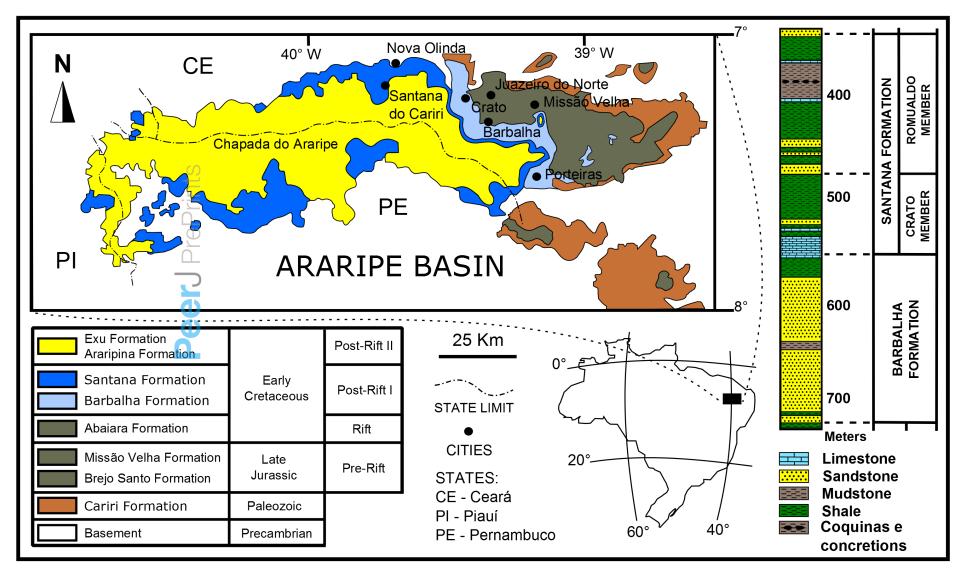


Figure 1. The Araripe Basin locality, the stratigraphic columns, units and chronology. (Adapted and modified from Coimbra *et al.*, 2002; Vianna & Neumann, 2002; Assine, 2007)

Table 2(on next page)

Measures of the new specimens

Table 2. Values of the measures of the three specimens. Legend: NP - Not present. Dimensions are in mm

Specimen	Width	Length	Larger barb	Minor barb	Calamus	Rachis	Rachis thickness
GP/2E-7853	12,36	16,14	8,65	4,85	NP	9,43	0,49
GP/2E-7854	12,76	19,00	17,83	4,30	0,24	12,03	0,49
GP/2E-8771	15,63	33,50	16,45	4,12	NP	29,35	0,03
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Table 2.	Values of the m	easures of the three	specimens.	Legend: NP	• - Not present.	. Dimensions	are in mm
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Figure 2(on next page)

Samples (Feathers and fish)

Figure 2. Fossilized feathers and fish of the Santana Formation. (A) GP/2E-7853; (B) GP/2E-7854; (C) GP/2E-8771. (D) The Dastilbe fossil fish of the GP/2E-7853 specimen; (E-F) External molt of the calamus of GP/2E-7854; (E) Photograph of the the umbilicus proximallis; (F) Interpretative drawing of the calamus. Legend: CL - Calamus; BI - Isolated Barbule; VX -Vexillum (vanes); RQ - Rachis. Scale bars: Scale bar: (A, E-F) 2 mm; (B-C) 5 mm; (D) 10 mm. PeerJ reviewing

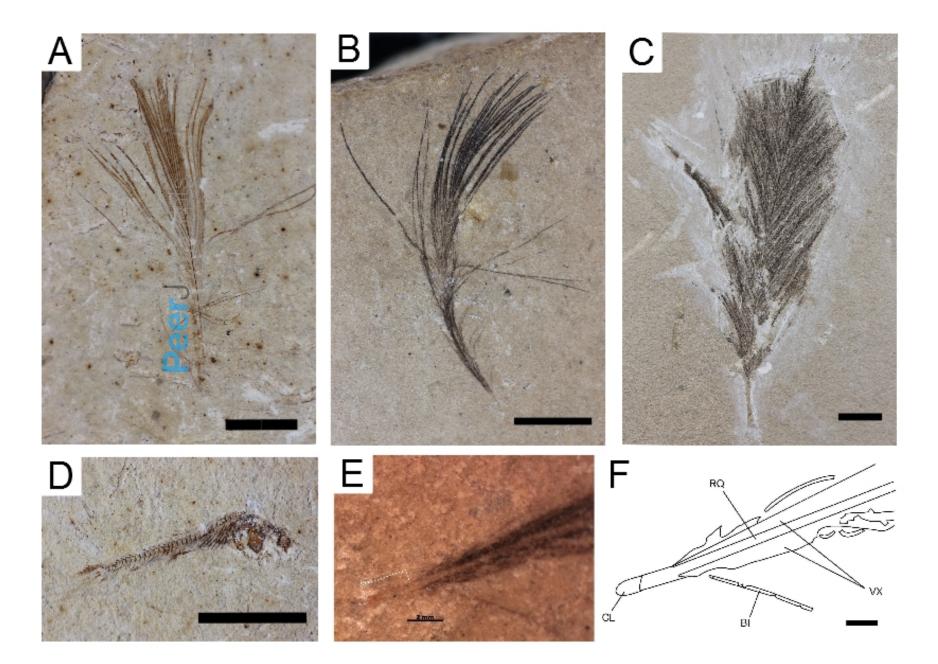


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Table 3(on next page)

Feather portions calculation

Table 3. Difference in percentage between portions of the feathers compared to the maximum length. Legend: ND - No data available.

STOLICTUDE	PERCENTAGE					
STRUCTURE	GP/2E-7853	GP/2E-7854	GP/2E-8771			
Larger barb	46,41	6,16	50,90			
Minor barb	69,95	77,37	87,70			
Calamus	ND	1,26	ND			
Rachis 4	41,57	36,68	12,39			
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Table 3. Difference in percentage between portions of the feathers compared to the maximum length. Legend: ND - No data available.

Table 4(on next page)

Taxonomic assignment

Table 4. Classification of the described feathers. PeerJ reviewing

Table 4. Classification of the described feathers.

SPECIMEN	MORPHOTYPE	EVOLUTIONARY- DEVELOPMENTAL MODEL	MORPHOTYPE MODEL PRESENT OF THE FOSSIL RECORD	
GP/2E-7853	Downy feathers	IIIb	Morphotype 4	
GP/2E-7854	2		1 51	
GP/2E-8771	Semiplume (Contour feather)	IIIa+b	Morphotype 6	
Peenda Pe				

Figure 3(on next page)

Simplified Cladogram of dinosauria and feathers

Figure 3. Simplified cladogram of Dinosauria with the distribution of feathers according to the fossil record. Despite its more ancient origin, it was only in maniraptoriformes that moderntype feathers (plumulaceous and pennaceous feathers) have arisen (Based in Xu & Guo, 2009; Clarke, 2013; Godefroit et al., 2013; Han et al., 2014; Koshchowitz et al., 2014).

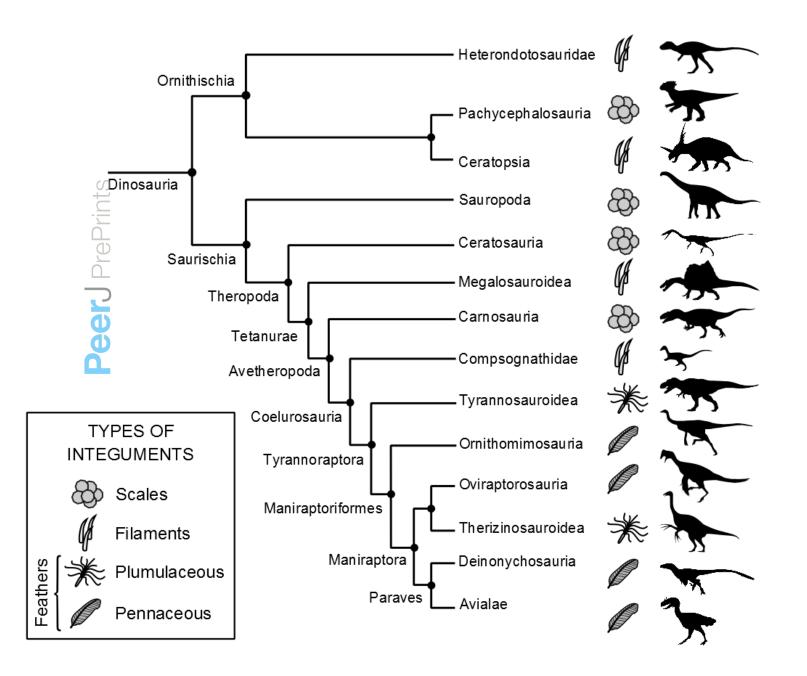


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