



The Cenozoic history of the Antarctic ice sheet

Arjen Stroeven

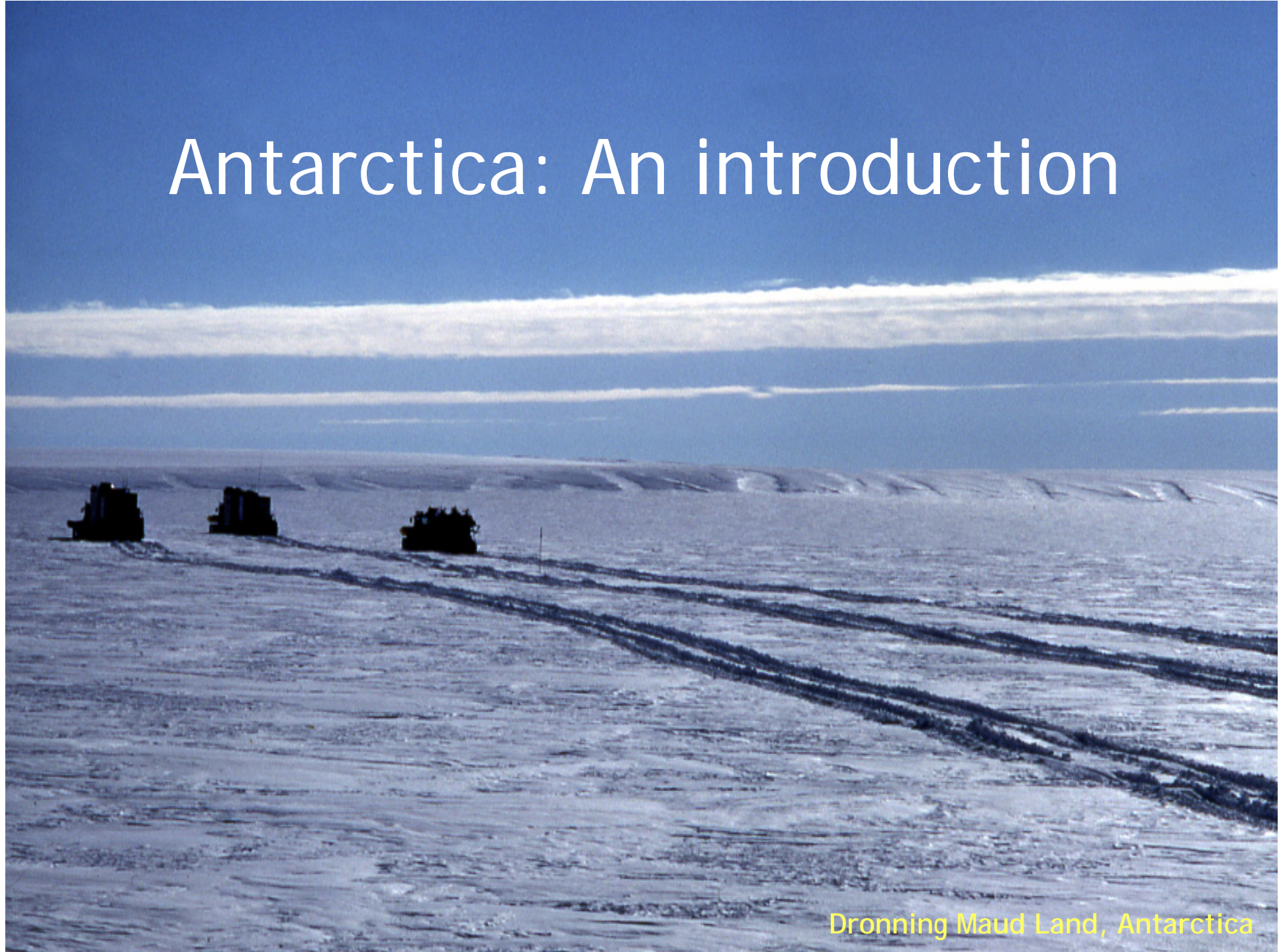
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Structure of the lecture

- Antarctica: An introduction
- Antarctic geological history
 - plate tectonics
- Antarctic glacial history
 - how deduced?
 - what have we learned?
- Antarctic geomorphology
 - Dry Valleys
- Ice sheet modelling

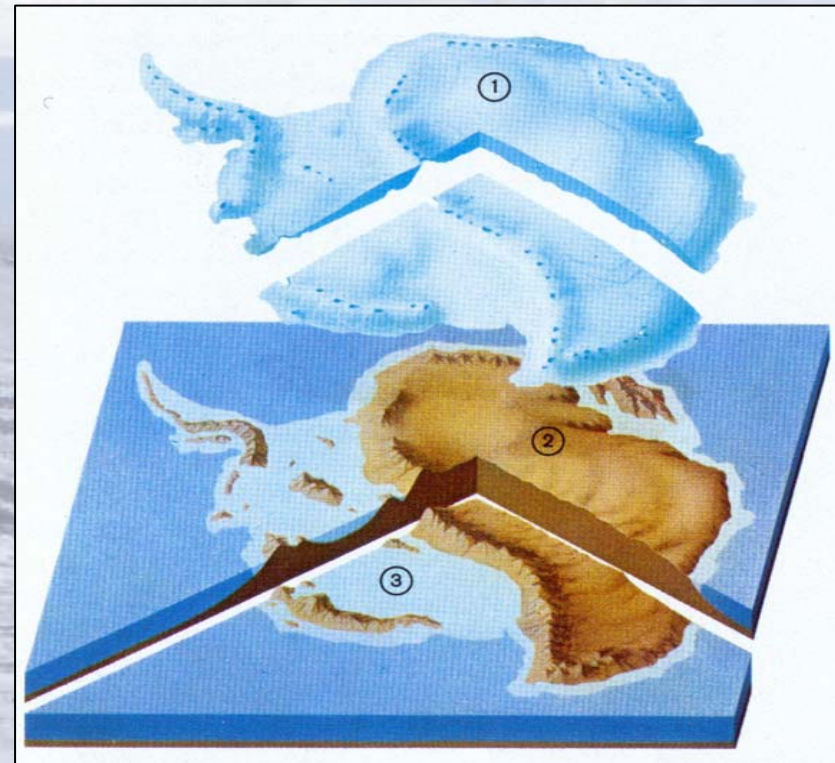
Antarctica: An introduction



Dronning Maud Land, Antarctica

Antarctica: An introduction

- East Antarctica (2):
continental
-66 m water equivalent
- West Antarctica (3):
marine
-6 m water equivalent
- Average thickness (1):
~ 2,079 m (IPCC, 2001)

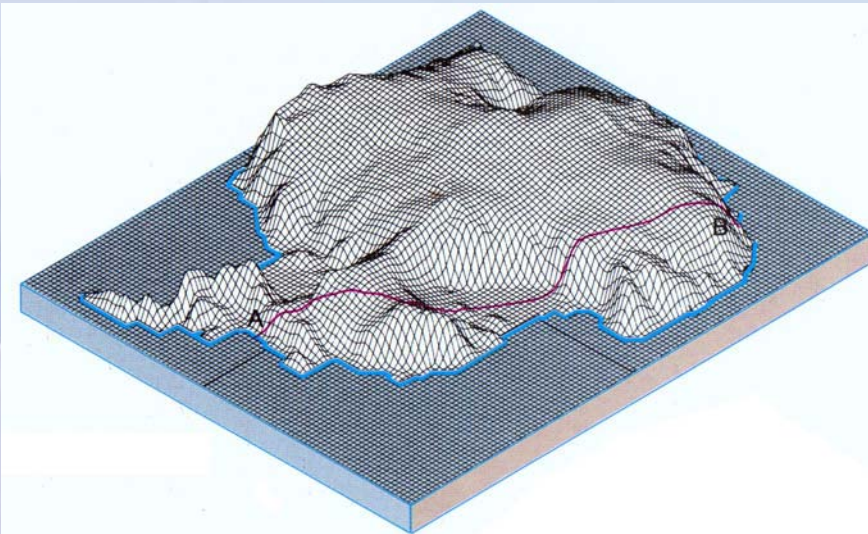


Antarctica: An introduction

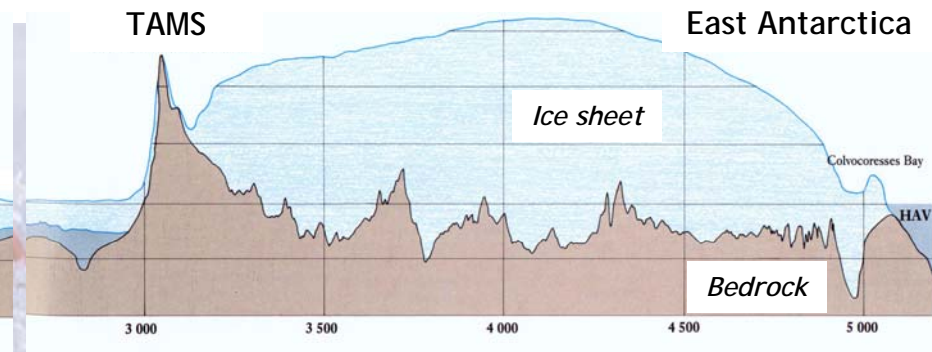
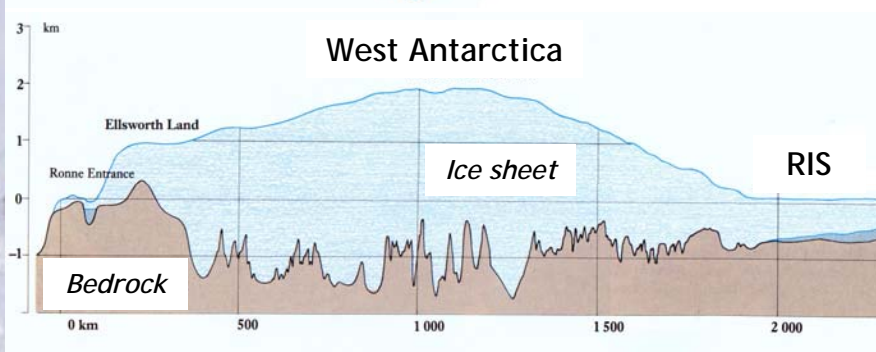


Dronning Maud Land, Antarctica

Antarctica: An introduction



- Both ice sheets rest on bedrock below sea level - potentially sensitive situation
- Ice exits Antarctica through ice streams and ice shelves

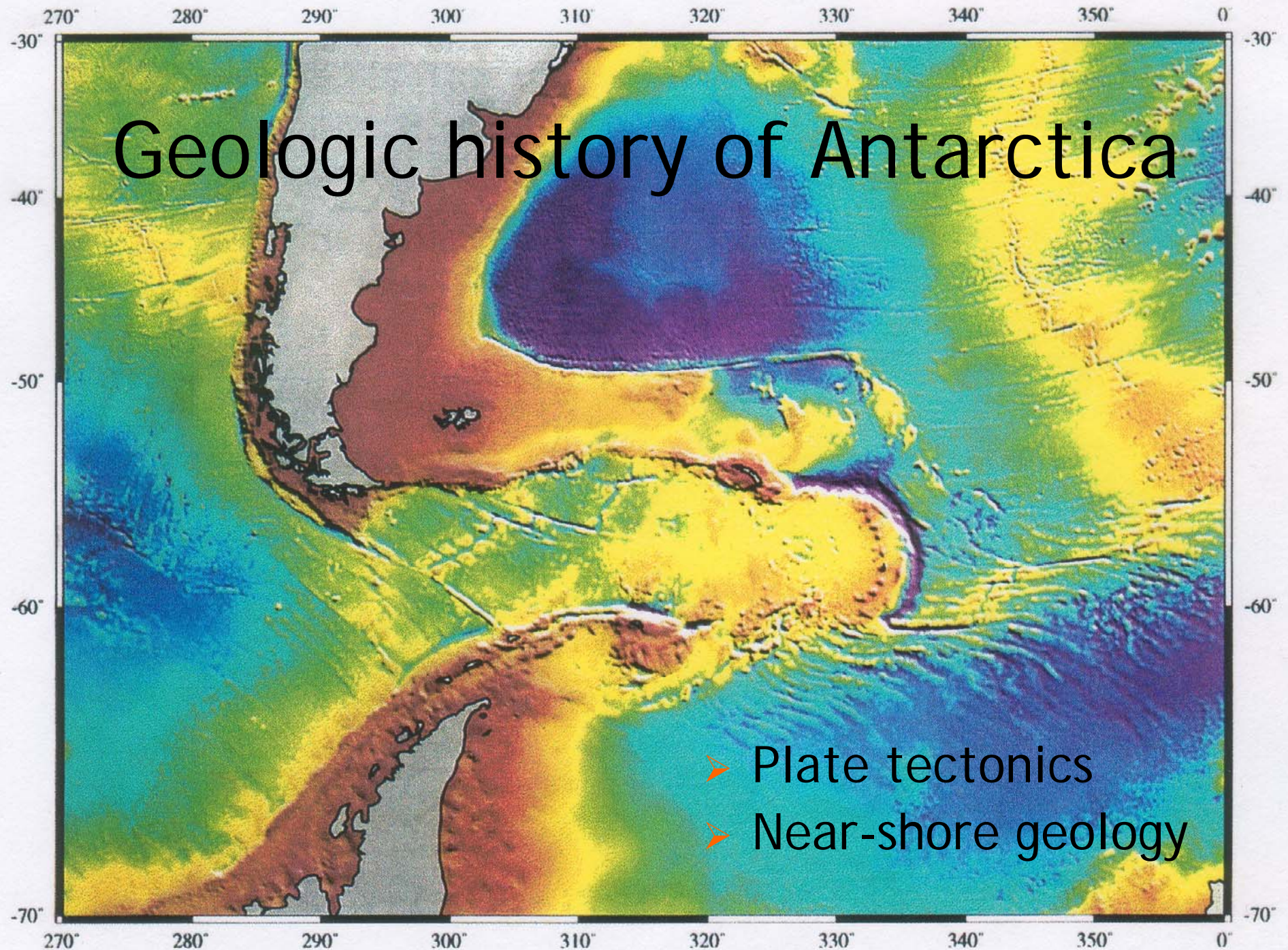


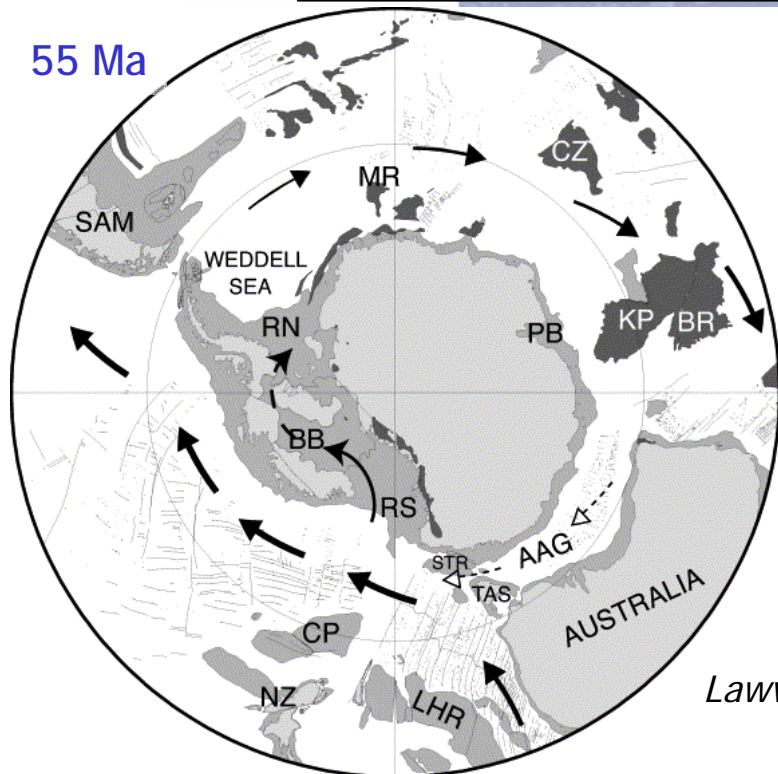
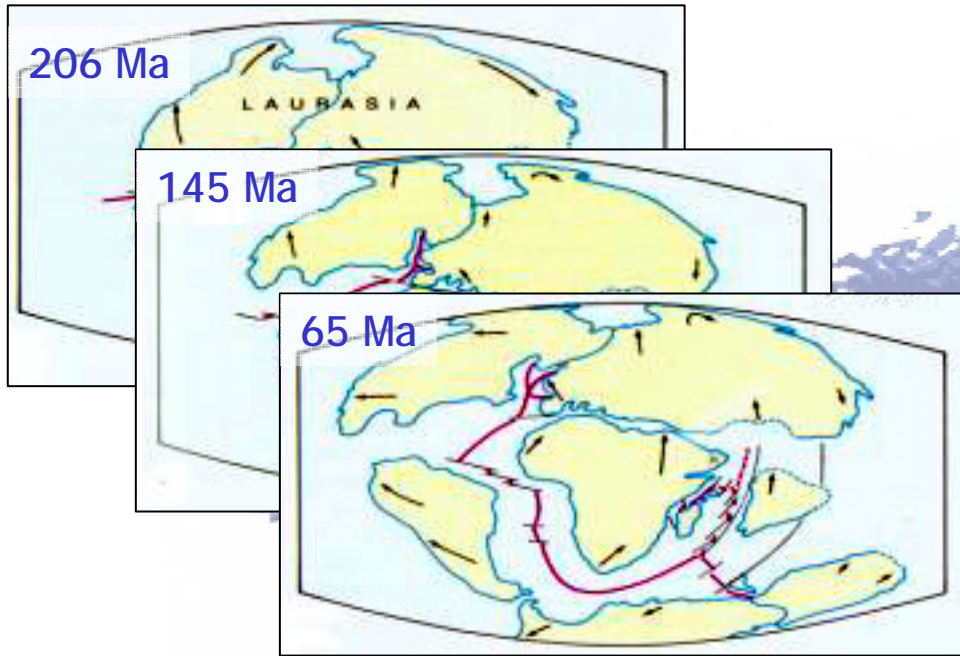
May (ed.) 1988, Dorling Kindersley Limited, London

- 
- Highest continent
average elevation: ~2 km
 - Coldest continent
record: -89,5 °C
 - Driest continent
1.63 cm water equivalent/year
 - Windiest continent
average annual: 80 km/hour

Air Devron Six Icefalls, Dry Valleys, Antarctica

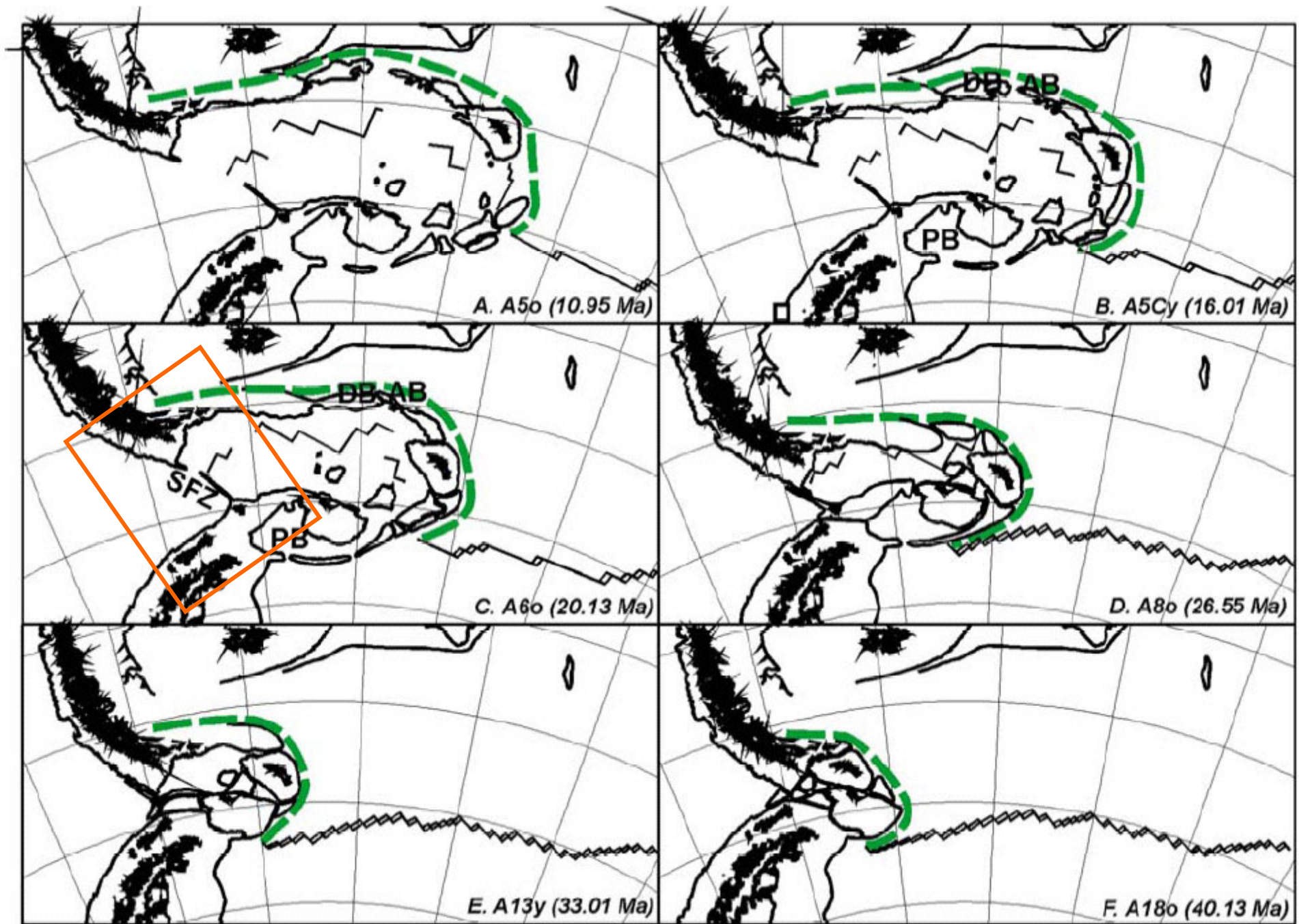
Geologic history of Antarctica



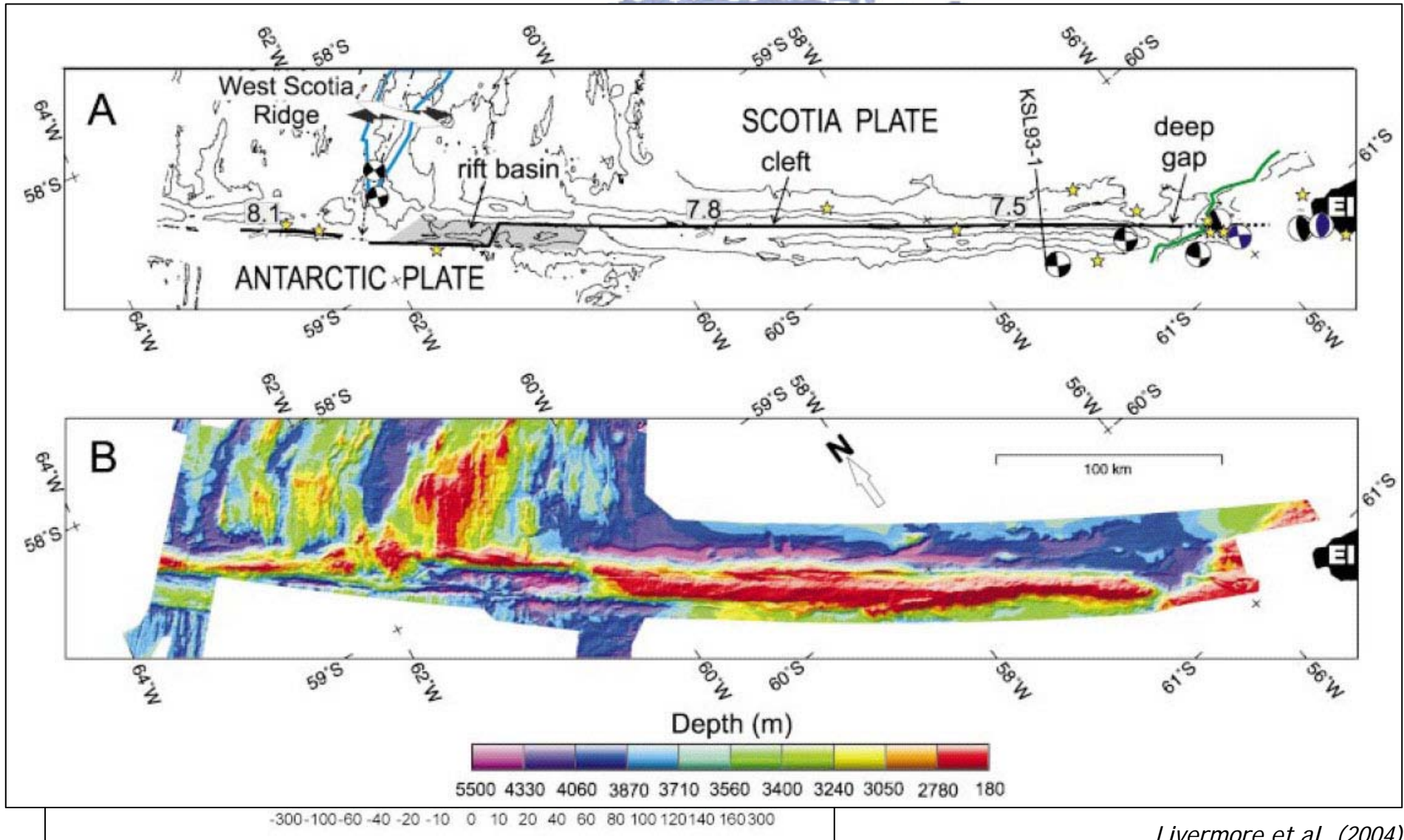


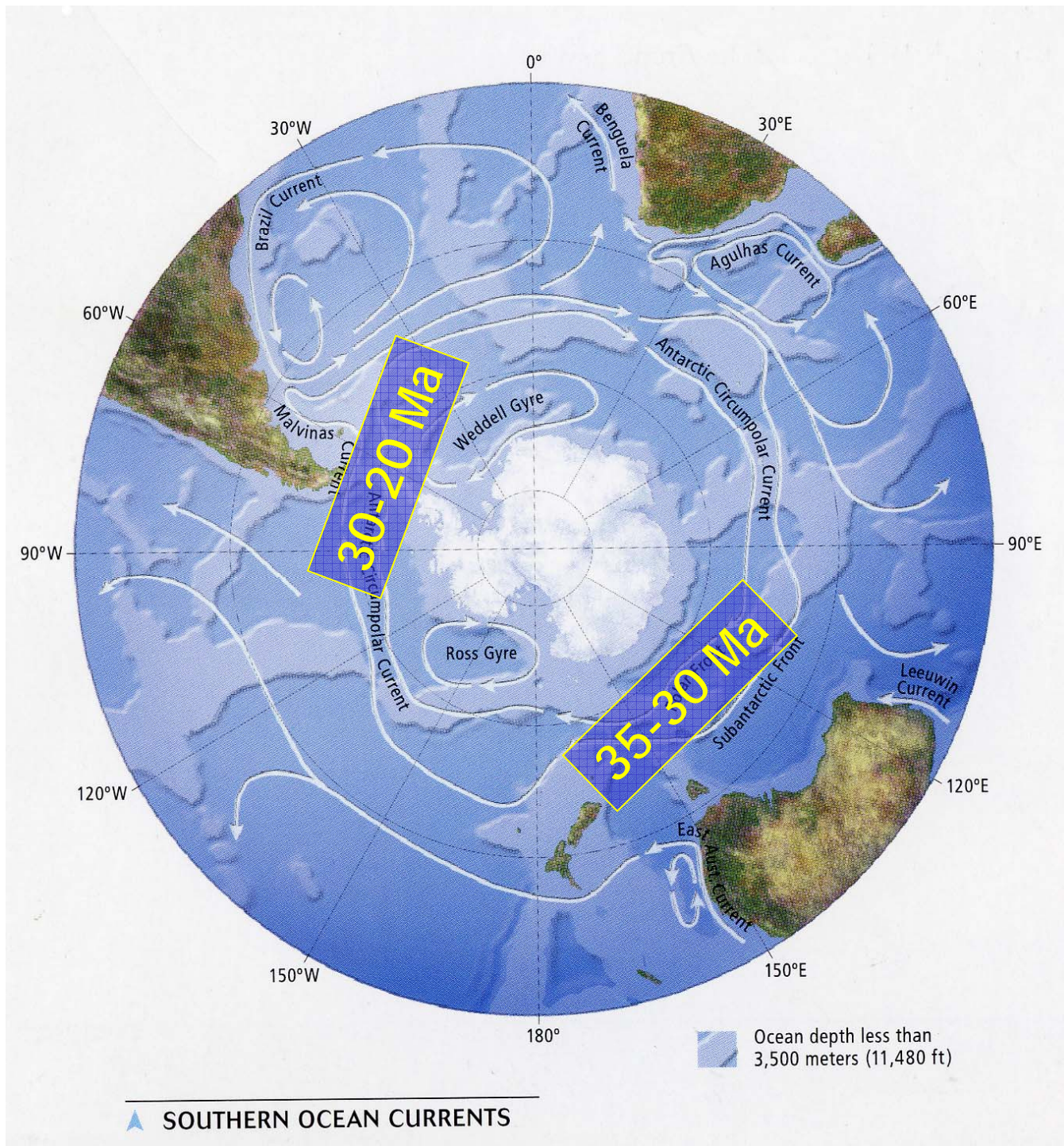
Lawver & Gahagan (2003)

ERA	PERIOD	EPOCH	YEARS AGO (millions)	
Cenozoic	Quaternary	Holocene	The present	
		Pleistocene	0.01	
	Tertiary	Pliocene	1.8	
		Miocene	5.3	
		Oligocene	24	
	Mesozoic	Cretaceous	Eocene	34
			Paleocene	55
Mesozoic	Jurassic		65	
			144	
			200	
Paleozoic	Triassic		248	
			290	
	Permian		354	
			417	
			443	
	Paleozoic	Carboniferous		490
				543
Paleozoic	Precambrian		3800	



Opening of the Drake Passage





McGonigal and Woodworth (2001)

Antarctic glacial history

- What are reconstructions based on?
 - Sediment cores (ocean)
 - sediment stratigraphy continental shelf
 - seismic studies
 - drill studies
 - Sea level reconstructions (coral)

Reconstructions of the Antarctic glacial history

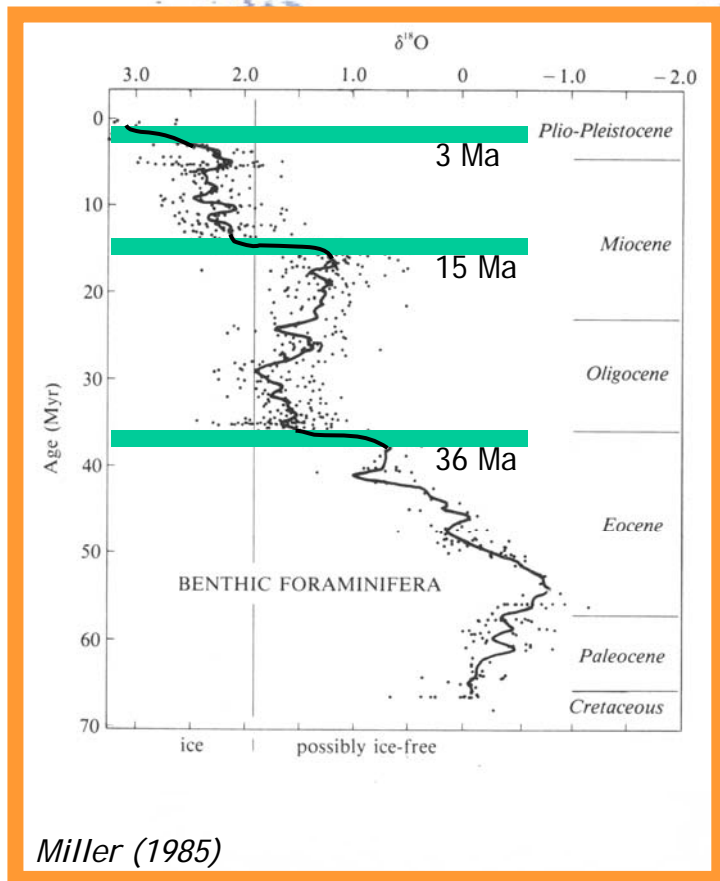
Far field evidence:

- indications from marine sediment cores
- indications from sea level data: far-field shelf stratigraphy & corals

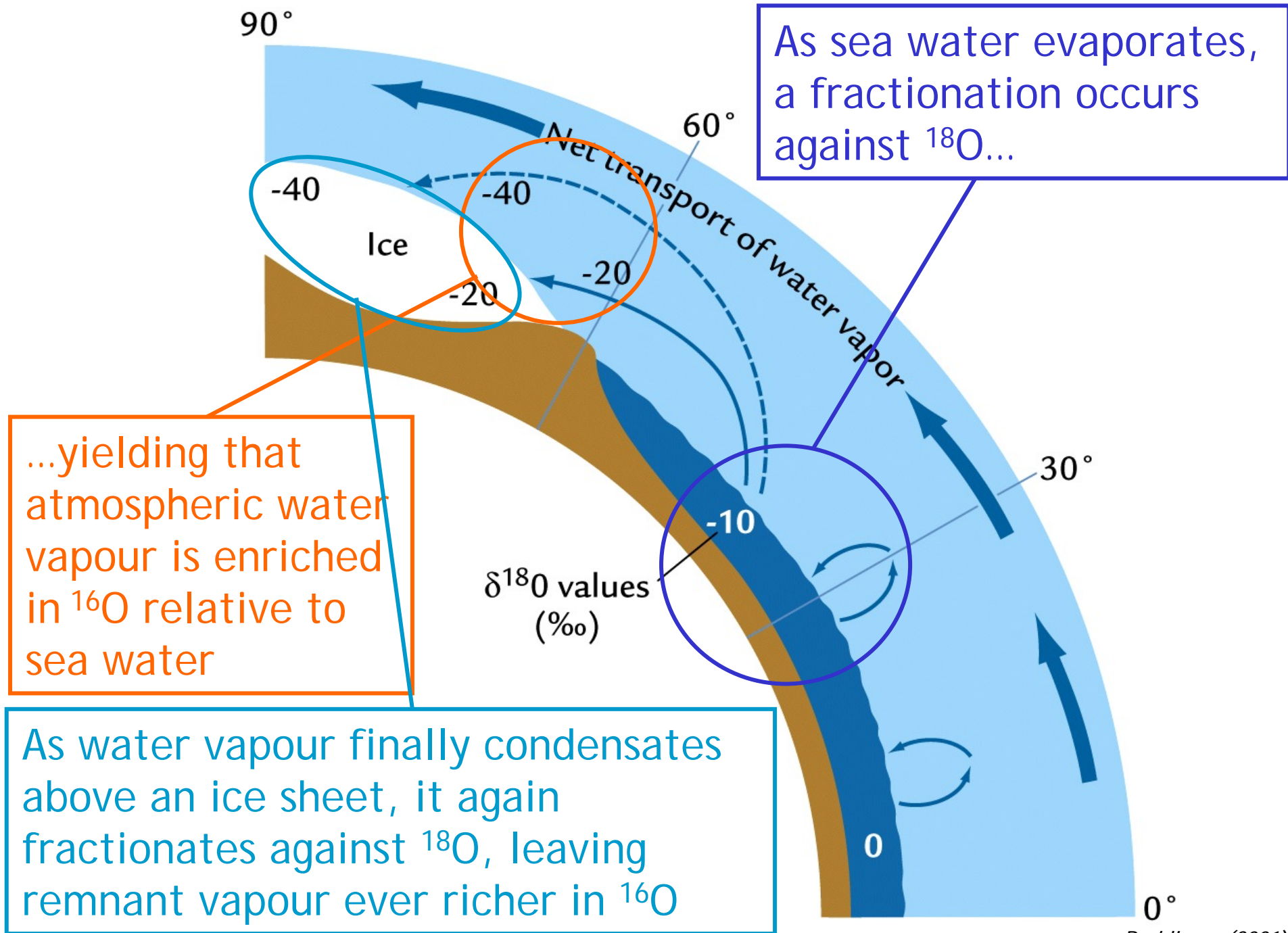
Near field evidence:

- indications from Antarctic shelf drill sites and stratigraphy
- indications from terrestrial sites: Mostly Pliocene
- Ice core records (not this lecture; 1 Ma presence)

Indications from marine cores



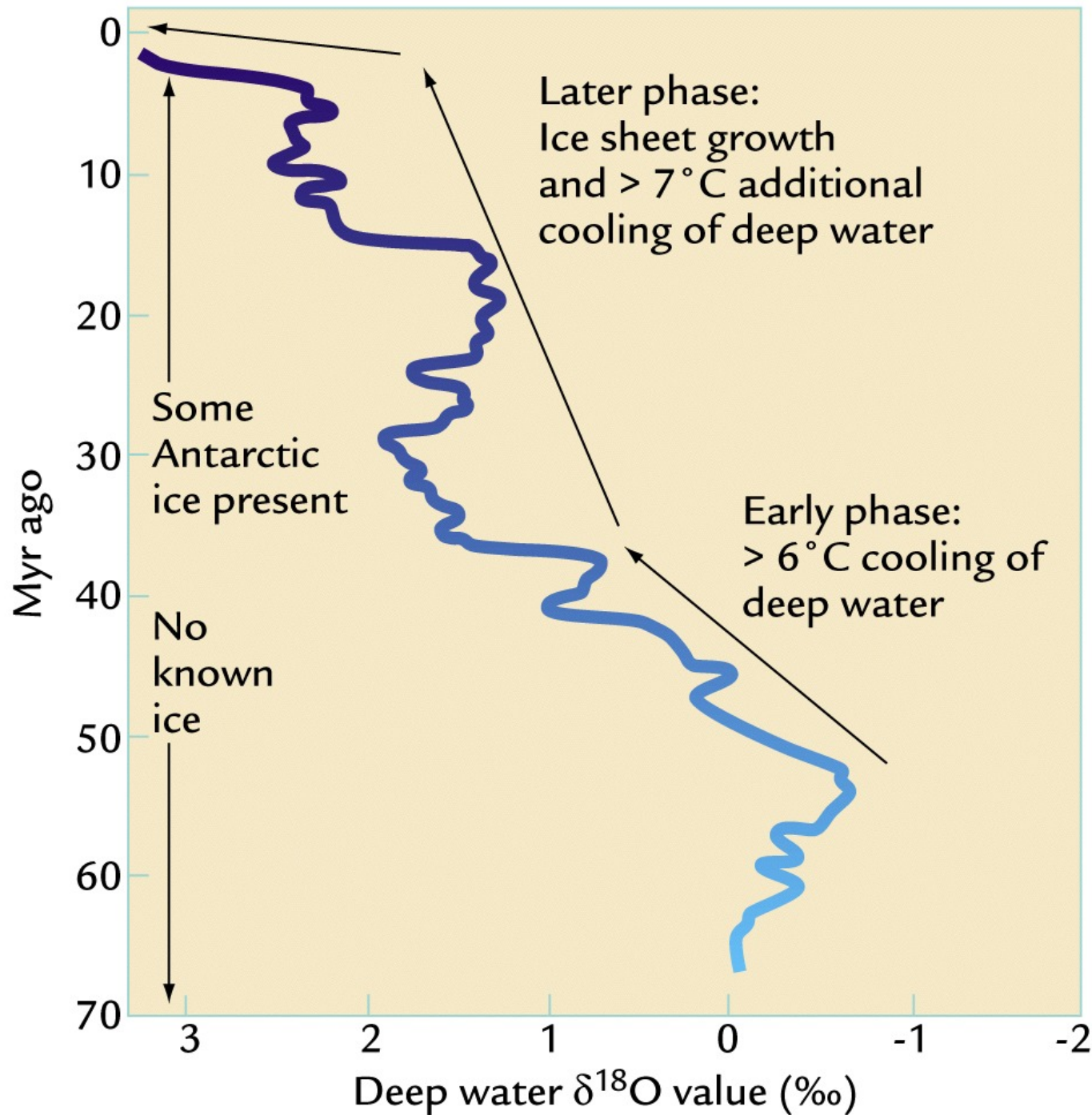
- Oxygen isotopes are derived from:
 - planktonic
 - benthic foraminifer
- Oxygen isotopes are interpreted as:
 - ocean temperature
 - ocean isotopic signature
 - ocean circulation (cancels out)
- This could mean that changes imply a change in either of these two remaining properties



As sea water evaporates, a fractionation occurs against ^{18}O ...

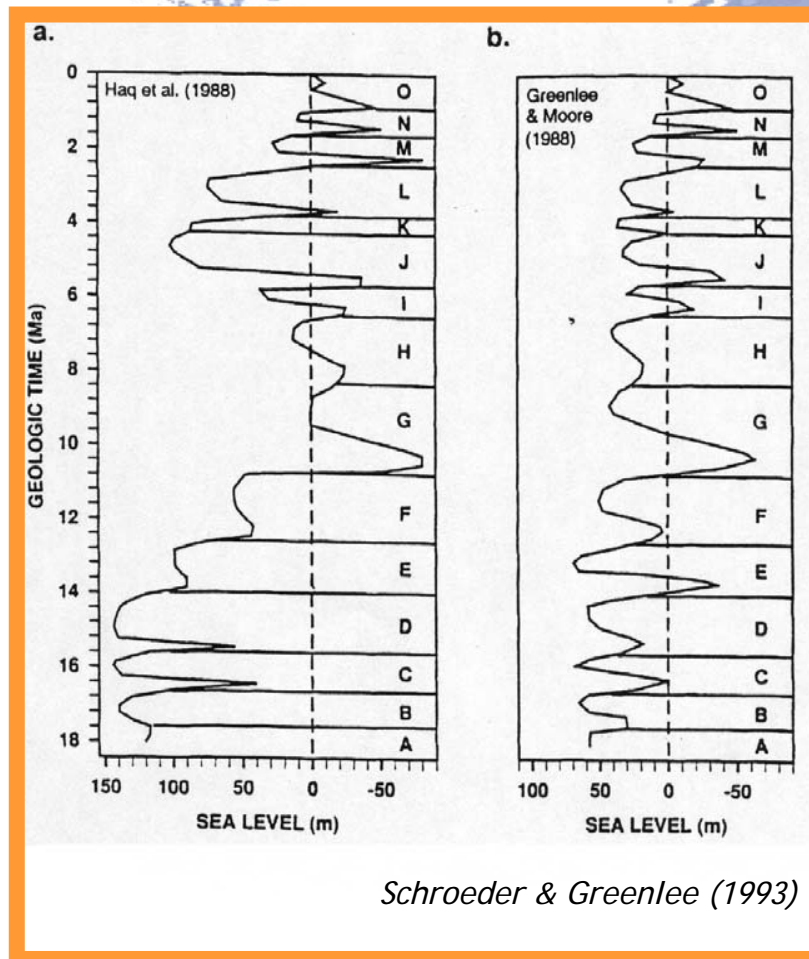
...yielding that atmospheric water vapour is enriched in ^{16}O relative to sea water

As water vapour finally condensates above an ice sheet, it again fractionates against ^{18}O , leaving remnant vapour ever richer in ^{16}O



About 1‰ $\delta^{18}\text{O}$ change during the Cenozoic is due to ice sheet growth, the rest is deep water temperature change

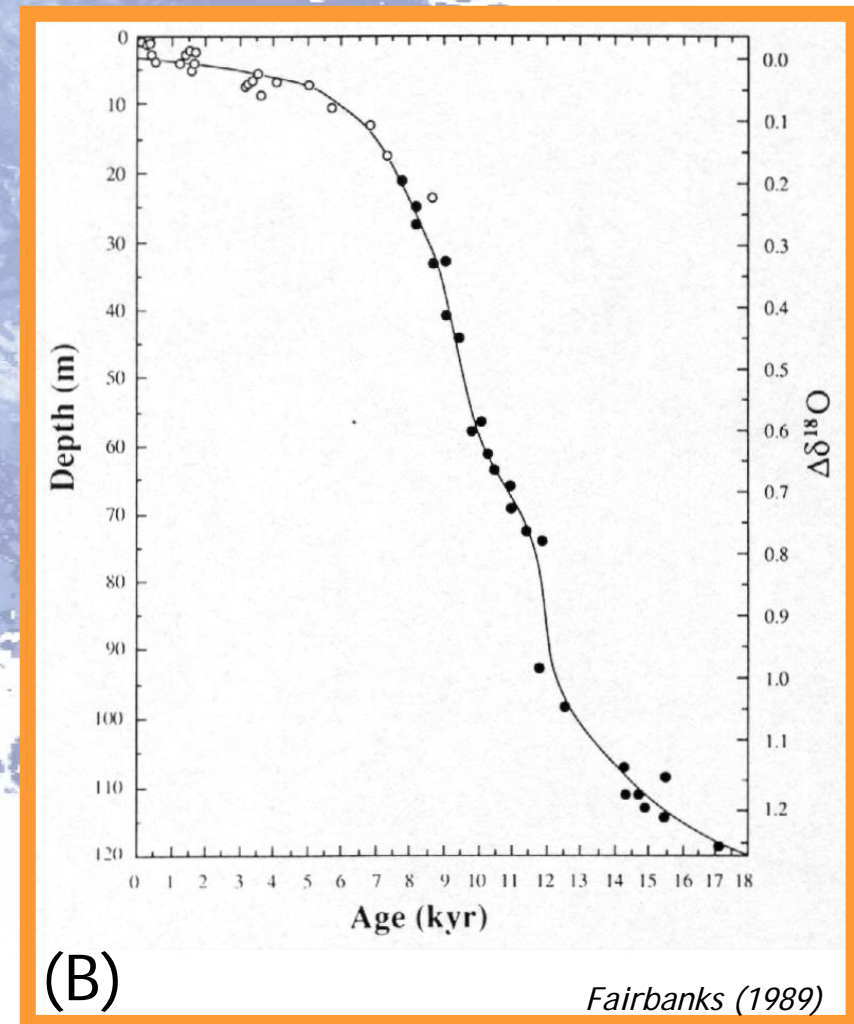
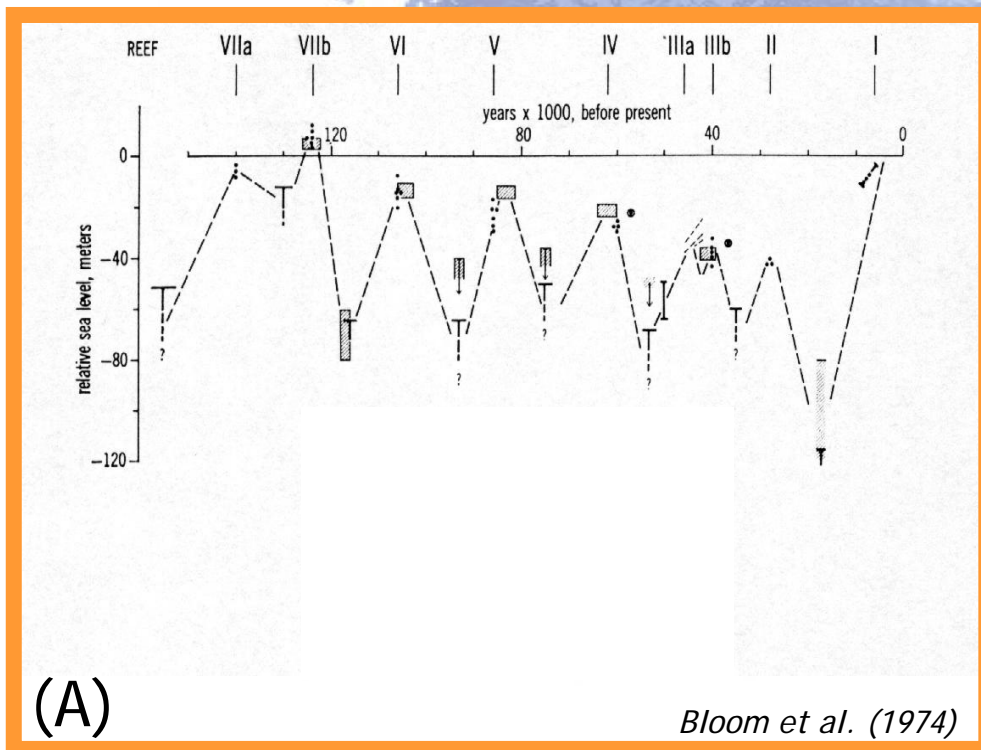
Indications from sea level data from the continental shelf



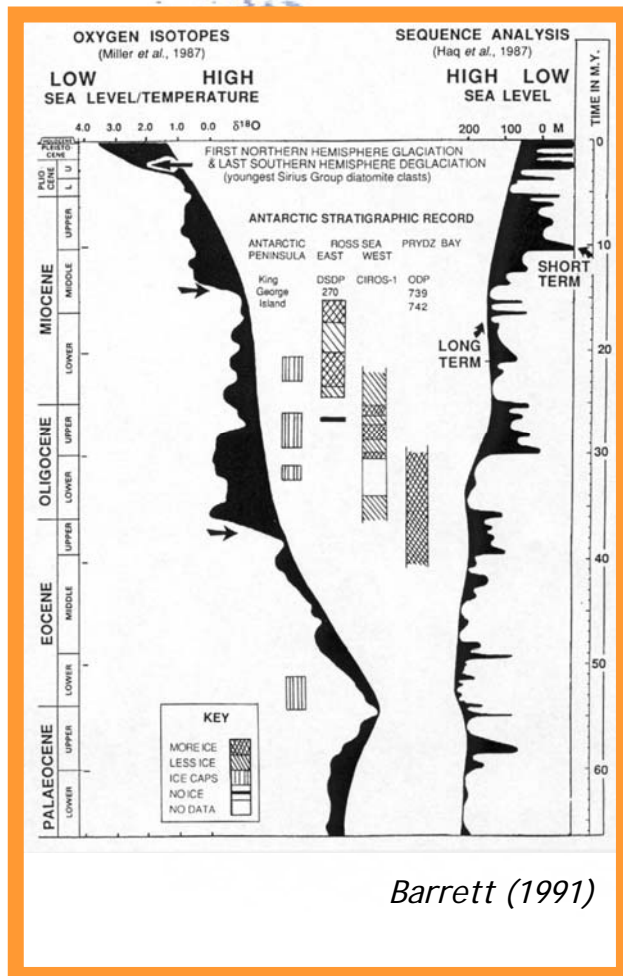
- Problems associated with sea level calculations:
 - amplitude of the signal (equated to ice volume)
 - generalisation, e.g. all of last 2 myr consists of two high "events", even though there were perhaps more than 20 glaciations.

Indications from sea level data from corals

- (A) Eemian interglacial-today
 - (B) late glacial
- global sea level history



Indications from sea level data: wrap-up

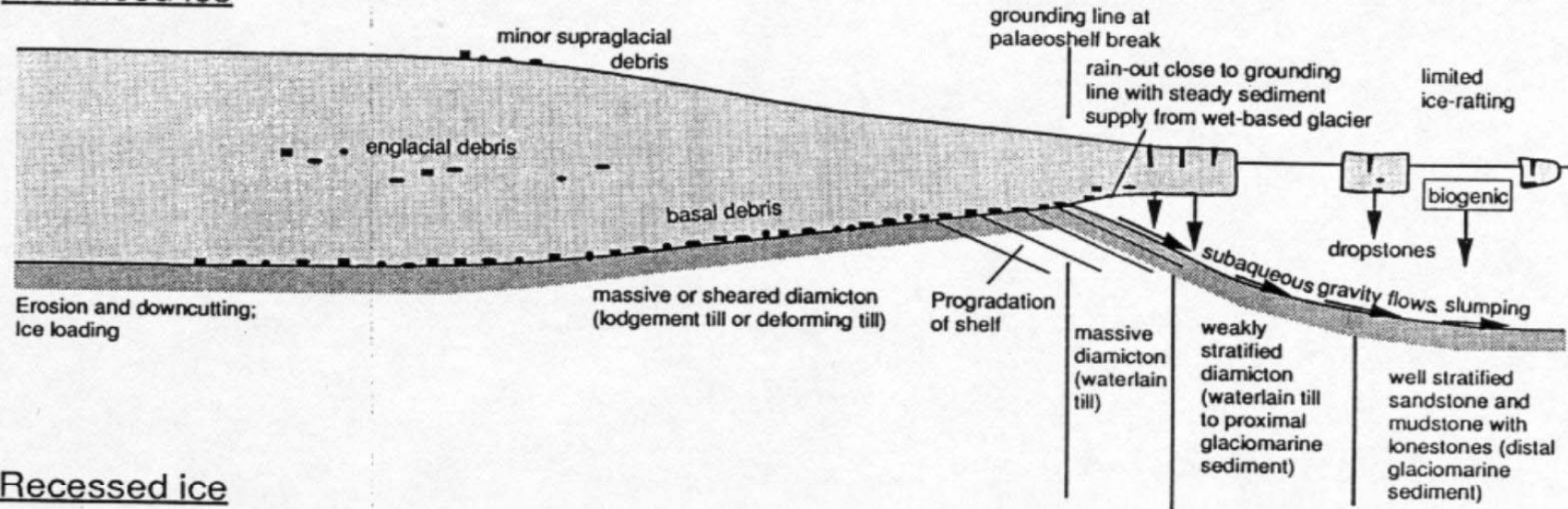


- Sea level changes are inferred from:
 - corals
 - depositional structures on the continental shelf, *i.e.*
 - transgression: horizontal structures
 - regression: erosion surfaces

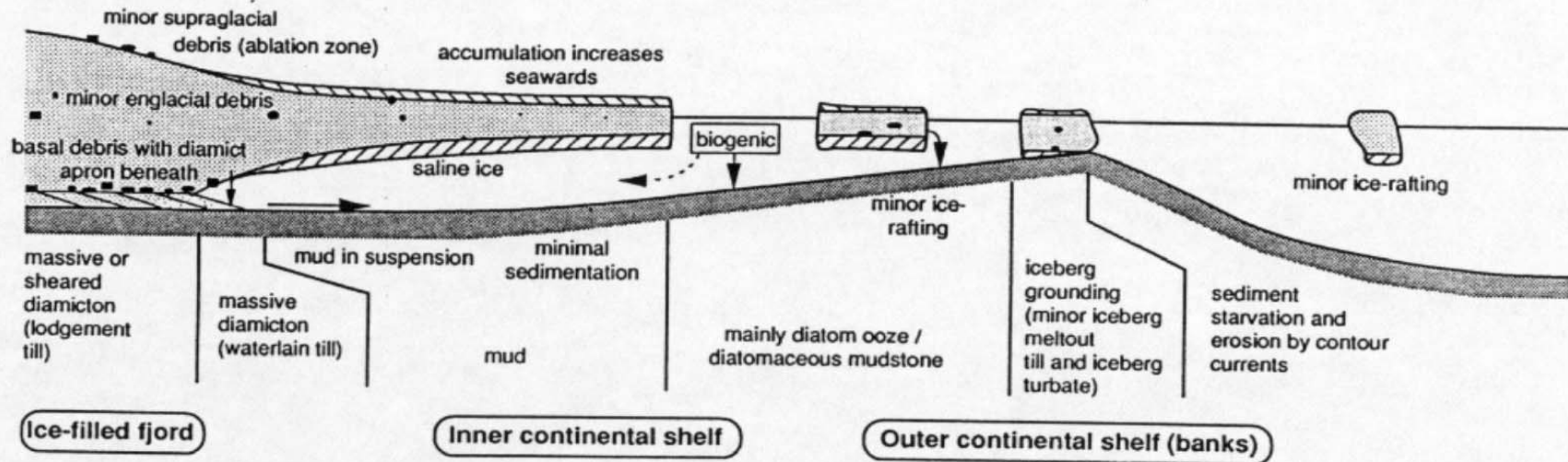
But remember:

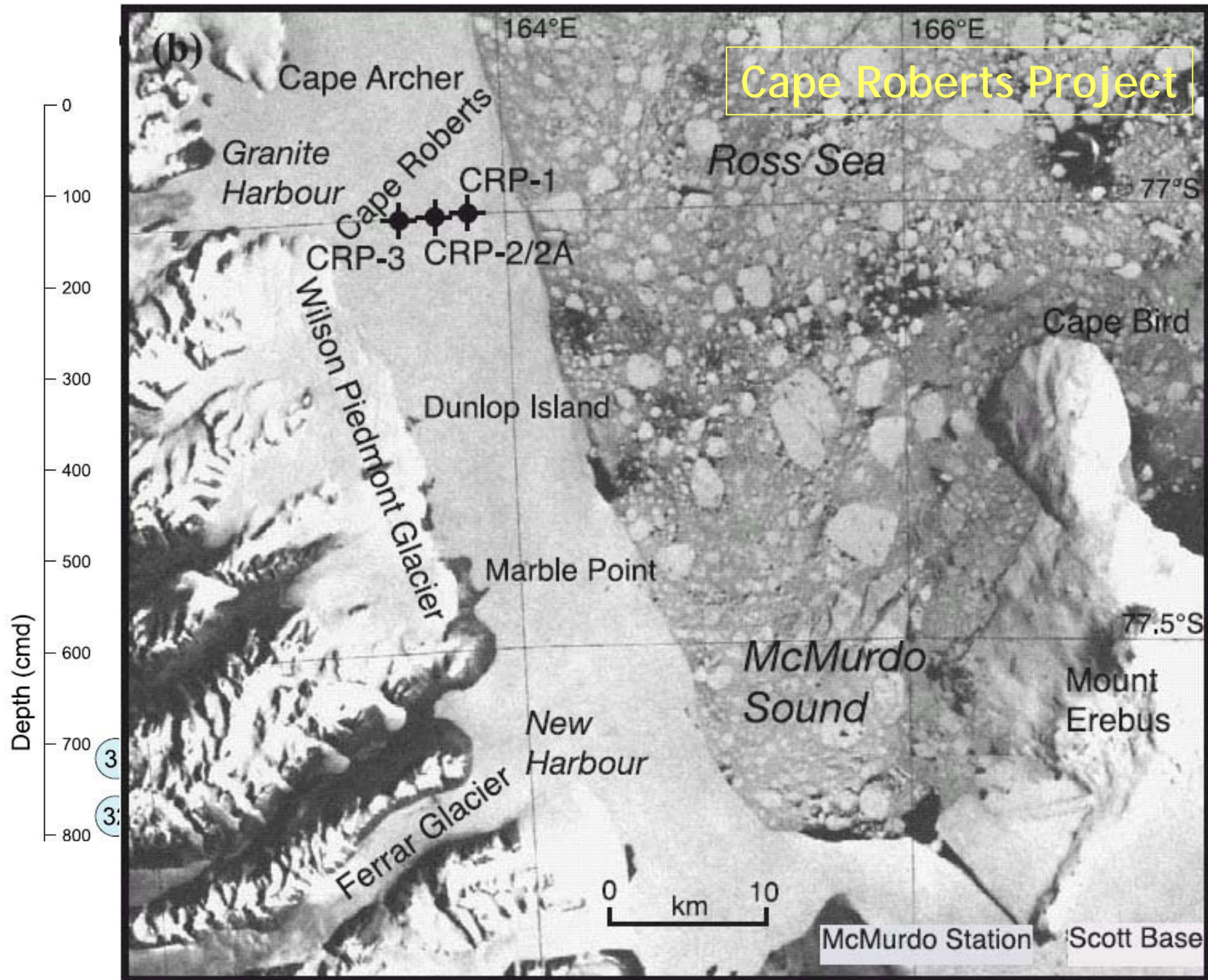
- Sea level changes are a function of:
 - tectonics (trends- global)
 - tectonics/isostasy (regional signature)
 - ice sheet volume (global signature)

(A) Advanced ice



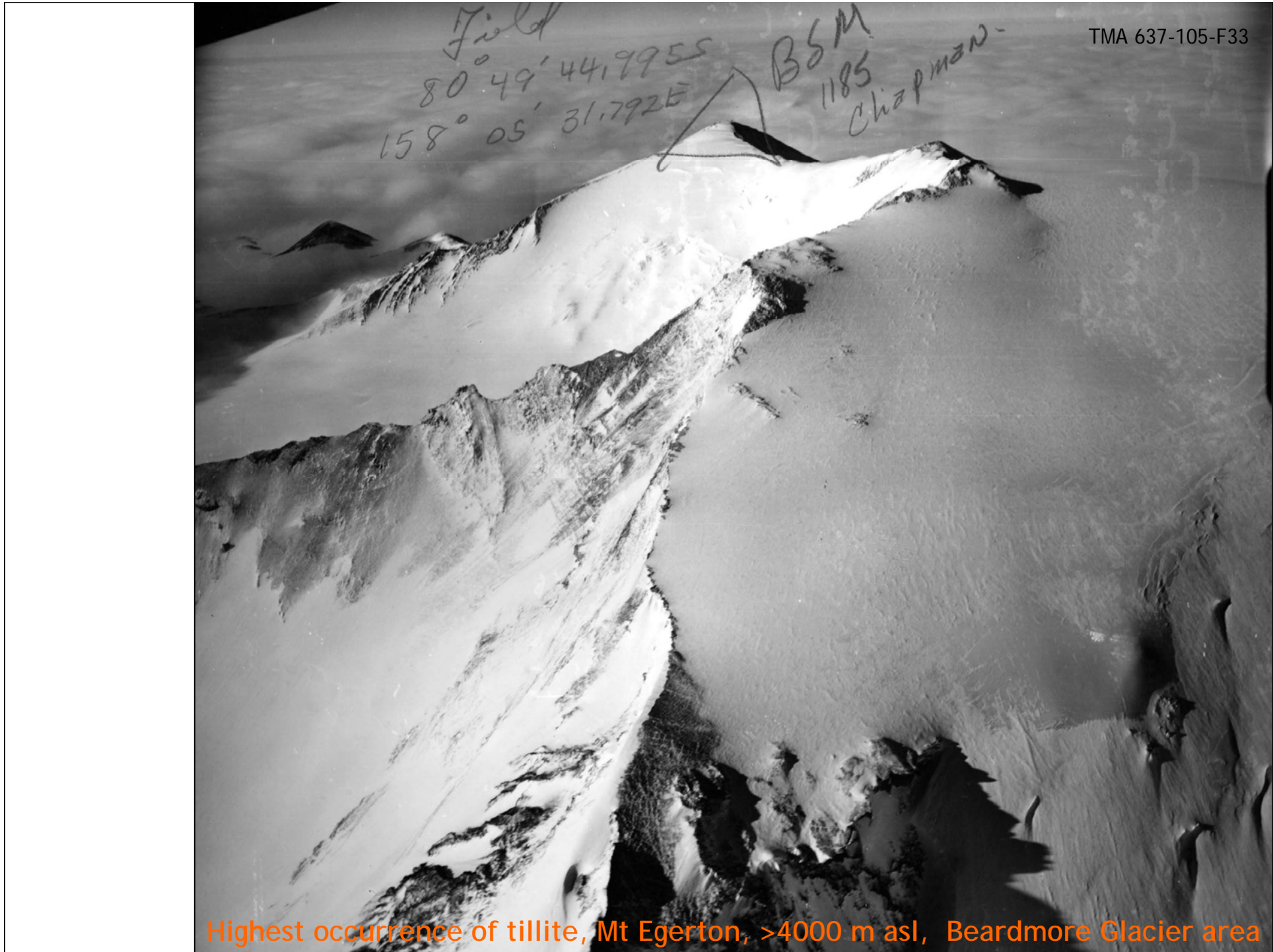
(B) Recessed ice





Indications from terrestrial sites: Near-coastal areas

- Evidence from vegetation remains:
 - was there a Pliocene community?
- Terrestrial glacial stratigraphy
 - Lambert Graben area
 - Transantarctic Mtns
- N. Peninsula evidence of MAT leading up to inception of ice sheets, last warm period early Eocene: 7-15 °C.
- Deposits dating back to Oligocene - e. Miocene, mostly using microfossil, ash stratigraphy & TCN dating evidence



Highest occurrence of tillite, Mt Egerton, >4000 m asl, Beardmore Glacier area

Indications from terrestrial sites: Transantarctic Mountains



Image by Alan Ashworth

Webb & Harwood (1993)

Mt. Fleming, Antarctica

- One Pliocene reconstruction:
 - WAIS had melted away
 - EAIS strongly reduced

T. vulnifica

➤ The reconstruction is based on the occurrence of marine microfossils in tillite deposits in the TAMS

Tillite @ 2000 m asl on Mt Fleming

This is where diatoms were found



Mt Fleming, Antarctica



...in a pit....

...but only in the top ~5 cm layer (red), and they are wind-blown

A photograph of a dark, rocky slope in the foreground, leading to a body of water and snow-capped mountains in the background. The text is overlaid on the upper portion of the image.

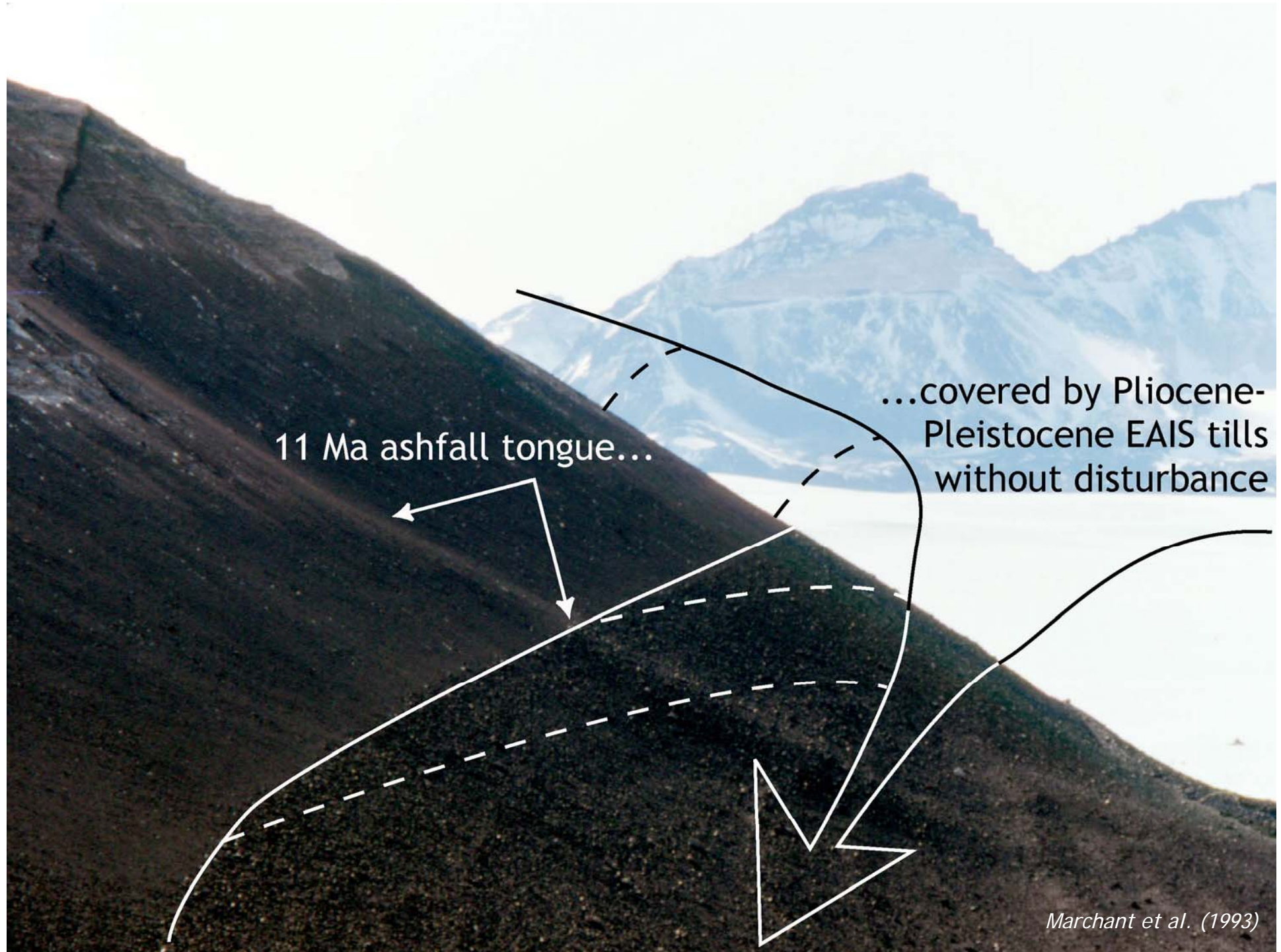
There are problems associated
with Pliocene warmth...

Taylor valley, Antarctica

...in Taylor Valley

11 Ma ashfall tongue...

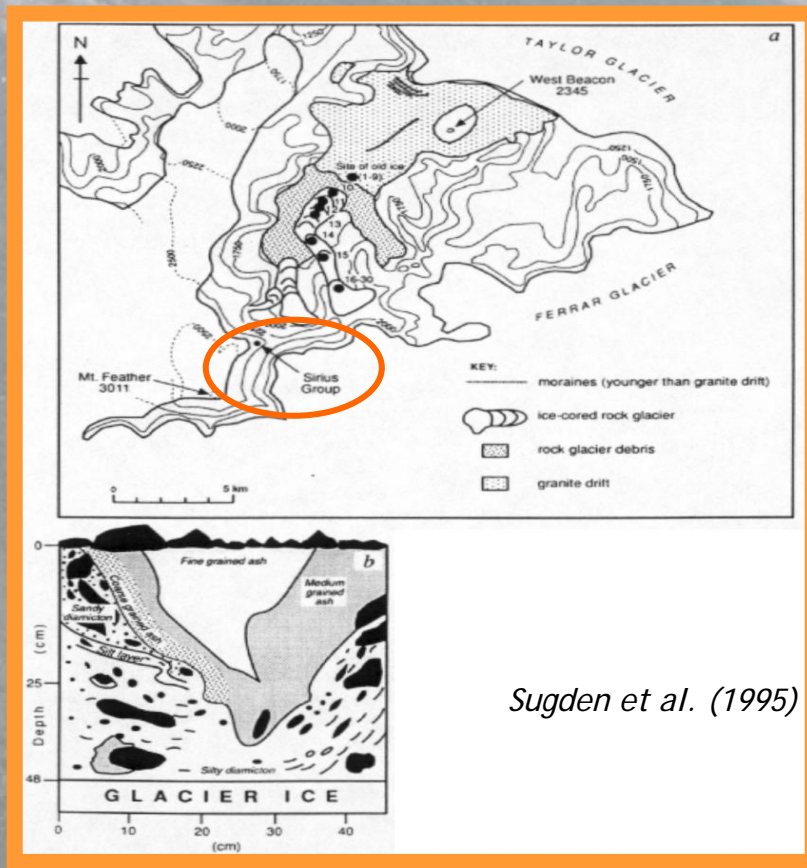




11 Ma ashfall tongue...

...covered by Pliocene-Pleistocene EAIS tills without disturbance

...and in Beacon Valley



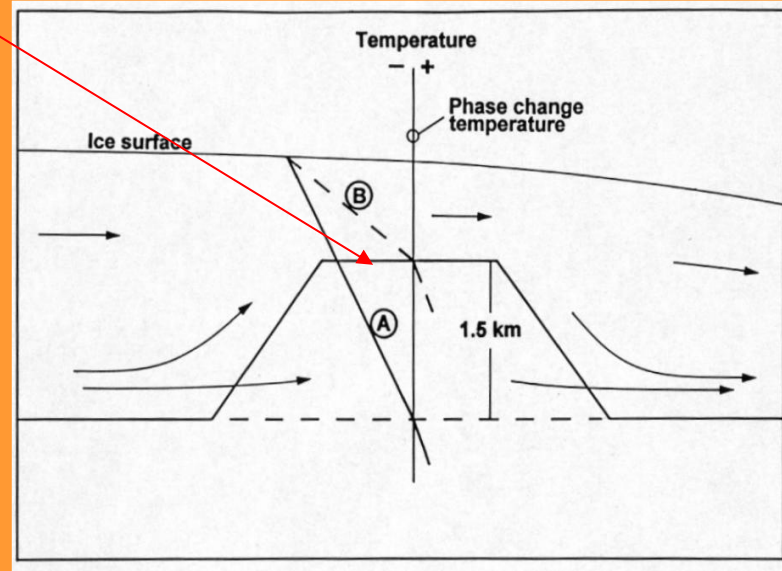
Sugden et al. (1995)

Current ash beds have been previously identified in the Middle Pleistocene. It has been suggested that the appearance of the ash beds in Beacon Valley implies that the deglaciation during the Pliocene.

In addition, these diatoms occur in odd topographic situations...

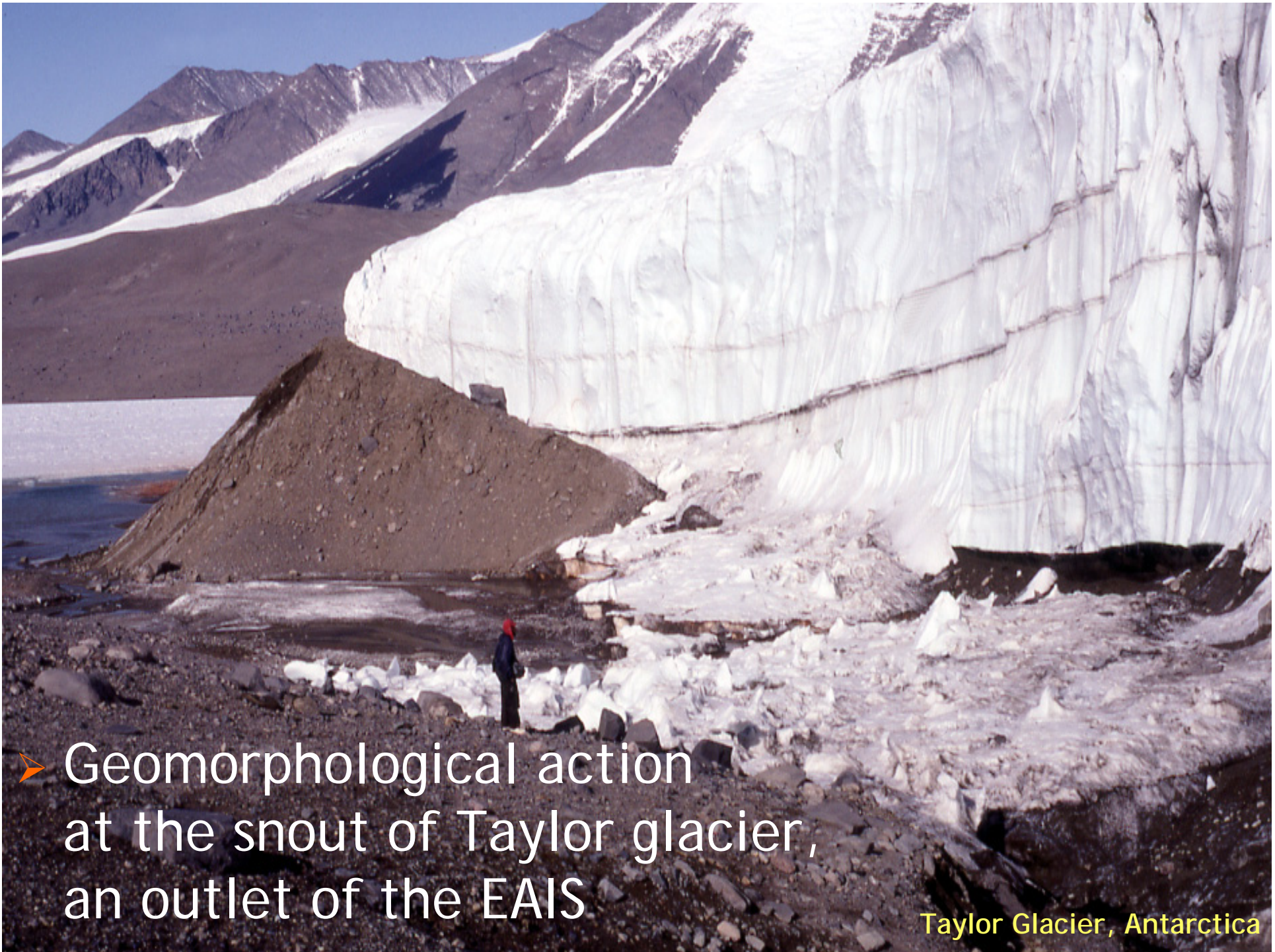
Tillite with microfossils

Foto: D. Marchant



Stroeven & Kleman (1999)

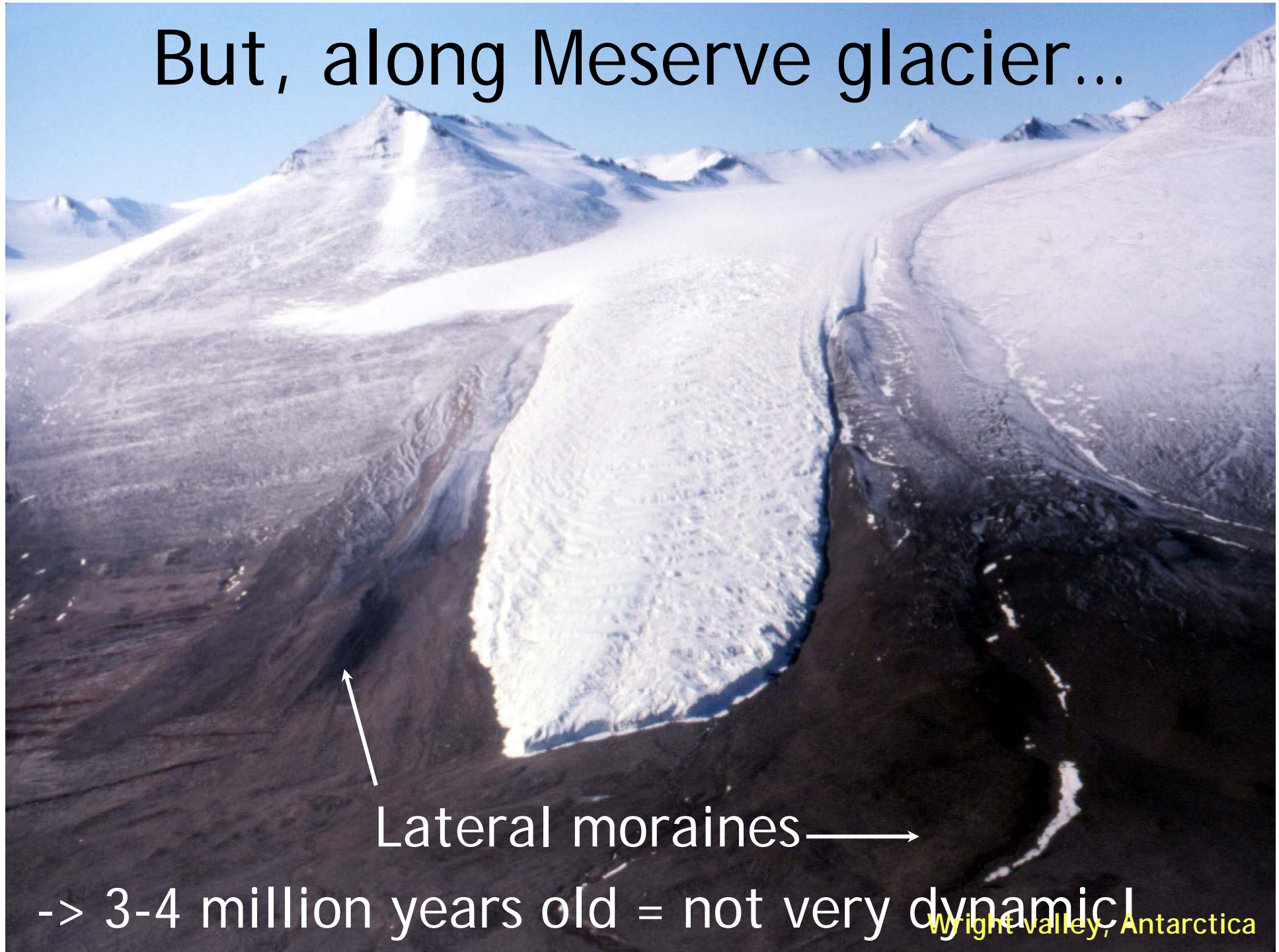
Mount Feather, Antarctica



- Geomorphological action at the snout of Taylor glacier, an outlet of the EAIS

Taylor Glacier, Antarctica

But, along Meserve glacier...



Lateral moraines

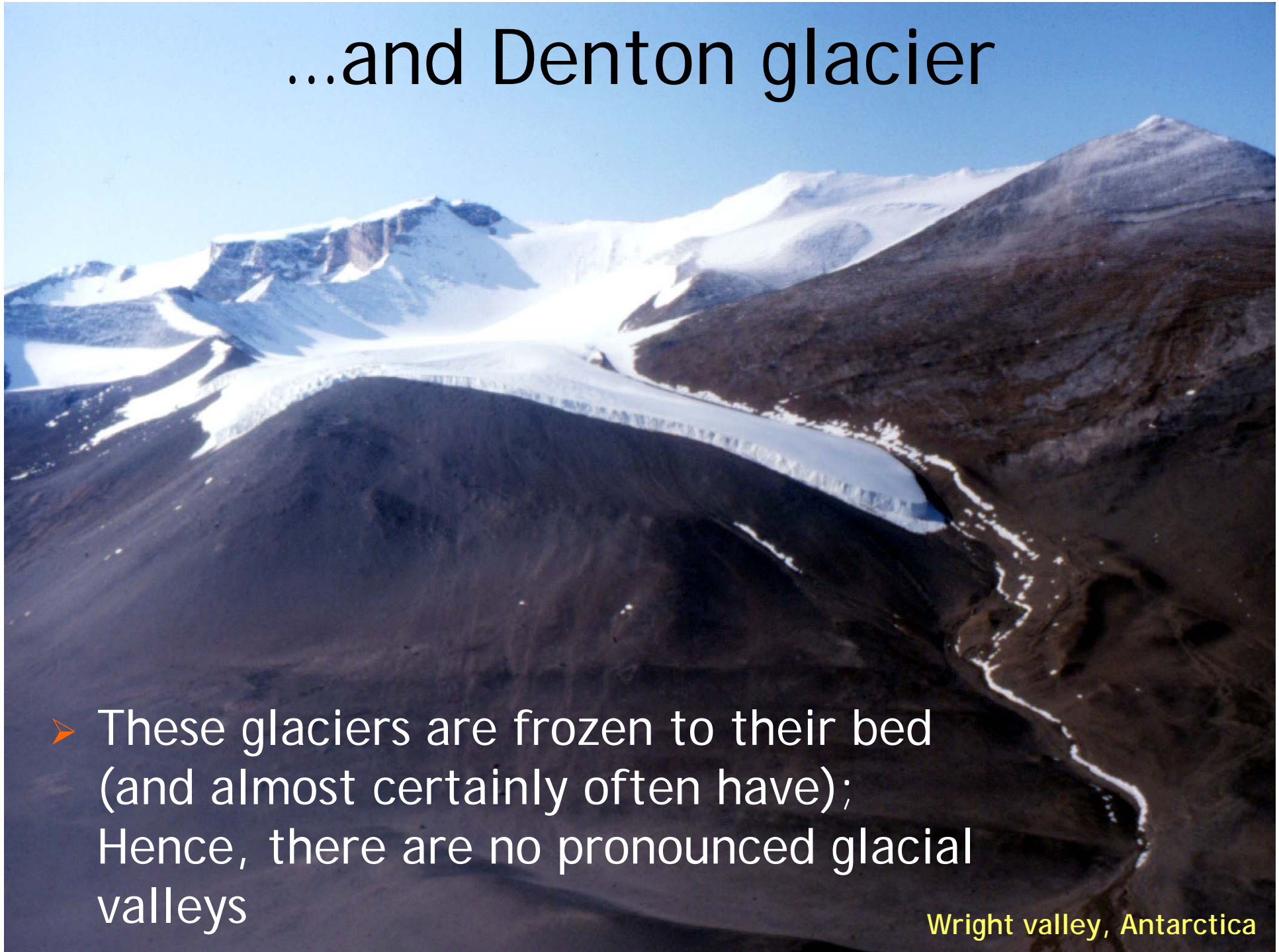
-> 3-4 million years old = not very dynamic

Wright valley, Antarctica

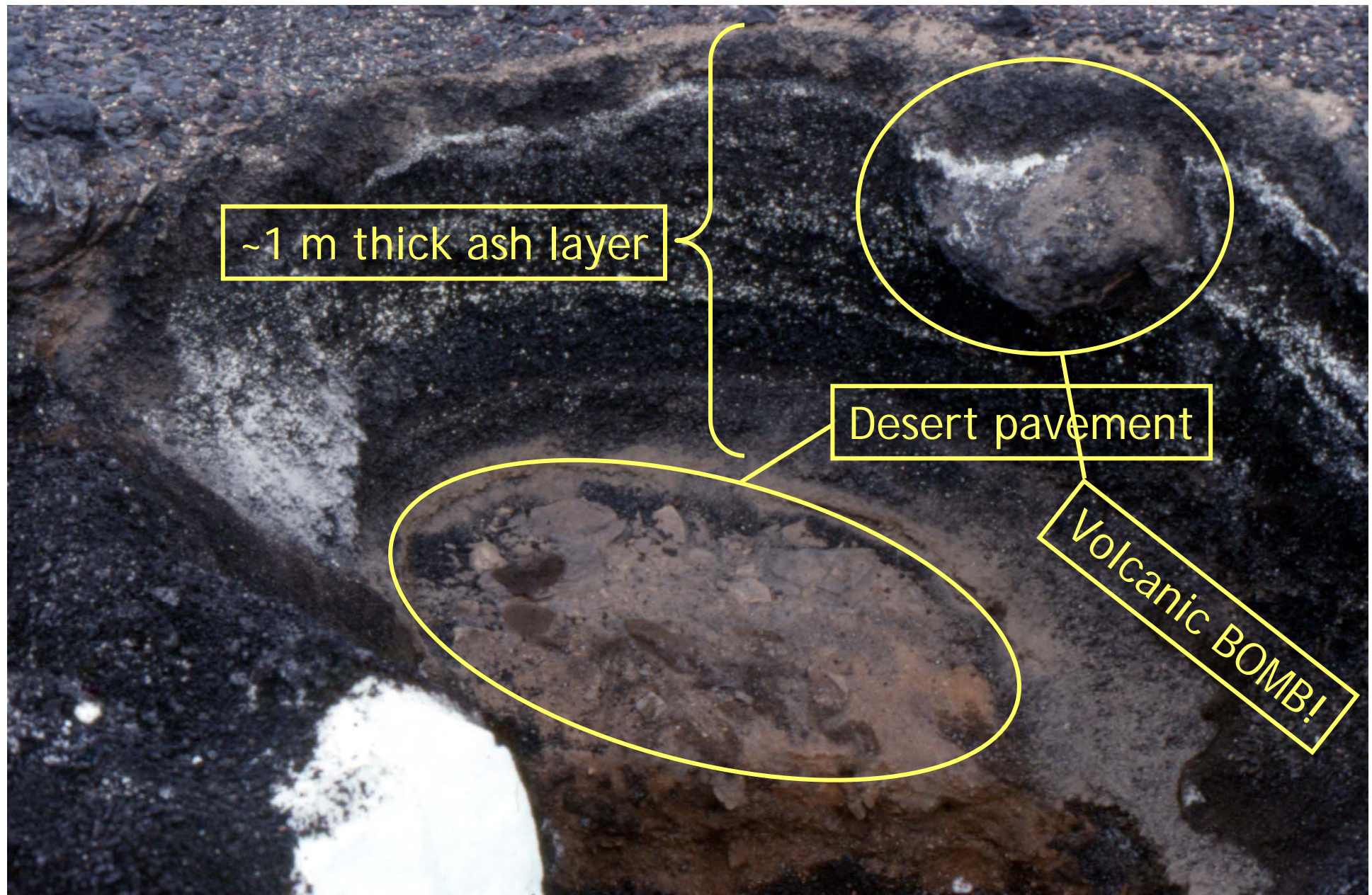
...and Denton glacier

- These glaciers are frozen to their bed (and almost certainly often have); Hence, there are no pronounced glacial valleys

Wright valley, Antarctica



Pliocene ashfall deposit in Taylor valley



Antarctic geomorphology wrap-up

- Pliocene lateral moraines (and volcanic cones)
- 8-15 million years old ashes
- Geomorphological comparison w/ USA
- Trunk valley in which they are situated is older than the Pliocene: so are *its* tills
- Valleys, ice and sediment underneath the ashes are older, tills on top were deposited by cold-based ice
- EAIS has not everywhere left a definitive imprint

TAMS glacial history...

		Pre-55 Ma	Post-rifting 55 Ma -15 Ma	15 Ma - present
Geographic setting		Continental interior	Uplifted margin close to open ocean	Similar to present
Elevation		Low	High	High
Precipitation climate		Dry	Moderately wet	Dry
Temperature		Cool temperate	Decreasing trend with superimposed oscillations → Extreme cold	
Expected process system	Semi-arid subaerial			
	Fluvial subaerial			
	Wet-bed glacial			
	Dry-bed glacial polar desert			
Relief		Low	Increasing	High
Suggested time of Sirius Group deposition		-----→		

Stroeven & Kleman (1999)

- Early Tertiary *cold temperate*
- 55->15 Ma *growing colder and drier*
- last 15 Ma *similar to today, cold and arid*

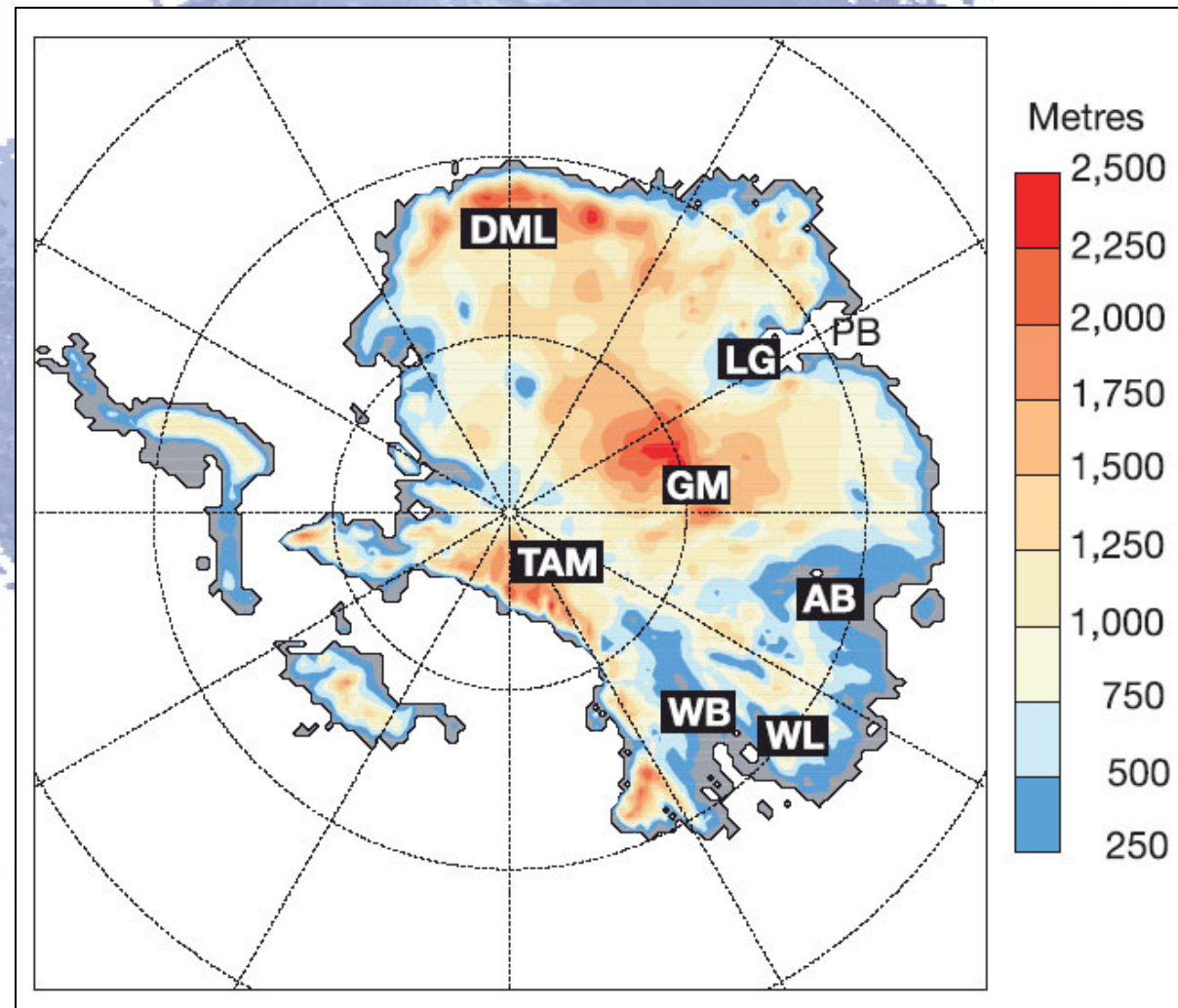
Wright valley, Antarctica

Ice sheet and climate modelling

- 34 Ma transition
- Pliocene warmth
- Cenozoic history
- DeConto & Pollard (2003a, b), *P³, Nature*
- Thorn & DeConto (2006), *P³*
- Huybrechts (1993): *Geogr. Ann. not shown*
- Oerlemans (2004), *P³*

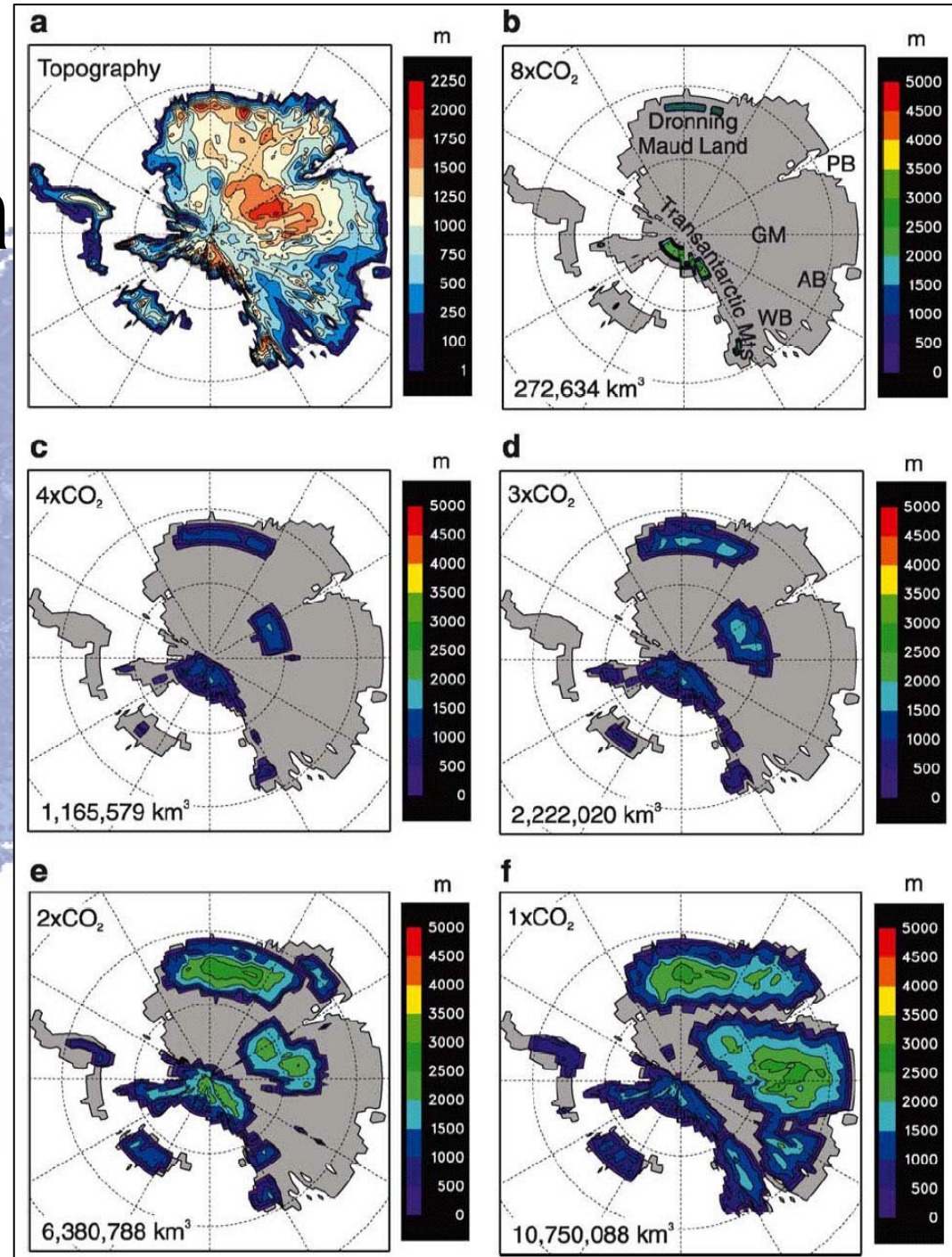
34 Ma glacial inception

Initial
topography:



34 Ma gla

Ice sheets run to steady state following five different CO₂ boundary conditions: Non-linear response



34 Ma glacial inception

Effect of Drake Passage upon the inception of ice sheets on Antarctica

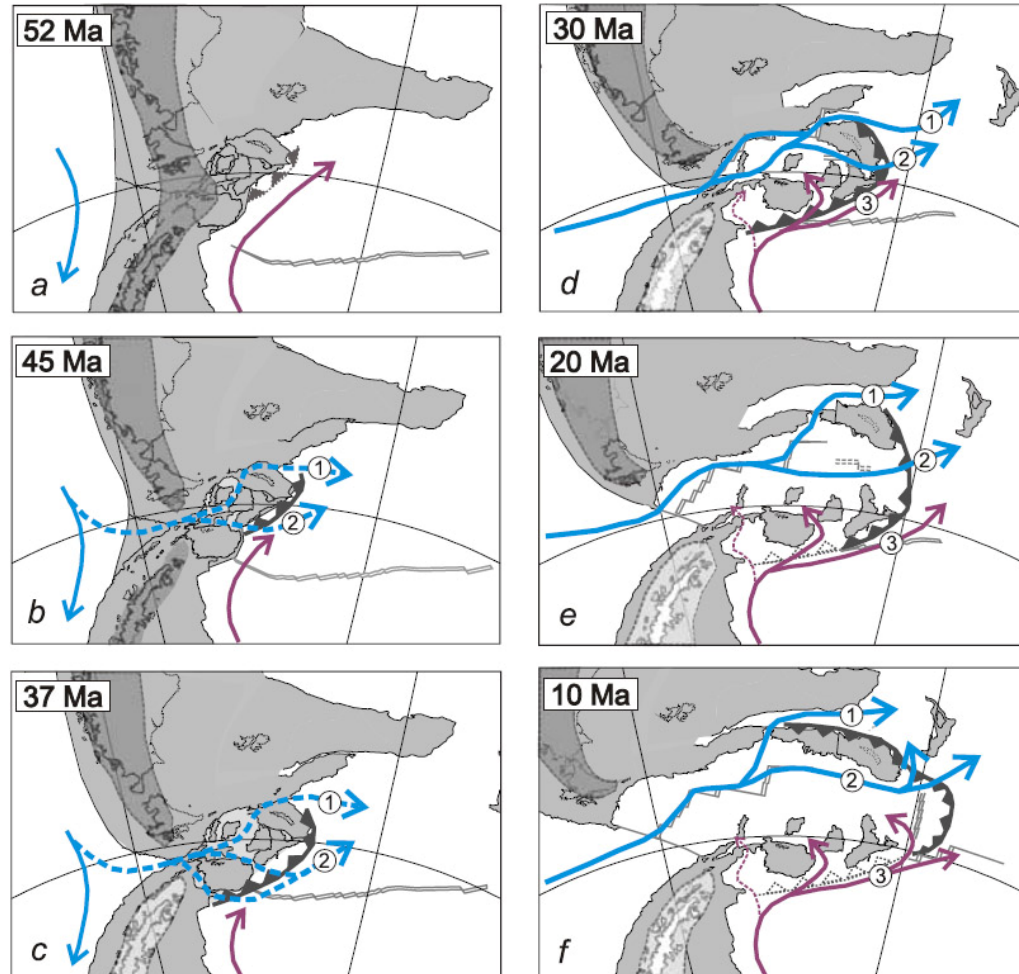
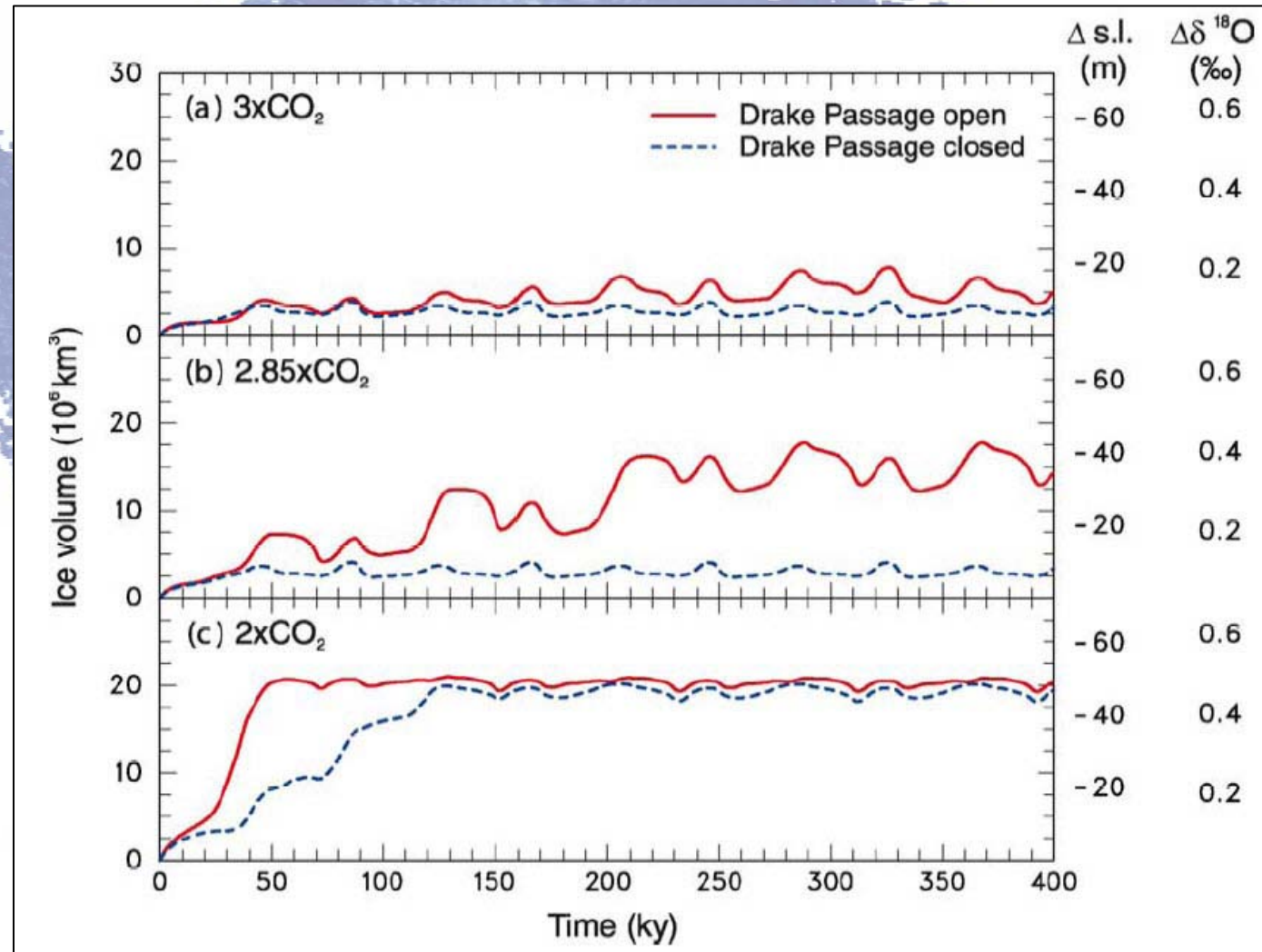


Fig. 2: Livermore et al. (2007)

34 Ma glacial inception

Effect of Drake Passage upon the inception of ice sheets on Antarctica



34 Ma

Effect of vegetation on 2-m surface temperatures on Antarctica for Evergreen and Tundra vegetation and ice sheet growth for 2 different CO₂ scenarios and 2 different orbital configurations

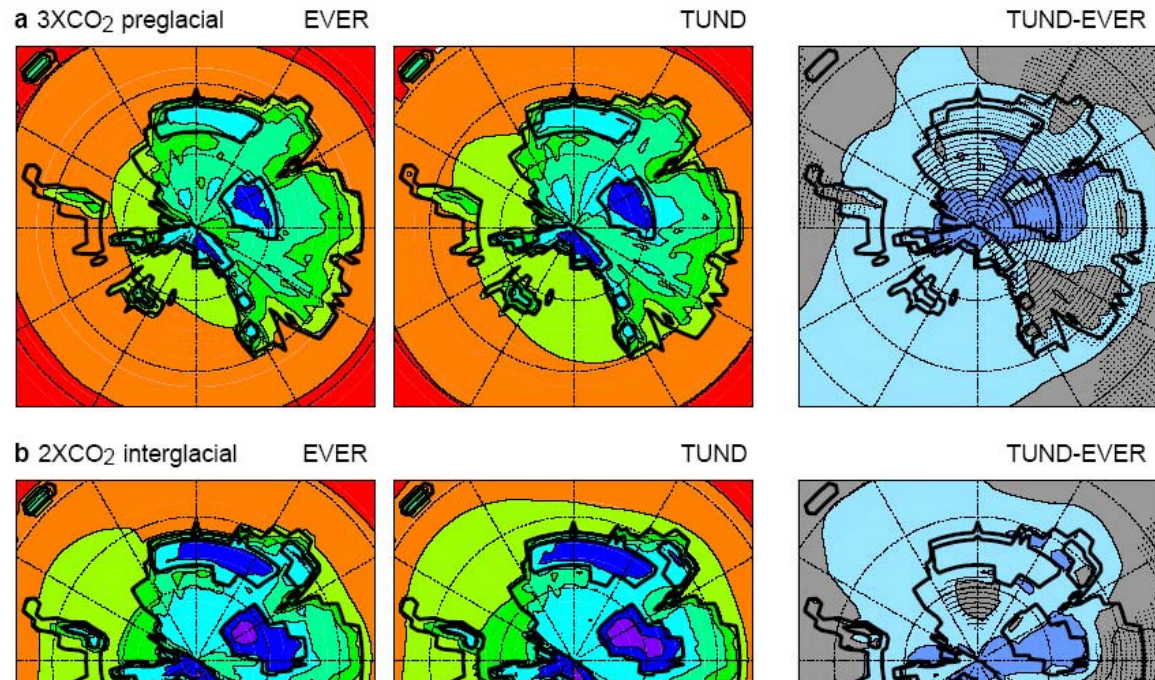


Table 2

Model simulations and relevant model inputs

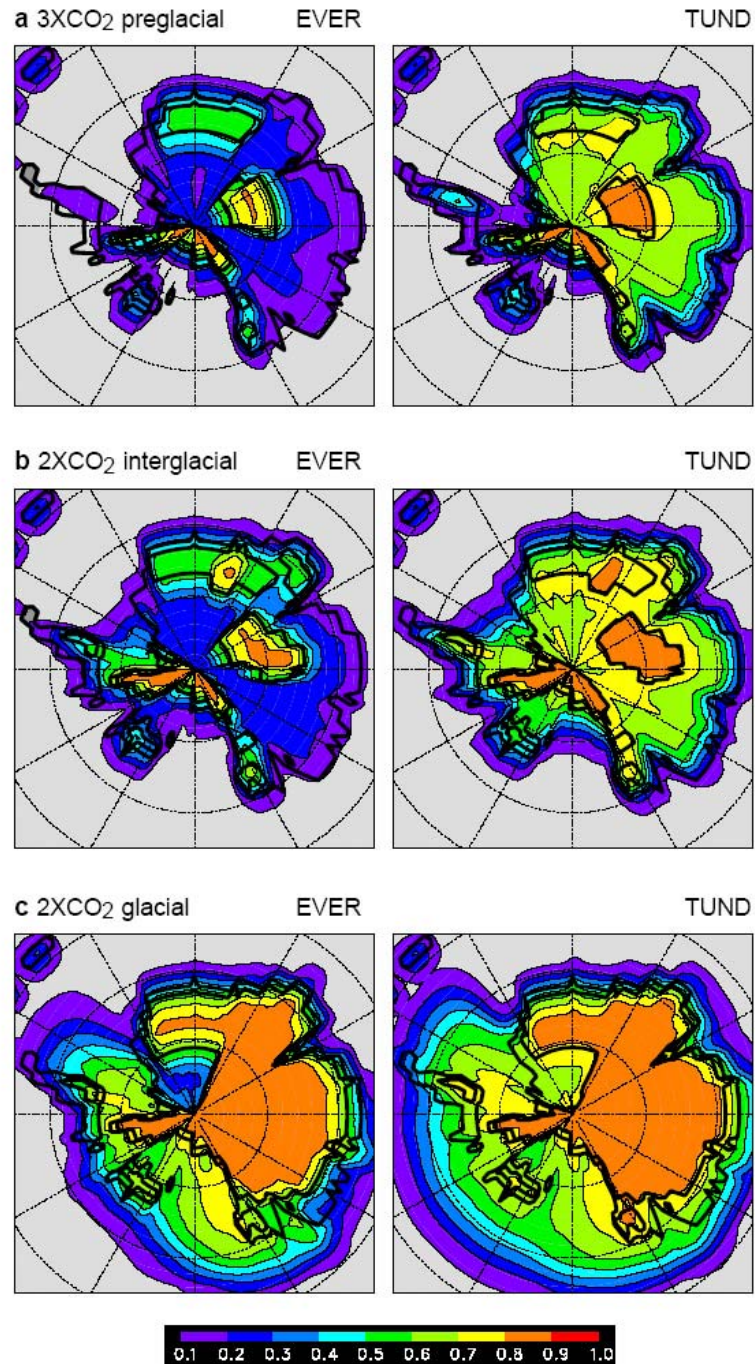
Model run	CO ₂ mixing ratio ^a	Orbital parameters (ecc., obl., pre.) ^b	Ice sheet geometry ^c	Antarctic vegetation ^d
1	2 × CO ₂	0.04, 23.5, 90	interglacial	TUND
2	2 × CO ₂	0.04, 23.5, 90	interglacial	EVER
3	2 × CO ₂	0.01, 22.5, 270.0	glacial	TUND
4	2 × CO ₂	0.01, 22.5, 270.0	glacial	EVER
5	3 × CO ₂	0.04, 23.5, 90	preglacial	TUND
6	3 × CO ₂	0.04, 23.5, 90	preglacial	EVER



Table 2: Thorn & DeConto (2006)

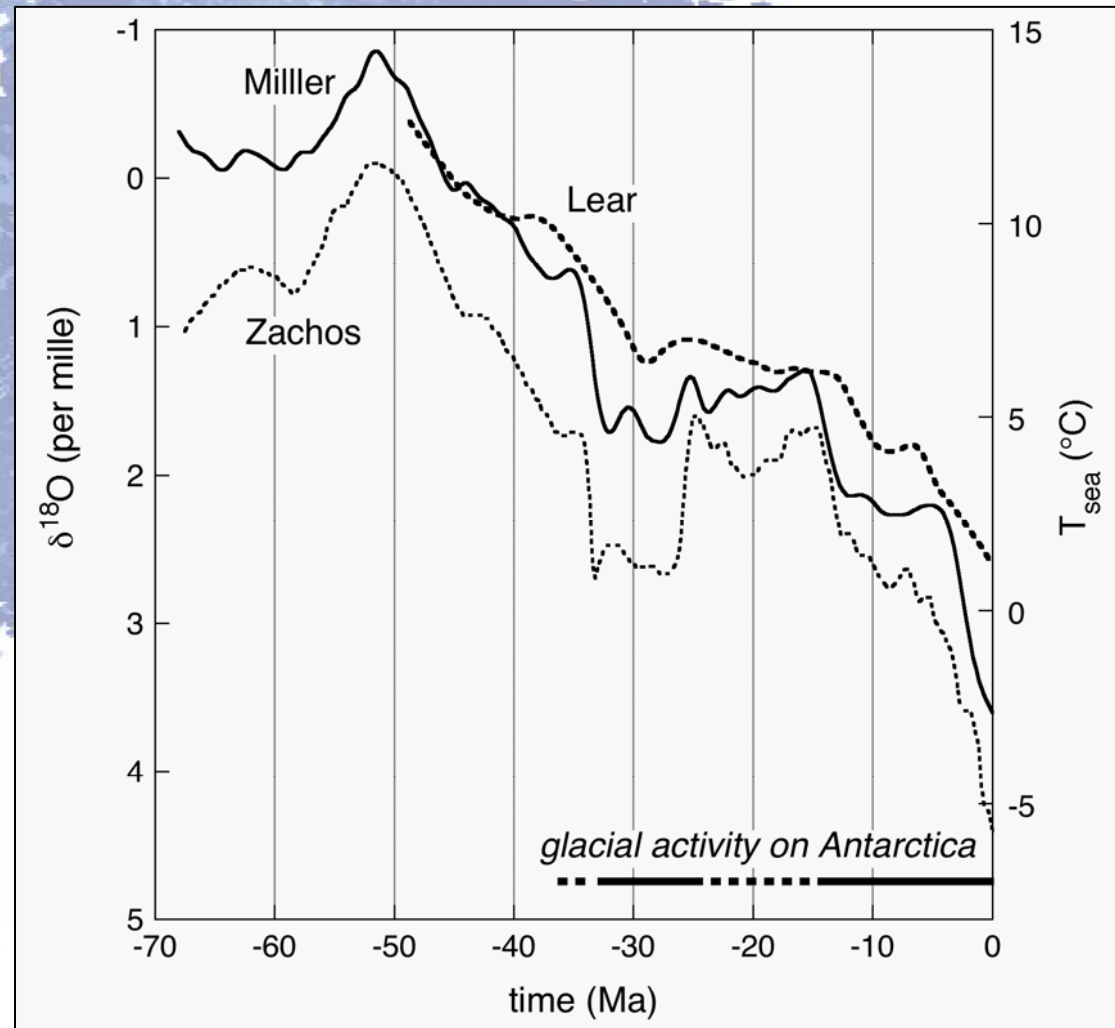
34 Ma glac

Effect of vegetation on surface albedo on Antarctica for Evergreen and Tundra vegetation and ice sheet growth for 2 different CO₂ scenarios and 2 different orbital configurations



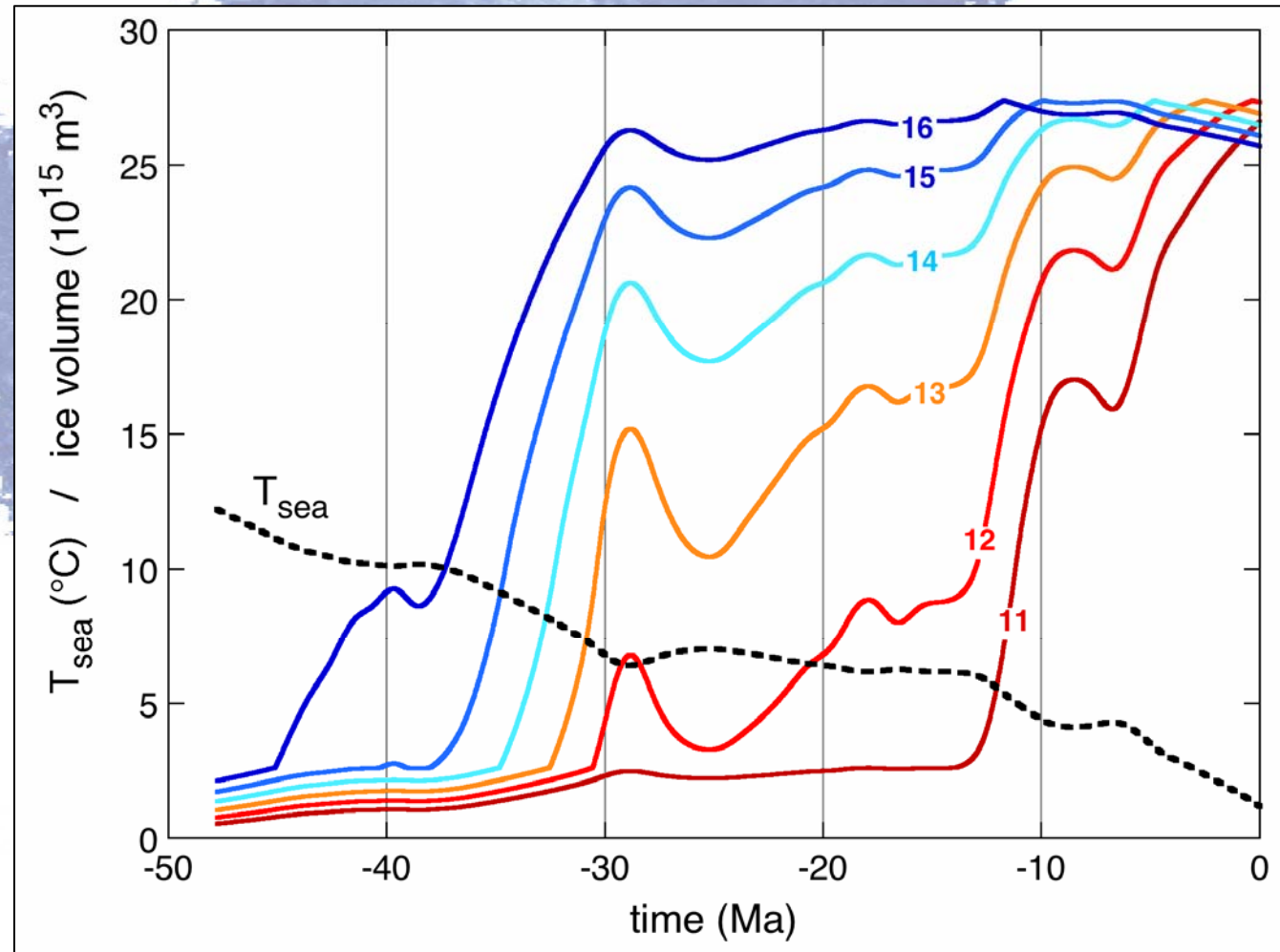
Cenozoic history

Forcing functions, Miller et al. (1987) and Zachos et al. (2001) deep sea temperatures (converted to Antarctic air temperatures) and Lear et al. (2000) sea surface temperatures.



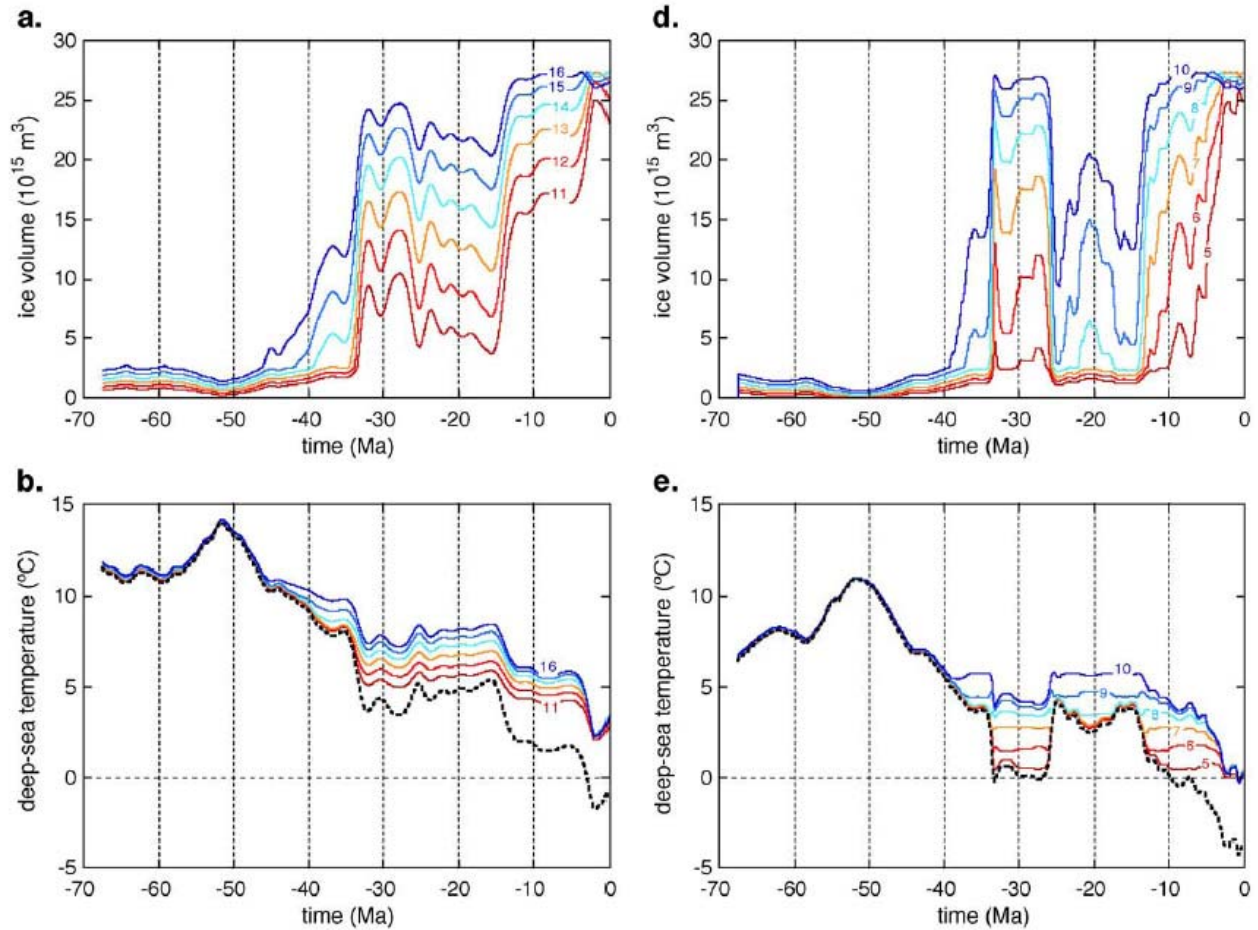
Cenozoic history

Strongly non-linear response of the Antarctic ice sheet for declining sea surface temperatures for different sensitivities



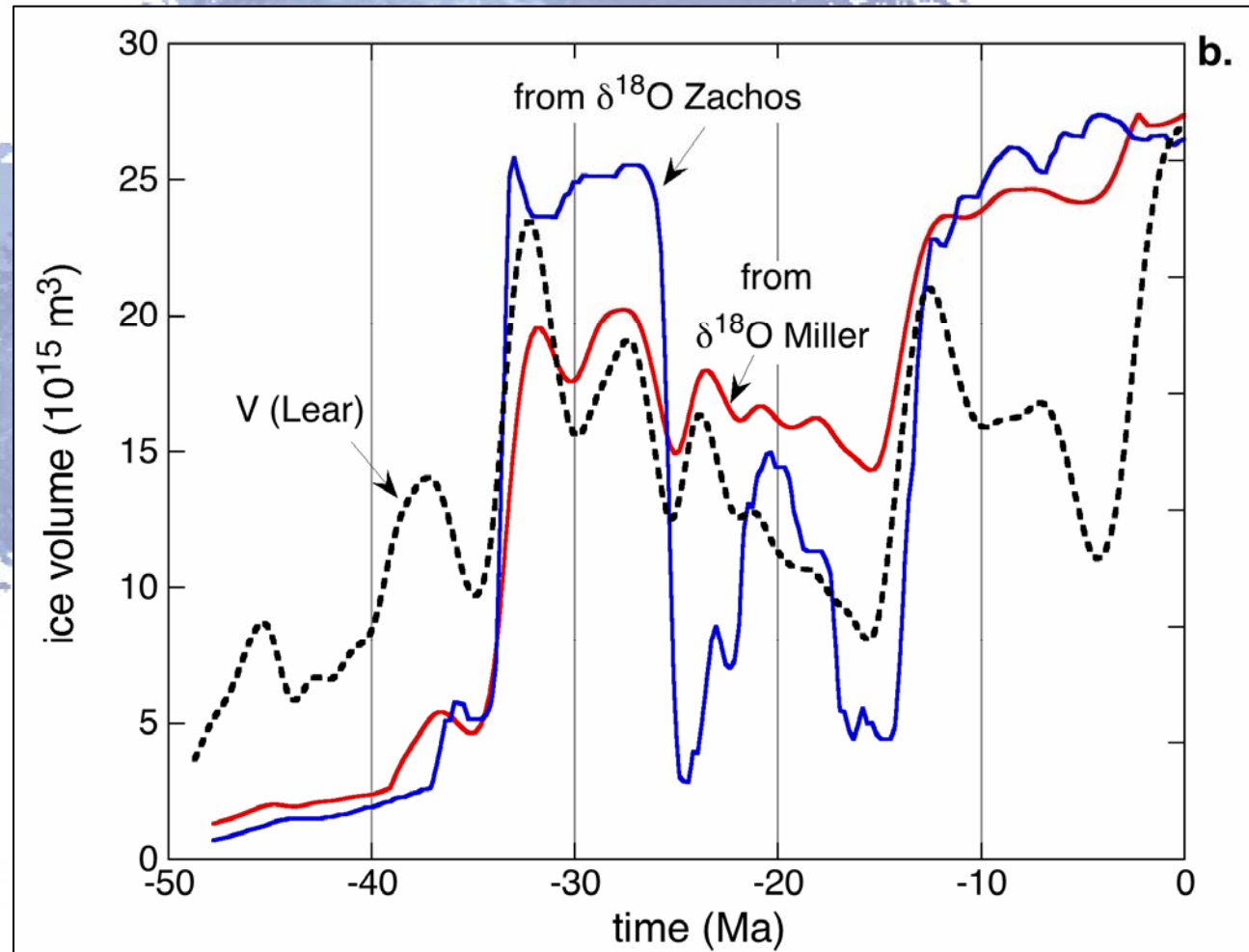
Cenozoic history

Strongly non-linear response of the Antarctic ice sheet using the Miller et al. 1987 forcing (a-b) or the Zachos et al. 2001 forcing (c-d).



Cenozoic history

Comparing Miller and Zachos best runs with Lear. Ice growth around 34 Ma for all 3 records and second/third pulses around 15 Ma and 3 Ma. Zachos decline in Miocene, Lear in early Pliocene.



Cenozoic history of Antarctica: Conclusions

Most important phases:

- Glaciation history spans the last 55 million years
- EAIS first formed ~34 million years ago (Ma); NLR
- Opening of Drake passage (?30-22 Ma)
-> resulted in gradual thermal isolation, which implies that its climate became less and less sensitive to extra-Antarctic forcing
- Formation of a cold-based EAIS of current proportions first after 15 Ma
- Inception of bipolar glaciation around 3 Ma affected Antarctica only through sea level falls

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