## 1988 <br> ANNUAL <br> TROPICAL CYCLONE REPORT



## JOINT TYPHOON WARNING CENTER GUAM, MARIANA ISLANDS

FRONT COVER: Contouring of visual satellite imagery enhances the cloud topography of Typhoon Hal (14W) with its large ragged eye and (at the right edge of the picture) the comma shape of Tropical Storm Irma (15W). These two tropical cyclones were active with Tropical Storm Jeff (16W) and Typhoon Uleki (01C) during the second week of September.
U. S. NAVAL OCEANOGRAPHY COMMAND CENTER JOINT TYPHOON WARNING CENTER COMNAVMARIANAS BOX 17 FPO SAN FRANCISCO, CA 96630

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DEFENSE TECHNICAL INFORMATION CENTER

The Annual Tropical Cyclone Report is prepared by the staff of the Joint Typhoon Warning Center (JTWC), a combined Air Force/Navy organization operating under the command of the Commanding Officer, U.S. Naval Oceanography Command Center/Joint Typhoon Warning Center, Guam. JTWC was established in April 1959 when USCINCPAC directed USCINCPACFLT to provide a single tropical cyclone waming center for the western North Pacific region. The operations of JTWC are guided by CINCPACINST 3140.1S.

The mission of the Joint Typhoon Waming Center is multi-faceted and includes:

1. Continuous monitoring of all tropical weather activity in the northern and southern hemispheres, from 180 degrees longitude westward to the east coast of Africa, and the prompt issuance of appropriate advisories and alerts when tropical cyclone development is anticipated.
2. Issuing warnings on all significant tropical cyclones in the above area of responsibility.
3. Determination of reconnaissance requirements for tropical cyclone surveillance and assignment of appropriate priorities.
4. Post-storm analysis of all significant tropical cyclones occurring within the westem North Pacific and North Indian Oceans, which includes an in-depth analysis of tropical cyclones of note and all typhoons.
5. Cooperation with the Naval Environmental Prediction Research Facility (NEPRF), Monterey, California, on the operational evaluation of tropical cyclone models and forecast aids, and the development of new techniques to support operational forecast scenarios.

Satellite imagery used throughout this report represents data obtained by the DMSP network. The personnel of Detachment 1, 1WW, collocated with JTWC at Nimitz Hill, Guam, coordinate the satellite acquisitions and tropical cyclone reconnaissance with the following units:

## Det 4, 20WS, Hickam AFB, Hawaii

Det 5, 20WS, Clark AB, Republic of the Philippines
Det 8, 20WS, Kadena AB, Japan
Det 15, 30WS, Osan AB, Korea
Air Force Global Weather Central, Offutt AFB, Nebraska

In addition, the Naval Oceanography Command Detachment, Diego Garcia, and Defense Meteorological Satellite Program (DMSP) equipped U.S. Navy ships have been instrumental in providing vital fixes of tropical cyclones in the Indian Ocean from satellite data.

Should JTWC become incapacitated, the Alternate Joint Typhoon Warning Center (AJTWC) located at the U.S. Naval Western Oceanography Center, Pearl Harbor, Hawaii, assumes warning responsibilities. Assistance in determining satellite reconnaissance requirements, and in obtaining the resultant data, is provided by Det 4, 20WS Hickam AFB, Hawaii.

Special thanks to: Navy Captain Carl W. Hoffman for his significant contributions and support; the men and women of the 27th Communications Squadron, Operating Location Charlie and the Operations department of the Naval Oceanography Command Center, Guam for their continuing support by providing high quality real-time satellite imagery; Marine Corps Air Station, Futenma, Japan for sharing their satellite imagery of STY Nelson (20W); the Pacific Fleet Audio-Visual Center, Guam for their assistance in the reproduction of satellite data for this report; to the Navy Publications and Printing Service Branch Office, Guam; the Royal Observatory Hong Kong for supporting synoptic data on TS Mamie (19W) and TY Ruby (23W); Dr. Bob Abbey and the Office of Naval Research for their technical support to this publication and support to the University of Hawaii for the Post Doctorate Fellow at JTWC; Dr. Mark Lander for his training efforts, suggestions and detailed work on CSUM; Dr. Greg Holland for his work with Major Joel Martin to develop a more representative periphery wind technique; the Australian Meteorological Service for the coastal radar reports on tropical cyclones via the AFTN to JTWC; Mr. Andy Chun and the CPHC for the data package on TY Uleki ( 01 C ); the Royal Navy and Royal Observatory Hong Kong for making possible the liaison visit by Lt. Cmdr. David Gray; National Weather Service Pacific Region for the startup of 24 -hour operations at Ponape, Truk, Majuro and Koror, the retrieval of the HANDAR observations from their computer in Hawaii for TY Roy ( 01 W ), and Mr. Tom Yoshida for the development of the Tropical Cyclone Reporting Form for post-storm support to JTWC; First Weather Wing for expediting the delivery of the Alden receiver for PACDIGS; and the Federated States of Micronesia for endorsing the AMOS installation on Faraulep Atoll.

Note: Appendix IV contains information on how to obtain past issues of the Annual Tropical Cyclone Report (titled Annual Typhoon Report prior to 1980).

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## CONTRACTIONS

| ABIO | Significant Tropical <br> Weather Advisory for the <br> Indian Ocean | CLIPER | Climatology and Persistence <br> Technique | HPAC | Mean of XTRP and CLIM <br> Techniques (Half Persis- <br> tence and Climatology) |  |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| ABPW | Significant Tropical <br> Weather Advisory for the <br> Western Pacific Ocean | CNOC | Centimeter <br> Commander Naval <br> Oceanography Command | ICAO | HR(S) | Hour(s) |
| ACFT | Aircraft | COSMOS | Cyclops Objective Steering <br> Model Output Statistics | INTT | Inganization Civil Aviation |  |



Intentionally left blank.

## CHAPTER I - OPERATIONAL PROCEDURES

## 1. GENERAL

The Joint Typhoon Warning Center (JTWC) provides a variety of routine services to the organizations within its area of responsibility, including:
a. SIGNIFICANT TROPICAL WEATHER ADVISORIES

Issued daily, these products describe all tropical disturbances and assess their potential for further development during the advisory period.

## b. TROPICAL CYCLONE FORMATION ALERTS

Issued when synoptic or satellite data indicate development of a significant tropical cyclone, in a specified area, is likely.

## c. TROPICAL CYCLONE WARNINGS

Issued periodically throughout each day, for significant tropical cyclones, giving forecasts of position and intensity of the system.

## d. PROGNOSTIC REASONING MESSAGES

Issued with each warning for tropical depressions, tropical storms, typhoons and super typhoons in the western North Pacific; these messages discuss the rationale behind the most recent JTWC warnings.

JTWC's customers determine the content of JTWC's products according to their changing requirements. Therefore, the spectrum of routine services is subject to change from year to year. Such changes are usually the result of deliberations held at the Annual Tropical Cyclone Conference.

## 2. DATA SOURCES

## a. COMPUTER PRODUCTS

A standard array of numerical and statistical guidance are available from the USN Fleet Numerical Oceanography Center (FNOC) at Monterey, California. FNOC products are received through the Naval Environmental Data Network (NEDN), the Naval Environmental Satellite Network (NESN) and by microcomputers connected to mainframe computers via military and commercial telephone lines.

## b. CONVENTIONAL DATA

This data set is comprised of land-based, ship surface and upper-air observations recorded at, or within six hours of, synoptic times. It incorporates cloud-motion winds derived twice a day from satellite imagery and commercial and military Aircraft Reports (AIREPS) of enroute meteorological observations, within six hours of synoptic times. There has been an effort to increase the frequency and use of AIREPS to describe the synoptic situation in otherwise data sparse regions. Additional conventional data sources include three Automated Meteorological Observing Station (AMOS) sites on the islands of Saipan and Rota in the Mariana Islands, and Faraulep in the Caroline Islands. The conventional data is hand plotted and analyzed in the tropics for the surface/gradient and 200 mb levels. These analyses are prepared twice daily from 0000Z and 1200 Z synoptic data. Also, FNOC supplies JTWC with computer streamline analyses and prognoses at the $925 \mathrm{mb}, 850 \mathrm{mb}, 700 \mathrm{mb}, 500$ $\mathrm{mb}, 400 \mathrm{mb}, 250 \mathrm{mb}$ and 200 mb levels from 0000 Z and 1200 Z synoptic data.

## c. AIRCRAFT RECONNAISSANCE

Aircraft of opportunity are a valuable for providing meteorological data at flight level around the periphery of tropical cyclones and in
describing the environment away from tropical cyclones that frequently affect tropical cyclone motion. With their airborne radar, they can remotely sense the inner rainbands and core of the tropical cyclone. Flight safety considerations preclude the use of transient aircraft for tropical cyclone penetration.

## d. SATELLITE RECONNAISSANCE

Meteorological satellite imagery recorded at USAF/USN ground sites and USN ships supply day and night coverage in JTWC's area of responsibility. Interpretation of these satellite data provide tropical cyclone positions and estimates of current and forecast intensities (Dvorak, 1984).

## e. RADAR RECONNAISSANCE

During 1988, as in previous years, landbased radar coverage was utilized extensively, when available. Once a tropical cyclone moved within the range of land-based radar sites, their reports were essential for determination of small-scale movement. Use of radar reports during 1988 is discussed in Chapter II.

## f. DRIFTING METEOROLOGICAL BUOYS

In July 1988, nine drifting meteorological buoys were deployed in the western North Pacific to increase the availability of synoptic data, in an often data sparse region. The Naval Oceanography Command provided the funds for procurement and deployment of the buoys and data acquisition equipment. The Naval Oceanographic Office, and JTWC, planned and coordinated the buoy deployments. The 21st Tactical Airlift Squadron at Clark Air Base in the Republic of the Philippines deployed the buoys from C-130 aircraft. All of the buoys provided sea-level pressure, air and sea surface temperatures. Six of the Tropical

Ocean Global Atmosphere (TOGA) buoys were retrofitted to provide wind speed and direction. Five buoys were deployed at seven-degree longitude intervals along three degrees North latitude, just south of the Caroline Islands. One of these buoys failed to activate on deployment. (Ironically, a buoy from the 1987 deployment remained active, trapped in an atoll near the failed buoy.) Three buoys were deployed in the Philippine Sea, and the remaining buoy was deployed east of Guam.

JTWC acquired the drifting buoy data directly through its Local User Terminal (LUT). The buoys transmit data to the Tiros-N polar orbiting satellites, which in turn store and relay the data to JTWC's LUT. Four to six observations per day are available from each buoy via direct readout. The stored data is dumped to Service ARGOS in Washington, D.C., where it is passed to the National Data Buoy Center (NDBC) for quality control. JTWC receives the data from NDBC via the Automated Weather Network (AWN). The buoys' positioning systems, wind speed, air pressure and temperature sensors provided accurate data. However, the wind direction data on the modified TOGA's were unreliable. JTWC's internal quality control program computed coarse corrections in an attempt to salvage wind directions, but the results were only of marginal use for the manual surface/gradient level streamline analyses.

The buoys furnished data on Super Typhoon Nelson (20W), Typhoons Odessa (21W), Pat (22W) , Ruby (23W), Skip (24W), Tess (25W), and Tropical Storms Lee (18W) and Val (26W). The tracks of the drifting buoys can be seen in Figure 1-1. Three buoys ceased operating after washing ashore in the Philippine Islands during October, November and December. By the end of 1988, five of nine buoys from the 1988 deployment and one buoy from the 1987 deployment remained operational.


## 3. COMMUNICATIONS

a. AUTOMATED DIGITAL NETWORK (AUTODIN)

AUTODIN is used for dissemination of warnings, alerts and other related bulletins to Department of Defense installations. These messages are relayed for further transmission over Navy Fleet Broadcasts, and Coast Guard CW (continuous wave Morse Code) and voice broadcasts. Inbound message traffic for JTWC is received via AUTODIN addressed to NAVOCEANCOMCEN GQ or DET 1, 1WW NIMITZ HILL GQ.

## b. AUTOMATED WEATHER NETWORK (AWN)

The AWN provides weather data over the Pacific Meteorological Data System (PACMEDS). Operational for JTWC in April 1988, the PACMEDS allows Pacific-Theater agencies to receive weather information at 1200 baud, which is an upgrade from the older 75 baud teletype systems, eliminating data backlogs. The system provides the large volume of meteorological reports necessary to satisfy JTWC requirements. Weather bulletins prepared by JTWC are inserted into the AWN circuit through the Nimitz Hill Naval Telecommunications Center (NTCC), which is controlled by the Naval Communications Area Master Station (NAVCAMS) Western Pacific located on Guam.

## c. AUTOMATIC VOICE NETWORK (AUTOVON)

AUTOVON is a world-wide general purpose switched voice network for the Department of Defense. The network provides a rapid and vital voice link for JTWC to communicate tropical cyclone information to customers. The AUTOVON telephone number for JTWC is 344-4224.

## d. NAVAL ENVIRONMENTAL DATA NETWORK (NEDN)

The NEDN continues to provide processed and raw environmental data from FNOC to JTWC, and is a communication line for requesting and receiving forecast aids via FNOC's mainframe computers.

## e. TIME-SHARING NETWORK (TYMNET)

The use of TYMNET through the Automated Tropical Cyclone Forecast (ATCF) system started in 1987 and aided JTWC's shift away from exclusive dependence on the Naval Environmental Display Station (NEDS) for processing and transmission of tropical cyclone forecast aids, and for receiving environmental fields and raw data. The use of the ATCF microcomputers has improved both the speed of handling and the quality control of these data.

## f. DEFENSE DATA NETWORK (DDN)

The DDN is a Department of Defense network of communication lines/links to exchange data files. Because the DDN has links, or gateways, to non-military information networks, it has demonstrated an excellent potential for interfacing with the research community. It was used routinely to transmit SSM/I data from AFGWC to JTWC.

## g. TELEPHONE FACSIMILE (TELEFAX)

TELEFAX provides the capability to rapidly scan and transmit, or receive, documents over commercial telephone lines or AUTOVON. The TELEFAX is used to disseminate tropical cyclone advisories and warnings to key agencies on Guam and in special situations the other Micronesian Islands. Inbound documents for JTWC are received via commercial telephone at (671) 477-6186. If inbound through AUTOVON, the Guam AUTOVON
operator 322-1110 can transfer the call to the commercial number 477-6186.

## h. NAVAL ENVIRONMENTAL SATELLITE NETWORK (NESN)

The NESN's primary function is to pass satellite data between FNOC and regional centers. It can provide a limited back-up for the NEDN.

## 4. DATA DISPLAYS

a. NAVAL ENVIRONMENTAL DISPLAY STATION (NEDS)

The NEDS is the mainstay for producing displays and hard copies of FNOC environmental products. However, it now serves as a backup for the transmission and receipt of FNOC's objective forecast aids and JTWC's weather messages.

## b. AUTOMATED TROPICAL CYCLONE FORECAST SYSTEM (ATCF)

Increased usage of microcomputers in the ATCF has shortened the processing time and improved the quality control of weather bulletins; especially the warnings. The ATCF is still a step away from direct interface with NTCC for AWN and AUTODIN message transmissions, but the Joint Typhoon Warning Center Automation Project (JTWC-AP) will make this needed interface a reality.

## c. PACIFIC DIGITAL INFORMATION GRAPHICS SYSTEM (PACDIGS)

The PACDIGS is a new communications circuit that was expanded to include JTWC in 1988. Air Force Global Weather Central (AFGWC) at Omaha, Nebraska provides a standard set of numerical products to the PACDIGS circuit.

## d. NAVAL SATELLITE DISPLAY SYSTEM (NSDS)

The NSDS functions primarily for display of FNOC stored satellite imagery and can provide limited back up for the NEDN, via the NESN.

## 5. ANALYSES

A composite surface/gradient-level ( 3000 ft ( 914 m )) manual analysis of the JTWC area of responsibility is accomplished daily on the 0000 Z and 1200 Z conventional data. Analysis of the wind field using streamlines is stressed for tropical and subtropical regions. Analysis of the pressure field outside the tropics is accomplished routinely by the Naval Oceanography Command Center Operations watch team and is used by JTWC in conjunction with their analysis of the tropical wind fields.

A composite upper-tropospheric manual streamline analysis is accomplished daily at 0000 Z and 1200 Z , utilizing rawinsonde data from 300 mb through 100 mb , winds obtained from satellite-derived cloud motion analysis, and AIREPS (taken plus or minus three hours of chart valid time) at or above $31,000 \mathrm{ft}(9,449$ $\mathrm{m})$. Wind and height data are used to generate a representative analysis of tropical cyclone outflow patterns, mid-latitude steering currents, and features that may influence tropical cyclone intensity. All charts are hand-plotted in the tropics to provide all available data as soon as possible to the Typhoon Duty Officer (TDO). These charts are augmented by computerplotted charts for the final analysis.

Computer analyses for the 925,850 , $700,500,400,250$ and 200 mb levels are available from the 0000 Z and 1200 Z data base. Additional sectional charts at intermediate synoptic times and auxiliary charts, such as station-time plot diagrams and pressure-change charts, are also analyzed during periods of significant tropical cyclone activity.

A Hovmöller Trough-and-Ridge Diagram for 500 mb at $40^{\circ}$ North and $30^{\circ}$ South latitudes for the entire hemisphere is produced daily to provide a quick look at trough and ridge progression with time.

## 6. FORECAST AIDS

The following objective tropical cyclone forecasting techniques were employed during 1988 (a description of each technique is presented in Chapter V):
a. MOVEMENT
(1) EXTRAPOLATION
(2) CLIMATOLOGY
(3) HPAC (Half Persistence - Half Climatology Blend)
(4) CLIPER (Climatology and Persistence Technique)
(5) COSMOS (Model Output Statistics)
(6) CSUM (Statistical Dynamic Model)
(7) OTCM (Dynamic Model)
(8) TAPT (Empirical)
(9) TYAN78 (Analog)
b. INTENSITY
(1) CLIMATOLOGY
(2) DVORAK (Empirical)
(3) HOLLAND/MARTIN
(Empirical)

## 7. FORECAST PROCEDURES

## a. INITIAL POSITIONING

The warning position is the best estimate of the center of the surface circulation at synoptic time. It is estimated from an analysis of all fix information received up to one and one-half hours after synoptic time. This analysis is based on a semi-objective weighting of fix information based on the historical accuracy of the fix platform and the meteorological features used for the fix. The interpolated warning position reduces the weighting of any single fix and results in a more consistent movement and a warning position that is more representative of the larger-scale circulation. If the fix data are not available due to reconnaissance platform malfunction or communication problems, synoptic data or extrapolation from previous fixes are used.

## b. TRACK FORECASTING

A preliminary forecast track is developed based on an evaluation of the rationale behind the previous warning and the guidance given by the most recent set of objective techniques and numerical prognoses. This preliminary track is then subjectively modified based on the following considerations:
(1) The prospects for recurvature or erratic movement are evaluated. This determination is based primarily on the present and forecast positions and amplitudes of the middle-tropospheric, mid-latitude troughs and ridges as depicted on the latest upper-air analysis and numerical forecasts.
(2) Determination of the best steering level is partly influenced by the maturity and vertical extent of the tropical cyclone. For mature tropical cyclones located south of the subtropical ridge axis, forecast changes in speed of movement are closely correlated with anticipated changes in the intensity or relative position of the ridge. When steering currents are relatively weak, the tendency for tropical
cyclones to move northward due to internal forces is an important consideration.
(3) Over the 12 - to 72 -hour (12- to 48-hour in the southern hemisphere) forecast period, speed of movement during the early forecast period is usually biased towards persistence, while the later forecast periods are biased towards objective techniques. When a tropical cyclone moves poleward, and toward the mid-latitude steering currents, speed of movement becomes increasingly more biased toward a selective group of objective techniques capable of estimating acceleration.
(4) The proximity of the tropical cyclone to other tropical cyclones is closely evaluated to determine if there is a possibility of interaction.

A final check is made against climatology to determine whether the forecast track is reasonable. If the forecast deviates greatly from one of the climatological tracks, the forecast rationale will be reappraised.

## c. INTENSITY FORECASTING

Heavy reliance is placed on the empirically derived Dvorak (1984) model for forecasting tropical cyclone intensity. Other techniques used for forecasting intensity are extrapolation of synoptic wind and pressure data and climatology. An evaluation of the entire synoptic situation is made, including the location of major troughs and ridges, the position and intensity of any nearby Tropical Upper-Tropospheric Troughs (TUTTs), the vertical and horizontal extent of the tropical cyclone's circulation and the extent of the associated upper-level outflow pattern. Each intensity forecast is affected by the accompanying forecast track and environmental influences along that track; such as, terrain, vertical wind shear and extratropical weather features.

## d. WIND-RADII FORECASTING

A new wind profile and steering diagnostic is being used at JTWC. The technique is the result of efforts by Dr. G. J. Holland (Office of Naval Research contractor) and Maj. J. Martin. The technique adapts an earlier work (Holland, 1980) and specifically addresses the need for realistic $30-, 50$ - and $100-\mathrm{kt}$ wind radii around tropical cyclones. It solves equations for basic gradient wind relations within the tropical cyclone area, using input parameters obtained from enhanced infrared satellite imagery. For the first time, diagnoses also include asymmetric areas of winds caused by tropical cyclone movement. Size and intensity parameters are also used to diagnose internal steering components of tropical cyclone motion known collectively as "Beta-drift". The Holland/Martin wind radii technique replaces the more general Huntley (1980) technique.

## 8. WARNINGS

Tropical cyclone warnings are issued when a closed circulation is evident and maximum sustained winds are forecast to increase to 34 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) within 48 -hours, or if the tropical cyclone is in such a position that life or property may be endangered within 72hours. Warnings may also be issued in other situations if it is determined that there is a need to alert military or civil interests to threatening tropical weather conditions.

Each tropical cyclone warning is numbered sequentially and includes the following information: the position of the surface center; estimate of the position accuracy and the supporting reconnaissance (fix) platforms; the direction and speed of movement during the past six hours (past 12 -hours in the southern hemisphere); the intensity and radial extent of over $30-, 50$-, and $100-\mathrm{knot}$ surface winds, when applicable. At forecast intervals of $12-, 24-, 48-$, and 72 -hours (12-, 24 -, and $48-$ hours in the southern hemisphere), information on the tropical cyclone's anticipated position, intensity and wind radii are also provided. Vectors indicating the mean direction and mean
speed between forecast positions are also included in all warnings.

Warnings in the western North Pacific and North Indian Oceans are issued every six hours valid at standard times: 0000Z, 0600Z, 1200 Z and 1800 Z (every 12-hours: 0000Z, 1200 Z or $0600 \mathrm{Z}, 1800 \mathrm{Z}$ in the southern hemisphere). All warnings are released to the communications network no earlier than synoptic time and no later than synoptic time plus two and one-half hours, so that recipients will have a reasonable expectation of having all warnings "in hand" by synoptic time plus three hours (0300Z, 0900Z, 1500 Z and 2100 Z ).

Warning forecast positions are later verified against the corresponding "best track" positions (obtained during detailed post-storm analyses to determine the most probable path and intensity of the cyclone). A summary of the verification results for 1988 is presented in Chapter V.

## 9. PROGNOSTIC REASONING MESSAGES

This plain language message is intended to provide meteorologists with the reasoning behind the latest forecast. For tropical depressions, tropical storms, typhoons and super typhoons in the western North Pacific Ocean, prognostic reasoning messages are transmitted following each warning. This is a change from 1987, when prognostic reasoning messages for western North Pacific tropical storms, typhoons and super typhoons were transmitted after the 0000 Z and 1200 Z warnings, or whenever the previous forecast reasoning was no longer valid.

In addition to this message, prognostic reasoning information, applicable to all customers, is provided in the remarks section of warnings when significant forecast changes are made or when deemed appropriate by the TDO.

## 10. TROPICAL CYCLONE FORMATION ALERTS

Tropical Cyclone Formation Alerts (TCFAs) are issued whenever interpretation of satellite imagery and other meteorological data indicate that the formation of a significant tropical cyclone is likely. These Alerts will specify a valid period not to exceed twenty-four hours and must either be cancelled, reissued, or superseded by a tropical cyclone warning prior to the expiration of the valid time.

## 11. SIGNIFICANT TROPICAL WEATHER ADVISORIES

This product contains a general, nontechnical description of all tropical disturbances in JTWC's area of responsibility (AOR) and an assessment of their potential for further (tropical cyclone) development. In addition, all tropical cyclones in warning status are briefly discussed. Two separate messages are issued daily and are valid for a 24 -hour period. The Significant Tropical Weather Advisory for the western Pacific Ocean (ABPW PGTW) covers the area east of $100^{\circ}$ East longitude to the dateline and is issued by 0600Z. The Significant Tropical Weather Advisory for the Indian Ocean (ABIO PGTW) covers the area west of $100^{\circ}$ East longitude to the coast of Africa and is issued by 1800 Z . These are reissued whenever the situation warrants. For each suspect area, the words "poor", "fair", or "good" are used to describe the potential for development. "Poor" will be used to describe a tropical disturbance in which meteorological conditions are currently unfavorable for development; "fair" will be used to describe a tropical disturbance in which the meteorological conditions are favorable for development but significant development has not commenced; and "good" will be used to describe the potential for development of a tropical disturbance covered by a Tropical Cyclone Formation Alert.

## CHAPTER II - RECONNAISSANCE AND FIXES

## 1. GENERAL

The Joint Typhoon Warning Center depends on reconnaissance to provide necessary, accurate, and timely meteorological information in support of advisories, alerts and warnings. JTWC relies primarily on three reconnaissance platforms: aircraft, satellite, and radar. In data rich areas, synoptic data are also used to supplement the above. Optimum use of all available reconnaissance resources is obtained through the Selective Reconnaissance Program (SRP); various factors are considered in selecting a specific reconnaissance platform including capabilities and limitations, and the tropical cyclone's threat to life and property both afloat and ashore. A summary of reconnaissance fixes received during 1988 is included in Section 6 of this chapter.

## 2. RECONNAISSANCE AVAILABILITY

## a. AIRCRAFT

Due to budgetary constraints, 1987 was the final year for dedicated aircraft weather reconnaissance in the western North Pacific. The thrust in 1988 was to increase the frequency and reliability of commercial/military airways meteorological reports, thus enhancing synoptic analysis, particularly in data sparse regions.

Limited aircraft of opportunity were available in the western North Pacific, in 1988, for use as synoptic track missions. Aircraft of opportunity can provide direct measurements of standard pressure-level heights, temperature and flight-level wind data. These data, plus the use of airborne radar, can provide the forecaster vital information on changing tropical cyclone characteristics.
b. SATELLITE

Satellite fixes from Air Force/Navy ground sites and Navy ships provide day and night coverage in JTWC's area of responsibility. Interpretation of this satellite imagery provides tropical cyclone positions and estimates of current and forecast intensities through the Dvorak technique.

## c. RADAR

Land-based radar remotely senses and maps precipitation within tropical cyclones in the proximity (usually within $175 \mathrm{~nm}(324 \mathrm{~km})$ ) of radar sites in the Republic of the Philippines, Taiwan, Hong Kong, Japan, South Korea, Kwajalein and Guam. In 1987 the USAF upgraded the radars at Yongsan AB,' South Korea; Yokota AB, Japan; Kadena AB, Okinawa, Japan; and Andersen AFB, Guam. (The upgrade included increased range, continuous clockwise or counterclockwise scan, a range height indicator to an altitude of 21 km ( 13 nm ) in $1 \mathrm{~km}(0.6 \mathrm{~nm}$ ) intervals, a digital video integrator/processor, range normalization, a color enhanced digital remote scope and local area/operations area mapping program.) These new radars are a welcome improvement to the existing network. The next upgrade will be the arrival of the first next generation Doppler radars in the early 1990's.

## d. SYNOPTIC

JTWC also determines tropical cyclone positions based on the analysis of the surface/gradient-level synoptic data. These positions were helpful in situations where the vertical structure of the tropical cyclone was weak or accurate surface positions from aircraft or satellite were not available.

## 3. AIRCRAFT RECONNAISSANCE SUMMARY

There were no vortex fix or investigative missions flown into western North Pacific
tropical cyclones in 1988. A synoptic track and airborne radar fix were provided on Typhoon Roy (01W) in January by aircraft of opportunity. These data described the mid-level steering flow and center location.

## 4. SATELLITE RECONNAISSANCE SUMMARY

The USAF provides satellite reconnaissance support to JTWC through the DMSP Tropical Cyclone Reporting Network (DMSP Network), which consists of tactical sites and a centralized facility. Tactical DMSP sites monitoring DMSP, NOAA and geostationary satellite data are located at Nimitz Hill, Guam; Clark AB, Republic of the Philippines; Kadena AB, Okinawa, Japan; Osan AB, Republic of Korea; and Hickam AFB, Hawaii. These sites provide a combined coverage that includes most of JTWC's area of responsibility in the western North Pacific, from near the dateline westward to the Malay Peninsula. For the remainder of its AOR, JTWC relies on the AFGWC to provide coverage using stored satellite data. The Naval Oceanography Command Detachment, Diego Garcia, furnishes interpretation of NOAA polar orbiting coverage in the central Indian Ocean and USN ships equipped for direct satellite readout contribute supplementary support.

AFGWC, located at Offutt AFB, Nebraska, is the centralized member of the DMSP network. In support of JTWC, AFGWC processes stored imagery from DMSP and NOAA spacecraft. Stored imagery is recorded onboard the spacecraft as they pass over the earth and later down-linked to AFGWC via a network of command readout sites and communication satellites. This enables AFGWC to obtain the coverage necessary to fix all tropical cyclones within JTWC's AOR. AFGWC has the primary responsibility to provide tropical cyclone reconnaissance over the entire Indian Ocean, southwest Pacific, and the area near the dateline in the western North Pacific Ocean. Additionally, AFGWC can be tasked to provide tropical cyclone support in the entire western North Pacific as backup to
coverage routinely available in that region.
The hub of the DMSP network is Detachment 1, First Weather Wing (Det 1, 1WW), colocated with JTWC at Nimitz Hill, Guam. Based on available satellite coverage, Det 1, 1WW is responsible for coordinating satellite reconnaissance requirements with JTWC and tasking the individual network sites for the necessary tropical cyclone fixes, current intensity estimates and forecast intensities. When a particular satellite pass is selected to support the development of JTWC's next tropical cyclone warning, two sites are tasked to fix the tropical cyclone from the same pass. This "dual-site" concept provides the necessary redundancy to virtually guarantee JTWC a satellite fix on the tropical cyclone.

The network provides JTWC with several products and services. The main service is one of monitoring the AOR for indications of tropical cyclone development. If an area exhibits potential for development, JTWC is notified. Once JTWC issues either a Tropical Cyclone Formation Alert or warning, the network is tasked to provide three products: tropical cyclone positions, current intensity estimates and forecast intensities. Each satellite tropical cyclone position is assigned a Position Code Number (PCN) to indicate the accuracy of the fix position. The PCN is determined by the availability of visible landmarks in the image that can be used as references for precise gridding and the degree of organization of the tropical cyclone's cloud system (Table 2-1).

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Figure 2-1. Dvorak code for communicating estimates of current and forecast intensity derived from satellite data. In the example, the current "T-number" is 3.5 , but the current intensity is 4.5 . The cloud system has weakened by 1.5 "T-numbers" since the previous evaluation conducted 24 -hours earlier. The plus (+) symbol indicates an expected reversal of the weakening trend or very little further weakening of the tropical cyclone during the next 24 -hour period.

Starting in 1987, Detachment 1, First Weather Wing increased the number of estimates of the tropical cyclone's current intensity from two to four per day once a tropical cyclone formation alert or tropical cyclone warning was issued. Current intensity estimates and 24 -hour intensity forecasts are made using the Dvorak technique (NOAA Technical Report NESDIS 11) for both visual and enhanced infrared imagery (Figure 2-1).

Figure 2-2 shows the status of operational polar orbiting spacecraft. Two DMSP spacecraft were operational in 1988. The year began with one operational DMSP satellite, the 19543 (F8) spacecraft. After overheating forced a temporary shutdown on 3 December 1987, the Special (passive) Sensor, Microwave Imager (SSM/I) on the F8 spacecraft was reactivated in mid-January 1988. The 20542 (F9) DMSP satellite was launched 3 February as a replacement for the 18541 (F7) satellite, which failed 17 October 1987. The thermal channel used for intensity estimates began to degrade shortly after launch and was of marginal use at year's end. The NOAA 10 spacecraft performed

well throughout the year. The NOAA 11 was launched 24 September and replaced the aging NOAA 9 satellite on 8 November.

During 1988, data from the DMSP network was the primary input to warnings and best tracks in the western North Pacific. This increased emphasis on satellite data resulted in almost all the warnings being based on satellite reconnaissance.

The DMSP network provided JTWC with a total of 2,044 satellite fixes on 27 tropical cyclones in the western North Pacific Ocean. In addition, 117 fixes were made on tropical cyclones in the North Indian Ocean and 1144 in the southern hemisphere. A comparison of those fixes and their corresponding best tracks is shown in Tables 2-2A and 2-2B. (Note: Those fixes which were out of limits when compared with the best track are not included.) The network also provided an additional 224 fixes on tropical disturbances which did not develop into significant tropical cyclones. The standard relationship between tropical cyclone "T-number", maximum surface wind speed (Dvorak, 1984) and minimum sealevel pressure (Atkinson and Holliday, 1977) is outlined in Table 2-3.

TABLE 2-2A
MEAN DEVLATION (NM) CE ALL SAITHTITE DERIVED TROPICAL CYCLONE POSITIONS FRCM JTWC BRST TRACK POSITIONS IN TEIE HESTIERN NORTH PACTFIC AND NORTH INDIAN OCEANS (NUNBIMR OF CASES IN PARENIHESES)

WESTERN NORTH PACIFIC OCEAN
NORTH INDIAN OCEAN

| PCN | 1978-1987 AVERAGE |  | 1988 AVERAGE |  | 1980-1987 AVERAGE |  | 1988 AVERAGE |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1 | 14.2 | (1737) | 13.4 | (78) | 16.9 | (42) | 7.2 | (2) |
| 2 | 15.6 | (2582) | 13.7 | (337) | 16.9 | (9) | 11.3 | (4) |
| 3 | 21.2 | (2488) | 18.5 | (82) | 24.8 | (34) | 24.5 | (8) |
| 4 | 22.2 | (2047) | 16.7 | (281) | 45.1 | (21) | 25.2 | (5) |
| 5 | 37.5 | (4294) | 30.2 | (185) | 38.3 | (313) | 37.4 | (30) |
| 6 | 38.4 | (5023) | 30.9 | (973) | 40.5 | (417) | 38.7 | (64) |
| $1 \& 2$ | 15.0 | (4319) | 13.6 | (415) | 17.0 | (51) | 9.9 | (6) |
| 3\&4 | 21.7 | (4535) | 17.1 | (363) | 32.5 | (55) | 24.8 | (13) |
| 5\&6 | 38.0 | (9317) | 30.8 | (1158) | 39.6 | (730) | 38.3 | (94) |
| TOTAL | 28.4 | (18171) | 24.5 | (1936) | 37.8 | (836) | 35.2 | (113) |

TABLE 2-2B

PCN 1985-1987 AVERAGE 1988 AVERAGE

| 1 | 16.6 | $(82)$ | 15.2 | $(21)$ |
| :--- | ---: | ---: | ---: | ---: |
| 2 | 16.1 | $(442)$ | 17.9 | $(122)$ |
| 3 | 35.9 | $(112)$ | 25.4 | $(13)$ |
| 4 | 27.0 | $(408)$ | 27.1 | $(130)$ |
| 5 | 40.8 | $(474)$ | 39.2 | $(74)$ |
| 6 | 36.4 | $(2938)$ | 40.1 | $(713)$ |
| $1 \& 2$ | 16.2 | $(524)$ | 17.5 | $(143)$ |
| $3 \& 4$ | 28.9 | $(520)$ | 26.9 | $(143)$ |
| $5 \& 6$ | 37.0 | $(3412)$ | 40.0 | $(787)$ |
| TOTALS | 33.6 | $(4456)$ | 35.3 | $(1073)$ |


| TABLE 2-3 | TROPICAL CYCLONE INTENSITY NUMBER $\begin{aligned} & 0.0 \\ & 0.5 \\ & 1.0 \\ & 1.5 \\ & 2.0 \\ & 2.5 \\ & 3.0 \\ & 3.5 \\ & 4.0 \\ & 4.5 \\ & 5.0 \\ & 5.5 \\ & 6.0 \\ & 6.5 \\ & 7.0 \\ & 7.5 \\ & 8.0 \end{aligned}$ | WIND SPEED $<25$ 25 25 25 30 35 45 55 65 77 90 102 115 127 140 155 170 | (KI) <br> AND <br> D <br> IP) <br> MSLP <br> (NW PACIEIC) |
| :---: | :---: | :---: | :---: |

## 5. RADAR RECONNAISSANCE SUMMARY

Twelve of the twenty-seven significant tropical cyclones in the western North Pacific during 1988 passed within range of land-based radar with sufficient cloud pattern organization to be fixed. The land-based radar fixes that were obtained and transmitted to JTWC totaled 430. (Only one radar fix was obtained by aircraft of opportunity.)

The WMO radar code defines three categories of accuracy: good (within 10 km ( 5 nm ), fair (within $10-30 \mathrm{~km}(5-16 \mathrm{~nm}$ ), and poor (within $30-50 \mathrm{~km}(16-27 \mathrm{~nm}$ ). Of the 428 radar fixes encoded in this manner, 169 were good, 120 were fair, and 139 were poor. Compared to JTWC's best track, the mean vector deviation for land-based radar sites was $19 \mathrm{~nm}(35 \mathrm{~km})$. Excellent support from the radar network through timely and accurate radar fix positioning allowed JTWC to track and forecast tropical cyclone movement through even the most difficult erratic tracks.

The availability of data from radar sites in the Republic of Philippines was of concern. In 1988 these radar sites provided a valuable but limited number of reports on tropical cyclones. Reports were received from only two stations, in contrast to five in 1987. As in previous years, no radar reports were received on North Indian Ocean or southern hemisphere tropical cyclones.

## 6. TROPICAL CYCLONE FIX DATA

A total of 2,474 fixes on twenty-seven western North Pacific tropical cyclones and 117 fixes on five North Indian Ocean tropical cyclones were received at JTWC. Table 2-4A and Table 2-4B delineate the number of fixes per platform for each individual tropical cyclone for the western North Pacific and North Indian Oceans respectively. Season totals and percentages are also indicated. (Table 2-4C provides the same information for the South Pacific and South Indian Oceans.)

| TABLE $2-4 A$HESTHPN NORIP PACIFIC |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |
| WESTERN NORTH PACIEIC |  | SATEILITE | RADAR | TOTA * |
| TY SUSAN | (01W) | 155 | 60 | 215 |
|  | (02W) | 75 | 35 | 110 |
| TD 03W | (03W) | 29 | 0 | 29 |
| TY THAD | (04K) | 96 | 26 | 122 |
| TS VANESSA | (05W) | 76 | 8 | 84 |
| TY WARREN | (06\%) | 141 | 30 | 171 |
| TS AGNES | (07W) | 33 | 0 | 33 |
| TS BIILS | (08W) | 43 | 0 | 43 |
| TS CLARA | (09W) | 36 | 0 | 36 |
| TY DOYLE | (10W) | 96 | 0 | 96 |
| TS ELSIE | (11W) | 39 | 0 | 39 |
| TY FABIAN | (12W) | 81 | 0 | 81 |
| TS GAY | (13W) | 21 | 0 | 21 |
| TY HAL | (14W) | 110 | 18 | 128 |
| TY ULEKI | (01C) | 62 | 0 | 62 |
| TS IRMA | (15W) | 51 | 0 | 51 |
| TS JEFF | (16W) | 36 | 0 | 36 |
| TS KIT | (17W) | 51 | 9 | 60 |
| TS LEE | (18W) | 64 | 8 | 72 |
| TS MAMIE | (19W) | 48 | 0 | 48 |
| STY NELSON | (20W) | 125 | 184 | 309 |
| TY ODESSA | (21W) | 104 | 0 | 104 |
| TY PAT | (22W) | 81 | 9 | 90 |
| TY RUBY | (23W) | 140 | 42 | 182 |
| TY SKIP | (24W) | 133 | 0 | 133 |
| TY TESS | (25W) | 55 | 0 | 55 |
| TS VAL | (26W) | 63 | 1 | 64 |
| TOTALS |  | 2044 | 430 | 2474 |
| PERCENTAGE OF TOTALS |  | 83\% | 17\% | 100\% |
| * NO AIRCRAFT OR SYNOPTIC EIXES WERE RECEIVED |  |  |  |  |

TABLE 2-4B
NORITR INDIAN OCHPA FIX PIATFORM SUNGARY FOR 1988

TROPICAL CYCLONE SATEMTITE*

| TC 01A | 20 |
| :--- | :--- |
| TC 02B | 15 |
| TC 03B | 11 |
| TC 04B | 55 |
| TC 05B | 16 |

TOTAL NUMBER OF FIXES 117

* NO SYNOPTIC FIXES WERE RECEIVED


Intentionally left blank.

## CHAPTER III - SUMMARY OF WESTERN NORTH PACIFIC AND NORTH INDIAN OCEAN TROPICAL CYCLONES

## 1. GENERAL

During the calendar year 1988, JTWC issued warnings on 27 tropical cyclones in the western North Pacific - one super typhoon, 12 typhoons, 13 tropical storms and one tropical depression. This includes Typhoon Uleki ( 01 C ), which initially developed in the central North Pacific (Table 3-1). The total number of western North Pacific tropical cyclones is lower than the climatological mean of 30.7, and two above the 1987 total (Table 3-2). Five tropical cyclones - one of typhoon and four of tropical storm intensity - developed in the North Indian Ocean. This is average. The climatological mean is 4.7. During 1988, warnings were issued on a total of 32 northern hemisphere tropical cyclones. A chronology of western North Pacific and North Indian Ocean tropical cyclones is provided in Figure 3-1.

For the year, there were 114 "warning days" in the western North Pacific. A warning day is defined as a day during which JTWC issued warnings on at least one tropical cyclone. A "one-cyclone day" refers to a day when we were warning on only one tropical cyclone. A "two-cyclone day" refers to a day when we warned on two different tropical cyclones simultaneously. A "three-cyclone day" means JTWC was warning on three tropical cyclones at once. Considering only the western North Pacific, there were 15 two-cyclone days and four three-cyclone days (Table 3-3). When

North Indian Ocean tropical cyclones are included, there were 128 warning days of which 16 were two-cyclone days and four were threecyclone days. There were no four-cyclone or five-cyclone days. Thus, JTWC was in northern hemisphere warning status 35 percent of the year; we were in a multiple-cyclone situation (that is, warning on two or more tropical cyclones at the same time) for 20 days or six percent of the year.

JTWC issued 471 warnings on 27 western North Pacific tropical cyclones and 44 warnings on five North Indian Ocean tropical cyclones, for a grand total of 515 warnings. There were 33 initial Tropical Cyclone Formation Alerts issued for western North Pacific disturbances and four for the North Indian Ocean. Twenty-six western North Pacific and four North Indian Ocean tropical cyclones developed subsequent to the issuance of an Alert. Only one western North Pacific tropical cyclone-Tropical Storm Elsie (11W)regenerated, and an Alert was issued prior to regeneration. Typhoon Uleki (01C) was passed to JTWC from CPHC while in warning status, so JTWC did not issue an Alert (Table 3-4). In the western North Pacific, the false alarm rate was 21 percent and the mean lead time (to issuance of the first warning) was 8.5 -hours. For the North Indian Ocean, the false alarm rate was zero and the mean lead time was 5.1 hours. An Alert was not issued for Tropical Cyclone 02B.


TABLE 3-2 LEGEND


The criteria used in Table 3-2 are as follows:

1. If a tropical cyclone was first wamed on during the last two days of a particular month and continued into the next month for longer than two days, then that system was attributed to the second month.
2. If a tropical cyclone was warned on prior to the last two days of a month, it was attributed to the first month, regardless of how long the system lasted.
3. If a tropical cyclone began on the last day of the month and ended on the first day of the next month, that system was attributed to the first month. However, if a tropical cyclone began on the last day of the month and continued into the next month for only two days, then it was attributed to the second month.

| TABLE 3-2 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| WESTERN NORTH PACIFIC TROPICAI CYCLONE DISTRIBUTIO |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| YEAR | JAN | FEB | MAR | APR | MAY | JuN | JUL | AUG | SEP | $\propto$ | NOV | DEC | TOTALS |  |
| 1959 | 0 | 1 | 1 | 1 | 0 | 1 | 3 | 8 | 9 | 3 | 2 | 2 | 31 |  |
|  | 000 | 010 | 010 | 100 | 000 | 001 | 111 | 512 | 423 | 210 | 200 | 200 | 177 | 7 |
| 1960 | 1 | 0 | 1 | 1 | 1 | 3 | 3 | 9 | 5 | 4 | 1 | 1 | 30 |  |
|  | 001 | 000 | 001 | 100 | 010 | 210 | 210 | 810. | 041 | 400 | 100 | 100 | 198 | 3 |
| 1961 | 1 | 1 | 1 | 1 | 4 | 6 | 5 | 7 | 6 | 7 | 2 | 1 | 42 |  |
|  | 010 | 010 | 100 | 010 | 211 | 114 | 320 | 313 | 510 | 322 | 101 | 100 | 20111 | 11 |
| 1962 | 0 | 1 | 0 | 1 | 3 | 0 | 8 | 8 | 7 | 5 | 4 | 2 | 39 |  |
|  | 000 | 010 | 000 | 100 | 201 | 000 | 512 | 701 | 313 | 311 | 301 | 020 | 246 | 9 |
| 1963 | 0 | 0 | 1 | 1 | 0 | 4 | 5 | 4 | 4 | 6 | 0 | 3 | 28 |  |
|  | 000 | 000 | 001 | 100 | 000 | 310 | 311 | 301 | 220 | 510 | 000 | 210 | 196 | 3 |
| 1964 | 0 | 0 | 0 | 0 | 3 | 2 | 8 | 8 | 8 | 7 | 6 | 2 | 44 |  |
|  | 000 | 000 | 000 | 000 | 201 | 200 | 611 | 350 | 521 | 331 | 420 | 101 | 2613 | 5 |
| 1965 | 2 | 2 | 1 | 1 | 2 | 4 | 6 | 7 | 9 | 3 | 2 | 1 | 40 |  |
|  | 110 | 020 | 010 | 100 | 101 | 310 | 411 | 322 | 531 | 201 | 110 | 010 | 2113 | 6 |
| 1966 | 0 | 0 | 0 | 1 | 2 | 1 | 4 | 9 | 10 | 4 | 5 | 2 | 38 |  |
|  | 000 | 000 | 000 | 100 | 200 | 100 | 310 | 531 | 532 | 112 | 122 | 101 | 2010 | 8 |
| 1967 | 1 | 0 | 2 | 1 | 1 | 1 | 8 | 10 | 8 | 4 | 4 | 1 | 41 |  |
|  | 010 | 000 | 110 | 100 | 010 | 100 | 332 | 343 | 530 | 211 | 400 | 010 | 2015 | 6 |
| 1968 | 0 | 1 | 0 | 1 | 0 | 4 | 3 | 8 | 4 | 6 | 4 | 0 | 31 |  |
|  | 000 | 001 | 000 | 100 | 000 | 202 | 120 | 341 | 400 | 510 | 400 | 000 | 207 | 4 |
| 1969 | 1 | 0 | 1 | 1 | 0 | 0 | 3 | 3 | 6 | 5 | 2 | 1 | 23 |  |
|  | 100 | 000 | 010 | 100 | 000 | 000 | 210 | 210 | 204 | 410 | 110 | 010 | 136 | 4 |
| 1970 | 0 | 1 | 0 | 0 | 0 | 2 | 3 | 7 | 4 | 6 | 4 | 0 | 27 |  |
|  | 000 | 100 | 000 | 000 | 000 | 110 | 021 | 421 | 220 | 321 | 130 | 000 | 1212 | 3 |
| 1971 | 1 | 0 | 1 | 2 | 5 | 2 | 8 | 5 | 7 | 4 | 2 | 0 | 37 |  |
|  | 010 | 000 | 010 | 200 | 230 | 200 | 620 | 311 | 511 | 310 | 110 | 000 | 2411 | 2 |
| 1972 | 1 | 0 | 1 | 0 | 0 | 4 | 5 | 5 | 6 | 5 | 2 | 3 | 32 |  |
|  | 100 | 000 | 001 | 000 | 000 | 220 | 410 | 320 | 411 | 410 | 200 | 210 | 228 | 2 |
| 1973 | 0 | 0 | 0 | 0 | 0 | 0 | 7 | 6 | 3 | 4 | 3 | 0 | 23 |  |
|  | 000 | 000 | 000 | 000 | 000 | 000 | 430 | 231 | 201 | 400 | 030 | 000 | 129 | 2 |
| 1974 | 1 | 0 | 1 | 1 | 1 | 4 | 5 | 7 | 5 | 4 | 4 | 2 | 35 |  |
|  | 010 | 000 | 010 | 010 | 100 | 121 | 230 | 232 | 320 | 400 | 220 | 020 | 1517 | 3 |
| 1975 | 1 | 0 | 0 | 1 | 0 | 0 | 1 | 6 | 5 | 6 | 3 | 2 | 25 |  |
|  | 100 | 000 | 000 | 001 | 000 | 000 | 010 | 411 | 410 | 321 | 210 | 002 | 146 | 5 |
| 1976 | 1 | 1 | 0 | 2 | 2 | 2 | 4 | 4 | 5 | 0 | 2 | 2 | 25 |  |
|  | 100 | 010 | 000 | 110 | 200 | 200 | 220 | 130 | 410 | 000 | 110 | 020 | 1411 | 0 |
| 1977 | 0 | 0 | 1 | 0 | 1 | 1 | 4 | 2 | 5 | 4 | 2 | 1 | 21 |  |
|  | 000 | 000 | 010 | 000 | 001 | 010 | 301 | 020 | 230 | 310 | 200 | 100 | 118 | 2 |
| 1978 | 1 | 0 | 0 | 1 | 0 | 3 | 4 | 8 | 4 | 7 | 4 | 0 | 1132 | 2 |
|  | 010 | 000 | 000 | 100 | 000 | 030 | 310 | 341 | 310 | 412 | 121 | 000 | 1513 | 4 |
| 1979 | 1 | 0 | 1 | 1 | 2 | 0 | 5 | 4 | 6 | 3 | 2 | 3 | 28 |  |
|  | 100 | 000 | 100 | 100 | 011 | 000 | 221 | 202 | 330 | 210 | 110 | 111 | 149 | 5 |
| 1980 | 0 | 0 | 1 | 1 | 4 | 1 | 5 | 3 | 7 | 4 | 1 | 1 | 28 |  |
|  | 000 | 000 | 001 | 010 | 220 | 010 | 311 | 201 | 511 | 220 | 100 | 010 | 159 | 4 |
| 1981 | 0 | 0 | 1 | 1 | 1 | 2 | 5 | 8 | 4 | 2 | 3 | 2 | 29 |  |
|  | 000 | 000 | 100 | 010 | 010 | 200 | 230 | 251 | 400 | 110 | 210 | 200 | 1612 | 1 |
| 1982 | 0 | 0 | 3 | 0 | 1 | 3 | 4 | 5 | 6 | 4 | 1 | 1 | 28 |  |
|  | 000 | 000 | 210 | 000 | 100 | 120 | 220 | 500 | 321 | 301 | 100 | 100 | 197 | 2 |
| 1983 | 0 | 0 | 0 | 0 | 0 | 1 | 3 | 6 | 3 | 5 | 5 | 2 | 25 |  |
|  | 000 | 000 | 000 | 000 | 000 | 010 | 300 | 231 | 111 | 320 | 320 | 020 | 1211 | 2 |
| 1984 | 0 | 0 | 0 | 0 | 0 | 2 | 5 | 7 | 4 | 8 | 3 | 1 | 30 |  |
|  | 000 | 000 | 000 | 000 | 000 | 020 | 410 | 232 | 130 | 521 | 300 | 100 | 1611 | 3 |
| 1985 | 2 | 0 | 0 | 0 | 1 | 3 | 1 | 7 | 5 | 5 | 1 | 2 | 27 |  |
|  | 020 | 000 | 000 | 000 | 100 | 201 | 100 | 520 | 320 | 410 | 010 | 110 | 179 | 1 |
| 1986 | 0 | 1 | 0 | 1 | 2 | 2 | 2 | 5 | 2 | 5 | 4 | 3 | 27 |  |
|  | 000 | 100 | 000 | 100 | 110 | 110 | 200 | 410 | 200 | 320 | 220 | 210 | 198 | 0 |
| 1987 | 1 | 0 | 0 | 1 | 0 | 2 | 4 | 4 | 7 | 2 | 3 | 1 | 125 |  |
|  | 100 | 000 | 000 | 010 | 000 | 110 | 400 | 310 | 511 | 200 | 120 | 100 | 186 | 1 |
| 1988 | 1 | 0 | 0 | 0 | 1 | 3 | 2 | 5 | 8 | 4 | 2 | 1 | 27 |  |
|  | 100 | 000 | 000 | 000 | 100 | 111 | 110 | 230 | 260 | 400 | 200 | 010 | 1412 | 1 |
| (1959-1988) |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| AVG | 0.6 | 0.3 | 0.6 | 0.7 | 1.2 | 2.1 | 4.5 | 6.2 | 5.7 | 4.5 | 2.8 | 1.4 | 30.6 |  |
| CASES | 17 | 9 | 18 | 22 | 37 | 63 | 134 | 185 | 172 | 136 | 83 | 43 | 919 |  |



## 2. WESTERN NORTH PACIFIC TROPICAL CYCLONES

Distinguishing features of the 1988 western North Pacific tropical cyclone season were the low number of super typhoons (1), the short average lifespan of the tropical cyclones, the aclimatic location of the monsoon trough and an active Tropical Upper-Tropospheric Trough (TUTT). The northward displacement of the monsoon trough during the summer and early fall and the active TUTT may have accounted for the relatively large number of tropical cyclones that had anomalous tracks. The normal lifespan of a tropical cyclone in the western North Pacific usually exceeds four warning days. This year JTWC encountered a large number of tropical cyclones (13) that were in warning status for four days or less. The short lifespans led to a relatively low number of warnings and forecast verifications.

## JANUARY THROUGH MAY

Typhoon Roy (01W) was only the second typhoon in the past twelve years to develop in the western North Pacific during January. The typhoon's near miss of Guam resulted in the most destruction since Super Typhoon Pamela (1976) struck the island. After Roy (01W) there was a long break in activity until the end of May. The synoptic pattern during the last week of May was anomalous, with low-level southwesterlies extending across the northern Philippine Sea into the northern Marianas and southern Bonin Islands. Surface pressures in the monsoon trough were 4 to 5 mb below normal. Cyclonic vortices in the trough were transitory until Typhoon Susan (02W) formed off the coast of Luzon.

## JUNE

As Susan (02W) moved northeastward, Tropical Depression 03W developed in the enhanced low-level southwesterly monsoonal flow left behind Susan (02W). Then Tropical Depression 03W moved into a subsidence area over China and dissipated. A two week hiatus in
tropical cyclone activity followed. Then Typhoon Thad (04W) formed in the eastern Carolines. It tracked over 2000 nm ( 3704 km ) during its lifetime, recurving just east of the island of Luzon and passing $80 \mathrm{~nm}(148 \mathrm{~km})$ southeast of Okinawa. With Thad (04W) weakening over water to the north, Tropical Storm Vanessa (05W) generated to the south in the Philippine Sea. It was the first "straightrunner" of the year. Vanessa (05W) tracked across the Philippine Islands and into the South China Sea before dissipating over southern China.

## JULY

Almost two weeks passed after Vanessa's (05W) demise before Typhoon Warren (06W) developed in the eastem Caroline Islands. Warren (06W) was the second tropical cyclone of the year to threaten Guam. Warren ( 06 W ) was the second "straight-runner" of the year and maintained a west-northwestward track during almost its entire lifetime. The system skirted northern Luzon prior to making landfall in southeastern China. Tropical Storm Agnes (07W) followed a week later and was noteworthy for several reasons. It was the last of only two tropical cyclones to develop in July, a month that normally averages five systems and played a major role in the changing synoptic pattern during late July. When Warren (06W) dissipated, the monsoon trough remained much farther north than normal. Agnes (07W) formed in the area of lower pressures southeast of Japan where the monsoon trough merged with a mid-latitude low pressure system to the northeast. Agnes (07W) followed the path of least resistance and accelerated northnortheastward along the trough axis.

## AUGUST

Once Agnes (07W) went extratropical the monsoon trough underwent a major readjustment. It now stretched eastward from the Gulf of Tonkin, across the South China Sea, through the Luzon Strait and abruptly terminated near Okinawa. Tropical Storm Bill
(08W) consolidated rapidly at the eastern end of the monsoon trough, brushed by the island of Okinawa and reached a peak intensity of 45 kt ( $23 \mathrm{~m} / \mathrm{sec}$ ) before making landfall near Shanghai, China. Bill (08W) remained well organized even after making landfall, and caused widespread destruction and loss of life in China. The other four tropical cyclones that developed in August (Clara (09W), Doyle (10W), Elsie (11W) and Fabian (12W)) all formed north of $20^{\circ}$ North latitude. Tropical Storm Clara (09W) began in the easterly trade winds as an area of weakly organized convection $540 \mathrm{~nm}(1,000 \mathrm{~km})$ north of Wake Island. Clara (09W) initially tracked westward, then abrubtly changed direction toward the north. Throughout its short lifespan, the system was consistently hindered by vertical wind shear and only peaked at an intensity of 45 kt ( $23 \mathrm{~m} / \mathrm{sec}$ ). Typhoon Doyle (10W) also fell into the track category of "other" due to its erratic behavior. Initially, Doyle (10W) moved rapidly toward the south-southwest and looped before tracking northeastward. To make the forecasts more complicated, Doyle (10W) interacted with a TUTT cell while maintaining typhoon intensity. Once Doyle (10W) was extratropical, Tropical Storm Elsie (11W) and Typhoon Fabian (12W) formed from persistent convection in the monsoon trough. Both displayed erratic movement during their early stages and underwent binary interaction before transitioning into extratropical systems.

## SEPTEMBER

With Elsie (11W) and Fabian (12W) going extratropical, Tropical Storm Gay (13W) generated in the monsoon trough 420 nm ( 778 km ) east of Okinawa and attained a peak intensity of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$. It took the path of least resistance and tracked up the trough to the northeast. Gay (13W) was short-lived. It was followed by a more normal synoptic situation, where the monsoon trough shifted equatorward. The monsoon trough later expanded, stretching from Vietnam to $175^{\circ}$ East longitude, and spawned seven tropical cyclones. As Gay (13W) dissipated east of Japan and Uleki (01C)
churned across the Central Pacific, Typhoon Hal (14W) formed just west of Wake Island. Its development was aided by upper-level divergence from a TUTT cell to its north. Hal (14W) combined with Typhoon Uleki (01C), Tropical Storm Irma (15W), and later with Tropical Storm Jeff (16W) to create two separate three-storm situations. In the meantime, Typhoon Uleki ( 01 C ) became the third hurricane in the past thirty years to form in the central North Pacific and cross the international dateline while in a warning status. Tropical Storms Irma (15W) and Jeff (16W) developed in Hal's (14W) strong low-level southwesterly inflow. As Hal (14W), with a large ragged eye, tracked northward, Irma (15W) and Jeff (16W) followed and were sheared away. Once Hal (14W) went extratropical east of Japan, Tropical Storm Kit (17W) developed in the monsoon trough 240 $\mathrm{nm}(444 \mathrm{~km})$ east of the Philippine Islands. Kit (17W) was a "straight-runner" and tracked over the northern tip of Luzon. It made landfall over southern China, causing loss of life and property damage. While Kit (17W) was moving into Luzon, Tropical Storm Lee (18W) was slowly developing. Lee (18W) tracked over $1,300 \mathrm{~nm}(2408 \mathrm{~km}$ ) during a four day period as an identifiable area of convection before the first warning was issued. It then moved northwestward before recurving to the northeast and tracking within $45 \mathrm{~nm}(83 \mathrm{~km}$ ) southeast of Okinawa. Tropical Storm Mamie (19W) formed in tandem with Kit (17W) and was the second significant tropical cyclone to develop in the South China Sea. Mamie (19W) had an anomalous track. After a prolonged southwestward movement, it made a sharp turn and moved northward towards Hong Kong.

## OCTOBER

Once Lee (18W) and Mamie (19W) were gone, there was a five day break before Super Typhoon Nelson (20W). It was the only super typhoon of 1988 . The tropical cyclone initially moved westward towards the Philippine Islands, then west-northwestward as it tracked along the southwestern side of the subtropical
ridge. A break in the ridge northwest of Nelson (20W) was identified and recurvature was correctly forecast. Nelson (20W) rapidly deepened for two days and reached super typhoon intensity shortly before recurvature. It then recurved and threatened Okinawa. Later, as the system became extratropical and accelerated toward the northeast, it also threatened Japan. While Nelson (20W) was weakening and accelerating, Typhoon Odessa (21W) began as an area of convection superimposed on broad low-level easterly tradewinds 600 nm ( 1111 km ) south-southeast of Japan. During its first two days, Odessa (21W) moved rapidly to the west-northwest at a speed of 17 to 18 kt ( 32 to $33 \mathrm{~km} / \mathrm{hr}$ ). It began a gradual recurvature and moved toward the cooler, drier polar air mass from the Asian continent. Initially, interaction with cold air was expected to weaken the tropical cyclone. Instead, Odessa (21W) intensified into a midget typhoon, peaking at an intensity of 90 kt ( $46 \mathrm{~m} / \mathrm{sec}$ ). At the same time Tropical Storm Pat (22W) formed equatorward of $10^{\circ}$ North latitude. The system tracked westward and attained a peak intensity of 75 kt ( $39 \mathrm{~m} / \mathrm{sec}$ ) prior to making landfall over central Luzon. Pat (22W) then moved west-northwestward across the South China Sea, tracking over the island of Hainan Dao and dissipating 30 nm ( 56 km ) northeast of Hanoi, Vietnam. As Pat (22W) was winding down in the South China Sea, Typhoon Ruby (23W) was intensifying in the Philippine Sea. Ruby (23W) was the fifth tropical cyclone to track across the Philippine Islands and the third system to affect Vietnam in 1988. It tracked in a west-northwestward direction throughout its lifetime. The system reached a peak intensity of 125 kt ( $64 \mathrm{~m} / \mathrm{sec}$ ) shortly before making landfall in the Philippines. The result was extensive damage and loss of life. In the Philippines alone, at least 300 persons were killed and over 470,000 were left homeless.

Ruby (23W) passed 65 nm ( 120 km ) northnortheast of Manila, causing the strongest winds at Clark Air Base since 1978. Ruby (23W) then tracked into the South China Sea. Later, flash flooding from the dissipating system's torrential rainshowers resulted in over 100 deaths, hundreds of thousands homeless and widespread destruction of crops in Vietnam.

## NOVEMBER THROUGH DECEMBER

November marked a change in the synoptic flow pattern as the northeast monsoon became well established across the South China Sea and southeastern Asia. Easterly tradewinds dominated the Philippine Sea north of the nearequatorial trough. After a one week respite, Typhoon Skip (24W) appeared. It was a "straight-runner" and covered over $2,000 \mathrm{~nm}$ ( 3704 km ) during its nine day lifetime. Skip (24W) tracked through the Philippine Islands and into the South China Sea. The system caused widespread damage to crops in the Philippines, killed over 100 persons and left over 600,000 homeless. Typhoon Tess (25W) formed in the near-equatorial trough before Skip ( 24 W ), but was slow to intensify. It was the only tropical cyclone to track across southern Vietnam this year. After Skip (24W) and Tess (25W), a break in tropical cyclone activity occurred until the third week of December. Following a massive outbreak of polar air from Asia, the southern Philippine Sea filled with convection and a near-equatorial trough formed. Tropical Storm Val (26W), which developed in the trough, proved difficult to position and hard to forecast. While decelerating from 25 kt ( 46 $\mathrm{km} / \mathrm{hr}$ ), Val (26W) peaked at an intensity of 55 $\mathrm{kt}(28 \mathrm{~m} / \mathrm{sec})$. Finally, the low-level circulation separated from the deep convection and was carried to the southwest along the edge of the winter monsoon.

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## TYPHOON ROY (01W)

Typhoon Roy was the first significant tropical cyclone of 1988 in the western North Pacific. It formed as a "twin" (Figure 3-01-1) with its southern hemisphere counterpart, Tropical Cyclone 07P (Anne). During a period of eleven days in January, Roy made a 4000 nm
( 7408 km ) westward trek, caused significant damage on Kwajalein Atoll and the islands of Guam and Rota, crossed the Philippine Islands and dissipated over the South China Sea. Typhoon Roy's close approach to Guam resulted in the most destruction since Super


Figure 3-01-1. Typhoon Roy (01W) with its southern hemisphere counterpart, Tropical Cyclone 07P (Anne) (081957Z January NOAA infrared imagery).

Typhoon Pamela (1976).
Prior to tropical cyclone genesis, above normal sea surface temperature anomalies and greater than normal cloudiness persisted in the central Pacific Ocean. Roy began in this area of increased cloudiness southeast of the Marshall Islands, with persistent convection first noted on the Significant Tropical Weather Advisory at 060600 Z . The suspect area was mirrored by another area of persistent convection in the southern hemisphere, which developed into Tropical Cyclone 07P (Anne). A band of anomalous low-level equatorial westerlies was located between the two cloud systems. Gradient-level wind reports from Tarawa (WMO 91610) in the Kiribati Islands during early January consistently indicated moderate westerly winds. (Climatic windflow at Tarawa for January is east-northeasterly at 12 kt ( 6 $\mathrm{m} / \mathrm{sec}$ ).)

By 7 January, Roy's cloud organization had improved and JTWC issued a Tropical Cyclone Formation Alert at 072000Z. Satellite intensity technique estimates of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ combined with synoptic reports of 30 kt ( 15 $\mathrm{m} / \mathrm{sec}$ ) surface winds and a 997 mb surface pressure from Majuro (WMO 91376) prompted the issuance of the first warning on Tropical Depression 01W at 080000Z. (Tropical Cyclone 07P (Anne) in the southern hemisphere reached tropical storm intensity 12 -hours earlier). As Tropical Depression 01W moved north of Majuro, the island experienced maximum sustained winds of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) with gusts to $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec}$ ), and several buildings suffered minor structural damage.

Satellite reconnaissance continued to detect deepening of the system and Tropical Depression 01W was upgraded to a tropical storm at 080600Z. Roy (Figure 3-01-1) passed $35 \mathrm{~nm}(65 \mathrm{~km})$ south of Kwajalein Atoll at 081800Z. Kwajalein Island (WMO 91366) reported maximum sustained winds of 48 kt ( 25 $\mathrm{m} / \mathrm{sec}$ ) with a peak gust of $57 \mathrm{kt}(29 \mathrm{~m} / \mathrm{sec})$, a minimum sea-level pressure (MSLP) of 992 mb and light-to-moderate structural damage. Ebeye


Figure 3-01-2. Roy (01W) near peak intensity ( 100823 Z January DMSP enhanced infrared imagery).

Island just $4 \mathrm{~nm}(7 \mathrm{~km})$ to the north experienced moderate-to-severe structural damage, one death and loss estimates of five million dollars. Both islands had 20 to $22 \mathrm{ft}(6.1$ to 6.7 m ) surf and low-lying areas were flooded. Using their weather radar, meteorologists on Kwajalein were the first to detect the formation of Roy's eye at. 081000 Z . Later at 091200 Z , a satellite estimate of $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ resulted in the upgrade to typhoon intensity.

While at a forward speed of 22 kt (41 $\mathrm{km} / \mathrm{hr}$ ) at 101200Z, Roy (Figure 3-01-2) reached a peak intensity of $115 \mathrm{kt}(59 \mathrm{~m} / \mathrm{sec}) 510 \mathrm{~nm}$ ( 945 km ) east-southeast of Guam . The typhoon was embedded in a moderate mid-tropospheric east-southeasterly flow south of the subtropical ridge axis, as indicated by aircraft reports at 500 mb (Figure 3-01-3). Then Typhoon Roy slowed to $12 \mathrm{kt}(22 \mathrm{~km} / \mathrm{hr})$ as it approached Guam (Figure 3-01-3). Detachment 2, 20 Weather


Figure 3-01-3. 100000 Z January 500 mb streamline analysis with additional reports along the aircraft track to the north of Roy. The aircraft reports are from point A (100057Z) to point B (100340Z).


Figure 3-01-4. Spiral rainband echoes define the typhoon's eye at 112310 Z January 120 nm ( 222 km ) east-southeast of the radar site. The weather radar (at point $A$ ) is located at Andersen Air Force Base on the northeastem tip of Guam (photo courtesy of MSgt Robert W. Yates and Detachment 2 , 20th Weather Squadron).

Figure 3-01-5. 110838Z January digitized radar display of Roy's rain shield. Note the island of Rota (call sign PGRO) lies within the ragged eye. The call sign PGUA marks the location of Andersen Air Force Base on Guam and PGSN the airfield on Saipan (photo courtesy of MSgt Robert W. Yates and Detachment 2, 20th Weather Squadron).



Figure 3-01-6. At 110838 Z January the Andersen Air Force Base weather radar display paints 37,000 feet ( 11.3 km ) rain echo tops in the outer eye wall cloud. The radar returns in the lower left of the picture and closest to the radar site are attenuated due to heavy rain (photo courtesy of MSgt Robert W. Yates and Detachment 2, 20th Weather Squadron).

Squadron at Andersen Air Force Base on Guam first detected the eye on radar at 111930 Z (Figure 3-01-4). These weather radar data were instrumental in tracking Typhoon Roy's center, as it made its closest point of approach (Figures 3-01-5 and 3-01-6) 32 nm ( 59 km ) north of Guam at 120930Z. Wind estimates near the center were $110 \mathrm{kt}(57 \mathrm{~m} / \mathrm{sec})$. However, Andersen Air Force Base on the northeastern tip of Guam measured maximum sustained winds of $66 \mathrm{kt}(34 \mathrm{~m} / \mathrm{sec}$ ) with peak gusts to $98 \mathrm{kt}(50$ $\mathrm{m} / \mathrm{sec}$ ) as the eye wall passed just to the north. Buildings, particularly those on the northern part of Guam, sustained light-to-moderate structural damage (Figure 3-01-7). Crops and vegetation on Guam suffered extensive damage, with estimates of losses as high as 23.5 million dollars. As a credit to the disaster preparedness team, no severe injuries or loss of life were


Figure 3-01-7. Roy almost went on a shopping spree as indicated by the structural damage to the Andersen Air Force Base commissary on Guam.


Figure 3-01-8. Strong winds rearranged these parked vehicles.
reported on Guam (Figure 3-01-8).
In comparison, the island of Rota, 40 nm ( 74 km ) north-northeast of Guam, suffered the heaviest damage. At 120724Z, shortly before eye passage, Rota's automated weather observing equipment reported maximum sustained winds of $71 \mathrm{kt}(37 \mathrm{~m} / \mathrm{sec})$ with peak gusts to 104 kt ( $54 \mathrm{~m} / \mathrm{sec}$ ). Because of apparent communication problems, no further data were received until the 120905 Z report of 60 kt ( 31 $\mathrm{m} / \mathrm{sec}$ ) with gusts to $89 \mathrm{kt}(46 \mathrm{~m} / \mathrm{sec}$ ). Residents of Rota described the eye passage as a marked lessening of wind speed and clearing skies from 120730Z to 120810 Z . Concurrently, a microbarograph trace from the Naval Oceanography Command Detachment, Agana, located on central Guam indicated a minimum sea-level pressure of 979 mb from 120800 Z to 121000Z (Figure 3-01-9). A large percentage of the homes on Rota were destroyed and the remainder damaged. Four minor injuries were reported, which resulted when a flying roof impacted another building where people had sought shelter. Numerous coconut trees were downed and all crops destroyed. With an estimated 95 percent of the utility poles knocked down, lack of power and potable water completely disrupted the community.

After moving through the southern Marianas, Roy continued to slow. Earlier analysis of 500 mb aircraft reports revealed a mid-tropospheric anticyclone east of the Philippine Islands with ridging extending to the northeast of Roy's center. As Typhoon Roy approached the Mariana Islands, it apparently responded to the weaker mid-level steering flow and decelerated. A weakness in the subtropical ridge was located almost due north of Guam . It was initially thought that Roy would weaken the subtropical ridge and ultimately recurve. However, this did not happen. Instead, the lower tropospheric ridge built, as reflected by 700 mb pressure-height rises at Iwo Jima (WMO 47971). In turn the typhoon accelerated to the southwest. (By this time, Roy's maximum sustained winds had weakened to $90 \mathrm{kt}(46 \mathrm{~m} / \mathrm{sec})$. This intensity was maintained until reaching the mountainous terrain of southern Luzon.)

At 141800 Z , Roy returned to a more westward course along the southern edge of the subtropical ridge and increased its speed of 20 kt ( $37 \mathrm{~km} / \mathrm{hr}$ ). From 160000 Z to 170000 Z , Roy tracked across southern Luzon. The mountains and increased vertical wind shear further


Figure 3-01-9. Microbarograph (pressure) trace of Typhoon Roy ( 01 W ) recorded at the Naval Oceanography Command Detachment, Agana, Guam indicates a minimum sea-level pressure of 979 mb from 120800 Z to 121000 Z January.
weakened Roy from $75 \mathrm{kt}(39 \mathrm{~m} / \mathrm{sec})$ to 40 kt ( $21 \mathrm{~m} / \mathrm{sec}$ ). Once in the South China Sea, Roy's interaction with the low-level northeasterly flow
of the winter monsoon spawned gales, but dissipation was imminent. At 171200 Z , Tropical Storm Roy was downgraded to a
tropical depression and the final warning followed at 180000Z.

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## TYPHOON SUSAN (02W)

Susan was a short-lived typhoon with maximum sustained winds of $75 \mathrm{kt}(39 \mathrm{~m} / \mathrm{sec})$. Slow to exit the the South China Sea, it threatened the southeast of coast of China, then churned across the southern tip of Taiwan and rapidly weakened.

The synoptic pattern during the fourth week of May was anomalous with low-level southwesteries extending across the northern Philippine Sea into the northern Marianas and southern Bonin Islands (Figure 3-02-1). Surface pressures in the monsoonal trough, that was north of this southwesterly flow, were 4 to 5 mb below normal. Cyclonic vortices that formed in the trough were transitory until 28 May when a persistent circulation formed off
the northwest coast of Luzon (see Figure 3-022 ). Initially the cenvection was displaced equatorward of the the low-level circulation center by vertical wind shear, but within a day the cloudiness became more centralized. The cloud system as a whole then appeared to isolate itself from the surrounding zone of maximum cloudiness. The Significant Tropical Weather Advisory was reissued at 290200 Z May to include this suspect monsoon depression. Although the upper-level outflow was restricted in the north and west, the amount of central convection and organization continued to increase, prompting JTWC to issue a Tropical Cyclone Formation Alert at 300200 Z May. The first warning on Tropical Storm Susan at 300600 Z followed from a satellite


Figure 3-02-1. Surface/gradient analysis ( 260000 Z May) shows the anomalous southwesterly flow extending eastward from the South China Sea.


Figure 3-02-2. Susan as a tropical disturbance (280055Z May DMSP visual imagery).
analysis wind estimate of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$.

At warning time Susan's initial position was $65 \mathrm{~nm}(120 \mathrm{~km})$ west of northern Luzon. Past movement had been erratic because the low-level circulation was located within the larger monsoonal trough. For forecast movement the tropical cyclone was near the axis of the subtropical ridge and recurvature was favored by the Typhoon Acceleration Prediction Technique (TAPT) (Weir, 1982). However, TAPT guidance identified the 200 mb northwesterly flow as unfavorable for rapid acceleration. The initial track forecasts were correct based on this guidance and Susan recurved and moved slowly to the northeast.

Susan intensified rapidly after recurvature. At 310600 Z Susan was upgraded to a typhoon based upon satellite intensity estimates. The sustained winds increased to 75 kt ( $39 \mathrm{~m} / \mathrm{sec}$ ) at 010600 Z June (see Figure 3-023). Now packing its most dangerous winds,


Figure 3-02-3. Typhoon Susan at maximum intensity in the Luzon Strait (011109Z June DMSP infrared imagery).

Susan accelerated towards the northeast and the southern tip of Taiwan. Aloft, a mid-level trough in the polar westerlies was advancing across eastern China. The trough became more meridional as it approached Susan. A combination of acceleration along-track, terrain effects (induced by the rugged mountains of Taiwan) and increasing vertical shear stripped away Susan's deep central convection, leaving behind an exposed low-level circulation center (Figure 3-02-4). The typhoon was downgraded to tropical storm intensity at 021200 Z and further to a tropical depression at 030000 Z . The final warning was issued at 030600 Z . Twelve hours later the residual low-level vortex was no longer discernible on satellite imagery or in the synoptic data.

In retrospect, the majority of the damage
to the island of Luzon, Republic of the Philippines resulted from heavy rains, not winds. A landslide triggered by these rains in Olongapo City, $50 \mathrm{~nm}(93 \mathrm{~km})$ northwest of Manila, led to one death. In Manila another landslide killed five people. Flooding closed the main roads in Manila, disrupted travel and caused the loss of millions of prawns and lobsters from fishponds. Also, a tornado destroyed 18 homes outside of Manila.

Although Susan passed about 10 nm (18 km ) south of the island of Okinawa, Japan at 022200 Z , the system had rapidly weakened and the peak wind recorded at Kadena Air Base was $41 \mathrm{kt}(21 \mathrm{~m} / \mathrm{sec})$ with $47 \mathrm{kt}(24 \mathrm{~m} / \mathrm{sec})$ at Naha. No deaths or injuries were reported by authorities on Okinawa. No reports of damage were received from Taiwan.


Figure 3-02-4. The residual exposed low-level circulation of Tropical Depression 02W (Susan) (030057Z June DMSP visual imagery).


## TROPICAL DEPRESSION 03W

On 3 June the remains of Typhoon Susan (02W) sped northeastward and left behind the low-level southwest monsoonal flow which terminated abruptly in the northwestern Philippine Sea. Within a day the enhanced convection in the northwestern Philippine Sea acquired convective banding and cyclonic rotation. A Tropical Cyclone Formation Alert documented this event at 040200 Z . The
convection consolidated near the low-level circulation center and the first warning on Tropical Depression 03W followed at 040600 Z based on a satellite intensity estimate of 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) surface winds. The satellite imagery (Figure 3-03-1) shows Tropical Depression 03W near its maximum intensity of 30 kt ( 15 $\mathrm{m} / \mathrm{sec}$ ).


Figure 3-03-1. Tropical Depression 03W near peak intensity (040037Z June DMSP visual data).

Shortly after the first warning was issued, the central convection collapsed and further intensification ceased. In Figure 3-03-2 note that the system center is basically free of deep convection with only remnants of high cloud debris evident. The banding feature is displaced to the south and east. The next
daytime visual imagery (Figure 3-03-3) reveals low-level stratiform cloudiness filling the center. The deep convection is well removed from the center with the exception of one transitory cumulonimbus. The final warning was issued on 051200 Z June, when it became apparent that the system was dissipating.


Figure 3-03-2. Only high-level cloud debris are in evidence over the center of the low-level circulation ( 041318 Z June DMSP infrared imagery).

In retrospect Tropical Depression 03W's failure to mature and achieve tropical storm intensity may be related to its track. In contrast to Typhoon Susan (02W), which traveled northeastward along a zone of increased cloudiness, Tropical Depression 03W took a west-northwesterly track into the cloud minimum area that had settled across southeastern China and the northern South China Sea. Bao (1981) developed a hypothesis
for forecasting typhoon movement based on satellite observed cloudiness which suggested that tropical cyclones move into, or along, areas of preexisting enhanced cloudiness - or conversely, tropical cyclones don't move into areas of minimum cloudiness. If they do, there is a price. Unless the tropical cyclone is large enough to modify the ambient environment, which is unfavorable, dissipation will result.


Figure 3-03-3. Stratiform low-level cloudiness fills the center of TD 03W (042302Z June NOAA visual imagery).


## TYPHOON THAD (04W)

Thad was the second of three tropical cyclones to develop during June and the third typhoon of 1988. Typhoon Thad tracked over $2000 \mathrm{~nm}(3704 \mathrm{~km}$ ) during its lifetime, recurved just east of the island of Luzon in the Republic of the Philippines and passed $80 \mathrm{~nm}(148 \mathrm{~km})$ southeast of the island of Okinawa, Japan before dissipating over water. The recurvature forecast was complicated by a complex interaction of the tropical cyclone with upperlevel synoptic features.

After Tropical Depression 03W dissipated during the first week of June, there was a two week hiatus in tropical cyclone activity in the western North Pacific and low-
level westerly flow established itself across the southern Philippine Sea. Thad began in the zone of increased cyclonic shear between this westerly flow and the easterly tradewinds 300 $\mathrm{nm}(556 \mathrm{~km})$ south of Ulithi Atoll in the western Caroline Islands. The disturbance was first mentioned on the Significant Tropical Weather Advisory at 180600Z. Initially Thad's intensification may have been slowed by increased upper-level wind shear across the system, caused by the unfavorable location of an intense Tropical Upper-Tropospheric Trough (TUTT) low to its northeast. However, as the disturbance's central convection consolidated, the separation between Thad's upper-level circulation center and the TUTT low lessened.


Figure 3-04-1. Thad, shortly before reaching tropical storm intensity. Note the vigorous TUTT cell (at A) to the northeast of the tropical disturbance (191914Z June NOAA infrared imagery).

The system's upper-level outflow pattern improved and a Tropical Cyclone Formation Alert followed at 190800Z (Figure 3-04-1). At 200000 Z , satellite intensity analysis indicated a T-number of 2.5 , corresponding to maximum sustained surface winds of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ (Figure 3-04-2) and the Alert was upgraded to Tropical Storm Thad.

Throughout this period of gradual intensification, Thad was embedded in the flow south of the lower tropospheric subtropical ridge axis and moved northwestward, except for one excursion - the "stair-step" jog in the track from 200000Z to 201800Z. Afterward, Thad continued to track northwestward, intensifying for two more days, until reaching


Figure 3-04-2. Thad intensifies as it approaches the island of Luzon (212128Z June NOAA visual imagery).
the westernmost end of the subtropical ridge where it began to recurve at 221800 Z . The system reached its peak intensity and was upgraded to typhoon intensity at 220600 Z , based on a satellite intensity estimate of 70 kt ( $36 \mathrm{~m} / \mathrm{sec}$ ) (Figure 3-04-3). Thad then developed a central cold cover (Dvorak, 1984) and further development was arrested.

At 221800Z, Typhoon Thad recurved in response to the approaching mid-level trough in the westerlies aloft over eastern China. The forecast for this event was complicated by the failure of JTWC's dynamic forecast aid, the One Way (Interactive) Tropical Cyclone Model (OTCM), to change from a persistent northwestward track (Figure 3-04-3). At 231200 Z , the decision was made to disregard the objective forecast guidance and forecast recurvature based on synoptic data analyses.

This decision was correct, but in retrospect, the timing was 18 -hours late.

From 240000 Z to 250000 Z , Thad underwent rapid weakening as it tracked to the northeast and entered a region of increased vertical wind shear. At 240000Z, Typhoon Thad was downgraded to tropical storm intensity and six hours later the system weakened to a $50 \mathrm{kt}(26 \mathrm{~m} / \mathrm{sec})$ intensity as it made a closest point of approach of 80 nm (148 km ) southeast of the island of Okinawa, Japan. Both Kadena Air Base and Naha airport reported wind speeds below 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) during Thad's passage. With dissipation over water underway, Thad was downgraded to a tropical depression at 250000 Z and the final warning issued. No reports of damage were received.


Figure 3-04-3. OTCM forecast guidance (dashed lines) from 210000 Z to the point of recurvature at 221800 Z basically held Thad to a northwestward track through each 72 -hour period, in contrast to the best track (solid line).


## TROPICAL STORM VANESSA (05W)

The third of three significant tropical cyclones to develop during the month of June, Vanessa was the first "straight-runner" of the year in the western North Pacific. It tracked across the Philippine Islands and into the South China Sea before dissipating over southern China.

As Thad (04W) moved northeastward and weakened over the Philippine Sea east of the island of Okinawa, Vanessa was first detected at 241200 Z by satellite imagery analysts 125 nm ( 232 km ) east of Koror in the western Caroline Islands. A flare-up of convection at 241800 Z resulted in an increase in high clouds. The upper-level outflow began to show organization (Figure 3-05-1). At 250600 Z , the tropical disturbance was described
on JTWC's Significant Tropical Weather Advisory as an area of persistent convection 55 $\mathrm{nm}(102 \mathrm{~km})$ southeast of Koror. Synoptic data indicated a well organized low-level circulation embedded in the near-equatorial trough. Satellite imagery revealed a Tropical UpperTropospheric Trough (TUTT) low located 250 $\mathrm{nm}(463 \mathrm{~km})$ northeast of the disturbance's lowlevel circulation. The upper-level cold low interrupted the disturbance's upper-level outflow in its northeast quadrant. At 260440Z, a Tropical Cyclone Formation Alert was issued when satellite imagery revealed increased convection. The TUTT low had weakened and passed north of the low-level circulation, resulting in divergent upper-level flow across the disturbance.


Figure 3-05-1. Vanessa as a tropical disturbance (250028Z June DMSP visual imagery).

The first warning followed at 261200Z, when the tropical disturbance was upgraded to Tropical Depression 05W, based on continued improvement in the system's organization and convection. Satellite intensity analysis indicated surface wind speeds of 30 to 35 kt ( 15 to $18 \mathrm{~m} / \mathrm{sec}$ ). The system continued to intensify. At 270000Z, Vanessa (Figure 3-05-2) was again upgraded, this time to a tropical storm, based on a satellite intensity analysis of
$35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ surface winds. With winds of 45 kt ( $28 \mathrm{~m} / \mathrm{sec}$ ), Vanessa made landfall over the Republic of the Philippines, at 270600Z, between the islands of Samar and Mindanao.

The tropical cyclone tracked rapidly across the central Philippine Islands, as a weak tropical storm, and entered the South China Sea at 280200 Z . Vanessa continued its rapid movement as it tracked across the South China


Figure 3-05-2. Nearing peak intensity, Vanessa approaches the central Philippine Islands (270130Z June DMSP visual imagery).

Sea. Despite increased vertical wind shear, Vanessa tenaciously resisted weakening until 290000 Z. Over the next six hours, the deep convection was stripped away from the lowlevel circulation center (Figure 3-05-3) and the final warning followed at 290600Z. Vanessa
made landfall just west of Macao on the south coast of China at 290800 Z. Nearby land stations reported $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec}$ ) maximum surface winds. No reports of major damage were received.


Figure 3-05-3. Vanessa's exposed low-level circulation, shortly before the system made landfall near Macao (290802Z June NOAA visual imagery).


## TYPHOON WARREN (06W)

Typhoon Warren was the first of two significant tropical cyclones to develop during the month of July, and the fourth tropical cyclone of the year to reach typhoon intensity. Warren was a "straight-runner" and maintained a west-northwestward direction of movement during almost its entire life span.

The tropical disturbance that eventually developed into Typhoon Warren was first mentioned on the Significant Tropical Weather Advisory at 110600 Z . The suspect area had persisted as a poorly organized area of convection for 12 -hours in the eastern Caroline Islands. Its potential for further development was assessed as "poor." However, better convective organization and improved upperlevel outflow raised the potential for development into a significant tropical cyclone to "fair" and JTWC reissued the Advisory at 111230 Z . Increased convection prompted a Tropical Cyclone Formation Alert at 120530Z. The system was now 180 nm ( 333 km ) southeast of Guam and headed for the island. Satellite intensity analysis estimated sustained surface winds of 25 kt ( $13 \mathrm{~m} / \mathrm{sec}$ ).

The appearance of a central dense overcast (CDO) on the 121237 Z satellite imagery led the satellite analyst to increase the intensity estimate of surface winds to 30 kt ( 15 $\mathrm{m} / \mathrm{sec}$ ). From 121200 Z to 121500 Z , the system's CDO increased in size. Because of the disturbance's steady development and its proximity to Guam, JTWC issued an abbreviated warning for the tropical depression at 121600 Z : the detailed warning followed at 121800 Z . The center of the system passed 55 $\mathrm{nm}(102 \mathrm{~km})$ to the south of Guam at 130000 Z .

After returning to a more westerly track, Tropical Depression 06W was upgraded to Tropical Storm Warren at 130600Z. (Postanalysis showed that Warren probably attained tropical storm intensity earlier at 130000Z.) At 141800Z, Warren reached typhoon intensity. This intensification process continued and
peaked at $115 \mathrm{kt}(58 \mathrm{~m} / \mathrm{sec})$ in the Philippine Sea 300 nm ( 556 km ) east of Luzon (Figure 3-06-1). During this same two day period as the winds doubled in intensity, Warren also doubled its forward speed to 15 kt ( $28 \mathrm{~km} / \mathrm{hr}$ ).

While Warren tracked across the Philippine Sea, the One Way (Interactive) Tropical Cyclone Model (OTCM) outlook began to take the track northward into the


Figure 3-06-1. Typhoon Warren at peak intensity ( 162247 Z July NOAA infrared imagery).


Figure 3-06-2. A comparison of OTCM 72-hour guidance with JTWC's best track. Note OTCM's systematic strong northward bias.


Figure 3-06-3. Seventy-two hour forecasts tended to be north of the best track and more conservative than the OTCM guidance.
subtropical ridge. OTCM is JTWC's operational dynamic aid and in general provides the best performance of all the objective aids. This numerical guidance moved Warren north and east of the island of Taiwan and eventually suggested recurvature within 48 - to 72 -hours (Figure 3-06-2). However, JTWC's synoptic data analyses and mid-level prognoses maintained the subtropical ridge to the north. This was reinforced by satellite observed persistent minimum cloudiness, which was associated with subsidence and ridging, to the north of Warren. Consequently, recurvature was not forecast (Figure 3-06-3).

From 171800 Z to 180000 Z , Typhoon Warren weakened as it skirted the northern coast of Luzon with damage to rice and corn crops in northern Luzon estimated to be $\$ 10$
million.
After its brush with Luzon, the tropical cyclone maintained typhoon intensity until making landfall at 190600Z (Figure 3-06-4) near the city of Shantou in southeastern China. China's official media reported 17 killed and 153 injured by Warren. Additionally, in the province of Guangdong in southeastern China, over 13,000 homes were destroyed and over 150,000 homes damaged. At 200000Z, JTWC issued its final warning on Tropical Storm Warren with the system well inland and northwest of Hong Kong. The remnants of Warren's low-level circulation continued tracking across southern China for another day before merging with a weak summer front, enhancing cloudiness and precipitation.


Figure 3-06-4: Tropical Storm Warren after making landfall (190745Z July NOAA visual imagery).


## TROPICAL STORM AGNES (07W)

Agnes was of note in several respects. It was the second of only two tropical cyclones to develop in July, a month that normally averages five. It played a part in the major shift of the synoptic pattern in the western North Pacific in the latter part of July and later became a small, but vigorous, extratropical cyclone.

As Typhoon Warren (06W) moved into southern China on 19 July, lower tropospheric ridging and fair weather prevailed over eastern China and the Philippine Sea. Once Warren
(06W) dissipated, the monsoon trough, instead of maintaining its climatological position across the northern Philippine Islands and southern Philippine Sea, remained over Asia. Southeast of Japan, an area of disturbed weather with lower than normal sea-level pressures and enhanced convection generated in the lower tropospheric troughing (Figure 3-07-1). By 27 July, a closed circulation and supporting southwesterly monsoonal flow at 700 mb had developed along this trough (Figure 3-07-2).


Figure 3-07-1. 700 mb analysis at 190000 Z July with Typhoon Warren ( 06 W ) (point W) and troughing (line X to X ) southeast of Japan.


Figure 3-07-2. 270000 Z July 700 mb analysis with a closed cyclonic circulation (at point X ) and deep west-southwesterly monsoonal flow.

At first, Agnes appeared as a monsoon depression with a low-level cyclonic circulation and deep convection displaced to the south by vertical shear aloft. This was discussed in the Significant Tropical Weather Advisory at 260600 Z . The suspect area drifted southeastward in the trough and had a poor potential for development into a significant tropical cyclone due to the unfavorable vertical wind shear of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$.) When the vertical wind shear dropped to $20 \mathrm{kt}(10 \mathrm{~m} / \mathrm{sec})$ at 280600 Z and satellite imagery indicated increased upper-level outflow and deep convection, the system's potential for significant development was upgraded to "fair". A Tropical Cyclone Formation Alert was issued at 281500 Z based on a satellite intensity estimate of 25 kt (13 $\mathrm{m} / \mathrm{sec}$ ). The continued increase in the system's deep convection and overall organization led to the first waming at 290000 Z .

Initially, Agnes was forecast to intensify, separate from the monsoon trough and track to the northwest. However, the monsoon trough, which was farther north than normal, merged with a mid-latitude low pressure system to the northeast of Japan. Agnes followed the path of least resistance and accelerated to the northnortheast along this trough axis. The tropical cyclone was upgraded to tropical storm intensity at 291200 Z and reached its peak intensity (Figure 3-07-3) of 45 kt ( $23 \mathrm{~m} / \mathrm{sec}$ ) 12hours later. The loss of persistent central convection at 301200 Z resulted in the issuance of the final waming at 301800 Z .

Unfavorable vertical shear from the strong southwesterlies aloft increased and Agnes (Figure 3-07-4) accelerated northnortheastward at more than 30 kt ( $56 \mathrm{~km} / \mathrm{hr}$ ). Although the system appeared to be

extratropical at 310000 Z , ships in the right front quadrant (relative to the forward motion) of this hybrid system reported maximum sustained surface winds of $60 \mathrm{kt}(31 \mathrm{~m} / \mathrm{sec})$ at 310600 Z and $50 \mathrm{kt}(26 \mathrm{~m} / \mathrm{sec})$ at 311200 Z . Herbert and Poteat (1975) address this type of system, or subtropical cyclone, where translational speeds greater than $20 \mathrm{kt}(37 \mathrm{~km} / \mathrm{hr}$ ) are added to the
intensity estimate determined from the cloud signature. Even though the final warning was issued when the system was at 32 degrees North, Agnes stubbornly maintained some of its tropical characteristics well into the midlatitudes. No reports of casualties or damages were received.


Figure 3-07-4. The remnants of Tropical Storm Agnes during transition into an extratropical system ( 301829 Z July NOAA infrared imagery).


## TROPICAL STORM BILL (08W)

Tropical Storm Bill was the first of five tropical cyclones to develop during the month of August. It formed in the Philippine Sea, brushed by the island of Okinawa and reached a peak intensity of 45 kt ( $23 \mathrm{~m} / \mathrm{sec}$ ) before making landfall near Shanghai. Bill remained well organized over China, causing widespread destruction and loss of life. It exemplifies the serious impact that is possible from a system of
tropical storm intensity that doesn't dissipate rapidly after moving over land.

On the first day of August the monsoon trough (Figure 3-08-1) stretched from the Gulf of Tonkin across the South China Sea to the Luzon Strait and extended northeastward along the Japanese Islands. By 5 August (Figure 3-08-2) this synoptic feature had undergone a


Figure 3-08-1. The 010000Z August 925 mb winds and streamlines show the monsoon trough extending across the South China Sea and northeastward.


Figure 3-08-2. The 925 mb winds and streamlines for 051200 Z , when compared with Figure 3-08-1, reveal the shift of the monsoon trough into China and its abrupt termination near the island of Okinawa, Japan.
major readjustment. Now it was oriented west-to-east over southern China and terminated abruptly near the island of Okinawa, Japan. Excess relative cyclonic vorticity, low-level convergence and associated enhanced convection persisted near the eastern end of the trough. With the environment favorable for tropical cyclogenesis, Bill rapidly consolidated
and at 050430Z the first Tropical Cyclone Formation Alert was issued on the system. Early on 6 August the cloudiness associated with the tropical disturbance (Figure 3-08-3) separated from the maximum cloud zone. As the tropical cyclone moved away from the axis of the monsoon trough, it came under the steering influence of the subtropical ridge to the


Figure 3-08-3. Bill as a tropical disturbance south of the island of Okinawa ( 060009 Z August DMSP visual imagery).
northeast. The track, which was initially to the north, became northwestward. Consolidation continued, but the pace slowed, and at 060330 Z a second Alert was issued. The center of the low-level circulation passed just southwest of the southern tip of Okinawa at 061500 Z . Kadena Air Base (WMO 47931) reported a minimum sea-level pressure of 997 mb , sustained surface winds of $20 \mathrm{kt}(10 \mathrm{~m} / \mathrm{sec})$ and a peak surface gust of $40 \mathrm{kt}(21 \mathrm{~m} / \mathrm{sec})$. Earlier,

Kadena's $061200 Z$ upper-air sounding indicated a layer between the 875 mb and 550 mb levels of 35 to 40 kt ( 18 to $21 \mathrm{~m} / \mathrm{sec}$ ) winds, which supported the peak gusts.

Slow intensification progressed and the first warning ( 070000 Z ) on Tropical Storm Bill followed a satellite intensity estimate of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) (Figure 3-08-4). At 070600Z, Bill reached a peak intensity of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$ and


Figure 3-08-4. Bill, shortly after being upgraded to tropical storm intensity ( 070131 Z August DMSP visual imagery).
sustained it until making landfall $120 \mathrm{~nm}(222$ km ) south of Shanghai.

Once overland, instead of rapidly dissipating, Bill retained its convection and organization (Figure 3-08-5). At 080000Z, the intensity was downgraded to a tropical depression and the final warning issued. The remains of this cloud system tracked to the northwest and could still be identified on satellite imagery two days later.

Bill's slow dissipation resulted in extensive damage and loss of life in China. Torrential rains led to widespread flooding and local topographic effects may have produced wind gusts as high as $70 \mathrm{kt}(35 \mathrm{~m} / \mathrm{sec})$. When it was all over, at least 110 people had perished and numerous bridges, dams and watercraft were damaged or destroyed. News releases from China reported "the worst economic loss from a storm in 30 years."


Figure 3-08-5. Tropical Storm Bill continued to be a very well organized system, despite having already been over land for several hours ( 072156 Z August DMSP visual imagery).

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## TROPICAL STORM CLARA (09W)

Tropical Storm Clara was the second of five significant tropical cyclones to develop during August. Although hindered by vertical wind shear, Clara proved to be very persistent. Even after the central deep convection was stripped away and the final warning issued, the residual cyclonic vorticity could be identified four days later as a spiral of low-level stratocumulus on the satellite imagery.

Clara was originally detected on satellite imagery as an area of weakly organized convection about $540 \mathrm{~nm}(1,000 \mathrm{~km})$ north of

Wake Island and it was mentioned on the Significant Tropical Weather Advisory on 070600Z August. Easterly flow along the southern edge of the subtropical ridge steered the disturbance westward. Over the next 16hours the convection persisted and became more organized. The satellite intensity analysis at 072126 Z indicated a shearing-type cloud pattern with sustained surface winds of 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) and an exposed low-level circulation defined by cumulus lines spiraling inwards. The first Tropical Cyclone Formation Alert followed at 072330Z. A second Alert was issued at


Figure 3-09-1. Tropical Storm Clara, shorlly after the first warning (102201Z August NOAA visual imagery).

081600 Z for procedural, not meteorological, reasons. However, the disturbance experienced stronger vertical wind shear, and its deep convection was displaced more than $90 \mathrm{~nm}(167$ km ) north of the circulation center. Based on this, the second Alert was canceled at 090600 Z .

Increased central convection on 10 August and an evaluation of satellite imagery and synoptic data, which indicated 20 to 30 kt ( 10 to $15 \mathrm{~m} / \mathrm{sec}$ ) sustained surface winds, resulted in a third Alert at 100100 Z . As the disturbance continued to move westward, its upper-level outflow improved and its central convection increased. The first warning on

Tropical Storm Clara was issued at 101915Z. Sustained surface winds at that time were estimated to be 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) (Figure 3-091). Due to the close proximity to regions of stronger vertical wind shear, Clara was not forecast to intensify beyond the 40 to 45 kt ( 21 to $23 \mathrm{~m} / \mathrm{sec}$ ) range.

As ridging to the west increased, Clara made a small counterclockwise loop while maintaining its 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) intensity. Clara reached a peak intensity of 45 kt ( 23 $\mathrm{m} / \mathrm{sec}$ ) at 110000 Z and began moving northeastward (Figure 3-09-2) at 110600Z. Analysis of satellite imagery on 12 August


Figure 3-09-2. Clara near peak intensity (110513Z August NOAA visual imagery).
indicated Clara had weakened to 30 kt (15 $\mathrm{m} / \mathrm{sec}$ ) sustained surface winds as its main convection was displaced southeast of the lowlevel circulation center. The final warning on the system was issued at 120000 Z . The remnants of Clara tracked northward,
northwestward and, then again, northward for the next four days in response to steering flow around the subtropical ridge to its east. By 16 August, the remains of the system were no longer identifiable. No damage reports relating to Clara were received.


## TYPHOON DOYLE (10W)

Typhoon Doyle was the third of five tropical cyclones and the first of two typhoons to occur in the western North Pacific during August. In keeping with the aclimatological trend for the month, Doyle was no exception. It formed north of 20 degrees North latitude, moved south-southwestward and looped before tracking to the northeast.

On 12 August, as Tropical Storm Clara (09W) moved northward and weakened, a portion of the Central Pacific high moved in from the east to fill the void. The high pushed southwestward across the dateline. During the adjustment process, a low-level cloud vortex
appeared along the leading edge of the flow and west of the remnants of Hurricane Fabio (08E). The unusual south-southwestward track of this vortex appeared to be related to the steering provided by a lower-tropospheric anticyclone to the north. The Significant Tropical Weather Advisory at 130600 Z mentioned the vortex when deep convection became associated with the low-level cyclonic circulation. Throughout the night of the 13 th and early morning hours of the 14th, the convection became centralized. This prompted a Tropical Cyclone Formation Alert at 140130Z. Visual and infrared imagery (Figure 3-10-1) at 140743 Z implied a well developed low-level inflow. About this time


Figure 3-10-1. Satellite imagery of the suspect area (Doyle). Low-level circulation is implied by the lines of cumulus surrounding the center. The left picture is visual and the right is enhanced infrared (140743Z August DMSP visual and infrared imagery).
the system (Doyle) executed a small counterclockwise loop and began tracking westnorthwestward. The potential for the system to develop remained good and a second Tropical Cyclone Formation Alert was issued at 150130Z.

The satellite intensity estimate of 40 kt ( $21 \mathrm{~m} / \mathrm{sec}$ ) maximum surface winds was followed at 151200 Z by the first warning, when the system was $96 \mathrm{~nm}(178 \mathrm{~km})$ east-northeast of Wake Island. For the 24 -hour period from

151800 Z to 161800 Z (Figure $3-10-2$ ) the intensity increased from 50 to 115 kt ( 26 to 59 $\mathrm{m} / \mathrm{sec}$ ). This was the equivalent of a sixty millibar pressure fall and met the criteria (Dunnavan, 1981) for explosive deepening. Although Doyle, which was close to becoming a compact typhoon, passed $55 \mathrm{~nm}(102 \mathrm{~km})$ north of Wake Island (WMO 91245) at 152100 Z , the island only experienced gusts to $40 \mathrm{kt}(21 \mathrm{~m} / \mathrm{sec})$. As a result, the low-lying island incurred only minor damage.


Figure 3-10-2. Time/intensity comparison of satellite intensity estimates for the Guam satellite site and the best track. Notice the extended period of explosive deepening of $30 \mathrm{mb} / 12$-hours from 151800 Z (point A) to 161800 Z (point B).

Nearing the western periphery of the mid-level subtropical ridge, Doyle peaked in intensity at 161800Z (Figure 3-10-3) and assumed a northward track at 170000 Z . To complicate the track forecasts, a TUTT cell stalled, then appeared to dissipate to the north. The main effect of the TUTT cell was to shield the tropical cyclone from strong mid- and upper-level westerlies. As a result, expected acceleration along the track didn't take place and Doyle's speed was never greater than 12 kt
( $22 \mathrm{~km} / \mathrm{hr}$ ). Doyle's track followed lower pressures and heights present between the subtropical ridge to the southeast and another high cell to the northwest centered near $42^{\circ}$ North latitude.

After gradual weakening, Doyle was forecast to become extratropical as the convection began to move into the northeast semicircle at 180900 Z. However, some central convection remained until 22 August. Doyle


Figure 3-10-3. Satellite imagery of Typhoon Doyle at maximum intensity. The medium sized eye ranged from 20 to 30 nm ( 37 to 57 km ) in diameter ( 1622122 August DMSP visual imagery).
dissipated over colder water as the system slowed and moved northeastward.

Doyle fell into the "other" track category for several reasons: rapid south-southwestward movement for 24 -hours, looping and interaction with a TUTT cell while at typhoon intensity. Normally, tropical cyclone objective forecast guidance does not perform well for aclimatic
systems. Figure 3-10-4 compares the final best track and the performance of the two best performing aids, CSUM and OTCM, up to the major track change at 170000 Z . Although OTCM had the lowest mean forecast errors at 72-hours of all the aids, it was slow in catching the major track change to the northeast. CSUM had the same problem predicting this track change.


Figure 3-10-4. Graphical display of selected objective aids performance (OTCM and CSUM).



## TROPICAL STORM ELSIE (11W) AND TYPHOON FABIAN (12W)

Elsie and Fabian were the last of five significant tropical cyclones to develop in August. They both initially displayed erratic movement as immature systems within the larger monsoon trough, interacted as a binary pair (Brand, 1968) and, later, became extratropical.

Both systems were first detected on 24 August and mentioned on the Significant Tropical Weather Advisory as areas of persistent convection in the monsoon trough, that was anomalously far north and extended southeastward from Japan. Typical of formative vortices in the larger monsoon trough, their tracks were less than straight forward. Fabian's track wobbled around until 28 August, when it
sped off towards the east on a possible collision course with Elsie. In the interim, Fabian required three Tropical Cyclone Formation Alerts: the first at 271430 Z based on a 30 kt ( 15 $\mathrm{m} / \mathrm{sec}$ ) satellite estimate of surface winds, a second at 281430 Z and the third at 291230 Z when the tropical disturbance retained its potential for development and failed to weaken. Finally, Fabian's convection consolidated (Figure 3-11/12-1), resulting in a satellite intensity estimate of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$ and the first warning at 300000 Z .

In the meantime, Elsie, which was located at the southeastern end of the monsoon trough, was in an area of stronger low-level convergence and associated cloudiness, but


Figure 3-11/12-1. Fabian at tropical storm intensity is on the left. Tropical Storm Elsie is on the right (301140Z August DMSP visual imagery).
suffered from lack of central convection. Better organization at 280600 Z with estimated surface winds of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ from the satellite analysis led to a Tropical Cyclone Formation Alert at 280900 Z . The first warning followed at 281200 Z based on satellite intensity estimates of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec}$ ). At this point Elsie's course, which had been southeastward, abruptly changed to northeastward. As a consequence, the monsoon trough also began to shift northward. This reorientation of the trough axis is the most probable explanation for Fabian's unusual track to the east.

Figure 3-11/12-3. For the period 281200Z August to 010000 Z September the plotted positions of Elsie and Fabian relative to the midpoint (reference Figure 3-11/12-2) at point C describe the motion of this binary system.

At 281200 Z , Elsie and Fabian were separated by $1070 \mathrm{~nm}(1982 \mathrm{~km})$. The two systems closed on each other until only a 400 nm ( 741 km ) separation remained three days later. Figure 3-11/12-2 shows the two best tracks with the locus of midpoints (point A to point B) between each system at six hourly intervals. Plotting the locus of midpoints at one point (C) in Figure 3-11/12-3 and replotting the relative locations of Elsie and Fabian, the cyclonic motion of the binaries (Dong and Neumann, 1983) is captured.

Figure 3-11/12-2. Tracks for Elsie and Fabian for the period 281200Z August to 010000 Z September. For each respective six hourly position for the two systems a midpoint is plotted. The locus of these midpoints stretches from point $\mathbf{A}$ to point $\mathbf{B}$.


While this interaction was taking place, Elsie's convection weakened and the forecast intensity estimate dropped below 35 kt ( 18 $\mathrm{m} / \mathrm{sec}$ ). This prompted an amendment of the 291800 Z warning, which had called for dissipation in 48-hours over water with remarks that "signs for regeneration would be monitored." It appears that Elsie stabilized as a weak tropical storm and assumed a track to the northwest. The remnants of Elsie were not expected to flare up again due to the binary interaction with Fabian, which reduced the separation between the two and increased the vertical wind shear across Elsie from Fabian's
outflow. However, the central convection returned (Figure 3-11/12-4) and a Tropical Cyclone Formation Alert - the second for Elsie was issued at 302300Z. Elsie's compact reorganization resisted the unfavorable conditions aloft and the Alert was almost immediately upgraded to a regenerated tropical storm at 310000 Z .

Once Elsie was past the closest point of approach to Fabian, it changed course to the northeast, accelerated and rapidly became extratropical. The final warning - the second for Elsie - was issued at 311800 Z . In contrast,


Figure 3-11/12-4. The remnants of Elsie, on the right, undergo a rapid reorganization. The second Tropical Cyclone Formation Alert was issued upon receipt of this picture. Fabian is to the left of Elsie (302240Z August DMSP visual imagery).

Fabian intensified during this encounter and reached typhoon intensity (Figure 3-11/12-5) at 010600 Z just prior to making an abrupt track change to the north six hours later. Typhoon Fabian reached a peak intensity of 75 kt ( 39 $\mathrm{m} / \mathrm{sec}$ ) at 020600 Z . By 030000 Z , the onset of
rapid acceleration and stronger upper-level westerlies led to a loss of convective organization. Fabian was downgraded to tropical storm intensity and finalled at 030600 Z . No reports of damage or loss of life were received for these two tropical cyclones.


Figure 3-11/12-5. Fabian nearing peak intensity. Elsie has moved off the picture to the northeast of Fabian (312220Z:August DMSP visual imagery).

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## TROPICAL STORM GAY (13W)

Gay was a short-lived (two day) tropical cyclone. Only six warnings were issued on the system before it dissipated over water. The disturbance that eventually became Gay formed in the monsoon trough $420 \mathrm{~nm}(778 \mathrm{~km}$ ) east of Okinawa. It was first mentioned on the Significant Tropical Weather Advisory at 020600 Z and rated as having a "fair" potential for significant development. The system's convection rapidly increased in amount and became better organized when its low-level circulation moved beneath the divergent area ahead of an approaching upper-level trough. After synoptic data ( 020600 Z ) indicated the disturbance had sustained surface winds in the

20 to 30 kt ( 10 to $15 \mathrm{~m} / \mathrm{sec}$ ) range and satellite wind estimates of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ), a Tropical Cyclone Formation Alert was issued at 021040 Z . The first warning ( 021800 Z ) on Tropical Storm Gay followed. Gay moved northeastward at speeds of 14 kt ( $26 \mathrm{~km} / \mathrm{hr}$ ), or more, and reached a peak intensity of 45 kt ( 23 $\mathrm{m} / \mathrm{sec}$ ) sustained surface winds (Figure 3-13-1). After peaking, the tropical cyclone tracked into an environment of strong vertical wind shear, which separated the system's deep convection from its low-level circulation. The final warning was issued on Gay at 040000 Z while it was dissipating over water 290 nm ( 537 km ) southeast of Tokyo.


Figure 3-13-1. Tropical Storm Gay at peak intensity (030605Z September NOAA visual imagery).


## TYPHOON HAL (14W)

Typhoon Hal was the second of eight significant tropical cyclones and the first of two typhoons to form in the western North Pacific during September. Hal combined with Typhoon Uleki (01C), Tropical Storm Irma (15W), and later with Tropical Storm Jeff (16W) to create the first three-storm situations of 1988 in the western North Pacific.

On 5 September the remnants of Gay (13W) dissipated east of Japan. In the central North Pacific Uleki (01C) churned westnorthwestward from Hawaii and a large tropical upper-tropospheric trough (TUTT) low was situated northwest of Wake Island. Hal formed just west of Wake Island as a tropical disturbance induced by this TUTT low. At 050600 Z , the system was first mentioned on the Significant Tropical Weather Advisory. Over
the next two days, the disturbance moved westward along the southern side of the subtropical ridge and became more organized. This growth prompted a Tropical Cyclone Formation Alert at 070430Z. Hal continued to organize. At 080000 Z , the Alert was superseded by the first warning on Tropical Depression 14W, then upgraded ( 081200 Z ) to Tropical Storm Hal (Figure 3-14-1), when satellite intensity analysis indicated sustained surface winds of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ). Initially, Hal tracked west-southwestward, but eventually settled into a west-northwestward track as the subtropical ridge to its north weakened slightly. Hal continued to intensify and, at 101800Z, reached typhoon intensity.

Earlier at 101200 Z, when Hal was 120 $\mathrm{nm}(222 \mathrm{~km})$ northeast of Maug in the northern


Figure 3-14-1. Hal shortly after being upgraded to a tropical storm ( 082128 Z September NOAA visual imagery).

Marianas, the tropical cyclone started to decelerate and track to the southwest in response to stronger ridging to its north and west. After Typhoon Hal reached its peak intensity of $105 \mathrm{kt}(54 \mathrm{~m} / \mathrm{sec})$ at 111200 Z , it continued onward and passed over Maug, which is uninhabited, at 111800 Z . On Guam (WMO 91212), $395 \mathrm{~nm}(732 \mathrm{~km}$ ) to the south, the enhanced southwesterly inflow into Hal brought brisk surface winds with gusts to 40 kt ( 20 $\mathrm{m} / \mathrm{sec}$ ). Power outages and minor property damage were reported on the islands of Guam
and Saipan.
With a mid-latitude trough creating lower pressure-heights in the subtropical ridge north of the typhoon, Hal's direction of track changed to the north-northwest. Japan braced for the possibility of being affected by three tropical cyclones: Typhoon Hal (Figure 3-14-2), plus Tropical Storms Irma (15W) and Jeff (16W), which had developed southeast and southwest, respectively. During the next three days, Hal weakened, developed a large 60 nm


Figure 3-14-2. Typhoon Hal, at peak intensity. Also visible are Tropical Storms Irma (15W) and Uleki (01C) (122141Z September NOAA visual imagery).
(111 km) diameter eye and continued moving north-northwestward. At 150000 Z , Hal (Figure 3-14-3) approached 32 degrees North latitude, started to recurve and accelerate. Typhoon Hal further weakened to $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ and made its closest point of approach - 195 nm ( 361 km ) - to Tokyo, Japan at 151800 Z . High surf caused several deaths and injuries along the coastal areas near Tokyo.

As Hal moved off to the northeast, its central convection was stripped away from its low-level circulation center by strong midlatitude westerlies. When the final warning was issued at 170000 Z , the system had weakened to 45 kt ( $23 \mathrm{~m} / \mathrm{sec}$ ), increased forward speed to 32 kt ( $59 \mathrm{~km} / \mathrm{hr}$ ) and was extratropical.


Figure 3-14-3. Typhoon Hal, Tropical Storms Irma (15W) and Jeff (16W) (150936Z September NOAA infrared imagery).


Uleki was only the third hurricane (Sarah (1967) and Peke (1987) were the previous two) in the past thirty years to form in the eastern North Pacific Ocean and cross the international dateline while in a warning status. Uleki tracked over $3,300 \mathrm{~nm}(6,105 \mathrm{~km})$ during its eighteen day life span.

Uleki was first detected by the Central Pacific Hurricane Center (CPHC) at 281800 Z August. During the next four days, Uleki tracked westward and intensified. CPHC's warnings were disseminated to military customers by the Naval Western Oceanography Center (NWOC). At 291800Z, Uleki had reached tropical storm intensity and was upgraded to a hurricane at 310000 Z . The tropical cyclone slowed its forward motion and
continued to intensify until 2 September, when it reached a peak intensity of $110 \mathrm{kt}(57 \mathrm{~m} / \mathrm{sec}$ ). As Uleki approached the Hawaiian Islands, weather reconnaissance aircraft joined with satellite reconnaissance to watch the hurricane. At peak intensity, the direction of movement changed from west-northwestward to northward. Uleki headed directly towards the island of Oahu. The hurricane approached to within $270 \mathrm{~nm}(500 \mathrm{~km}$ ) of Honolulu at 040000 Z before changing course to the westnorthwest and accelerating. The tropical cyclone began a weakening trend as it entered a shearing environment, and the upper-level outflow in the western semicircle became restricted. Uleki continued to move westnorthwestward and approached the international dateline (Figure 3-01C-1). It appeared that


Figure 3-01C-1. Hurricane Uleki heads towards the international dateline. Note the distinct shadow on the eye wall caused by the low sun angle. The Hawaiian Islands are visible to the east of Uleki. Photo courtesy of the National Weather Service Forecast Office, Honolulu, Hawaii (071846Z September GOES West visual imagery).


Figure 3-01C-2. Uleki, shortly after the final warning (130839Z September NOAA infrared imagery).

Uleki would be exiting CPHC and NWOC's area of responsibility, and entering JTWC's area of responsibility after the 080000 Z warning.

Warning responsibility was transferred for the 080600 Z warning and JTWC issued its first warning on Uleki. The system was redesignated Typhoon Uleki. At this time the tropical cyclone had an intensity of 90 kt ( 46 $\mathrm{m} / \mathrm{sec}$ ). Uleki pressed onward to the westnorthwest along the southern edge of a subtropical ridge, and gradually slowed. At 100600 Z , the speed of movement had dropped from $15 \mathrm{kt}(28 \mathrm{~km} / \mathrm{hr})$ to $6 \mathrm{kt}(11 \mathrm{~km} / \mathrm{hr})$. The typhoon had entered the weak 700 mb steering flow in an area between two anticyclones in the subtropical ridge. With a mid-latitude trough approaching from the west, Uleki was forecast
to recurve during the next 24 - to 48 -hours. The trough caused the tropical cyclone to "step climb" to the north-northwest, but was not able to bring about recurvature. Uleki returned to a northwestward track and weakened in response to increased vertical wind shear and entrainment of low-level cooler air. At 130000Z, strong vertical wind shear associated with a second trough caused the tropical cyclone to weaken rapidly and be downgraded to a tropical storm. Satellite imagery showed a long, narrow plume of cirrus streaming from Uleki to the northeast. The final warning was issued at 130600 Z (Figure 3-01C-2) and at 140000Z, all of Uleki's deep convection had been sheared away to the northeast. The low-level circulation center persisted over water until 15 September (Figure 3-01C-3).


Figure 3-01C-3. Uleki dissipating over water (142247Z September DMSP infrared imagery).

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## TROPICAL STORM IRMA (15W) AND TROPICAL STORM JEFF (16W)

Irma and Jeff, circulations spawned by enhanced inflow into Typhoon Hal (14W), never achieved typhoon intensity. Both were part of multiple tropical cyclone outbreaks of 12 to 16 September and sheared away when Hal (14W) moved northward through the subtropical ridge.

By the second week of September Hal (14W), which started earlier as a tropical uppertropospheric trough (TUTT) induced system, had matured south of the subtropical ridge and developed a large supporting low-level
southwesterly inflow. (This inflow was separated from, and not the normal eastward extension of, the Asian southwest monsoon.) Jeff formed at the extreme western end and Irma at the extreme eastern end of this southwesterly inflow.

Compared to Jeff, Irma got a head start in central convection and was first noted on the Significant Tropical Weather Advisory at 110000 Z . Although the outflow from Hal (14W) to the west streamed across the area, Irma persisted. This increased convective


Figure 3-15/16-1. Irma about four hours before the first warning (111946Z September DMSP visual imagery).
activity (Figure 3-15/16-1) required a Tropical Cyclone Formation Alert at 110920 Z and a first warning, based on a satellite intensity estimate of $40 \mathrm{kt}(21 \mathrm{~m} / \mathrm{sec})$, at 120000 Z .

Meanwhile, Jeff consolidated and was included on the 120600 Z Advisory. Organization of the convection continued and an Alert followed at 121200 Z . Another Alert was issued at 131030 Z before a $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ satellite
surface wind estimate precipitated the first warning at 140000 Z . Jeff's relatively slow development was related to its upper-level outflow being severely restricted in the northeast quadrant by the larger outflow from Hal (14W) to the northeast. This shear in the vertical, in fact, confined Jeff's central convection to the southern half of its low-level circulation for the lifetime of the system.


Figure 3-15/16-2. Jeff at peak intensity. Hal (14W) is at top right. Jeff's outflow is severely restricted to the north and east (142051Z September DMSP visual imagery).

Briefly, Jeff (Figure 3-15/16-2) reached a peak intensity of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$ on 14 September before gradually weakening. In contrast, Ima, which was aided by troughing in Hal's (14W) upper-level outflow, attained 55 kt ( $28 \mathrm{~m} / \mathrm{sec}$ ) at 131200 Z half a day earlier and maintained that intensity through 0000 Z on 15 September. Later, both Jeff and Irma were finalled within six hours of each other - Irma (Figure 3-15/16-3) at 151800 Z and Jeff at 160000 Z - as Hal (14W) moved northward through the subtropical ridge.

The relationship between the three tropical cyclones and the subtropical ridge is of interest. Earlier on 12 September, as Hal (14W) tracked to the north, the subtropical ridge segmented into two cells. These high pressure cells worked to narrow and restrict the low latitude inflow into Hal (14W). This appears to have affected the relative positions between Jeff, Irma and Hal (14W) (Figure 3-15/16-4). Initially, at 130000 Z , the baseline (from A to B) between Jeff and Irma is relatively long compared to the height (from C to D ) between Hal (14W) and the baseline. However, at


Figure 3-15/16-3. The low-level circulation is all that remains of Irma ( 152228 Z September DMSP visual imagery).

160000 Z the baseline has decreased to almost half its previous length and the height has more than doubled. These triangles (Figure 3-15/164) suggest a subtle tertiary interaction between Hal (14W) and the two smaller tropical cyclones in an almost non-divergent flow.

This multiple cyclone activity resulted in three sets of warnings being issued from 140000 Z to 160000 Z. Earlier, from $120000 Z$ to 130600 Z , Uleki ( 01 C ), Hal (14W) and Irma had required three sets of warnings. No damage reports were received for Jeff and Irma.


Figure 3-15/16-4. The tracks of Hal (14W), Irma and Jeff. Compare the length of the baseline (from A to B) between Jeff and Irma and the height, which is measured along the vertical ( $C$ to $D$ ) from the baseline to Hal (14W), at 130000 Z with the second triangle at 160000 Z . Note the relative adjustment of the three as Hal (14W) moves to the north.

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## TROPICAL STORM KIT (17W)

Tropical Storm Kit was the sixth of eight significant tropical cyclones in September. It was a "straight runner" to the northwest and made landfall on the south coast of China. Kit caused loss of life and significant property damage in southeastern China.

The tropical cyclone was first detected
on satellite imagery on 18 September 300 nm ( 556 km ) east of Manila. The disturbance rapidly developed in the eastward extension of the monsoon trough and immediately became the subject of a Tropical Cyclone Formation Alert at $182230 Z$ (Figure 3-17-1). Increased deep convection in the banding feature, improved outflow aloft, plus a satellite intensity


Figure 3-17-1. Kit as a tropical disturbance ( $190051 Z$ September DMSP visual imagery).
estimate of sustained 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) surface winds, dictated the upgrade to Tropical Depression 17W at 090600Z.

Even though Kit tracked across the northern tip of Luzon, it continued to intensify. At 191800Z, another upgrade was needed this time to tropical storm intensity. After being over land for six hours, it once again moved over open waters. The system developed a strong low-level inflow from the southwest and improved its upper-level outflow to the
southeast through southwest (Figure 3-17-2). A day later, at 210600 Z , Kit reached its peak intensity of $60 \mathrm{kt}(31 \mathrm{~m} / \mathrm{sec})$.

After the intensity peaked, the tropical storm approached the coast of southern China and weakened. The final warning was issued on Kit at 220000 , when it made landfall 120 nm ( 222 km ) northeast of Hong Kong. Press releases from China indicated widespread flooding, loss of electrical power and at least three lives lost in the Guangdong province.


Figure 3-17-2. Kit at peak intensity (to the left) and Tropical Storm Lee (18W) (to the lower right of the picture) (211003Z September DMSP infrared imagery).

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## TROPICAL STORM LEE (18W)

Tropical Storm Lee was the seventh of eight significant tropical cyclones to occur during September. Lee had a formative period of over four days, and was tracked over 1300 $\mathrm{nm}(2408 \mathrm{~km})$ as an identifiable area of convection before the first warning was issued.

On 16 September Tropical Storms Irma (15W) and Jeff (16W) had just been finalled, Typhoon Hal (14W) was recurving east of Japan and a new area of persistent convection was mentioned on the Significant Tropical Weather Advisory at 160600 Z . This persistent convection was tropical upper-tropospheric trough (TUTT) induced (Sadler, 1979) and was
superimposed on the broad low-level easterly flow. A steady westward movement was noted for the next four days, during which time there was little change in the poorly organized convection. At 200600 Z the disturbance was upgraded to a "fair" suspect area due to improved organization. A Tropical Cyclone Formation Alert followed at 201730 Z after the system continued to show improved organization and intensification to 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) sustained surface winds, based on a satellite analysis estimate. Another estimate of 35 kt (18 $\mathrm{m} / \mathrm{sec}$ ) sustained surface winds followed at 210000 Z and prompted the first warning on Tropical Storm Lee (Figure 3-18-1). Lee


Figure 3-18-1. Tropical Storm Lee just after the first warning. Satellite intensity analysis indicated a T-number of 2.5 , corresponding to sustained surface winds of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) ( 210012 Z September DMSP visual imagery).


Figure 3-18-2. Shortly before recurvature, Lee shows the effects of increased vertical wind shear. The low-level circulation center is partially exposed to the northeast of the centeral dense overcast ( $232304 Z$ September DMSP visual imagery).
tracked west-northwestward along the southwestern side of the subtropical ridge for the next 24 -hours, and acquired its maximum intensity of $55 \mathrm{kt}(28 \mathrm{~m} / \mathrm{sec})$ at 211800 Z . The forecast was for Lee to continue on its northwestward track around the periphery of the 700 mb subtropical ridge.

Visual satellite imagery (Figure 3-18-2) on 23 September showed a partially exposed low-level circulation center, as Lee encountered increasing vertical wind shear. When night arrived, Lee's poorly defined deep convection provided targets for remote sensing - bright cold tops on the satellite infrared and rain echoes for the radar. Beginning at 231100Z, radar position reports from Ishigaki Jima (WMO 47918) confirmed the movement of the rain echoes to the northwest, which paralleled

Kit's (17W) earlier track into southeastern China. However, remarks the following morning on the relocated 240000 Z warning summed it up: "Visual satellite pictures indicate Lee has an exposed low-level circulation that has been moving northeastward. Tropical Storm Lee has recurved earlier than expected and should now pass about 55 nm east of Okinawa."

This forecast was accurate and Lee's closest point of approach was $45 \mathrm{~nm}(83 \mathrm{~km})$ southeast of the island of Okinawa at 240400 Z . Lee started to lose its convective organization and showed signs of becoming extratropical at 240600 Z . The final warning was issued at 241200Z. Lee became an extratropical low at 242100 Z and continued moving rapidly northeastward.

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## TROPICAL STORM MAMIE (19W)

Mamie was the second significant tropical cyclone of 1988 to develop in the South China Sea. It formed in tandem with Tropical Storm Kit (17W) in the monsoon tough and was slow to develop. Typical of monsoon depressions, Mamie proved to be a particularly difficult system to locate and forecast. Large fluctuations in its central convection within the larger synoptic scale trough contributed to this difficulty.

The disturbance that would become Tropical Storm Mamie formed in September, in the monsoon trough, 600 nm ( 1111 km ) westnorthwest of Kit (17W). Kit (17W) was at the eastern end of the monsoon trough. Mamie was first mentioned on the Significant Tropical Weather Advisory at 190600Z. During the next 18-hours, the disturbance moved southwestward
at $10 \mathrm{kt}(19 \mathrm{~km} / \mathrm{hr})$, most probably in response to binary interaction with Kit (17W). As Mamie became better organized, its deep convection began the first of several flare-ups. This prompted the first Tropical Cyclone Formation Alert at 200230Z, after satellite intensity analysis indicated 30 kt ( $15 \mathrm{~m} / \mathrm{sec}$ ) sustained surface winds. The system (Figure 3-19-1) continued southwestward until 210000 Z , when it made a sharp turn to the east. The persistent central convection and potential for development required a second Alert at 210030Z.

Later, at 211800 Z, the disturbance was headed northeastward along the trough axis the opposite direction from its initial track, two days earlier. Apparently Mamie's track to the northeast was the result of increased low-level


Figure 3-19-1. Mamie, the subject of a Tropical Cyclone Formation Alert, is on the left. To the right of Mamie is Tropical Storm Kit (17W) (200802Z September NOAA visual imagery).
southwesterly inflow into Kit (17W), which was now over water between Luzon and the southeastern coast of China. Mamie maintained its overall convective organization and a third Alert was issued at 220030Z. Based on a ship report of southeasterly surface winds of 40 $\mathrm{kt}(20 \mathrm{~m} / \mathrm{sec})$ and a minimum sea-level pressure of 991 mb , plus a satellite intensity analysis of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ), the first warning on Tropical Storm Mamie followed at 220600Z. Mamie
continued up the monsoon trough axis towards Kit (17W), which had just made landfall in southeastern China. As vertical wind shear increased aloft over Mamie, JTWC issued its final warning at 230000 Z . Mamie's remnants, and associated gales, then moved northnortheastward and dissipated on the coast of China northeast of Hong Kong. No reports of damage or loss of life were received.

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## SUPER TYPHOON NELSON (20W)

Nelson was the first significant tropical cyclone of October and the only super typhoon of 1988. It developed in the Philippine Sea in the monsoon trough. The super typhoon recurved and threatened the Ryukyu Islands and the main Japanese Islands of Kyushu and Honshu.

In late September, after the multiple outbreak of Tropical Storms Kit (17W), Lee (18W) and Mamie (19W), there was a week long lull in tropical cyclone activity. In the meantime polar air pushed southward across the Asian mainland and Japanese Islands. The
monsoon trough had returned to its normal climatic location along with the maximum cloud zone. The disturbance that would later become Super Typhoon Nelson was first detected in this maximum cloud zone 200 nm ( 370 km ) southwest of Guam by satellite reconnaissance. The Significant Tropical Weather Advisory, that is normally issued at $0600 Z$ each day, was reissued at $301400 Z$ September to include this area of suspect cloudiness. Nelson developed within the monsoon trough and began steadily organizing (Figure 3-20-1). A noticeable increase in central convection led to the issuance of a


Figure 3-20-1. Nelson as a tropical disturbance (010016Z October DMSP visual imagery).

Tropical Cyclone Formation Alert at 0900 Z on 1 October. The first warning followed at 011200 Z , based on a satellite intensity estimate of 45 kt ( $23 \mathrm{~m} / \mathrm{sec}$ ) sustained surface winds. (Post analysis indicates the disturbance, most probably, achieved tropical storm intensity earlier at 010600Z.) Nelson initially moved westward towards the Philippine Islands, and then west-northwestward as it tracked along the periphery of the subtropical ridge.

Only 24 hours after the initial warning was issued, a satellite intensity estimate of 65 kt ( $33 \mathrm{~m} / \mathrm{sec}$ ) winds resulted in an upgrade to typhoon status at 021200Z. At 022100Z, a 15 $\mathrm{nm}(28 \mathrm{~km})$ diameter eye first became visible on satellite imagery. (The eye persisted until 7 October.) Nelson continued to rapidly intensify
and reached super typhoon intensity at 040600 Z (Figure 3-20-2). The normal rate of intensification (Dvorak, 1984) is one T-number per day. From 020000 Z to 041200 Z , Nelson developed more rapidly than normal (Figure 3-20-3). Conversion (Atkinson and Holliday, 1977) of intensity to minimum sea-level pressure indicates a fall from 991 to 898 mb 93 mb in 60 -hours - and sustained rapid intensification (Holliday and Thompson, 1979)(Figure 3-20-4). On 4 October, Nelson slowed and tracked through an area where, according to climatology (Annual Typhoon Report, 1970), a large number of tropical cyclones reach super typhoon intensity (Figure 3-20-5). The typhoon's intensity peaked at 140 $\mathrm{kt}(72 \mathrm{~m} / \mathrm{sec})$ at 041200 Z .


Figure 3-20-2. Super Typhoon Nelson near peak intensity displays a well defined 20 $\mathrm{nm}(37 \mathrm{~km})$ diameter eye ( 040709 Z October NOAA visual imagery).


Figure 3-20-3. Analysis of intensity with time shows Super Typhoon Nelson's rapid intensification from 020000Z to 041200Z October. Note the peak intensity of $140 \mathrm{kt}(72 \mathrm{~m} / \mathrm{sec}$ ) persisted from 041200Z to 051200 Z October.


Figure 3-20-4. Nelson's final best track is superimposed upon the areas where tropical cyclones rapidly intensified during summer and early fall ( 20 June - 16 October) for the years 1956 to 1976 (Holliday and Thompson, 1979).


Figure 3-20-5. Nelson's final best track superimposed upon climatic areas of super typhoon occurrence. Areas of first super typhoon intensity include number of occurrances from the period 1959 to 1970 (Annual Typhoon Report, 1970).

Packing the most intense winds of any tropical cyclone for the year, Nelson rounded the western end of the subtropical ridge at a speed of 6 kt ( $11 \mathrm{~km} / \mathrm{hr}$ ), and slowly accelerated northeastward. In addition to satellite reconnaissance, a total of 177 radar position reports greatly aided the accurate tracking of the
typhoon's recurvature and subsequent acceleration. Moving along the edge of the modifying polar air, Nelson weakened and was downgraded to typhoon intensity at 060000 Z . The tropical cyclone passed 85 nm ( 157 km ) southeast of the island of Okinawa at 060930 Z (Figures 3-20-6 and 3-20-7). The maximum


Figure 3-20-6. Nelson's primary rainband and eye as viewed by the radar at Kadena Air Base, Okinawa at 061258 Z. Dots have been added to enhance the subtle edge of the rain echoes (photograph courtesy of Detachment 8, 20th Weather Squadron).


Figure 3-20-7. Microbarograph trace from Kadena Air Base, Okinawa. The time of the lowest minimum sea-level pressure of 29.02 inches Hg coincides with Nelson's closest approach to the island (barograph trace courtesy of Detachment 8, 20th Weather Squadron).
sustained winds reported by Detachment 8, 20th Weather Squadron at Kadena Air Base on Okinawa were $38 \mathrm{kt}(20 \mathrm{~m} / \mathrm{sec})$, with a peak gust of 59 kt ( $30 \mathrm{~m} / \mathrm{sec}$ ). Close by, maximum sustained winds of $40 \mathrm{kt}(21 \mathrm{~m} / \mathrm{sec})$, with a peak gust of $64 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ were reported by the Marine Corps Air Station at Futenma. The rainfall totals recorded on Okinawa ranged from 7.30 inches ( 18.54 cm ) at Futenma to 8.35 inches ( 21.21 cm ) at Kadena Air Base.

Nelson continued to weaken, move
northeastward and accelerate (Figure 3-20-8). As it lost its persistent central convection, the typhoon transitioned to an extratropical system $190 \mathrm{~nm}(352 \mathrm{~km})$ southeast of Tokyo, Japan at 081500 Z . By this time it was moving at a speed of $28 \mathrm{kt}(52 \mathrm{~km} / \mathrm{hr}$ ). Extratropical Nelson retained winds of typhoon intensity and moved rapidly northeastward. The final warning was issued at 081800Z. The remnants of Nelson were identifiable on satellite imagery for the next two days. No reports of significant damage were received.


Figure 3-20-8. Nelson during its weakening stage (072235Z October NOAA visual imagery).


## TYPHOON ODESSA (21W)

Odessa was the second of four significant tropical cyclones to occur during October. Slow to develop, it was tracked for three and a half days by satellite before the first warning was issued. After recurvature, Odessa rapidly intensified into a midget typhoon, despite interaction with a frontal system.

On 8 October, as Super Typhoon Nelson (20W) was weakening and accelerating to the northeast in higher latitudes, Odessa began as an area of convection superimposed on broad low-level easterly tradewinds 460 nm ( 852 km ) south-southeast of Minami Tori-Shima. The persistence of this convective area was first mentioned on the Significant Tropical Weather Advisory at 080600Z. After two days of faster than normal - 17 to 18 kt ( 32 to $33 \mathrm{~km} / \mathrm{hr}$ ) movement to the west-northwest, surface winds
in the area increased to $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$. At 102100Z, the Significant Tropical Weather Advisory was reissued to address this increase (Figure 3-21-1) and a Tropical Cyclone Formation Alert followed at 102300 Z . A satellite intensity estimate of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ prompted the first warning at 111200 Z .

At 121800 Z , Odessa was moving northnorthwestward toward the cooler, drier polar air that was spilling off the Asian mainland. As interaction with this air mass commenced, the tropical cyclone began tracking to the northeast and intensifying. Satellite intensity analysis at 131106 Z indicated sustained surface winds of $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ and Odessa was upgraded to a typhoon at 131200Z. Initially, the interaction with the cold front was expected to weaken the tropical cyclone; instead Odessa intensified into


Figure 3-21-1. Odessa as a tropical disturbance (110021Z October DMSP infrared imagery).


Figure 3-21-2. Typhoon Odessa displays a small eye before attaining its peak intensity (141044Z October NOAA visual imagery).
a midget typhoon. At 141200 Z , the intensity peaked at 90 kt ( $46 \mathrm{~m} / \mathrm{sec}$ ) (Figure 3-21-2).

Loss of organization and deep convection started to be evident at 151200 Z . With extratropical transition underway and the
low-level circulation exposed, the final warning was issued at 161800 Z . The extratropical circulation (Figure 3-21-3) made a counterclockwise loop on 17 October before moving off to the northeast along the frontal zone.


Figure 3-21-3. Odessa's well organized low-level circulation persists (170627Z October NOAA visual imagery).


## TYPHOON PAT (22W)

Pat was the third of four typhoons to develop during October. Unlike most of the western North Pacific tropical cyclones in 1988 that preceded, it formed equatorward of 10 degrees North latitude.

The tropical cyclone was first detected on satellite imagery and the Significant Tropical Weather Advisory was reissued at 160630 Z to include the low-level cyclonic circulation, which was located 300 nm ( 556 km ) south of Guam. Initially, synoptic and satellite data comparison indicated the maximum convection was in a convergent zone south of the low-level circulation. From 16 to 18 October, Pat slowly developed over the warm Philippine Sea and moved through an area of relatively low vertical
wind shear. By 18 October the convection had organized and the first Tropical Cyclone Formation Alert was issued at 180300 Z . Surface synoptic data indicated a minimum sealevel pressure (MSLP) of about 1002 mb and winds of 25 to 30 kt ( 13 to $15 \mathrm{~m} / \mathrm{sec}$ ). Satellite imagery continued to show an uneven distribution of deep convection with significantly more convection in the system's eastern semicircle. Pat's slow development required a second Alert at 181530 Z . Improving upper-level outflow and increasing central convection prompted the first warning at 181800Z.

Then Pat (Figure 3-22-1) assumed a more westerward course along the edge of the


Figure 3-22-1. Pat crossing the Philippine Sea (190514Z October NOAA visual imagery).
modifying polar high to the north. The tropical cyclone's convection and organization increased again, which required an upgrade to typhoon intensity at 200000 Z . Shortly before making landfall on the island of Luzon, Pat reached its peak intensity of 75 kt ( $39 \mathrm{~m} / \mathrm{sec}$ ). Maintaining typhoon intensity as it crossed central Luzon, the system (Figure 3-22-2) passed 75 nm (139 km ) north of Clark Air Base at 201800Z . The base, due to the sheltering effect of the nearby mountain ranges, only recorded peak gusts of $21 \mathrm{kt}(11 \mathrm{~m} / \mathrm{sec})$.

After entering the South China Sea, Pat (Figure 3-22-3) pressed on to the westnorthwest and sustained minimal typhoon intensity. It crossed the South China Sea and was downgraded to a tropical storm, as interaction with Hainan Dao occurred. Once across the island, the weakening system moved into the Gulf of Tonkin and entered northern Vietnam. It passed 30 nm ( 56 km ) northeast of Hanoi and dissipated inland.


Figure 3-22-2. Typhoon Pat entering the South China Sea (202152Z October DMSP infrared imagery).


Figure 3-22-3. Typhoon Pat approaches Hainan Dao (220011Z October NOAA visual imagery).


## TYPHOON RUBY (23W)

The last of four typhoons to develop in the western North Pacific during October, Ruby became the fifth tropical cyclone to hit the Philippine Islands in 1988.

On 20 October, as Typhoon Pat (22W) approached the Philippine Islands, Ruby formed to the east in the Philippine Sea. The Significant Tropical Weather Advisory was reissued at 201800 Z to include this new disturbance. Increased central convection and organization warranted a Tropical Cyclone Formation Alert at 210430 Z and the first warning at 211200 Z .

Ruby assumed the track of a "straight runner" and continued to intensify. At

240600Z, as it neared land, Ruby developed a $15 \mathrm{~nm}(28 \mathrm{~km})$ diameter eye and reached its peak intensity of $125 \mathrm{kt}(64 \mathrm{~m} / \mathrm{sec})$ at 241200 Z . The eye persisted for twelve hours before Ruby tracked into the mountainous terrain of central Luzon (Figure 3-23-1).

Like most tropical cyclones that track over the Philippine Islands, Ruby weakened significantly as it moved across Luzon; however, it was near super typhoon intensity shortly before it made landfall. The result was widespread damage and loss of life. More than three hundred people were killed, including over 150 who drowned when the ferry DONA MARILYN capsized at sea 300 nm ( 556 km ) southeast of Manila, and over 470,000 people


Figure 3-23-1. Ruby, shortly before reaching its peak intensity and making landfall over central Luzon (241000Z October DMSP infrared imagery).


Figure 3-23-2. Convective bands brought prolonged torrential rains, high winds and hazardous surf conditions to westem Luzon, as Ruby slowly entered the South China Sea ( 250048 Z October DMSP visual imagery).
were left homeless. Also the freighter JET ANN FIVE sank near Bohol Island in the southern Philippines after encountering rough seas from Typhoon Ruby.

Entering the South China Sea at 250300Z, Ruby (Figure 3-23-2) slowed. As a result of Ruby's slow departure from Luzon, Subic Bay Naval Base and Clark Air Base received their worst weather following the passage of the typhoon's center. Bands of wind and rain from the southwest slammed into western Luzon, causing torrential downpours and strong, gusty winds. Peak gusts recorded were $69 \mathrm{kt} \mathrm{( } 36 \mathrm{~m} / \mathrm{sec}$ ) at Subic Bay (Figure 3-$23-3)$ and $46 \mathrm{kt}(24 \mathrm{~m} / \mathrm{sec})$ at Clark Air Base.
(These were the strongest winds reported at Clark Air Base since Super Typhoon Rita (1978) produced gusts to $58 \mathrm{kt}(30 \mathrm{~m} / \mathrm{sec})$.)

Ruby then began to track toward the island of Hainan and made landfall at 280400Z. Interaction with the mountainous terrain of Hainan caused the tropical cyclone to weaken. The final warning was issued at 281800 Z , after satellite imagery indicated the absence of central convection. Although Ruby's circulation dissipated over water, heavy rainshowers caused flash floods in northern Vietnam that killed at least 100 people, left thousands homeless and destroyed over 300,000 tons of agricultural produce.


Figure 3-23-3. Ruby's high winds caused widespread damage. This tree toppled into a housing unit. (Photo courtesy of the Naval Oceanography Command Facility, Cubi Point, Republic of the Philippines.)


The first of two significant tropical cyclones to develop during November, Skip was a classic "straight runner" that covered over $2,000 \mathrm{~nm}(3,704 \mathrm{~km})$ during its ten day lifetime. This typhoon was especially damaging to the Philippine Islands because it followed close behind Typhoons Ruby (23W) and Tess (25W).

On the first of November the northeast monsoon was well established across the South China Sea and southeastern Asia. Easterly tradewinds dominated the Philippine Sea north of the near-equatorial trough and a disturbance, that was to become Typhoon Tess (25W), was bringing more rain and wind to the central Philippine Islands. The next day Skip began as
an area of convection in the monsoon trough about 360 nm ( 667 km ) southwest of the island of Guam. After the convection had persisted for a day, the disturbance was listed on the Significant Tropical Weather Advisory at 030600Z. Visual satellite imagery indicated a well defined low-level cyclonic circulation immediately west of the curved band of convection (Figure 3-24-1). A satellite intensity analysis estimate of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec}$ ) surface winds precipitated the first Tropical Cyclone Formation Alert at 030700Z. The disturbance continued to develop and moved out of the Alert box, necessitating the issuance of a second Alert at 031400Z. Four hours later, after a satellite intensity estimate of $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$,


Figure 3-24-1. Skip, shortly before the first Tropical Cyclone Formation Alert ( $030642 Z$ November NOAA visual imagery).
the first warning was issued. Skip (Figure 3-242) tracked westward and intensified, peaking in intensity at $125 \mathrm{kt}(64 \mathrm{~m} / \mathrm{sec})$ at 060600 Z .

As the typhoon approached the central Philippine Islands, it began to weaken and accelerate. The tropical cyclone reached a forward speed of 21 kt ( $39 \mathrm{~km} / \mathrm{hr}$ ), as it tracked across the island of Mindoro. The typhoon's trek through the Philippine Islands caused significant damage and loss of life. At least 104 people were killed by mudslides, floods and flying debris and another 95 persons were listed as missing. In all, Skip left over 600,000 homeless, caused extensive damage to coconut,
rice and sugar crops, and widespread disruption of power and communication lines. Numerous watercraft were reported lost, missing or aground.

After weakening over the Philippine Islands, Skip slowed as it entered the South China Sea at 071800Z. During the next four days, Skip pushed west-northwestward along the southern side of the narrow subtropical ridge. At 100600 Z the typhoon was downgraded to a tropical storm and further weakening led to the final warning at 110000 Z . The remnants of Skip (Figure 3-24-3) drifted into the Gulf of Tonkin and dissipated.


Figure 3-24-2. Typhoon Skip at peak intensity. Note the eye feature and symmerry of deep convection ( 061041 Z November NOAA infrared imagery).


Figure 3-24-3. Skip maintains its organization, shortly after the final waming was issued (110157Z November DMSP visual imagery).


## TYPHOON• TESS (25W)

The last of two significant tropical cyclones to occur during November, Tess developed slowly for three days before the first warning. Tess was the only tropical cyclone to track across southern Vietnam during 1988.

Tess' persistent area of convection was first mentioned on the Significant Tropical Weather Advisory at 010600 Z . For the next
three days the disturbance tracked southwestward along the edge of the deep northeasterly flow of the winter monsoon. Once across the rugged Philippine Islands and over open water in the Sulu Sea, the tropical cyclone (Figure 3-25-1) became more organized and required a Tropical Cyclone Formation Alert at 031730Z. A satellite intensity estimate of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$ prompted the first warning


Figure 3-25-1. As a disturbance, Tess showed improved convection and organization as it entered the Sulu Sea (032221Z November DMSP visual imagery).
on Tropical Depression 25W (Figure 3-25-2) at 040000Z.

Almost immediately after the first warning, the track became westward. The most probable explanation for this change appeared in the low-level northeasterly gradient flow. The pressure gradient between the winter high and the lower pressure associated with Tess had sustained a persistent flow of at least 30 kt ( 15 $\mathrm{m} / \mathrm{sec}$ ) upstream of the tropical cyclone since 1 November. This upstream pressure gradient relaxed on 4 November and the gales clustered around Tess.

Along with this track change came
intensification, as the system crossed Palawan Island and entered the South China Sea. At 040600 Z , satellite intensity estimates indicated $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec}$ ) surface winds and the tropical depression was upgraded to tropical storm intensity. The system (Figure 3-25-3) reached its peak intensity of $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$ at 051200Z.

As Tess approached the coast of southern Vietnam, it began to weaken. The tropical cyclone was downgraded to tropical storm intensity and finalled at 060600 Z . The remnants of Tess continued to track westward across the Mekong river delta. No reports of damage or loss of life were received.


Figure 3-25-2. Tess, shortly after the first waming was issued. The cloudiness at the picture's lower right is associated with Skip (24W) (040053Z November DMSP visual imagery).


Figure 3-25-3. Tess just before reaching its peak intensity in the South China Sea ( 051054 Z November DMSP infrared imagery).


## TROPICAL STORM VAL (26W)

Tropical Storm Val (26W) was the final significant tropical cyclone of 1988 and the only storm to develop during December. It proved difficult to position and hard to forecast. In less than 24 -hours, Val decelerated from a 25 kt ( 46 $\mathrm{km} / \mathrm{hr}$ ) to quasi-stationary. Also, the system's cirrus outflow prevented timely detection of shearing and decoupling of the low-level circulation center from the central convection.

On 13 December a massive outbreak of polar air started to push southeastward from Asia across the Philippine Sea. By 18 December the leading edge of the air mass stretched from the southern Philippine Islands to the northem Marianas and northeastward. As the major thrust of the cold air diminished, the southern Philippine Sea filled with deep convection and a near-equatorial trough formed. Multiple low-level cyclonic circulations
appeared in the trough with nearby gales to the north due to the strong northeast monsoon. Finally, the convection began to consolidate in the western Caroline Islands and the area was initially mentioned at 210600 Z on the Significant Tropical Weather Advisory. The convection continued to organize and a Tropical Cyclone Formation Alert was issued at 220700 Z based on satellite imagery that indicated an increase in upper-level organization. Plus, surface synoptic reports revealed pressures as low as 1005 mb and 20 kt ( $10 \mathrm{~m} / \mathrm{sec}$ ) westerly winds to the south of the circulation center.

A second Alert was issued at 221800 Z to cover the unusually rapid movement - 25 kt ( $46 \mathrm{~km} / \mathrm{hr}$ ) - of the circulation to the west. Still, the area (Figure 3-26-1) developed, and the upper-level outflow and surface circulation


Figure 3-26-1. Val with its major convective band (230429Z December NOAA visual imagery).
improved. The first warning followed close behind with a valid time at 221800 Z based on a satellite intensity analysis estimate of 30 kt ( 15 $\mathrm{m} / \mathrm{sec}$ ). Then Val (26W) began to decelerate and intensify. It reached a peak intensity of 55 $\mathrm{kt}(28 \mathrm{~m} / \mathrm{sec}$ ) at 240000 Z (Figure 3-26-2).

As the intensity peaked and forward motion ground to a halt on 24 December, high cloudiness obscured the low-level circulation
center. The deep central convection and upperlevel circulation center, which were the targets for remote sensing, tracked northward. The shallow system (Figure 3-26-3) continued to weaken and the final warning was issued at 260000 Z . The dissipating low-level circulation center accelerated to the southwest along the eastern boundary of the northeast monsoon. No reports of damage or loss of life were received.


Figure 3-26-2. Tropical Storm Val at maximum intensity ( 240114 Z December DMSP visual imagery).


Figure 3-26-3. Val's exposed low-level circulation center appears at the southern edge of the central cloud mass ( 250054 Z December DMSP visual imagery).

## 3. NORTH INDIAN OCEAN TROPICAL CYCLONES

Five significant tropical cyclones developed in the North Indian Ocean during 1988. This was average and well below 1987's all-time record of eight tropical cyclones. The only tropical cyclone to form in the Arabian Sea, Tropical Cyclone 01A, was also the only cyclone of the spring transition season. In contrast, the other four tropical cyclones were all part of the fall transition season and occurred
in the Bay of Bengal. The most damaging of these, Tropical Cyclone 04B, was one of the most intense to strike the Ganges River delta in this century. The spring and fall in the North Indian Ocean are periods of transition between major climatic controls - the summer, or southwest monsoon, and the winter, or northeast monsoon - and the most favorable seasons for tropical cyclone formation. Tables 3-5 and 3-6 provide a summary of information for 1988 and comparison with earlier years.

| TABLE 3-5 1988 SIGNIFICANT TRCOPICAL CYCIONESNORTH INDIAN OCEAN |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |
| TROPICAL CYCLONE | PERIOD OE WARNING | NUMBER OF WARNINGS ISSUED | $\begin{aligned} & \text { MAXIMUM } \\ & \text { SURFACE } \\ & \text { HINDS-KT (M/S) } \end{aligned}$ | $\begin{aligned} & \text { ESTIMATED } \\ & \text { MSIP - MB } \end{aligned}$ |
| TC 01A | 10 JUN - 12 JUN | 10 | 35 (18) | 996 |
| TC 02B | 18 ОСT - 19 ОСT | 3 | 35 (18) | 996 |
| TC 03B | 18 NOV | 3 | 55 (28) | 984 |
| TC 04B | 24 NOV - 29 NOV | 22 | 110 (57) | 933 |
| TC 05B | 07 DEC - 08 DEC | 6 | 45 (23) | 991 |
|  | TOTAL | 44 |  |  |



WARNINGS: NUMBER OF CALENDAR WARNING DAYS: 14
THERE WERE NO CALENDAR WARNING DAYS WITH TWO OR MORE TROPICAL CYCLONES.

## 4. NORTH INDIAN OCEAN CLIMATOLOGY

This climatology* of peak tropical cyclone intensity for the North Indian Ocean was prepared from data in the 1971 through 1988 Annual Tropical Cyclone Reports and Annual Typhoon Reports. JTWC became responsible for the Bay of Bengal east of $90^{\circ}$ East longitude on 4 June 1971. Table 3-7 lists the number of significant tropical cyclones by peak intensity and month for the Bay of Bengal. JTWC became responsible for the Arabian Sea in 1975; thus the Arabian Sea data are from 1975 through 1988. Table 3-8 lists the number of significant tropical cyclones by peak intensity and month for the Arabian Sea.

Note: data for tropical cyclones that passed from the Bay of Bengal into the Arabian Sea were included in the statistics for both basins when the peak intensity was 35 kt ( 18 $\mathrm{m} / \mathrm{sec}$ ) or greater in both basins. If the peak intensity was $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec})$ or greater and the tropical cyclone transitioned from one month to another, it was included in each month's statistics. Tropical cyclones with less than 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ) were not considered. The data set for 1971 through 1974 in the Bay of Bengal is incomplete, because the AOR at that time only included tropical cyclones that developed or tracked east of $90^{\circ}$ East longitude. These years are included since they provide limited additional data on the peak intensities.

* Climatology prepared by Capt John Rogers,
USAF.

| $\begin{gathered} \text { Peak } \\ \text { Intensity } \\ (\mathbf{k t}) \end{gathered}$ | TABLE 3-7 BAY OF BENGAL (1971-1988) |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Jan | Feb | Mar | Apr | May | Jun | Jul | Aug | Sep | Oct | Nov | Dec | Total |
| 125 | I |  |  |  | 11 |  |  | , |  |  |  |  | 1 |
| 120 | 1 |  |  |  |  |  |  | I |  |  |  |  | 0 |
| 115 |  |  |  |  |  |  |  |  |  |  | 1 |  | 1 |
| 110 |  |  |  |  |  |  |  |  |  |  | 1 |  | 1 |
| 105 | 1 |  |  |  | 1 |  |  | 1 |  |  |  |  | 0 |
| 100 | 1 |  |  |  | 1 |  |  | 1 |  | 1 |  |  | 1 |
| 95 | ! |  |  |  | I |  |  | I |  |  | 1 |  | 1 |
| 90 |  |  |  |  |  |  |  |  |  |  | 1 |  | 1 |
| 85 |  |  |  | 1 | 1 |  |  |  |  |  | 1 |  | 3 |
| 80 | 1 |  |  |  | 1 |  |  | 1 | 1 |  |  |  | 1 |
| 75 | 1 |  |  |  | 11 |  |  | 1 |  |  | 2 | 2 | 5 |
| 70 | 1 |  |  |  | I |  |  | I | 1 |  | 1 |  | 2 |
| 65 |  |  |  |  |  |  |  |  |  |  | 1 |  | 1 |
| 60 |  |  |  |  | 3 |  |  |  |  |  | 1 | 1 | 5 |
| 55 | 1 | 1 |  |  | 1 | 2 |  |  |  |  | 5 | 1 | 9 |
| 50 | 1 |  | 1 | 1 | 11 |  |  | 1 |  | 5 | 3 | 1 | 11 |
| 45 | 1 |  |  |  | I |  |  | I | - | 2 | 1 | 1 | 5 |
| 40 | 1 |  |  |  |  |  |  |  | 1 | 4 |  |  | 6 |
| 35 | 11 |  |  |  | 1 | , |  | 1 |  | 2 |  | 3 | 16 |
| Total | 13 | 1 | 10 | 2 | 17 | 2 | 0 | 10 | 1 | 14 | 19 | 9 | 60 |


| Peak$\begin{gathered}\text { Intensity } \\ \text { (kt) }\end{gathered}$ | TABLE 3-8 ARABIAN SEA (1975-1988) |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Jan | Feb | Mar | Apr | May | Jun | Jul | Aug | Sep | Oct | Nov | Dec | Total |
| 95 |  |  |  |  | 1 |  |  |  |  |  |  |  | 1 |
| 90 |  |  |  |  |  |  |  |  |  |  | 1 |  | 1 |
| 85 |  |  |  |  |  |  |  |  |  |  |  |  | 0 |
| 80 |  |  |  |  |  |  |  |  |  | 1 | 1 |  | 2 |
| 75 |  |  |  |  | , |  |  |  |  |  |  |  | 0 |
| 70 |  |  |  |  |  |  |  |  |  |  | 1 |  | 1 |
| 65 |  |  |  |  |  |  |  |  |  |  |  |  | 0 |
| 60 |  |  |  |  |  | 1 |  |  |  | 1 | 1 | 1 | 4 |
| 55 |  |  |  |  |  |  |  |  | 1 |  |  |  | 1 |
| 50 |  |  |  |  | 1 | 2 |  |  |  | 1 |  |  | 4 |
| 45 |  |  |  |  | 1 |  |  | 1 |  |  | 1 | 1 | 4 |
| 40 |  |  |  |  |  | 1 |  |  |  |  | 1 |  | 2 |
| 35 |  |  |  |  |  | 1 |  |  |  |  | 2 |  | 3 |
| Total | 0 | 0 | 0 | 0 | 3 | 5 | 0 | 1 | 1 | 3 | 8 | 2 | 23 |




## TROPICAL CYCLONE 01A

Tropical Cyclone 01A was the first and only significant tropical cyclone to develop in the North Indian Ocean during the spring transition season. Due to persistent upper-level cloudiness (Figure 3-01A-1) the system proved difficult to position, track and forecast. Tropical Cyclone 01A was initially identified as an area of convection about $240 \mathrm{~nm}(444 \mathrm{~km})$ south of

Bombay, India on 8 June by the Air Force Global Weather Central. It was first mentioned on the Significant Tropical Weather Advisory at 081800Z. After satellite imagery indicated a central dense overcast, increased convection and upper-level organization, a Tropical Cyclone Formation Alert was issued at $091430 Z$. The system's organization continued


Figure 3-01A-1. The bright, cold cloudiness persisted over and masked the low-level circulation (100530Z June DMSP infrared imagery).
to improve and JTWC issued its first warning at 100000 Z when satellite imagery and synoptic data indicated an intensity of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ). Finally, vertical wind shear exposed the lowlevel circulation center. As a result, the 110000 Z warning was amended and relocated -
the circulation center appeared 240 nm ( 407 km ) east-southeast of the upper-level circulation center (Figure 3-01A-2). Unfavorable conditions aloft continued and at 120600 Z the final warning was issued, as the tropical cyclone dissipated over water.


Figure 3-01A-2. The low-level circulation is exposed. Note the dust blowing seaward from coastal areas to the north (110511Z June DMSP visual imagery).

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## TROPICAL CYCLONE 02B



Figure 3-02B-1. The tropical disturbance was first detected 135 nm ( 250 km ) southwest of Rangoon, Burma and the Significant Tropical Weather Advisory was reissued at 170800 Z to include it. The first waming followed at 181200 Z as the cyclone reached a peak intensity of 35 kt ( $18 \mathrm{~m} / \mathrm{sec}$ ). Once onshore, the final warning was issued at 190000 Z . Press releases cited at least 35 deaths, more than 1000 injuries and an estimated 1500 missing. The fishing fleet was particularly hard hit and there were press reports of $15 \mathrm{ft}(4.6 \mathrm{~m})$ waves and winds as high as $65 \mathrm{kt}(33 \mathrm{~m} / \mathrm{sec})$. The above photo shows Tropical Cyclone 02B at the coast of Bangladesh (182219Z October NOAA visual imagery).



Figure 3-03B-1. The second significant tropical cyclone to develop in the Bay of Bengal during the fall transition season, Tropical Cyclone 03B was the first to make landfall over Burma. By 12 November the presence of Tropical Cyclone 01S in the South Indian Ocean strengthened the low-level westerlies along the equator and across the southem portion of the Bay of Bengal. As Tropical Cyclone 01S moved west-southwestward, a disturbance organized north of the equator. The Significant Tropical Weather Advisory was reissued at 150600 Z to describe the circulation. The system tracked northward around the western edge of a subtropical anticyclone. At 172230Z, a Tropical Cyclone Formation Alert was issued after satellite intensity analysis indicated sustained surface winds of $30 \mathrm{kt}(15 \mathrm{~m} / \mathrm{sec})$. The disturbance's organization continued to improve and, at 180000Z, the first warning was issued. The final waming followed 12 -hours later as Tropical Cyclone 03B made landfall on the coast of Burma. The enhanced infrared picture above shows the tropical cyclone during intensification (180309Z November DMSP infrared imagery).


## TROPICAL CYCLONE 04B

Tropical Cyclone 04B was the second of two significant tropical cyclones to develop in the Bay of Bengal during November. This cyclone (Figure 3-04B-1) was one of the most intense to strike Bangladesh and eastern India in this century.

The formation of Tropical Cyclone 04B was preceded by a sustained surge of the northeast winter monsoon and low-level convergence across the Malay Peninsula. Beginning 19 November, prolonged, heavy rains occurred in northern Malaysia and southern Thailand. Flash flooding and mud-


Figure 3-04B-1. Tropical Cyclone 04B approaching the coast with maximum sustained winds of over $100 \mathrm{kt}(51 \mathrm{~m} / \mathrm{sec})$ ( 290251 Z November DMSP visual imagery).
slides from deforested hillsides swept across low-lying villages, killing at least 1000 people and rendering another 100,000 homeless. At 211800Z, Tropical Cyclone 04B consolidated in the Straits of Malacca and was first mentioned on the Significant Tropical Weather Advisory. The central convection, organization and satellite intensity estimates quickly increased and a Tropical Cyclone Formation Alert was issued at 231830 Z . The system tracked northwestward while remaining south of the subtropical ridge. The central convection continued to increase and organize, and the first warning followed at 240600 Z .

Shortly after attaining typhoon intensity, at 260000 Z , the system began tracking around
the western periphery of the mid-level subtropical ridge. The ridge was broad, which allowed the tropical cyclone to move northward for three days prior to making landfall over the delta of the Ganges River. During this period Tropical Cyclone 04B gradually intensified, reaching a peak of $110 \mathrm{kt}(57 \mathrm{~m} / \mathrm{sec})$ at the coast. As the cyclone swept inland, it ravaged the southern Bangladesh and northeastern India coastal zones, leaving at least 2000 people dead, 6000 missing and almost three million homeless. Up to seventy percent of the crops ready for harvest were destroyed. Bangladesh, which was attempting to recover from earlier flooding during the summer, in which 1500 lives were lost and countless were left homeless, was particularly hard hit.

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## TROPICAL CYCLONE 05B



Figure 3-05B-1. The only significant tropical cyclone to form in the Bay of Bengal in December, Tropical Cyclone 05B was first detected as a poorly defined area of cloudiness in the southem Bay of Bengal. The Significant Tropical Weather Advisory was reissued at 061230 Z to include the disturbance as suspect for further development. The cloud system continued to develop and a Tropical Cyclone Formation Alert was issued at 062100Z. The first warning followed at 070000Z, after satellite intensity analysis indicated $35 \mathrm{kt}(18 \mathrm{~m} / \mathrm{sec}$ ) sustained surface winds. Once the peak intensity of $45 \mathrm{kt}(23 \mathrm{~m} / \mathrm{sec})$ was reached, the tropical cyclone began to weaken and the final waming was issued at 080600Z. The remnants of Tropical Cyclone 05B persisted and struck off westward across the Bay of Bengal for three additional days, before dissipating along the east coast of India north of the city of Madras. The photo above shows Tropical Cyclone 05B shortly before peak intensity (070332Z December DMSP visual imagery).

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## CHAPTER IV - SUMMARY OF SOUTH PACIFIC AND SOUTH INDIAN OCEAN TROPICAL CYCLONES

## 1. GENERAL

The JTWC area of responsibility (AOR) was expanded on 1 October 1980 to include the southern hemisphere from $180^{\circ}$ longitude westward to the east coast of Africa. Details on tropical cyclones in this region for July 1980 to June 1982 are contained in Diercks et. al. (1982). For the July 1982 through June 1984 period, reference the NOCC/JTWC TECH NOTE 86-1. As in earlier reports, data on tropical cyclones forming in, or moving into, the South Pacific Ocean east of $180^{\circ}$ longitude (NWOC's AOR) are included for completeness. JTWC provides the numbers for all South Pacific and South Indian Ocean tropical cyclones. The current convention (as stated in USCINCPACINST 3140.1S) for labelling tropical cyclones that develop in the South Indian Ocean (west of $135^{\circ}$ East longitude) is to add the suffix "S" to the assigned tropical cyclone number, while those originating in the South Pacific Ocean (east of $135^{\circ}$ East longitude) receive a " P " suffix. The " P " suffix also applies to significant tropical cyclones which form east of $180^{\circ}$ longitude in the South Pacific Ocean. Also, it should be noted that to encompass the southern hemisphere tropical cyclone season, which normally occurs from January through April, the limits of each southern hemisphere tropical cyclone year are defined as 1 July to 30 June. Thus, the 1988 southern hemisphere tropical cyclone year is from 1 July 1987 to 30 June 1988. This is in contrast to the labelling convention in the northern hemisphere, which is based on the calendar year (1 January to 31 December) to include the seasonal activity from May through December.

## 2. SOUTH PACIFIC AND SOUTH INDIAN OCEAN TROPICAL CYCLONES

Twenty-one significant tropical cyclones (Table 4-1 and Figure 4-1) occurred during the year (1 July 1987 through 30 June 1988) in the southern hemisphere AOR. This was lower than the twenty year average of 24.7 (Table 4-2) and significantly lower than the short term (19811985) average (Annual Tropical Cyclone Report, 1987) of 28.6. Five tropical cyclones occurred in the South Pacific Ocean, east of $165^{\circ}$ East longitude, which is very close to the long-term mean (Table 4-3). The Australian area ( $105^{\circ}$ to $165^{\circ}$ East longitude) accounted for only two tropical cyclones, compared to the twenty year climatological mean of 10.3 cyclones. Fourteen tropical cyclones developed in the South Indian Ocean, which is almost six more than the twenty year mean of 8.4 cyclones.

Caveat: Intensity estimates for southern hemisphere tropical cyclones are derived primarily from evaluation of satellite imagery (Dvorak, 1984) and from intensity estimates reported by other regional centers. Only in isolated cases are intensity estimates based on conventional surface observations. Estimates of minimum sea-level pressure are usually derived from the Atkinson and Holliday (1977) relationship between maximum sustained oneminute surface wind and minimum sea-level pressure (Table 4-4). This relationship has been shown to be representative for tropical cyclones in the western North Pacific and is also used by Australian region warning agencies to provide intensity estimates. However, these pressure estimates are usually based on wind intensities derived from interpretation of satellite imagery. Considerable caution should be exercised when using resultant pressure values in future tropical cyclone work.

| TABLE 4-1 SOUTH PACIFIC AND SOUTA INDIAN OCEANS |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |
| TROPICAL CYCLONE | PERIOD OF WARNING | NUMBER WARNINGS ISSUED | MAXIMUMSURFACEGINDS-KT_(M/SECI |  | ESTIMATED MSLP-MB |
| 015 - - - - | 01 Nov - 09 Nov | 20 | 55 | (28) | 984 |
| 02S - - - - | 24 NOV - 26 Nov | 5 | 40 | (21) | 994 |
| 03 S ARINY | 09 DEC - 14 DEC | 10 | 55 | (28) | 984 |
| 04P - - - - | 22 DEC - 23 DEC | 3 | 35 | (18) | 997 |
| 055 BERNANDRO | 27 DEC - 01 JAN | 11 | 35 | (18) | 997 |
| 06P AGI | 06 JAN - 07 JAN | 3 | 35 | (18) | 997 |
| 06P AGI * | 09 JAN - 14 JAN | 13 | 70 | (36) | 972 |
| 07P ANNE | 07 JAN - 14 JAN | 14 | 140 | (72) | 898 |
| 08 S CALIDERA | 13 JAN - 15 JAN | 5 | 65 | (33) | 976 |
| 09 S DOAZA | 23 JAN - 26 JAN | 7 | 55 | (28) | 984 |
| 09S DOAZA * | 28 JAN - 02 FEB | 11 | 115 | (59) | 927 |
| 10 S FREDERIC | 31 JAN - 02 FEB | 6 | 65 | (33) | 976 |
| 115 GWENDA ** | 08 FEB - 16 FEB | 16 | 90 | (46) | 954 |
| 12P CHARLIE | 21 FEB - 24 FEB | 9 | 45 | (23) | 991 |
| 12P CHARLIE * | 28 FEB - 29 FEB | 3 | 35 | (18) | 997 |
| 13P BOLA | 24 FEB - 04 MAR | 20 | 105 | (54) | 935 |
| 14 S - - - | 27 FEB - 02 MAR | 8 | 85 | (44) | 958 |
| 15P CILILA | 28 FEB - 03 MAR | 8 | 45 | (23) | 991 |
| 16S GASITAO | 16 MAR - 23 MAR | 16 | 130 | (67) | 910 |
| 17 S - - - | 17 MAR - 20 MAR | 7 | 45 | (23) | 991 |
| 18 S HELY | 27 MAR - 28 MAR | 3 | 40 | (21) | 994 |
| 19P DOVI | 09 APR - 15 APR | 12 |  | (36) | 972 |
| 205 IARISENA | 09 MAY - 10 MAY | 3 | 40 | (21) | 994 |
| $21 S$ - - - | 19 MAY - 20 MAY | 4 | 35 | (18) | 995 |
| TOTAL 217 |  |  |  |  |  |
| * REGENERATED |  |  |  |  |  |
| ** ALSO NAMED EZENINA |  |  |  |  |  |
| NOTE: NAMES OF SOUTHERN HEMISPHERE TROPICAL CYCLONES ARE GIVEN BY THE REGIONAL WARNING CENTERS (NANDI, BRISBANE, DARWIN, PERTH AND MAURITIUS) AND ARE APPENDED TO JTWC WARNINGS, WHEN AVAILABLE. |  |  |  |  |  |




| TABLE 4-3 |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: |
|  trCOPICAL CYCLONES EY OCTRAN BASIN |  |  |  |  |
| YEAB | $\begin{aligned} & \text { SOUTH INDIAN } \\ & \text { (GEST OF } 105^{\circ} \text { E) } \end{aligned}$ | AUSTRALIAN $\left(105^{\circ} E-165^{\circ} E\right)$ | SOUTH PACIFIC <br> (EAST OF $165^{\circ}$ EI | TOTAL |
| (1959-1978) |  |  |  |  |
| AVERAGE* | 8.4 | 10.3 | 5.9 | 24.7 |
| 1981 | 13 | 8 | 3 | 24 |
| 1982 | 12 | 11 | 2 | 25 |
| 1983 | 7 | 6 | 12 | 25 |
| 1984 | 14 | 14 | 2 | 30 |
| 1985 | 14 | 15 | 6 | 35 |
| 1986 | 14 | 16 | 3 | 33 |
| 1987 | 9 | 8 | 11 | 28 |
| 1988 | 14 | 2 | 5 | 21 |
| (1981-1988) |  |  |  |  |
| AVERAGE | 12.1 | 10.0 | 5.5 | 27.6 |
| TOTAL CASES | $5 \quad 97$ | 80 | 44 | 221 |
| * (GRAY, 19 | 979) |  |  |  |


| TABLE 4-4 |  |
| :---: | :---: |
| MAXCMUM SUSTATNUD SUREACE GINDS ARD EOUIVATMNT MINIMCM STRA-LEVEI PRESSURE (ATKINSON ARD HOUTIDAY, 1977) |  |
| MAXIMUM SUSTAINED SUREACE WIND (KTI) | MINIMUM SEA-IEVEL PRESSURE (MB) |
| 30. | 1000 |
| 35 | 997 |
| 40 | 994 |
| 45 | 991 |
| 50 | 987 |
| 55 | 984 |
| 60 | 980 |
| 65 | 976 |
| 70 | 972 |
| 75 | 967 |
| 80 | 963 |
| 85 | 958 |
| 90 | 954 |
| 95 | 948 |
| 100 | 943 |
| 105 | 938 |
| 110 | 933 |
| 115 | 927 |
| 120 | 922 |
| 125 | 916 |
| 130 | 910 |
| 135 | 906 |
| 140 | 898 |
| 145 | 892 |
| 150 | 885 |
| 155 | 879 |
| 160 | 872 |
| 165 | 865 |
| 170 | 858 |




Intentionally left blank.

## CHAPTER V - SUMMARY OF FORECAST VERIFICATION

## 1. ANNUAL FORECAST VERIFICATION

## a. WESTERN NORTH PACIFIC OCEAN

Verification of wamings at initial, 24-, 48- and 72-hour forecast positions was made against the final best track. The (scalar) forecast, along-track and cross-track errors (illustrated in Figure 5-1) were then calculated for each tropical cyclone and are presented in Tables $5-1 \mathrm{~A}, 5-1 \mathrm{~B}, 5-1 \mathrm{C}$ and $5-1 \mathrm{D}$, as appropriate. The frequency distributions of forecast errors in $30 \mathrm{~nm}(56 \mathrm{~km}$ ) increments for 24-, 48 -, and 72 -hour forecasts are in Figures 52 A through $5-2 \mathrm{C}$, respectively. A summation of the mean forecast errors, since 1971, is shown in Table 5-2A. Table 5-2B includes mean along-track and cross-track forecast errors for 1988. A comparison of the annual mean forecast errors for all tropical cyclones as compared to those tropical cyclones that reached typhoon intensity can be seen in Table $5-3$. The mean forecast errors for 1988 as compared to the nineteen previous years are graphed in Figure 5-3.

## b. NORTH INDIAN OCEAN

The positions given for warning times and those at the $24-, 48$-, and 72 -hour valid times were verified for tropical cyclones in the

Figure 5-1. Definition of cross-track error (XTE), along-track error (ATE) and forecast track error (FTE). In this example, the XTE is positive (to the right of the best track) and the ATE is negative (behind or slower than the best track).

North Indian Ocean by the same methods used for the western North Pacific. These error statistics should not be taken as representative of any trend due to the small sample number. Table 5-4 is the initial and forecast along-track and cross-track error summary for the North Indian Ocean. Table 5-5A contains a summary of the annual mean forecast errors for each year. Table 5-5B includes along-track and cross-track errors for 1988. Forecast errors are plotted in Figure 5-4 (Seventy-two hour forecast errors were evaluated for the first time in 1979). There were no verifying 72 -hour forecast in 1983 and 1985.

## c. SOUTH PACIFIC AND SOUTH INDIAN OCEANS

The positions given for warning times and those at the $24-, 48$-, and 72 -hour valid times were verified for tropical cyclones in the Southern Hemisphere by the same methods used for the western North Pacific. It should be noted that due to the lack of verifying groundtruth data, these error statistics should not be taken as representative of any trend. Table 56 A is the initial, forecast along-track and crosstrack error summary for the southern hemisphere. Table 5-6B has the number of warnings verified at each forecast period. Table 5-7A contains a summary of the annual mean forecast errors for each year. Table 5-7B includes alongand cross-track errors for 1988. Forecast errors are plotted in Figure 5-5.


|  |  |  |  |
| :---: | :---: | :---: | :---: |
| TROPI | AL CYCLONE | ERROR GNML | SAMPLE SIZE |
| (01w) | TY ROY | 13 | 41 |
| (02W) | TY SUSAN | 18 | 17 |
| (03W) | TD 03W | 19 | 6 |
| (04W) | TY THAD | 24 | 20 |
| (05W) | TS VANESSA | 27 | 11 |
| (06W) | TY WARREN | 19 | 30 |
| (07w) | tS AGNES | 33 | 8 |
| (08W) | TS BILL | 12 | 5 |
| (09W) | TS CLARA | 24 | 6 |
| (10W) | TY DOYLE | 14 | 24 |
| (11W) | TS EISIE | 28 | 6 |
| (11W) | TS ELSIE* | 27 | 4 |
| (12W) | ty fabian | 21 | 18 |
| (13W) | TS GAY | 33 | 6 |
| (14W) | TY HAL | 18 | 37 |
| (01C) | TY ULEKI | 17 | 21 |
| (15W) | TS IRMA | 17 | 16 |
| (16W) | TS JEFF | 15 | 9 |
| (17W) | TS KIT | 26 | 12 |
| (18W) | TS IEE | 21 | 15 |
| (19W) | TS MAMIE | 36 | 4 |
| (20w) | STY NELSON | 10 | 30 |
| (21w) | TY ODESSA | 16 | 22 |
| (22W) | TY PAT | 17 | 17 |
| (23W) | TY RUBY | 24 | 30 |
| (24W) | TY SKIP | 18 | 30 |
| (25W) | TY TESS | 14 | 10 |
| (26W) | TS VAL | 57 | 10 |
| TOTALS |  | 23 | 465 |
| * Regenerated |  |  |  |



| TABLE 5-1C |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | SUMPARY OF 48-HOUR FORECAST ERRRORS MESTERN NORTH PACIFIC OCTRAN 1988 SIGNIFICANT TROPICAL CYCTONIRS |  |  |  |  |  |  |  |
|  | TROPICAL CYCLONE |  | FORECASTERROR (NM) | ALONG-TRACK ERROR |  | CROSS-TRACK ERROR |  | $\begin{gathered} \text { SAMPLE } \\ \text { SIZE } \end{gathered}$ |
|  |  |  | MEAN* | MEDIAN | MEAN* | MEDIAN |  |
|  | (01W) | TY ROY |  | 246 | 196 | -62 | 119 | 22 | 33 |
|  | (02W) | TY SUSAN | 378 | 344 | ** | 106 | ** | 9 |
|  | (03W) | TD 03W | *** | *** | *** | *** | *** | 0 |
|  | (04W) | TY THAD | 274 | 233 | -223 | 123 | -81 | 12 |
|  | (05W) | TS VANESSA | 386 | 352 | ** | 158 | ** | 3 |
|  | (06W) | TY WARREN | 152 | 118 | -52 | 74 | 56 | 23 |
|  | (07W) | TS AGNES | 714 | 706 | ** | 108 | ** | 1 |
|  | (08W) | TS BILL | *** | *** | *** | *** | *** | 0 |
|  | (09W) | TS CLARA | *** | *** | *** | *** | *** | 0 |
|  | (10W) | TY DOYIE | 316 | 258 | -173 | 155 | -173 | 16 |
|  | (11W) | TS ELSIE | *** | *** | *** | *** | *** | 0 |
|  | R (11W) | TS ELSIE | *** | *** | *** | *** | *** | 0 |
|  | (12W) | TY FABIAN | 177 | 108 | ** | 124 | ** | 10 |
|  | (13W) | TS GAY | *** | *** | *** | *** | *** | 0 |
|  | (14W) | TY HAL | 249 | 194 | -122 | 130 | -71 | 29 |
|  | (01C) | TY ULEKI | 146 | 92 | 49 | 94 | 42 | 15 |
|  | (15W) | TS IRMA | 94 | 76 | ** | 45 | ** | 10 |
|  | (16W) | TS JEFF | 103 | 91 | ** | 48 | ** | 1 |
|  | (17W) | TS KIT | 120 | 68 | ** | 98 | ** | 4 |
|  | (18W) | TS LEE | 221 | 186 | ** | 89 | ** | 7 |
|  | (19W) | TS MAMIE | *** | *** | *** | *** | *** | 0 |
|  | (20W) | STY NELSON | 128 | 88 | 8 | 70 | -48 | 22 |
|  | (21W) | TY ODESSA | 245 | 217 | ** | 109 | ** | 10 |
|  | (22W) | TY PAT | 303 | 269 | ** | 110 | ** | 9 |
|  | (23W) | TY RUBY | 162 | 121 | -90 | 75 | 13 | 22 |
|  | (24W) | TY SKIP | $199$ | 164 | -84 | 85 | 53 | 21 |
|  | (25W) | TY TESS | 123 | 109 | ** | 53 | ** | 4 |
|  | (26W) | TS VAL | 768 | 75 | ** | 764 | ** | 1 |
|  | TOTALS |  | 216 | 170 | -100 | 103 | -14 | 262 |
|  | * THE MEAN WAS COMPUTED FROM ABSOLUTE VALUES. <br> ** THE MEDIAN WAS NOT COMPUTED FOR INSTANCES OF TEN CASES OR IESS. <br> *** FORECASTS WERE NOT ISSUED OR DID NOT VERIFY. <br> $\mathrm{R}=\mathrm{REGENERATED}$ <br> SEE TABLE 5-1B FOR EXPLANATIONS OF THE TERMS MEAN, MEDIAN, ALONG-TRACK ERROR AND CROSS-TRACK ERROR. |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |


| TABLE 5-1D |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | SUMAARY OF 72-HOUR FORECAST EMRRORS HIESTERN FORTH PACIFIC OCIRAN 1988 SIGNIFICANT TROPICAL CYCLORES |  |  |  |  |  |  |  |
| TROPICAI CYCIONE |  |  | FORECAST ERROR (NM) | $\begin{gathered} \text { ALONG-TRACK } \\ \text { ERROR } \end{gathered}$ |  | $\begin{aligned} & \text { CROSS-TRACK } \\ & \text { ERROR } \end{aligned}$ |  | $\begin{aligned} & \text { SAMPLE } \\ & \text { SIZE } \end{aligned}$ |
|  |  |  | MEAN* | median | MEAN* | MEDIAN |  |
|  | (01W) | TY ROY |  | 401 | 313 | -150 | 202 | 77 | 29 |
|  | (02W) | TY SUSAN | 423 | 385 | ** | 121 | ** | 5 |
|  | (03W) | TD 03W | *** | *** | *** | *** | *** | 0 |
|  | (04W) | TY THAD | 329 | 315 | ** | 75 | ** | 8 |
|  | (05W) | ts vanessa | *** | *** | *** | *** | *** | 0 |
|  | (06W) | TY WARREN | 246 | 193 | -106 | 113 | 96 | 16 |
|  | (07W) | ts Agnes | *** | *** | *** | *** | *** | 0 |
|  | (08W) | TS BILL | *** | *** | *** | *** | *** | 0 |
|  | (09W) | TS CLARA | *** | *** | *** | *** | *** | 0 |
|  | (10W) | TY DOYIE | 626 | 516 | -517 | 297 | -517 | 11 |
|  | (11W) | TS ELSIE | *** | *** | *** | *** | *** | 0 |
|  | R(11W) | ts elsie | *** | *** | *** | *** | *** | 0 |
|  | (12W) | ty fabian | 279 | 256 | ** | 91 | ** | 6 |
|  | (13W) | TS GAY | *** | *** | *** | *** | *** | 0 |
|  | (14W) | TY HAL | 355 | 298 | -124 | 151 | -104 | 25 |
|  | (01C) | ty uleki | 153 | 80 | 78 | 120 | 39 | 11 |
|  | (15W) | TS IRMA | 158 | 39 | ** | 135 | ** | 6 |
|  | (16W) | TS JEFF | *** | *** | *** | *** | *** | 0 |
|  | (17\%) | TS KIT | *** | *** | *** | *** | *** | 0 |
|  | (18W) | TS LEE | 265 | 236 | ** | 104 | ** | 3 |
|  | (19W) | TS MAMIE | *** | *** | *** | *** | *** | 0 |
|  | (20W) | STY NELSON | 148 | 92 | 15 | 103 | -79 | 18 |
|  | (21W) | TY ODESSA | 686 | 544 | ** | 398 | ** | 5 |
|  | (22W) | TY PAT | 436 | 366 | ** | 235 | ** | 5 |
|  | (23W) | TY RUBY | 210 | 112 | -105 | 156 | 28 | 18 |
|  | (24W) | TY SKIP | 260 | 202 | -194 | 143 | 138 | 17 |
|  | (25W) | TY TESS | *** | *** | *** | *** | *** | 0 |
|  | (26W) | TS VAL | *** | *** | *** | *** | *** | 0 |
| totals |  |  | 315 | 244 | -159 | 159 | -11 | 183 |
| * THE MEAN WAS COMPUTED FROM ABSOLUTE VALUES. <br> ** THE MEDIAN WAS NOT COMPUTED FOR INSTANCES OF TEN CASES OR less. *** NOT ENOUGH WARNINGS WERE ISSUED TO VERIFY THE FORECAST. <br> $\mathrm{R}=\mathrm{REGENERATED}$ |  |  |  |  |  |  |  |  |
| SEE TABLE 5-1B FOR EXPLANATIONS OF THE TERMS MEAN, MEDIAN, ALONG-TRACK ERROR AND CROSS-TRACK ERROR. |  |  |  |  |  |  |  |  |





TABIE 5-2A
ANNUAL MIGAN FCRECAST ERRRORS EOR THE WESTERN NORTH PACIFIC

|  | 24-HOUR |  | 48-HOUR |  | 72-HOUR |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| YEAR | FORECAST | RIGHT-ANGLE | FORECAST | RIGHT-ANGLE | FORECA | RIGHT-ANGLE |
| 1971 | 111 | 64 | 212 | 118 | 317 | 117 |
| 1972 | 117 | 72 | 245 | 146 | 381 | 210 |
| 1973 | 108 | 74 | 197 | 134 | 253 | 162 |
| 1974 | 120 | 78 | 226 | 157 | 348 | 245 |
| 1975 | 138 | 84 | 288 | 181 | 450 | 290 |
| 1976 | 117 | 71 | 230 | 132 | 338 | 202 |
| 1977 | 148 | 83 | 283 | 157 | 407 | 228 |
| 1978 | 127 | 75 | 271 | 179 | 410 | 297 |
| 1979 | 124 | 77 | 226 | 151 | 316 | 223 |
| 1980 | 126 | 79 | 243 | 164 | 389 | 287 |
| 1981* | 123 | 75 | 220 | 119 | 334 | 168 |
| 1982* | 113 | 67 | 237 | 139 | 341 | 206 |
| 1983* | 117 | 72 | 259 | 152 | 405 | 237 |
| 1984* | 117 | 66 | 233 | 137 | 363 | 231 |
| 1985* | 117 | 66 | 231 | 134 | 367 | 214 |
| 1986 | 121 | ** | 261 | ** | 394 | ** |
| 1987 | 107 | ** | 204 | ** | 303 | ** |
| 1988 | 114 | ** | 216 | ** | 315 | ** |

[^1]| TABLE $5-2 \mathrm{~B}$ 1988 MEAN FORECAST, ALONG-TRACK AND CF <br>  FOR THE FIESTERN NORTH PACIFIC OCRAN |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | FORECAST | $\begin{gathered} \text { ALONG-TRACK } \\ \text { ERROR } \end{gathered}$ |  | $\begin{gathered} \text { CROSS-TRACK } \\ \text { ERROR } \end{gathered}$ |  |
| FORECAST | ERROR | MEAN | MEDIAN | MEAN | MEDIAN |
| 24-HOUR | 114 | 85 | -45 | 58 | -9 |
| 48-HOUR | 216 | 170 | -100 | 103 | -14 |
| 72-HOUR | 315 | 244 | -159 | 159 | -11 |


| TABLE 5-3 ANNUAL MEAN FORECAST ERRORS (NI |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| IEAB |  | 24-HOUR <br> / TYPHOONS* |  | 48-HOUR <br> / TYPHCONS* |  | 72-HOUR <br> / TYPHOONS* |
| 1959 |  | 117** |  | 267** |  |  |
| 1960 |  | 177** |  | 354** |  |  |
| 1961 |  | 136 |  | 274 |  |  |
| 1962 |  | 144 |  | 287 |  | 476 |
| 1963 |  | 127 |  | 246 |  | 374 |
| 1964 |  | 133 |  | 284 |  | 429 |
| 1965 |  | 151 |  | 303 |  | 418 |
| 1966 |  | 136 |  | 280 |  | 432 |
| 1967 |  | 125 |  | 276 |  | 414 |
| 1968 |  | 105 |  | 229 |  | 337 |
| 1969 |  | 111 |  | 237 |  | 349 |
| 1970 | 104 | 98 | 190 | 181 | 279 | 272 |
| 1971 | 111 | 99 | 212 | 203 | 317 | 308 |
| 1972 | 117 | 116 | 245 | 245 | 381 | 382 |
| 1973 | 108 | 102 | 197 | 1.93 | 253 | 245 |
| 1974 | 120 | 114 | 226 | 218 | 348 | 357 |
| 1975 | 138 | 129 | 288 | 279 | 450 | 442 |
| 1976 | 117 | 117 | 230 | 232 | 338 | 336 |
| 1977 | 148 | 140 | 283 | 266 | 407 | 390 |
| 1978 | 127 | 120 | 271 | 241 | 410 | 459 |
| 1979 | 124 | 113 | 226 | 219 | 316 | 319 |
| 1980 | 126 | 116 | 243 | 221 | 389 | 362 |
| 1981 | 123 | 117 | 220 | 215 | 334 | 342 |
| 1982 | 113 | 114 | 237 | 229 | 341 | 337 |
| 1983 | 117 | 110 | 259 | 247 | 405 | 384 |
| 1984 | 117 | 110 | 233 | 228 | 363 | 361 |
| 1985 | 117 | 112 | 231 | 228 | 367 | 355 |
| 1986 | 121 | 117 | 261 | 261 | 394 | 403 |
| 1987 | 107 | 101 | 204 | 211 | 303 | 318 |
| 1988 | 114 | 107 | 216 | 222 | 315 | 327 |
| * FORECASTS WERE VERIFIED WHEN THE TROPICAL CYCLONE INTENSITIES WERE OVER 35 KT ( $18 \mathrm{M} / \mathrm{SEC}$ ). <br> ** FORECAST POSITIONS NORTH OF 35 DEGREES NORTH LATITUDE WERE NOT VERIFIED. |  |  |  |  |  |  |



Figure 5-3. Annual mean forecast errors (nm) for all significant tropical cyclones in the westem North Pacific.





Figure 5-4. Annual mean forecast errors ( nm ) for all significant tropical cyclones in the North Indian Ocean.

\begin{tabular}{|c|c|c|c|c|c|c|c|c|c|c|c|}
\hline \multicolumn{12}{|l|}{\multirow[t]{2}{*}{TABLE 5-6A

FORECAST ERROR SUMMARY
SOWHA PACIFIC AND SOUI}} <br>
\hline \& \& \& \& \& \& \& \& \& \& \& <br>

\hline \multirow[b]{3}{*}{TROPICAL CYCLONE} \& \multirow[t]{3}{*}{$$
\begin{aligned}
& \text { INITIAL } \\
& \text { POSIT } \\
& \text { ERROR } \\
& \hline
\end{aligned}
$$} \& \multirow[t]{3}{*}{24-HR FCST ERROR} \& \multicolumn{2}{|l|}{\multirow[t]{2}{*}{\[

$$
\begin{gathered}
\text { 24-HOUR } \\
\text { ALONG-TRACK }
\end{gathered}
$$

\]}} \& \multicolumn{2}{|l|}{\multirow[t]{2}{*}{\[

$$
\begin{aligned}
& \text { 24-HOUR } \\
& \text { CROSS-TRACK }
\end{aligned}
$$

\]}} \& \multirow[t]{3}{*}{48-HR FCST ERROR} \& \multicolumn{2}{|l|}{\multirow[t]{2}{*}{\[

$$
\begin{gathered}
\text { 48-HOUR } \\
\text { ALONG-TRACK }
\end{gathered}
$$

\]}} \& \multicolumn{2}{|l|}{\multirow[t]{2}{*}{\[

$$
\begin{gathered}
\text { 48-HOUR } \\
\text { CROSS-TRACK }
\end{gathered}
$$
\]}} <br>

\hline \& \& \& \& \& \& \& \& \& \& \& <br>
\hline \& \& \& MEAN* \& MEDIAN \& MEAN* \& MEDIAN \& \& Mrank \& Median \& MEANA \& MEDIAN <br>
\hline TC 01s \& 33 \& 125 \& 56 \& 9 \& 98 \& 24 \& 216 \& 148 \& 84 \& 137 \& 71 <br>
\hline TC 02S \& 21 \& 57 \& 54 \& ** \& 18 \& ** \& *** \& *** \& *** \& *** \& *** <br>
\hline TC 03S \& 34 \& 128 \& 88 \& ** \& 73 \& ** \& 273 \& 176 \& ** \& 164 \& ** <br>
\hline TC 04P \& 58 \& *** \& *** \& *** \& *** \& *** \& *** \& *** \& *** \& *** \& *** <br>
\hline TC 05s \& 39 \& 158 \& 125 \& ** \& 81 \& ** \& 517 \& 507 \& ** \& 98 \& ** <br>
\hline TC 06P \& 63 \& 242 \& 212 \& -212 \& 94 \& 50 \& 433 \& 364 \& ** \& 193 \& ** <br>
\hline TC 07P \& 25 \& 113 \& 59 \& -6 \& 78 \& -28 \& 198 \& 112 \& ** \& 139 \& ** <br>
\hline TC 08s \& 16 \& 60 \& 58 \& ** \& 13 \& ** \& 75 \& 60 \& ** \& 41 \& ** <br>
\hline TC 09S \& 37 \& 115 \& 76 \& -68 \& 59 \& -12 \& 233 \& 166 \& ** \& 115 \& ** <br>
\hline TC 10S \& 25 \& 179 \& 92 \& ** \& 140 \& ** \& 430 \& 390 \& ** \& 161 \& ** <br>
\hline тC 11s \& 26 \& 133 \& 89 \& -22 \& 79 \& -31 \& 283 \& 182 \& -41 \& 193 \& -42 <br>
\hline TC 12P \& 79 \& 206 \& 109 \& ** \& 55 \& ** \& 234 \& 183 \& ** \& 132 \& ** <br>
\hline TC 13P \& 26 \& 157 \& 122 \& -89 \& 102 \& 31 \& 337 \& 190 \& -66 \& 251 \& 19 <br>
\hline TC 145 \& 21 \& 111 \& 77 \& ** \& 67 \& ** \& 269 \& 184 \& ** \& 153 \& ** <br>
\hline TC 15P \& 58 \& 158 \& 121 \& ** \& 77 \& ** \& 558 \& 487 \& ** \& 222 \& * <br>
\hline TC 16S \& 32 \& 168 \& 117 \& -50 \& 112 \& -10 \& 299 \& 188 \& -56 \& 185 \& 4 <br>
\hline TC 17 S \& 21 \& 205 \& 108 \& ** \& 166 \& ** \& 375 \& 301 \& ** \& 160 \& ** <br>
\hline TC 18S \& 20 \& *** \& *** \& *** \& *** \& *** \& *** \& *** \& *** \& *** \& *** <br>
\hline TC 19P \& 22 \& 118 \& 95 \& ** \& 51 \& ** \& 292 \& 275 \& ** \& 84 \& ** <br>
\hline TC 20 S \& 18 \& 139 \& 11 \& ** \& 139 \& ** \& 277 \& 276 \& ** \& 17 \& ** <br>
\hline TC 21S \& 33 \& 207 \& 191 \& ** \& 66 \& ** \& *** \& *** \& *** \& *** \& *** <br>
\hline AVERAGES \& 34 \& 146 \& 98 \& -63 \& 83 \& 3 \& 290 \& 246 \& -20 \& 144 \& 13 <br>

\hline \multicolumn{12}{|l|}{\multirow[t]{2}{*}{| * the mean was computed from absolute values. |
| :--- |
| ** the median was not computed for instances of ten cases or iess. *** NOT ENOUGH WARNINGS WERE ISSUED TO VERIFY THE FORECAST. |
| SEE TABLE 5-1B FOR EXPLANATIONS OF THE TERMS MEAN, MEDIAN, ALONG-TRACK ERROR AND CROSS-TRACK ERROR. |}} <br>

\hline \& \& \& \& \& \& \& \& \& \& \& <br>
\hline
\end{tabular}

| TABLE 5-6B <br> SOOHH PACIFIC AND SOUMH INDIAN OCEAN NUMRDR OF RARNINGS |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| TRORICAI CYCLONE |  |  | $\begin{aligned} & \text { INITIAL } \\ & \text { POSITION } \end{aligned}$ | 24-HOUR FORECAST | 48-HOUR <br> FORECASI |
|  | TC-01S | ----- | 20 | 18 | 15 |
|  | TC-02S | ----- | 5 | 1 | 0 |
|  | TC-03S | ARINY | 10 | 7 | 6 |
|  | TC-04P | ----- | 2 | 0 | 0 |
|  | TC-05S | BERNANDRO | 11 | 9 | 1 |
|  | TC-06P | AGI | 3 | 1 | 0 |
|  | TC-06P | AGI* | 12 | 5 | 2 |
|  | TC-07P | ANNE | 14 | 6 | 3 |
|  | TC-08S | CALIDERA | 5 | 2 | 1 |
|  | TC-09S | DOAZA | 7 | 3 | 1 |
|  | TC-09S | DOAZA* | 11 | 5 | 2 |
|  | TC-10S | FREDERIC | 6 | 2 | 1 |
|  | TC-11S | GWENDA** | 16 | 7 | 3 |
|  | TC-12P | CHARLIE | 9 | 3 | 2 |
|  | TC-12P | CHARIIE* | 3 | 0 | 0 |
|  | TC-13P | BOLA | 20 | 9 | 4 |
|  | TC-14S | ----- | 6 | 3 | 1 |
|  | TC-15P | CILIA*** | 6 | 2 | 1 |
|  | TC-16S | GASITAO | 16 | 7 | 3 |
|  | TC-17S |  | 7 | 3 | 1 |
|  | TC-18S | HELY | 3 | 1 | 0 |
|  | TC-19P | DOVI | 12 | 5 | 2 |
|  | TC-20S | IARISENA | 3 | 1 | 0 |
|  | TC-21S | ----- | 3 | 1 | 0 |
|  | TOTALS |  | 210 | 101 | 49 |
| * REGENERATED <br> ** ALSO NAMED EZENINA <br> *** NWOC SYSTEM |  |  |  |  |  |


| TABLE 5-7A |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: |
|  | ANNOAL MEAN FORECAST ERRORS (NM) |  |  |  |
|  | SOUTH PACIFIC AND SOULH INDIAN OCBANS |  |  |  |
|  | 24-HOUR |  | 48-HOUR |  |
| YEAR | FORECAST | RIGHT-ANGLE | FORECAST | RIGHT-ANGLE |
| 1981 | 165 | 119 | 315 | 216 |
| 1982 | 144 | 91 | 274 | 174 |
| 1983 | 154 | 84 | 288 | 150 |
| 1984 | 133 | 73 | 231 | 124 |
| 1985 | 138 | 78 | 242 | 133 |
| 1986 | 133 | ** | 268 | ** |
| 1987 | 145 | ** | 280 | ** |
| 1988 | 146 | ** | 290 | ** |
| *** IN 1986, RIGHT-ANGLE ERROR WAS REPLACED BY CROSS-TRACK ERROR. SEE TABLE 5-1B FOR AN EXPLANATION OF CROSS-TRACK ERROR. |  |  |  |  |




Figure 5-5. Annual mean forecast errors (nm) for all significant tropical cyclones in the South Pacific and South Indian Oceans

## 2. COMPARISON OF OBJECTIVE TECHNIQUES

## a. GENERAL

Objective techniques used by JTWC are divided into five main categories:
(1) Extrapolation;
(2) Climatological and Analog Techniques;
(3) Model Output Statistics;
(4) Dynamic Models;
(5) Empirical and Analytical Techniques;

In September 1981, JTWC began to initialize its array of objective forecast techniques (described below) on the six-hour old preliminary best track position (an interpolative process) rather than the forecast (partially extrapolated) warning position, e.g. the 0600 Z warning is now supported by objective techniques developed from the 0000 Z preliminary best track position. This operational change has yielded several advantages:
(1) Techniques can now be requested much earlier in the warning development process, i.e. as soon as the track can be approximated by one or more fix positions after the valid time of the previous warning;
(2) Receipt of these techniques is virtually assured prior to the development of the next waming; and
(3) Improved (mean) forecast accuracy. This latter aspect arises because JTWC now has more reliable approximation of the short-term tropical cyclone movement. Further, since most of the objective techniques are biased towards persistence, this new procedure optimizes their performance and
provides more consistent guidance on shortterm movement, indirectly yielding a more accurate initial position estimate as well as lowering 24 -hour forecast errors.

## b. Description of Objective Techniques

(1) XTRP -- Forecast positions for 24 - and 48 -hours are derived from the extension of a straight line which connects the most recent and 12 -hour old preliminary best track positions.
(2) CLIM -- A climatological aid providing 24-, 48-, and 72-hour tropical cyclone forecast positions (and intensity changes in the western North Pacific) based upon the position of the tropical cyclone. The output is based upon data records from 1945 to 1981 for the western North Pacific Ocean and 1900 to 1981 for the North Indian Ocean.
(3) HPAC -- Forecast positions are generated from a blend of climatology and persistence. The 24-, 48- and 72-hour positions are equally weighted between climatology and persistence. Persistence is a straight line extension of a line connecting the current and 12 -hour old positions. Climatology is based on data from 1945 to 1981 for the western North Pacific Ocean and 1900 to 1981 for the North Indian Ocean.
(4) CLIPER -- A statistical regression technique based on climatology, current intensity, position and past movement. This technique is used as a crude measure of real forecast skill when verifying forecast accuracy.
(5) COSMOS -- A Model Output Statistics (MOS) routine based on the geostrophic steering at the $850-, 700-$, and $500-$ mb levels. The steering is derived from the HATTRACK point advection model run on Global prognostic fields from the FLENUMOCEANCEN's NOGAPS prediction system. The MOS forecast is then blended with the 6hour past movement to generate the forecast
track.
(6) Colorado State University Model (CSUM) -- A statistical-dynamic method developed by Matsumoto (1984) utilizes synoptic and persistence predictors by discretizing the forecast timeframe into three 24 -hour time steps. Climatology is incorporated into the forecast via a stratification scheme based on the position of the tropical cyclone relative to the 500 mb subtropical ridge. Depending on whether the tropical cyclone is south, on, or north of the ridge, three sets of regression equations are used to determine the north-south and east-west displacements
(7) One-way Interactive Tropical Cyclone Model (OTCM) -- A coarse-mesh, three-layer in the vertical, primitive equation model with a 205 km grid spacing over a 6400 x 4700 km domain. The model's fields are computed around a bogused, digitized cyclone vortex using FLENUMOCEANCEN's Numerical Variational Analysis (NVA) or NOGAPS prognostic fields for the specified valid time. The past motion of the tropical cyclone is compared to initial steering fields and a bias correction is computed and applied to the model. FLENUMOCEANCEN's NOGAPS global prognostic fields are used at 12-hour intervals to update the model's boundaries. The resultant forecast positions are derived by locating the 850 mb vortex at six-hour intervals to 72 -hours.
(8) TAPT -- An empirical technique which utilizes upper-tropospheric wind fields to estimate acceleration associated with the tropical cyclone's interaction with the midlatitude westerlies. It includes guidelines for the duration of acceleration, upper-limits, and probable path of the cyclone.
(9) TYAN -- An updated analog program which combines the earlier versions TYFN 75 and INJAN 74. The program scans a 30-year climatology with a similar history (within a specified acceptance envelope) to the current tropical cyclone. For the western North

Pacific Ocean, three forecasts of position and intensity are provided for $24-$, 48 -, and $72-$ hours: RECR - a weighted mean of all tropical cyclones which were categorized as "recurving" during their best track period; STRA - a weighted mean of all accepted tropical cyclones which were categorized as moving "straight" (westward) during their best track period: TOTL - a weighted mean of all accepted tropical cyclones, including those used in the RECR and STRA forecast. For the North Indian Ocean, a single (total) forecast track is provided for the 12 -hour intervals to 72 -hours.
(10) DVORAK -- An estimation of tropical cyclone's current and 24 -hour forecast intensity is made from interpolation of satellite imagery (DVORAK, 1984) and provided to the forecaster. These intensity estimates are used in conjunction with other intensity related data and trends to forecast tropical cyclone intensity.
(11) HOLLAND/MARTIN -- The technique adapts an earlier work (Holland, 1980) and specifically addresses the need for realistic $30-, 50-$ and $100-\mathrm{kt}$ wind radii around tropical cyclones. It solves equations for basic gradient wind relations within the tropical cyclone area, using input parameters obtained from enhanced infrared satellite imagery. For the first time, diagnosis also includes an asymmetric area of winds caused by tropical cyclone movement. Size and intensity parameters are also used to diagnose internal steering components of tropical cyclone motion known collectively as "Beta-drift". The Holland/Martin wind radii technique replaces the more general Huntley (1980) technique.

## c. Testing and Results

A comparison of selected techniques is included in Table 5-8 for all western North Pacific tropical cyclones, Table 5-9 for all North Indian Ocean tropical cyclones and Table 5-10 for the southern hemisphere. In these tables, " x-axis " refers to techniques listed vertically. For example (Table 5-8) in the 273 cases available for a (homogeneous) comparison, the
average forecast error at 24 -hours was 130 nm ( 241 km ) for TOTL and $137 \mathrm{~nm}(254 \mathrm{~km}$ ) for RECR. The difference of $7 \mathrm{~nm}(13 \mathrm{~km})$ is shown in the lower right. (Differences are not
always exact, due to computational round-off which occurs for each of the cases available for comparison).

TABLE 5-9

## 1988 ERROR STATISTICS FOR SELECTED OBIECTIVE TECHNIQUES <br> IN THE NORTHERN INDIAN OCEAN

## 24-HOUR MEAN FORECAST ERROR (NM)



## 48-HOUR MEAN FORECAST ERROR (NM)

|  | JTWC |  | OTCM |  | csum |  | HPAC |  | CLIM |  | XTRP |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| JTWC | $\begin{array}{r} 18 \\ 219 \end{array}$ | $\begin{array}{r} 219 \\ 0 \end{array}$ |  |  |  |  |  |  |  |  |  |  |  |
| OTCM | $\begin{array}{r} 18 \\ 227 \\ 17 \end{array}$ | $\begin{array}{r} 219 \\ 8 \\ 220 \end{array}$ | 20 211 <br> 19 | $\begin{array}{r} 211 \\ 0 \end{array}$ <br> 216 | 19 |  |  |  |  |  |  |  | - orricial mict ronecast <br> - owl-may tropical cyclowe hoozl <br> - COLORNDO STMTE UMIVLRSITY MCOLL <br> - toinl amalos (tyan 78 ) <br> - climitoiocy <br>  <br> - 12-bcon extrapolation |
|  | 576 | 356 | 539 | 323 | 539 |  |  |  |  |  |  |  |  |
| HPAC | 18 | 219 | 20 | 211 | 19 | 539 | 20 | 202 |  |  |  |  |  |
|  | 185 | -34 | 202 | -9 | 209 | -330 | 202 | 0 |  |  |  |  |  |
| CLIM | 18 | 219 | 20 | 211 | 19 | 539 | 20 | 202 | 20 | 162 |  |  |  |
|  | 141 | -78 | 162 | -49 | 165 | -374 | 162 | -40 | 162 |  |  |  |  |
| XTRP | 17 | 225 | 19 | 220 | 18 | 527 | 19 | 208 | 19 | 167 | 19 | 265 |  |
|  | 251 | 26 | 265 | 45 | 277 | -250 | 265 | 57 | 265 | 98 | 265 |  |  |

72-HOUR MEAN FORECAST ERROR (NM)

|  | JTWC |  | отСм |  |  | SUM | HPAC |  | CLIM |  | XTRP |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| JTWC | $\begin{array}{r} 12 \\ 409 \end{array}$ |  |  |  |  |  |  |  |  |  |  |  |
| отСм | $\begin{array}{r} 11 \\ 527 \end{array}$ | $\begin{aligned} & 426 \\ & 201 \end{aligned}$ | $\begin{array}{r} 11 \\ 527 \end{array}$ | $\begin{array}{r} 527 \\ 0 \end{array}$ |  |  |  |  |  |  |  |  |
| csum | $\begin{array}{r} 11 \\ 1043 \end{array}$ | $\begin{aligned} & 413 \\ & 630 \end{aligned}$ | $\begin{array}{r} 10 \\ 1034 \end{array}$ | $\begin{aligned} & 540 \\ & 494 \end{aligned}$ | $\begin{array}{r} 11 \\ 1043 \end{array}$ | $\begin{array}{r} 1043 \\ 0 \end{array}$ |  |  |  |  |  |  |
| HPAC | $\begin{array}{r} 12 \\ 298 \end{array}$ | $\begin{array}{r} 409 \\ -111 \end{array}$ |  | $\begin{array}{r} 527 \\ -216 \end{array}$ |  | $\begin{aligned} & 1043 \\ & -730 \end{aligned}$ | $\begin{array}{r} 12 \\ 298 \end{array}$ | 298 0 |  |  |  |  |
| CLIM | $\begin{array}{r} 12 \\ 231 \end{array}$ | $\begin{array}{r} 409 \\ -178 \end{array}$ |  | $\begin{array}{r} 527 \\ -292 \end{array}$ |  | $\begin{array}{r} 1043 \\ -810 \end{array}$ | $\begin{array}{r} 12 \\ 231 \end{array}$ | $\begin{aligned} & 298 \\ & -67 \end{aligned}$ | $\begin{array}{r} 12 \\ 231 \end{array}$ | $\begin{array}{r} 231 \\ 0 \end{array}$ |  |  |
| XTRP | 11 399 | 428 -29 |  | 557 -136 | 10 433 | -1037 | 11 399 | 312 87 | 11 399 | $\begin{aligned} & 237 \\ & 162 \end{aligned}$ | 11 399 | 399 |



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## CHAPTER VI - TROPICAL CYCLONE SUPPORT SUMMARY

# The Pocket Tropical Cyclone Model (PTCM) 

(Evans, J. L., Monash University, Australia and R. J. Miller, NAVENVPREDRSCHFAC)

The PTCM is a linear tropical cyclone motion prediction scheme incorporating the effects of large-scale environmental flow and west-northwestward propagation due to the rate of change of the Earth's vorticity over a given distance or (Beta-effect). The model is based on equations developed by Holland (1983) and has been operational in a modified form in the Australian region for a number of years. The current version of the model has been modified to accept FNOC data and produce up to 72-hour forecasts. The model is ready for testing.

The PTCM is being incorporated in the NEPRF Tropical Cyclone Forecast Simulation package which will be a resident program in the ATCF.

## Tropical Cyclone Motion Diagnostic Tool

(Chu, J. H., NAVENVPREDRSCHFAC)

Based on the barotropic vorticity equation, a set of "Typhoon Motion Equations" has been developed. These equations can be used to diagnose interactions between the tropical cyclone and its environment. The equations provide forecasters with information on the impact of environmental changes on tropical cyclone motion. The equations have been tested using input of theoretical or analytical fields. Plans have been made to adapt these equations to FNOC fields and incorporate them in the Tropical Cyclone Forecast Simulation Package.

## The Advanced Tropical Cyclone Model (ATCM)

(Hodur, R. M., NAVENVPREDRSCHFAC)

The ATCM was run in 1988 for use by JTWC forecasters. As in 1987, the ATCM again exhibited a strong northward bias. This bias was particularly evident with storms located deep in the tropics, despite numerous changes made after the 1987 tropical cyclone season. Also, erratic behavior was noticed on some forecasts of reasonably well-behaved storms. A description of problems discovered in the 1988 version of the ATCM follows.

The first problem concerns the data assimilation strategy used in the ATCM. Although the ATCM is not run until about six and a half hours after observation time, only those observations received up to three and a half hours after observation time were used. The ATCM used an assimilation cycle of 12hours, as opposed to 6-hours for NOGAPS. This meant that all late and off-time data was not used by the ATCM, resulting in degraded analyses. In 1989, the ATCM will use either a later data cut-off time or NOGAPS analyses. Second, the radiation parameter-ization in the ATCM caused excessive cooling of the atmosphere and produced noise in the model. Noise due to clouds appears to contribute the largest error to tropical cyclone track forecasts. An improved treatment of clouds, based on that used by the European Center, is being tested in the ATCM and will be implemented prior to the 1989 tropical cyclone season. Test runs of several 1988 storms, using NOGAPS analyses and no clouds, have shown dramatic improvement in forecasting tropical cyclone tracks. Finally, the ATCM exhibits a northward bias during the first 6 - to 12 -hours of a forecast. This occurs even without physical parameterizations and has been related to either biases in the initial fields or the structure of the bogus circulation. A similar tendency has been noted by the National Meteorological Center for tropical cyclones in the Atlantic Ocean. Experiments with simple basic flows are being performed to isolate the cause of this erroneous movement.

# Navy Tactical Applications Guide (NTAG), Volume 6 

(Fett, R. W., NAVENVPREDRSCHFAC)

An effort is now underway to develop a series of examples demonstrating the use of high quality satellite data for analysis and forecasting in the tropics. Data from polar orbiting and geostationary satellites are used to study the evolution of certain weather effects or of a particular weather phenomenon at a given time. These examples are organized and collated for publishing in the NTAG Volume 6. NTAG Volume 6, Part I, "Tropical Weather Analysis and Forecast Applications," was distributed in June 1986. Part II, "Tropical Cyclones," is scheduled for completion in January 1989. In March 1989, work will begin on the development of the "Tropical Cyclone Forecaster's Handbook."

## Automated Tropical Cyclone Forecasting System

## (Tsui, T. L., R. J. Miller, and A. J. Schrader, NAVENVPREDRSCHFAC)

The Automated Tropical Cyclone Forecasting (ATCF) system is an IBM PC compatible software package currently being developed for JTWC. The ATCF is designed to allow JTWC forecasters to graphically display tropical cyclone forecast information, merge and analyze synoptic wind fields, provide objective fix guidance, select optimum objective forecast aids, and expedite the issuance of tropical cyclone warnings. One great advantage of the ATCF is standardization of tropical cyclone forecasting procedures. During the course of preparing a tropical cyclone warning, forecasters will avoid neglecting decisional steps or available options. The ATCF automatically saves all tropical cyclone data, computes real-time and post storm statistics, and allows forecasters to randomly access any past storm data. A communications package included in the ATCF simplifies data transfer between JTWC and FNOC.

The ATCF was installed at JTWC in January 1988 and activated for operational use in June 1988. The system was also installed at the AJTWC (located at NWOC) during October 1988. The system software has been provided to OAO Corporation for inclusion in the JTWC Automation Project.

# North Pacific Tropical Cyclone Climatology 

## (Miller, R. J., T. L. Tsui and A. J. Schrader, NAVENVPREDRSCHFAC)

A tropical cyclone climatology for the North Pacific has been compiled and reviewed by EGPACOM and published by NEPRF. Data used for the western basin were taken from the JTWC Tropical Cyclone data base and covered a 40 year period from 1945 to 1984. Eastern basin data spanned a 34 year period from 1949 to 1982 and were obtained from the consolidated world-wide tropical cyclone data base at the National Climatic Data Center in Ashville, North Carolina. Tropical cyclones for both basins were sorted by day and month into twenty four 31-day overlapping periods. For each period, four charts are supplied: 1) actual storm paths; 2) mean storm paths; 3) average storm speed; and 4) storm constancy and frequency.

## EOF Post-Processing Forecast Technique

(Chu, J. H., R. J. Miller and T. L. Tsui, NAVENVPREDRSCHFAC)

NEPRF has adapted the Empirical Orthogonal Function (EOF) tropical cyclone post-processing forecast scheme on the FNOC computer system. This EOF technique, developed by the Naval Postgraduate School, objectively recognizes salient patterns of largescale horizontal wind fields relative to the center of a tropical cyclone. This information, in terms of EOF coefficients, is used via regression equations to modify tropical cyclone track forecasts produced by numerical models. Test results will be compiled after the 1988 western North Pacific tropical cyclone season.

## 1. GENERAL

Due to the rapid growth of microcomputers in the meteorological community and to save publishing costs, tropical cyclone track data (with best track, initial warning, 24-, 48and 72-hour JTWC forecasts) and fix data (satellite, aircraft, radar and synoptic) are now available separately upon request. The data will be in ASCII format on 5.25 inch "floppy" diskettes. The data sets are available on two diskettes. These include the western North Pacific Ocean (1 January - 31 December 1988) on one and North Indian Ocean (1 January - 31 December), South Western Pacific and South Indian Oceans (1 July 1987-30 June 1988) on the other. Agencies or individuals desiring these data sets should send the appropriate number of "floppy" diskettes (two if both data
sets are desired) to NAVOCEANCOMCEN/JTWC Guam with their request. When the request is received, the data will be copied onto your diskettes and returned with the explanation of the data formats. The use of floppy diskettes should facilitate the transfer of these rather large data files to your computer.

## 2. WARNING VERIFICATION STATISTICS

## a. WESTERN NORTH PACIFIC

This section includes verification statistics for each warning in the western North Pacific Ocean during 1988. Pre- and postwarning best track positions are not printed, but are available on floppy diskettes by request.

JTWC EORECAST TRACK AND INIENSITY ERRORS BY HARNING

| THPHOON | (01W) |  | O0h |  | 24h | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 13 | 123 | 246 | 401 |  |  |  |  |  |
|  |  |  | \# Cases | 41 | 37 | 33 | 29 |  |  |  |  |  |
| DTG | WH | BT IAT | BT ION | POS ER | 24 ER | 48 ER | 72 ER | BT WN | WW ER | 24.ER | 48 ER | 72 ER |
| 88010800 | 1 | 7.9 N | 171.5E | 29.7 | 119.1 | 84.8 | 127.2 | 30 | 0 | -10 | -20 | -35 |
| 88010806 | 2 | 8.0 N | 170.3 E | 32.0 | 80.3 | 109.0 | 197.7 | 35 | 0 | -5 | -25 | -25 |
| 88010812 | 3 | 8.1 N | 169.2 E | 13.4 | 78.0 | 213.1 | 251.8 | 45 | -5 | -5 | -40 | -20 |
| 88010818 | 4 | 8.2N | 168.OE | 24.7 | 122.4 | 253.7 | 245.3 | 50 | 0 | 0 | -30 | -5 |
| 88010900 | 5 | 8.2N | 166.7E | 6.0 | 142.4 | 293.3 | 295.2 | 55 | 0 | -10 | -30 | -10 |
| 88010906 | 6 | 8.3N | 165.3E | 13.3 | 165.8 | 270.1 | 296.4 | 60 | 0 | -20 | -25 | -10 |
| 88010912 | 7 | 8.3N | 163.7E | 23.7 | 225.0 | 299.5 | 325.6 | 65 | 0 | -35 | -15 | -5 |
| 88010918 | 8 | $8.4 N$ | 161.6E | 30.3 | 177.5 | 236.7 | 337.3 | 70 | -5 | -35 | -10 | -5 |
| 88011000 | 9 | 8.8 N | 159.4 E | 13.4 | 97.6 | 254.4 | 492.3 | 85 | -5 | -10 | 5 | 15 |
| 88011006 | 10 | 9.2 N | 157.3E | 8.4 | 63.1 | 290.0 | 519.5 | 95 | 0 | 5 | 15 | 15 |
| 88011012 | 11 | 9.9N | 155.1E | 21.4 | 130.2 | 349.8 | 608.9 | 115 | -10 | 10 | 15 | 20 |
| 88011018 | 12 | 10.6 N | 153.0E | 12.0 | 243.1 | 438.9 | 645.3 | 115 | -5 | 15 | 15 | 25 |
| 88011100 | 13 | 11.4 N | 151.5E | 25.2 | 209.8 | 391.6 | 652.9 | 115 | -5 | -10 | -10 | -5 |
| 88011106 | 14 | 11.7 N | 150.2E | 6.0 | 140.9 | 309.1 | 480.9 | 110 | -5 | -10 | -15 | 0 |
| 88011112 | 15 | 12.3N | 149.1E | . 0 | 18.5 | 101.2 | 281.9 | 110 | 0 | 0 | 0 | 0 |
| 88011118 | 16 | 12.9N | 147.9E | 5.8 | 11.6 | 64.3 | 296.1 | 105 | 5 | -10 | -5 | -5 |
| 88011200 | 17 | 13.1 N | 146.7E | 6.0 | 34.7 | 102.2 | 351.6 | 110 | -5 | -10 | 0 | -5 |
| 88011206 | 18 | 13.8 N | 145.6E | 11.7 | 36.7 | 198.8 | 318.2 | 110 | -5 | 0 | 10 | -5 |
| 88011212 | 19 | 14.0N | 144.3E | 11.6 | 141.1 | 124.2 | 117.7 | 110 | 0 | 0 | 5 | -10 |
| 88011218 | 20 | 14.6 N | 143.3 E | . 0 | 89.3 | 239.8 | 416.4 | 110 | 0 | 5 | 0 | -5 |
| 88011300 | 21 | 15.2N | 142.4E | 13.3 | 122.2 | 344.0 | 566.0 | 105 | 0 | 5 | 0 | 5 |
| 88011306 | 22 | 15.3N | 141.7E | 8.3 | 214.0 | 534.6 | 866.4 | 105 | -5 | 0 | -15 | 0 |
| 88011312 | 23 | 15.2N | 141.0 E | 11.6 | 225.7 | 479.6 | 739.8 | 100 | 0 | 5 | -10 | 20 |
| 88011318 | 24 | 14.7 N | 140.2E | 8.3 | 193.5 | 465.0 | 645.8 | 95 | 5 | 0 | -5 | 30 |
| 88011400 | 25 | 14.0N | 139.1 E | 13.3 | 234.4 | 502.5 | 659.6 | 90 | 0 | -10 | 0 | 25 |


| 88011406 | 26 | 13.2N | 137.4 E | 37.0 | 234.2 | 468.5 | 633.7 | 85 | 0 | -15 | 0 | 25 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88011412 | 27 | 12.6 N | 135.8 E | 8.4 | 64.7 | 46.3 | 76.0 | 85 | -5 | -20 | 0 | 0 |
| 88011418 | 28 | 12.3 N | 133.9 E | 11.7 | 77.0 | 24.7 | 50.0 | 90 | -10 | -15 | 10 | 5 |
| 88011500 | 29 | 12.4 N | 131.7E | 6.0 | 93.5 | 146.3 | 145.8 | 90 | 0 | -5 | 10 | 5 |
| 88011506 | 30 | 12.6 N | 129.7E | 6.0 | 117.0 | 169.2 | N/A | 90 | 0 | -5 | 15 | N/A |
| 88011512 | 31 | 12.8 N | 127.7E | . 0 | 110.8 | 145.8 | N/A | 90 | -5 | 10 | 15 | N/A |
| 88011518 | 32 | 13.0N | 125.9 E | 13.1 | 157.5 | 163.0 | N/A | 85 | -5 | 20 | 15 | N/A |
| 88011600 | 33 | 13.1N | 124.6E | 8.4 | 24.7 | 16.7 | N/A | 75 | -5 | 10 | 10 | N/A |
| 88011606 | 34 | 13.1 N | 123.3 E | 26.3 | 81.6 | N/A | N/A | 65 | 0 | 15 | N/A | N/A |
| 88011612 | 35 | 13.7N | 122.2E | 6.0 | 87.6 | N/A | N/A | 50 | 10 | 10 | N/A | N/A |
| 88011618 | 36 | 13.8 N | 121.0 E | 6.0 | 94.9 | $\mathrm{N} / \mathrm{A}$ | N/A | 40 | 5 | 10 | N/A | N/A |
| 88011700 | 37 | 13.4 N | 119.4 E | 8.4 | 96.2 | N/A | N/A | 40 | 0 | 5 | N/A | N/A |
| 88011706 | 38 | 13.7 N | 118.0E | 18.5 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88011712 | 39 | 13.8 N | 116.6E | 23.3 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |
| 88011718 | 40 | 14.0 N | 115.4 E | . 0 | $\mathrm{N} / \mathrm{A}$ | N/A | N/A | 25 | 0 | N/A | N/A | N/A |
| 88011800 | 41 | 13.9N | 114.5 E | 8.4 | N/A | N/A | N/A | 25 | 0 | N/A | N/A | N/A |





| Tropical | Storm Vanessa (05W) |  |
| :--- | ---: | ---: |
|  |  |  |
|  |  |  |
| DTG | W\# | BT LIAT |
| 88062612 | 1 | 7.6 N |
| 88062618 | 2 | 8.3 N |
| 88062700 | 3 | 9.1 N |
| 88062706 | 4 | 9.9 N |
| 88062712 | 5 | 10.7 N |
| 88062718 | 6 | 11.4 N |
| 88062800 | 7 | 12.3 N |
| 88062806 | 8 | 13.6 N |
| 88062812 | 9 | 15.2 N |
| 88062818 | 10 | 17.3 N |
| 88062900 | 11 | 19.4 N |


| 00 h | 24 h | 48 h | 72 h |
| ---: | ---: | ---: | ---: |
| 27 | 194 | 386 | $\mathrm{~N} / \mathrm{A}$ |
| 11 | 7 | 3 | 0 |

Average
\# Cases

| POS ER | $24 . E R$ | 48 EB | 22 ER | BT WN | WW ER | 24_WE | 48 WE | 72 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 34.8 | 228.1 | 486.2 | N/A | 30 | - | 10 | 5 | N/A |
| 6.0 | 125.9 | 322.2 | N/A | 35 | -5 | 0 | -10 | N/A |
| 13.3 | 170.0 | 352.1 | N/A | 40 | -5 | -10 | -5 | N/A |
| 23.6 | 194.1 | N/A | N/A | 45 | -10 | -10 | N/A | N/A |
| 23.6 | 186.2 | N/A | N/A | 35 | 0 | -10 | N/A | N/A |
| 30.0 | 227.1 | N/A | N/A | 35 | 0 | 0 | N/A | N/A |
| 42.6 | 226.7 | N/A | N/A | 40 | 0 | 15 | N/A | N/A |
| 63.1 | N/A | N/A | N/A | 40 | 0 | N/A | N/A | N/A |
| 29.0 | N/A | N/A. | N/A | 40 | 0 | N/A | N/A | N/A |
| 18.9 | N/A | N/A | N/A | 40 | 0 | N/A | N/A | N/A |
| 21.3 | N/A | N/A | N/A | 35 | 5 | N/A | N/A | N/A |



| 88071518 | 13 | 13.6 N | 134.9 E | 18.5 | 106.7 | 217.3 | 272.8 | 90 | -5 | -35 | -40 | -30 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | ---: | ---: | ---: | ---: |
| 88071600 | 14 | 14.2 N | 133.5 E | 25.1 | 115.2 | 190.9 | 359.3 | 90 | 0 | -20 | -5 | 10 |
| 88071606 | 15 | 14.8 N | 131.9 E | 34.7 | 140.0 | 190.3 | 277.3 | 95 | -5 | -15 | -5 | -10 |
| 88071612 | 16 | 15.2 N | 130.2 E | 16.7 | 80.0 | 109.6 | 145.9 | 100 | -5 | -30 | -20 | -5 |
| 88071618 | 17 | 15.9 N | 128.5 E | 8.3 | 48.3 | 97.4 | $\mathrm{~N} / \mathrm{A}$ | 110 | 0 | -25 | -25 | $\mathrm{~N} / \mathrm{A}$ |
| 88071700 | 18 | 16.5 N | 127.0 E | 18.9 | 36.4 | 90.9 | $\mathrm{~N} / \mathrm{A}$ | 115 | 0 | 20 | 20 | $\mathrm{~N} / \mathrm{A}$ |
| 88071706 | 19 | 16.9 N | 125.4 E | 5.7 | 34.5 | 68.7 | $\mathrm{~N} / \mathrm{A}$ | 110 | 5 | 5 | 15 | $\mathrm{~N} / \mathrm{A}$ |
| 88071712 | 20 | 17.8 N | 124.0 E | 5.7 | 87.3 | 73.9 | $\mathrm{~N} / \mathrm{A}$ | 110 | -5 | -25 | -20 | $\mathrm{~N} / \mathrm{A}$ |
| 88071718 | 21 | 18.4 N | 122.6 E | 5.7 | 32.8 | 6.0 | $\mathrm{~N} / \mathrm{A}$ | 105 | -10 | -20 | -10 | $\mathrm{~N} / \mathrm{A}$ |
| 88071800 | 22 | 18.8 N | 121.6 E | 16.5 | 45.6 | 73.8 | $\mathrm{~N} / \mathrm{A}$ | 95 | -10 | 15 | 5 | $\mathrm{~N} / \mathrm{A}$ |
| 88071806 | 23 | 19.3 N | 120.5 E | 13.3 | 54.3 | 126.1 | $\mathrm{~N} / \mathrm{A}$ | 90 | -10 | 25 | 5 | $\mathrm{~N} / \mathrm{A}$ |
| 88071812 | 24 | 20.0 N | 119.5 E | .0 | 44.7 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 85 | -5 | 20 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88071818 | 25 | 20.8 N | 118.1 E | 28.8 | 66.0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 80 | 0 | 30 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88071900 | 26 | 21.7 N | 117.0 E | 21.2 | 29.0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 75 | 5 | 15 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88071906 | 27 | 22.7 N | 116.3 E | 18.0 | 170.2 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 65 | 10 | 5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88071912 | 28 | 23.5 N | 115.5 E | 25.1 | 183.2 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 50 | 15 | 10 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88071918 | 29 | 24.2 N | 114.4 E | 24.0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 40 | 10 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88072000 | 30 | 25.0 N | 113.5 E | 21.0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 30 | 10 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |


| Tropical Storm Agnes (07W) |  | 00 h | $\frac{24 \mathrm{~h}}{}$ | $\frac{48 \mathrm{~h}}{}$ | $\frac{72 \mathrm{~h}}{}$ |
| ---: | :--- | ---: | ---: | ---: | ---: | ---: |
|  | Average | 33 | 322 | 714 | $\mathrm{~N} / \mathrm{A}$ |
|  | \# Cases | 8 | 5 | 1 | 0 |


| DTG | WH | BT LAT | BT LON | POS ER | 24EE | 48 ER | 72 ER | BT WN | Wh ER | 24. WE | 48. WE | 72 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88072900 | 1 | 20.9N | 140.7E | 36.4 | 295.8 | 714.1 | N/A | 30 | 0 | -5 | 5 | N/A |
| 88072906 | 2 | 21.8 N | 140.6E | 26.5 | 380.0 | N/A | N/A | 30 | 0 | 0 | N/A | N/A |
| 88072912 | 3 | 22.8 N | 140.8 E | 39.7 | 401.7 | N/A | N/A | 35 | -5 | -5 | N/A | N/A |
| 88072918 | 4 | 24.0 N | 141.7E | 81.4 | 427.2 | N/A | N/A | 40 | -10 | 0 | N/A | N/A |
| 88073000 | 5 | 25.7N | 142.4 E | 5.4 | 109.1 | N/A | N/A | 45 | -5 | 5 | N/A | N/A |
| 88073006 | 6 | 27.9N | 142.7 E | 29.1 | N/A | N/A | N/A | 40 | 0 | N/A | N/A | N/A |
| 88073012 | 7 | 29.9 N | 142.6 E | 13.1 | N/A | N/A | N/A | 35 | 5 | N/A | N/A | N/A |
| 88073018 | 8 | 31.7 N | 143.1 E | 35.5 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |


| Tropical Storm Bill (08W) |  | O0h | 24h | 48h | 72h |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | Average | 12 | 119 | N/A | N/A |
|  | * Cases | 5 | 3 | 0 | 0 |


| DTG | W违 | Bt LAT | BT ION | POS ER | $24 . \mathrm{ER}$ | 48 EB | 72 EB | BT.WN | WW ER | 24 WE | 48 WE | 72.WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88080700 | 1 | 27.4N | 125.6E | 10.7 | 111.2 | N/A | N/A | 35 | - | 5 | N/A | N/A |
| 88080706 | 2 | 28.2N | 124.2E | 6.0 | 123.5 | N/A | N/ $/$ A | 45 | 0 | 5 | N/A | N/A |
| 88080712 | 3 | 28.9N | 122.7 E | 6.0 | 123.8 | N/A | N/A | 45 | 0 | 15 | N/A | N/A |
| 88080718 | 4 | 29.7N | 121.0E | 18.7 | N/A | N/A | N/A | 40 | -5 | N/A | N/A | N/A |
| 88080800 | 5 | 30.7N | 119.4E | 20.8 | N/A | N/A | N/A | 30 | 0 | N/A | N/A | N/A |


| Tropical Storm Clara (09W) |  |  |  | $\begin{array}{r} 00 \mathrm{~h} \\ \hline 24 \\ 6 \end{array}$ | $\begin{array}{r} 24 \mathrm{~h} \\ 89 \\ 3 \end{array}$ | $\begin{array}{r} 48 h \\ N / A \\ 0 \end{array}$ | $\begin{array}{r} 72 \mathrm{~h} \\ \mathrm{~N} / \mathrm{A} \\ 0 \end{array}$ | BT WN | HM ER | 24 WE | 48 WE | 72 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average <br> * Cases |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |  |  |
| DTG | W | BT LAT | BT ION | POS ER | 24 ER | 48 ER | 72 ER |  |  |  |  |  |
| 88081018 | 1 | 28.7N | 160.1E | 10.5 | 107.6 | N/A | N/A | 40 | -5 | 0 | N/A | N/A |
| 88081100 | 2 | 28.5N | 160.3 E | 19.8 | 73.1 | N/A | N/A | 45 | -5 | 5 | N/A | N/A |
| 88081106 | 3 | 28.4N | 160.5E | 35.6 | 86.5 | N/A | N/A | 45 | 0 | 0 | N/A | N/A |
| 88081112 | 4 | 28.5N | 160.5 E | 18.8 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88081118 | 5 | 28.9N | 160.6 E | 16.9 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |
| 88081200 | 6 | 29.5N | 160.9E | 44.8 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |



| Tropical Storm Elsia (11W) |  | 00 h | 24 h | 48 h | 72 h |  |
| :---: | :---: | ---: | ---: | ---: | ---: | ---: |
|  |  | Average | 31 | 206 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
|  | Cases | 10 | 4 | 0 | 0 |  |


| DTG | W | BT LAT | BT ION | POS ER | 24_ER | 48 ER | 72 ER | BT WN | WW ER | 24.WE | 48 WE | 72 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88082812 | 1 | 21.9N | 158.2E | 5.6 | 362.3 | N/A | N/A | 35 | 0 | 20 | N/A | N/A |
| 88082818 | 2 | 22.3 N | 159.2 E | 47.5 | 205.2 | N/A | N/A | 35 | 0 | 15 | N/A | N/A |
| 88082900 | 3 | 23.1N | 159.9E | 82.0 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88082906 | 4 | 24.2N | 160.8 E | 40.7 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88082912 | 5 | 25.4N | 161.3 E | 16.2 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88082918 | 6 | 26.7 N | 161.6 E | 13.1 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88083100 | 7 | 32.6 N | 158.8 E | 26.0 | 132.7 | N/A | N/A | 45 | -5 | 0 | N/A | N/A |
| 88083106 | 8 | 34.0 N | 157.3E | 6.0 | 124.6 | N/A | N/A | 45 | 0 | 0 | N/A | N/A |
| 88083112 | 9 | 35.8 N | 156.0E | 25.9 | N/A | N/A | N/A | 40 | -5 | N/A | N/A | N/A |
| 88083118 | 10 | 37.1 N | 155.4 E | 51.6 | N/A | N/A | N/A | 35 | -5 | N/A | N/A | N/A |


| Typhoon | ian | (12W) |  | 00 h | 24h | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average <br> \# Cases | 21 | 141 | 177 | 279 |  |  |  |  |  |
|  |  |  |  | 18 | 14 | 10 | 6 |  |  |  |  |  |
| DTG | W考 | BT_LAT | BT LON | POS ER | 24_ER | 48 ER | 72.ER | BT WN | WW ER | 24 WE | 48 WE | 72. WE |
| 88083000 | 1 | 31.1 N | 144.8 E | 23.7 | 178.8 | 449.8 | 534.1 | 40 | 0 | 0 | -5 | -35 |
| 88083006 | 2 | 31.0 N | 145.5 E | 23.7 | 143.9 | 405.4 | 458.0 | 45 | -5 | 5 | -5 | -40 |
| 88083012 | 3 | 31.0 N | 146.4E | 11.9 | 72.7 | 84.6 | 312.0 | 50 | -5 | 5 | 0 | -5 |
| 88083018 | 4 | 31.1 N | 147.5E | 15.8 | 26.1 | 75.0 | 200.2 | 55 | 0 | 20 | 10 | -10 |
| 88083100 | 5 | 31.2 N | 148.6E | 13.1 | 36.4 | 92.1 | 46.0 | 55 | 0 | 15 | 5 | -5 |
| 88083106 | 6 | 31.4 N | 149.8 E | 15.8 | 115.8 | 182.1 | 125.8 | 55 | 0 | 0 | -10 | 0 |
| 88083112 | 7 | 31.4 N | 150.7 E | 19.5 | 146.2 | 117.3 | N/A | 55 | -10 | -30 | -45 | N/A |
| 88083118 | 8 | 31.4 N | 151.4E | 23.7 | 128.2 | 90.3 | N/A | 55 | -10 | -25 | -40 | N/R |
| 88090100 | 9 | 31.4 N | 152.2E | 30.7 | 103.9 | 119.3 | N/A | 60 | 0 | -25 | -35 | N/A |


| 88090106 | 10 | 31.4 N | 153.2E | 31.3 | 184.2 | 162.0 | N/A | 65 | -10 | -30 | -20 | N/A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88090112 | 11 | 31.7 N | 153.7E | 11.8 | 166.5 | N/A | N/A | 65 | -10 | -20 | N/A | N/A |
| 88090118 | 12 | 32.4 N | 153.8 E | 28.0 | 77.2 | N/A | N/A | 65 | -10 | -35 | N/A | N/A |
| 88090200 | 13 | 33.2 N | 153.8 E | 13.0 | 245.2 | N/A | N/A | 70 | 5 | 15 | N/A | N/A |
| 88090206 | 14 | 34.0 N | 153.7E | . 0 | 359.1 | N/A | N/A | 75 | 0 | 15 | N/A | N/A |
| 88090212 | 15 | 35.5 N | 153.9 E | . 0 | N/A | N/A | N/A | 75 | 0 | N/A | N/A | N/A |
| 88090218 | 16 | 36.8 N | 154.7E | 15.6 | N/A | N/A | N/A | 75 | -10 | N/A | N/A | N/A |
| 88090300 | 17 | 39.0 N | 156.9E | 4.7 | N/A | N/A | N/A | 70 | -5 | N/A | N/A | N/A |
| 88090306 | 18 | 41.3 N | 160.0E | 102.0 | N/A | N/A | N/A | 55 | -10 | N/A | N/A | N/A |

Tropical Storm Gay (13W)

| DTG | W\# | BT.LAT |
| :--- | ---: | ---: |
| 88090218 | 1 | 27.5 N |
| 88090300 | 2 | 28.4 N |
| 88090306 | 3 | 29.4 N |
| 88090312 | 4 | 30.5 N |
| 88090318 | 5 | 31.6 N |
| 88090400 | 6 | 32.7 N |


|  | $\frac{00 h}{}$ | 24 h | 48 h | 72 h |
| :--- | ---: | ---: | ---: | ---: |
| Average | 33 | 98 | N/A | N/A |
| \# Cases | 6 | 3 | 0 | 0 |



| 88091606 | 34 | $37.0 N$ | $146.6 E$ | 22.6 | N/A | N/A | N/A | 60 | -5 | N/A | N/A | N/A |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 88091612 | 35 | $38.4 N$ | $149.1 E$ | 38.1 | N/A | N/A | N/A | 55 | -5 | N/A | N/A | N/A |
| 88091618 | 36 | $39.7 N$ | $152.4 E$ | 23.8 | N/A | N/A | N/A | 50 | -5 | N/A | N/A | N/A |
| 88091700 | 37 | $40.4 N$ | $156.3 E$ | 37.7 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |


| Typhoon Olaki (01C) |  |  | Average \# Cases | $\begin{array}{r} \frac{006}{17} \\ 21 \end{array}$ | $\begin{array}{r} 24 h \\ 82 \\ 19 \end{array}$ | $\begin{array}{r} 48 \mathrm{~h} \\ \hline 146 \\ 15 \end{array}$ | $\begin{array}{r} 72 \mathrm{~h} \\ \hline 153 \\ 11 \end{array}$ | BT WN | WW ER | 24 WE | 48 WE | 72 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |  |  |
| DTG | W* | BT LAT | BT LON | POS ER | 24 ER | 48 ER | 72 ER |  |  |  |  |  |
| 88090806 | 1 | 25.2N | 179.8W | 6.0 | 39.4 | 63.6 | 78.7 | 90 | 0 | -10 | -20 | -20 |
| 88090812 | 2 | 25.6 N | 178.7 E | 18.8 | 90.7 | 226.0 | 332.9 | 90 | 0 | -10 | -20 | -15 |
| 88090818 | 3 | 26.1N | 177.3E | 17.2 | 27.8 | 105.3 | 122.7 | 90 | 0 | -10 | -20 | -15 |
| 88090900 | 4 | 26.7 N | 176.1E | 20.1 | 36.7 | 70.9 | 120.1 | 90 | -5 | -20 | -30 | -20 |
| 88090906 | 5 | 27.2 N | 175.2E | 8.0 | 93.8 | 139.5 | 212.2 | 90 | -5 | -20 | -20 | -20 |
| 88090912 | 6 | 27.5N | 174.3E | . 0 | 94.1 | 129.8 | 230.7 | 90 | 5 | -5 | -5 | -10 |
| 88090918 | 7 | 27.6N | 173.5E | 16.0 | 55.9 | 87.7 | 136.9 | 90 | 0 | -15 | -10 | -10 |
| 88091000 | 8 | 27.6N | 172.8E | 12.2 | 86.6 | 98.7 | 126.4 | 90 | -10 | -25 | -15 | -15 |
| 88091006 | 9 | 27.8 N | 172.2E | 18.8 | 99.5 | 125.8 | 65.2 | 90 | -15 | -10 | -5 | -5 |
| 88091012 | 10 | 28.2N | 171.4 E | 19.9 | 70.2 | 119.4 | 119.6 | 90 | -10 | 0 | 0 | 5 |
| 88091018 | 11 | 28.8N | 170.7E | 6.0 | 72.6 | 150.4 | 139.8 | 90 | -10 | -10 | -10 | 10 |
| 88091100 | 12 | 29.4 N | 170.3E | 10.5 | 97.4 | 302.0 | N/A | 90 | 0 | -10 | -15 | N/A |
| 88091106 | 13 | 29.8 N | 169.7 E | 26.0 | 151.4 | 398.9 | N/A | 80 | 10 | 0 | -5 | N/A |
| 88091112 | 14 | 30.2 N | 169.3E | 13.1 | 39.0 | 100.6 | N/A | 75 | 0 | 10 | 15 | N/A |
| 88091118 | 15 | 30.6 N | 169.0 E | 11.9 | 63.0 | 74.7 | N/A | 75 | 0 | 0 | 15 | N/A |
| 88091200 | 16 | 31.0 N | 168.6 E | 7.9 | 102.0 | N/A | N/A | 70 | 5 | -5 | N/A | N/A |
| 88091206 | 17 | 31.2 N | 168.1 E | 23.7 | 145.9 | N/A | N/A | 70 | 0 | 0 | N/A | N/A |
| 88091212 | 18 | 31.2 N | 167.6E | 48.3 | 123.7 | N/A | N/A | 65 | 0 | -5 | N/A | N/A |
| 88091218 | 19 | 31.2 N | 167.1 E | 24.5 | 73.9 | N/A | N/A | 65 | -5 | 0 | N/A | N/A |
| 88091300 | 20 | 31.3 N | 166.4 E | 26.3 | N/A | N/A | N/A | 60 | -5 | N/A | N/A | N/A |
| 88091306 | 21 | 31.6 N | 165.4 E | 26.3 | N/A | N/A | N/A | 55 | -5 | N/A | N/A | N/A |


| Tropical | Storm Irma | (15W) |  | OOh | 24h | 48h | 72h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 17 | 76 | 94 | 158 |  |  |  |  |  |
|  |  |  | \# Cases | 16 | 14 | 10 | 6 |  |  |  |  |  |
| DTG | W ${ }^{\text {\# }}$ | BT IAT | BTION | POS ER | 24 ER | 48 ER | 72 EB | BT WN | WW EB | 24.WE | 48.WE | 22 WE |
| 88091200 | 1 | 22.6N | 160.1 E | 21.1 | 91.9 | 60.9 | 84.7 | 40 | 0 | 5 | 0 | 0 |
| 88091206 | 2 | 23.2N | 159.9E | 12.0 | 80.5 | 78.2 | 93.9 | 45 | 0 | 5 | 5 | 5 |
| 88091212 | 3 | 23.6N | 159.6E | 25.1 | 52.0 | 24.1 | 57.8 | 45 | 0 | -5 | -10 | -5 |
| 88091218 | 4 | 23.9N | 159. OE | 17.5 | 54.4 | 140.0 | 210.0 | 45 | 0 | -5 | -10 | 0 |
| 88091300 | 5 | 24.4N | 158.4 E | 6.0 | 134.5 | 214.6 | 238.3 | 45 | 0 | 0 | -10 | 10 |
| 88091306 | 6 | 24.9N | 158.OE | 6.0 | 87.6 | 157.2 | 264.3 | 50 | 5 | 5 | 5 | 20 |
| 88091312 | 7 | 25.2N | 157.9E | 12.4 | 34.2 | 72.4 | N/A | 55 | 0 | 5 | 10 | N/A |
| 88091318 | 8 | 25.7N | 157.8E | 36.4 | 131.1 | 143.8 | N/A | 55 | 0 | 5 | 15 | N/A |
| 88091400 | 9 | 26.3 N | 157.6E | 16.1 | 57.6 | 13.1 | N/A | 55 | 0 | 10 | 25 | N/A |
| 88091406 | 10 | 26.6N | 157.4E | 17.2 | 45.9 | 36.4 | N/A | 55 | 0 | 15 | 30 | N/A |
| 88091412 | 11 | 27.0k | 157.0E | 18.8 | 67.8 | N/A | N/A | 55 | 0 | -10 | N/A | N/A |
| 88091418 | 12 | 27.3N | 156.7E | 5.3 | 39.8 | N/A | N/A | 55 | 0 | -5 | N/A | N/A |
| 88091500 | 13 | 27.6N | 156.4E | 8.0 | 69.2 | N/A | N/A | 55 | -5 | 0 | N/A | N/A |
| 88091506 | 14 | 28.3N | 156.0E | 36.7 | 124.0 | N/A | N/A | 50 | 0 | 15 | N/A | N/A |
| 88091512 | 15 | 28.9N | 155.7E | 15.9 | N/A | N/A | N/A | 45 | -5 | N/A | N/A | N/A |
| 88091518 | 16 | 29.8N | 155.6E | 19.7 | N/A | N/A | N/A | 40 | -5 | N/A | N/A | N/A |


| Tropical | orm | (16W) | Average <br> * Cases | $\begin{array}{r} \frac{00 h}{15} \\ 9 \end{array}$ | $\begin{array}{r} 24 \mathrm{~h} \\ 64 \\ 5 \end{array}$ | $\frac{48 \mathrm{~h}}{103}$ | $\begin{array}{r} 72 h \\ \mathrm{~N} / \mathrm{A} \\ 0 \end{array}$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| DTG | W4 | BT_LAT | BT ION | POS ER | 24 ER | 48 ER | 72 ER | BT WN | WW ER | 24.WE | 48 WE | 72 WE |
| 88091400 | 1 | 16.3 N | 136.7E | 21.4 | 77.0 | 103.3 | N/A | 35 | 0 | -5 | 10 | N/A |
| 88091406 | 2 | 17.1N | 137.5E | 12.9 | 62.0 | N/A | N/A | 40 | 5 | 5 | N/A | N/A |
| 88091412 | 3 | 18.0N | 138.2E | 8.3 | 79.6 | N/A | N/A | 45 | 0 | 10 | N/A | N/A |
| 88091418 | 4 | 19.0 N | 138.9E | 8.3 | 60.0 | N/A | N/A | 45 | 0 | 15 | N/A | N/A |
| 88091500 | 5 | 20.1N | 139.5E | 24.7 | 45.0 | N/A | N/A | 45 | -5 | 5 | N/A | N/A |
| 88091506 | 6 | 21.4 N | 140.0E | 30.4 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 88091512 | 7 | 22.8 N | 140.5 E | . 0 | N/A | N/A | N/A | 35 | 5 | N/A | N/A | N/A |
| 88091518 | 8 | 24.0N | 141.0E | 18.0 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88091600 | 9 | 25.3N | 141.5E | 13.2 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |





Super Typhoon Nelson (20W)

|  | 00 h | 24 h | 48 h | 72 h |
| :--- | ---: | ---: | ---: | ---: |
| Average | 10 | 64 | 128 | 148 |
| \# Cases | 30 | 26 | 22 | 18 |


| DTG | 且 | Bt lat | BT ION | POS ER | $24.8 B$ | 48 ER | 72 EB | BT WN | Wh EB | 24 WE | 48_WE | 72. WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88100112 | 1 | 12.2 N | 136.1E | 21.3 | 110.9 | 215.5 | 243.5 | 45 | 0 | - | -25 | -60 |
| 88100118 | 2 | 12.4 N | 135.0E | 24.0 | 133.7 | 191.1 | 190.9 | 45 | 0 | -5 | -40 | -60 |
| 88100200 | 3 | 12.8 N | 134.0E | 5.9 | 122.1 | 211.6 | 191.4 | 45 | 0 | -10 | -40 | -55 |
| 88100206 | 4 | 13.4N | 133.0E | 5.8 | 89.1 | 132.1 | 54.0 | 55 | -5 | -20 | -50 | -60 |
| 88100212 | 5 | 14.1N | 131.8 E | 8.4 | 37.4 | 63.8 | 58.3 | 65 | -10 | -15 | -50 | -50 |
| 88100218 | 6 | 14.7N | 130.6 E | 12.0 | 62.4 | 105.9 | 105.5 | 70 | -10 | -25 | -50 | -45 |
| 88100300 | 7 | 15.5 N | 129.4 E | 6.0 | 36.0 | 74.0 | 96.6 | 75 | 0 | -25 | -50 | -45 |
| 88100306 | 8 | 16.1N | 128.4 E | 8.3 | 61.6 | 56.0 | 107.1 | 90 | -10 | -35 | -50 | -45 |
| 88100312 | 9 | 16.8 N | 127.5E | 8.3 | 41.2 | 81.9 | 127.8 | 100 | -10 | -35 | -40 | -35 |
| 88100318 | 10 | 17.5N | 126.6E | 8.3 | 17.9 | 78.8 | 168.0 | 115 | -15 | -20 | -10 | -15 |
| 88100400 | 11 | 18.5 N | 125.8 E | 13.3 | 22.4 | 84.2 | 168.1 | 120 | -5 | -5 | 0 | 0 |
| 88100406 | 12 | 19.3N | 125.2E | 5.7 | 110.8 | 143.1 | 144.1 | 130 | -5 | -5 | -15 | -25 |
| 88100412 | 13 | 20.0 N | 124.7E | 5.6 | 30.0 | 123.9 | 162.9 | 140 | 0 | -5 | -5 | -10 |
| 88100418 | 14 | 20.7N | 124.4 E | 8.2 | 75.4 | 187.8 | 230.8 | 140 | 0 | 0 | -10 | -10 |
| 88100500 | 15 | 21.3 N | 124.4 E | . 0 | 68.1 | 172.9 | 99.4 | 140 | 0 | -5 | -15 | 5 |
| 88100506 | 16 | 21.9N | 124.7E | 8.2 | 71.9 | 141.6 | 197.2 | 140 | 0 | 0 | -25 | -5 |
| 88100512 | 17 | 22.5N | 125.1 E | 11.1 | 76.1 | 101.9 | 141.2 | 140 | -10 | -15 | -25 | -15 |
| 88100518 | 18 | 23.3N | 125.9E | 12.0 | 99.7 | 150.0 | 180.6 | 130 | -10 | -25 | -25 | -25 |
| 88100600 | 19 | 24.0 N | 126.7E | 11.0 | 60.9 | 82.9 | N/A | 125 | -10 | -15 | -20 | N/A |
| 88100606 | 20 | 24.6N | 127.7E | 13.2 | 49.7 | 75.5 | N/A | 115 | -5 | -25 | -20 | N/A |
| 88100612 | 21 | 25.3N | 128.9 E | 6.0 | 81.9 | 231.5 | N/A | 115 | -10 | -10 | -5 | N/A |
| 88100618 | 22 | 25.9N | 130.0 E | 13.2 | 33.5 | 115.0 | N/A | 115 | -15 | -10 | -5 | N/A |
| 88100700 | 23 | 26.6 N | 131.1 E | . 0 | 52.3 | N/A | N/A | 115 | -15 | 5 | N/A | N/A |
| 88100706 | 24 | 27.6 N | 132.2 E | 8.0 | 41.4 | N/A | N/A | 115 | -20 | -15 | N/A | N/A |
| 88100712 | 25 | 28.5N | 133.6 E | 8.0 | 24.5 | N/A | N/A | 100 | -10 | -5 | N/A | N/A |
| 88100718 | 26 | 29.5 N | 135.0 E | 10.4 | 56.5 | N/A | N/A | 90 | 0 | -5 | N/A | N/A |
| 88100800 | 27 | 30.4N | 136.5 E | 12.0 | N/A | N/A | N/A | 85 | -5 | N/A | N/A | N/A |
| 88100806 | 28 | 31.4 N | 138.3 E | 15.8 | N/A | N/A | N/A | 75 | -5 | N/A | N/A | N/A |
| 88100812 | 29 | 32.5 N | 140.6E | 26.0 | N/A | N/A | N/A | 70 | -5 | N/A | N/A | N/A |
| 88100818 | 30 | 33.6 N | 143.6E | 23.3 | N/A | N/A | N/A | 70 | -15 | N/A | N/A | N/A |


| Typhoon Odessa (21w) |  |  | Average <br> * Cases | $\begin{array}{r} \frac{00 \mathrm{~h}}{16} \\ 22 \end{array}$ | $\begin{array}{r} 24 \mathrm{~h} \\ 100 \\ 18 \end{array}$ | $\begin{array}{r} 48 \mathrm{~h} \\ \hline 245 \\ 10 \end{array}$ | $\begin{array}{r} 72 h \\ 686 \\ 5 \end{array}$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |  |  |
| DTG | W\# | BT LAT |  | BT LON | POS ER | 24 EB | 48 EB | 72 ER | BT WN | WW ER | 24 WE | 48 WE | 72 WE |
| 88101112 | 1 | 19.0 N | 132.5 E | 6.0 | 110.2 | 354.0 | 640.8 | 35 | 0 | 10 | -5 | -25 |
| 88101118 | 2 | 18.9 N | 131.3 E | 13.3 | 182.9 | 447.9 | 749.7 | 35 | 5 | 10 | 0 | -25 |
| 88101200 | 3 | 18.9N | 130.4 E | 6.0 | 191.3 | 475.2 | 778.0 | 40 | 5 | 20 | 25 | 0 |
| 88101206 | 4 | 19.2N | 129.6 E | 21.3 | 243.1 | 545.6 | 853.1 | 45 | 0 | -10 | -10 | -25 |
| 88101212 | 5 | 19.8 N | 129.4 E | . 0 | 57.4 | 216.0 | 408.6 | 45 | 0 | -10 | -25 | -25 |
| 88101218 | 6 | 20.5N | 129.2E | 12.7 | 17.7 | N/A | N/A | 45 | -5 | -35 | N/A | N/A |
| 88101300 | 7 | 21.1 N | 129.0E | 13.2 | 190.7 | N/A | N/A | 45 | -5 | -40 | N/A | N/A |
| 88101306 | 8 | 21.6 N | 128.9E | 5.6 | 39.1 | 48.6 | N/A | 65 | 5 | -25 | -60 | N/A |
| 88101312 | 9 | 21.9 N | 129.2E | 11.1 | 56.3 | N/A. | 1*/A | 65 | 0 | -45 | N/A | N/A |


| 88101318 | 10 | 22.1N | 129.5 E | 11.1 | 81.4 | N/A | N/A | 65 | 0 | -45 | N/A | N/A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88101400 | 11 | 22.4 N | 130.0 E | 5.5 | 60.0 | 58.2 | N/A | 65 | 5 | -40 | -55 | N/A |
| 88101406 | 12 | 23.0 N | 130.5 E | . 0 | 42.0 | 88.0 | N/A | 75 | -5 | -35 | -30 | N/A |
| 88101412 | 13 | 23.6 N | 130.9E | 8.1 | 37.6 | 91.4 | N/A | 90 | -10 | 5 | 15 | N/A |
| 88101418 | 14 | 24.2 N | 131.3 E | . 0 | 17.3 | 132.0 | N/A | 90 | 0 | 20 | 45 | N/A |
| 88101500 | 15 | 24.7 N | 131.8 E | . 0 | 36.9 | N/A | N/A | 90 | 0 | 0 | N/A | N/A |
| 88101506 | 16 | 25.2 N | 132.4 E | 5.4 | 93.6 | N/A | N/A | 90 | 0 | 5 | N/A | N/A |
| 88101512 | 17 | 25.6 N | 132.9E | 8.1 | 159.3 | N/A | N/A | 90 | 0 | 15 | N/A | N/A |
| 88101518 | 18 | 25.8 N | 133.3 E | . 0 | 191.4 | N/A | N/A | 90 | -5 | 20 | N/A | N/A |
| 88101600 | 19 | 26.1N | 133.4 E | 22.4 | N/A | N/A | N/A | 85 | 0 | N/A | N/A | N/A |
| 88101606 | 20 | 26.3 N | 133.1E | 74.0 | N/A | N/A | N/A | 75 | 0 | N/A | N/A | N/A |
| 88101612 | 21 | 26.4N | 132.8 E | 117.0 | N/A | N/A | N/A | 65 | -10 | N/A | N/A | N/A |
| 88101618 | 22 | 26.5 N | 132.4 E | 16.1 | N/A | N/A | N/A | 55 | -20 | N/A | N/A | N/A |




| 88102600 | 19 | 16.5 N | 117.1E | 36.0 | 130.6 | 164.8 | N/A | 65 | 5 | 15 | 35 | N/A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88102606 | 20 | 16.8 N | 116.6 E | 29.6 | 141.2 | 164.8 | N/A | 65 | 0 | 0 | 15 | N/A |
| 88102612 | 21 | 17.2N | 115.8 E | 50.1 | 131.5 | 176.5 | N/A | 65 | 0 | 5 | 20 | N/A |
| 88102618 | 22 | 17.6N | 115.0E | 22.9 | 77.0 | 222.5 | N/A | 65 | -5 | 0 | 15 | N/A |
| 88102700 | 23 | 17.9N | 114.3 E | 58.1 | 256.1 | N/A | N/A | 60 | 0 | -15 | N/A | N/A |
| 88102706 | 24 | 18.2 N | 113.3 E | 18.0 | 51.1 | N/A | N/A | 60 | 0 | 5 | N/A | N/A |
| 88102712 | 25 | 18.4 N | 112.5 E | 16.5 | 87.0 | N/A | N/A | 55 | 0 | 5 | N/A | N/A |
| 88102718 | 26 | 18.5 N | 111.7 E | 18.0 | 108.7 | N/A | N/A | 50 | 0 | 10 | N/A | N/A |
| 88102800 | 27 | 18.8 N | 110.9E | 20.8 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 88102806 | 28 | 18.7N | 110.1 E | 36.1 | N/A | N/A | N/A | 35 | 5 | N/A | N/A | N/A |
| 88102812 | 29 | 18.4N | 109.3E | 54.3 | N/A | N/A | N/A | 30 | 5 | N/A | N/A | N/A |
| 88102818 | 30 | 18.2N | 108.5E | 79.8 | N/A | N/A | N/A | 25 | 5 | N/A | N/A | N/A |



| Typhoon I | (25w) |  | 00h |  | 24h | 48h | 72 h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 14 | 57 | 123 | N/A |  |  |  |  |  |
|  |  |  | \# Cases | 10 | 8 | 4 | 0 |  |  |  |  |  |
| DTG | W\# | ET IAAT | BT ION | ROS ER | 24EB | 48 EB | 72.ER | BT WN | WN_EB | 24.WE | 48 WE | 72. We |
| 88110400 | 1 | 9.2 N | 119.3E | 18.8 | 59.7 | 183.0 | N/A | 30 | 0 | -20 | -25 | N/A |
| 88110406 | 2 | 9.3N | 118.5E | 23.7 | 87.7 | 181.9 | N/A | 35 | 0 | -5 | 0 | N/A |
| 88110412 | 3 | 9.4 N | 117.5E | 11.8 | 63.2 | 85.5 | N/A | 35 | 0 | -20 | 10 | N/A |
| 88110418 | 4 | 9.6 N | 116.6 E | 5.9 | 18.9 | 43.0 | N/A | 40 | 0 | -10 | 0 | N/A |
| 88110500 | 5 | 9.7 N | 115.8 E | 6.0 | 72.4 | N/A | N/A | 55 | 0 | 5 | N/A | N/A |
| 88110506 | 6 | 9.8 N | 114.5 E | 29.6 | 119.6 | N/A | N/A | 50 | 5 | 15 | N/A | N/A |


| 88110512 | 7 | 10.2N | 113.3E | 8.4 | 25.2 | N/A | N/A | 65 | 0 | 30 | N/A | N/A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88110518 | 8 | 10.5 N | 112.3 E | 13.4 | 13.2 | N/A | N/A | 65 | 0 | 15 | N/A | N/A |
| 88110600 | 9 | 10.8 N | 111.2 E | 8.4 | N/A | N/A | N/A | 65 | 0 | N/A | N/A | N/A |
| 88110606 | 10 | 11.0 N | 110.2E | 21.5 | N/A | N/A | N/A | 55 | 0 | N/A | N/A | N/A |


| Tropical | rm | (26W) | Average <br> \# Cases | $\begin{array}{r} 00 \mathrm{~h} \\ 57 \\ 10 \end{array}$ | $\begin{array}{r} 24 \mathrm{~h} \\ 228 \\ 6 \end{array}$ | $\begin{array}{r} 48 \mathrm{~h} \\ 768 \\ 1 \end{array}$ | $\begin{array}{r} 72 \mathrm{~h} \\ \mathrm{~N} / \mathrm{A} \\ 0 \end{array}$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| DTG | W | BT LAT | BT ION | POS ER | 24.88 | 48 ER | 72 ER | BT WN | WN ER | 24.WE | 48 WE | 72. WE |
| 88122218 | 1 | 11.9 N | 129.5 E | 21.5 | 128.9 | N/A | N/A | 30 | 0 | 0 | N/A | N/A |
| 88122300 | 2 | 13.0 N | 127.6E | 67.8 | 390.5 | 767.9 | N/A | 30 | 5 | -25 | -10 | N/A |
| 88122306 | 3 | 13.8 N | 126.2E | 42.9 | 367.7 | N/A | N/A | 35 | 0 | -25 | N/A | N/A |
| 88122312 | 4 | 14.3 N | 125.7E | 8.4 | 186.3 | N/A | N/A | 40 | -5 | -15 | N/A | N/A |
| 88122318 | 5 | 14.6 N | 125.3E | 6.0 | 107.0 | N/A | N/A | 45 | 5 | 25 | N/A | N/A |
| 88122400 | 6 | 14.7N | 125.0E | 47.9 | 187.7 | N/A | N/A | 55 | 0 | -5 | N/A | N/A |
| 88122406 | 7 | 14.8 N | 124.7E | 90.2 | N/A | N/A | N/A | 55 | 0 | N/A | N/A | N/A |
| 88122412 | 8 | 14.9 N | 124.5E | 93.5 | N/A | N/A | N/A | 50 | 0 | N/A | N/A | N/A |
| 88122418 | 9 | 14.9 N | 124.4 E | 88.8 | N/A | N/A | N/A | 45 | 0 | N/A | N/A | N/A |
| 88122500 | 10 | 15.0 N | 124.3E | 108.3 | N/A | N/A | N/A | 40 | 0 | N/A | N/A | N/A |

## b. NORTH INDIAN OCEAN

This section includes verification statistics for each warning in the North Indian

Ocean during 1988. Pre- and post- warning best track positions are not printed, but are available on floppy diskettes by request.

JHC FORECAST TRACK AND INTHNSITY ERRORS BY MARNING

Tropical Cyclone 01A

| DTG | W\# | BT LAT |
| :--- | ---: | ---: |
| 88061000 | 1 | 16.9 N |
| 88061006 | 2 | 16.8 N |
| 88061012 | 3 | 16.6 N |
| 88061018 | 4 | 16.3 N |
| 88061100 | 5 | 16.2 N |
| 88061106 | 6 | 16.1 N |
| 88061112 | 7 | 16.2 N |
| 88061118 | 8 | 16.3 N |
| 88061200 | 9 | 16.4 N |
| 88061206 | 10 | 16.6 N |

Tropical Cyclone 02B


| Tropical Cyclone 04B |  |  |
| :--- | ---: | ---: |
|  |  |  |
|  |  |  |
| DTG | WH | BT LAT |
| 88112406 | 1 | 10.0 N |
| 88112412 | 2 | 10.4 N |
| 88112418 | 3 | 10.7 N |
| 88112500 | 4 | 11.0 N |
| 88112506 | 5 | 11.1 N |
| 88112512 | 6 | 11.1 N |
| 88112518 | 7 | 11.2 N |
| 88112600 | 8 | 11.3 N |
| 88112606 | 9 | 11.6 N |
| 88112612 | 10 | 12.1 N |

$\begin{array}{lrrrr} & \frac{00 h}{} & 24 \mathrm{~h} & \frac{48 \mathrm{~h}}{} & \frac{72 \mathrm{~h}}{\mathrm{~N}} \\ \text { Average } & 61 & 188 & 484 & 2 \\ \text { \# Cases } & 10 & 6 & 2 & 0\end{array}$

| BT_LON | POS ER | 24ER | 48 EB | I2_ER | BT NN | WN ER | 24 ER |
| ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 69.4 E | 183.8 | 360.8 | 577.5 | $\mathrm{~N} / \mathrm{A}$ | 35 | 0 | -10 |
| 69.0 E | 86.4 | 253.1 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | -10 |
| 68.6 E | 81.3 | 210.2 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | -10 |
| 68.7 E | 84.0 | 251.7 | 390.1 | $\mathrm{~N} / \mathrm{A}$ | 35 | 0 | -10 |
| 69.0 E | 29.2 | 29.2 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | -10 |
| 69.5 E | 33.3 | 23.8 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | 0 |
| 69.8 E | 29.2 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | $\mathrm{~N} / \mathrm{A}$ |
| 70.2 E | 26.6 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | $\mathrm{~N} / \mathrm{A}$ |
| 70.4 E | 34.0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 35 | 0 | $\mathrm{~N} / \mathrm{A}$ |
| 70.7 E | 31.1 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 25 | 10 | $\mathrm{~N} / \mathrm{A}$ |


| 48 ER | $72.5 R$ |
| :---: | :---: |
| -10 | N/A |
| N/A | N/A |
| N/A | N/A |
| -5 | N/A |
| N/A | N/A |
| N/A | N/A |
| N/A | N/A |
| N/A | N/A |
| N/A | N/A |


|  | $00 h$ | 24 h | 48 h | 72 h |
| :--- | ---: | ---: | ---: | ---: |
| Average | 30 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| \# Cases | 3 | 0 | 0 | 0 |


| DTG | Hi* | BT LAT | ET ION | ROS ER | 24 EB | 48_EB | 72.ER | BT WN | WW ER | 24 ER | 48_ER | 72 ER |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88101812 | 1 | 19.6 N | 91.3E | 25.6 | N/A | N/A | N/A | 35 | 0 | N/A | N/A | N/A |
| 88101818 | 2 | 20.8N | 90.6E | 13.2 | N/A | N/A | N/A | 30 | 5 | N/A | ' $\mathrm{N} / \mathrm{A}$ | N/A |
| 88101900 | 3 | 22.2N | 89.9E | 53.1 | N/A | N/A | N/A | 30 | 5 | N/A | N/A | N/A |


|  | 00 h | 24 h | 48 h | 72 h |
| :--- | ---: | ---: | ---: | ---: |
| Average | 28 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| \# Cases | 3 | 0 | 0 | 0 |


| BT LON | POS ER | 24 ER | 48 ER | 72 ER | BT WN | WW ER | 24 EB | 48 EB | 22 ER |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 90.9 E | 17.1 | N/A | N/A | N/A | 40 | 0 | N/A | N/A | N/A |
| 91.7E | 34.5 | N/A | N/A | N/A | 50 | 0 | N/A | N/A | N/A |
| 93.0E | 33.5 | N/A | N/A | N/A | 55 | 0 | N/A | N/A | N/A |


|  | 00 h | 24 h | 48 h | 72 h |
| :--- | ---: | ---: | ---: | ---: |
| Average | 22 | 90 | 186 | 409 |
| \# Cases | 22 | 20 | 16 | 12 |


| 88112618 | 11 | 12.7 N | 88.3 E | 46.4 | 133.8 | 307.3 | 663.2 | 75 | 0 | 10 | 5 | -60 |
| :--- | :--- | :--- | :--- | :--- | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 88112700 | 12 | 13.2 N | 88.1 E | 30.0 | 56.7 | 216.4 | 508.8 | 75 | 5 | 5 | 0 | 20 |
| 88112706 | 13 | 13.7 N | 88.0 E | 13.3 | 25.0 | 96.3 | $\mathrm{~N} / \mathrm{A}$ | 75 | 0 | -15 | -45 | $\mathrm{~N} / \mathrm{A}$ |
| 88112712 | 14 | 14.2 N | 87.9 E | 29.4 | 44.1 | 151.7 | $\mathrm{~N} / \mathrm{A}$ | 80 | -5 | -20 | -55 | $\mathrm{~N} / \mathrm{A}$ |
| 88112718 | 15 | 14.7 N | 87.7 E | 35.3 | 41.8 | 147.8 | $\mathrm{~N} / \mathrm{A}$ | 85 | 0 | -10 | -30 | $\mathrm{~N} / \mathrm{A}$ |
| 88112800 | 16 | 15.3 N | 87.7 E | 36.7 | 69.0 | 168.4 | $\mathrm{~N} / \mathrm{A}$ | 90 | 0 | -15 | 0 | $\mathrm{~N} / \mathrm{A}$ |
| 88112806 | 17 | 15.8 N | 87.7 E | 8.3 | 115.3 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 90 | 0 | -15 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88112812 | 18 | 16.6 N | 87.8 E | 29.6 | 196.0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 90 | 0 | -30 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88112818 | 19 | 17.7 N | 88.0 E | 51.0 | 181.1 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 95 | -5 | -20 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88112900 | 20 | 19.0 N | 88.1 E | 18.0 | 101.6 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 100 | 0 | 5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88112906 | 21 | 20.4 N | 88.3 E | 12.8 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 105 | 5 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 88112912 | 22 | 21.8 N | 89.0 E | 6.0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | 110 | -25 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |



## c. SOUTHERN HEMISPHERE

This section includes verification statistics for each warning in the South Indian and western South Pacific Oceans from 1 July

1987 to 30 June 1988. Pre- and post- warning best track positions are not printed, but are available on floppy diskettes by request.

JTHC FORECAST TRACK AND INTINNSITY ERRORS BY WARNING


| Tropical Cyclone 02S |  |  | Average <br> * Cases | $\begin{array}{r} 00 \mathrm{~h} \\ 20 \\ 5 \end{array}$ | 24 h | $\begin{array}{r} 48 \mathrm{~h} \\ \mathrm{~N} / \mathrm{A} \\ 0 \end{array}$ | BT WN | HW ER | 24-WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | 57 |  |  |  |  |  |  |
|  |  |  | 1 |  |  |  |  |  |  |
| DTG | W* | BT LAT |  | BT ION | ROS ER | 24.ER |  |  |  |  | 48 ER |
| 87112406 | 1 | 10.75 |  | 41.5E | 16.8 | N/A | N/A | 35 | 0 | N/A | N/A |
| 87112418 | 2 | 11.5 s | 41.1 E | 26.7 | 56.8 | N/A | 40 | 5 | 10 | N/A |
| 87112506 | 3 | 12.2 S | 40.6 E | . 0 | N/A | N/A | 35 | 0 | N/A | N/A |
| 87112518 | 4 | 13.25 | 40.0E | 46.2 | N/A | N/A | 25 | 5 | N/A | N/A |
| 87112606 | 5 | 14.5 S | 39.6 E | 13.3 | N/A | N/A | 20 | 5 | N/A | N/A |


| Tropical Cyclone 035 |  |  | Average \# Cases | $\frac{00 h}{34}$ | 24h | 48h |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | 127 |  | 272 |  |  |  |  |
|  |  |  |  | 10 | 7 | 6 |  |  |  |  |
| DIG | W\# | BT IAT |  | BT ION | POS EB | 24_ER | 48 ER | BT WN | HN ER | 24.nE | 48 WE |
| 87120918 | 1 | 12.35 | 87.9E | 17.6 | 60.3 | 167.1 | 35 | 5 | 10 | 15 |
| 87121006 | 2 | 13.15 | 85.9E | 11.7 | 164.3 | 213.1 | 40 | 0 | -5 | 10 |
| 87121018 | 3 | 14.6 S | 84.8E | 24.0 | 172.5 | 270.7 | 50 | 0 | 5 | 15 |
| 87121106 | 4 | 16.15 | 84.4E | 35.1 | 91.0 | 203.7 | 55 | 5 | 15 | 10 |
| 87121118 | 5 | 16.85 | 84.0E | 16.6 | 70.7 | 299.0 | 55 | 10 | 5 | 5 |
| 87121206 | 6 | 17.85 | 83.1E | 12.9 | 140.6 | 481.2 | 55 | 0 | -10 | 0 |
| 87121218 | 7 | 18.55 | 82.3 E | 18.9 | 192.9 | N/A | 55 | -10 | -15 | N/A |
| 87121306 | 8 | 17.85 | 81.2E | 91.0 | N/A | N/A | 55 | -15 | N/A | N/A |


| 87121318 | 9 | $17.1 S$ | 79.3 E | 87.9 | N/A | N/A | 45 | -10 | N/A | N/A |
| :--- | ---: | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 87121406 | 10 | 17.1 S | 76.5 E | 24.7 | N/A | N/A | 35 | -10 | N/A | N/A |


|  |  | 00 h | 24 h | 48 h |
| ---: | ---: | ---: | ---: | ---: |
|  | Average | 58 | N/A | N/A |
|  | \# Cases | 2 | 0 | 0 |


| DTG | WH | BT LAT | BT LON | ROS ER | 24 ER | 48 ER | BT WN | WW ER | 24 WE | 48 WE |
| :--- | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 87122200 | 1 | $13.8 S$ | 179.3 E | 80.7 | N/A | N/A | 30 | 0 | N/A | N/A |
| 87122212 | 2 | $15.2 S$ | $180.8 E$ | 36.0 | N/A | N/A | 25 | 5 | N/A | N $/ \mathrm{A}$ |


| Tropical Cyclone 05S |  |  | Average * Cases | $\begin{array}{r} 00 \mathrm{~h} \\ 39 \\ 11 \end{array}$ | $\begin{array}{r} 24 b \\ 158 \\ 9 \end{array}$ | $\begin{array}{r} 48 \mathrm{~h} \\ 517 \\ 1 \end{array}$ | BT WN WW ER |  | 24 WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |
| DTG | W\# | BT IAT | BT L0N | POS ER | $24 . E R$ | 48 ER |  |  |  |  |
| 87122700 | 1 | 13.95 | 70.1 E | 33.4 | 288.2 | 516.6 | 35 | 0 | 0 | -5 |
| 87122 勺18 | 2 | 13.45 | 67.8 E | 29.8 | 264.3 | N/A | 35 | 0 | -5 | N/A |
| 87122806 | 3 | 13.75 | 65.6 E | 55.4 | 244.6 | N/A | 35 | 0 | -5 | N/A |
| 87122818 | 4 | 13.25 | 63.15 | 50.2 | 96.0 | N/A | 35 | 0 | -5 | N/A |
| 87122906 | 5 | 12.85 | 60.9 E | 42.9 | 48.0 | N/A | 35 | 0 | -5 | N/A |
| 87122918 | 6 | 13.45 | 59.2 E | 12.0 | 46.8 | N/A | 35 | 0 | -5 | N/A |
| 87123006 | 7 | 14.6 S | 57.7E | 35.4 | 169.1 | N/A | 35 | 0 | 0 | N/A |
| 87123018 | 8 | 14.95 | 56.2 E | 6.0 | 122.0 | N/A | 35 | 0 | 0 | N/A |
| 87123106 | 9 | 15.5 S | 55.8 E | 46.3 | 144.6 | N/A | 30 | 5 | 5 | N/A |
| 87123118 | 10 | 15.95 | 55.3 E | 88.7 | N/A | N/A | 30 | 0 | N/A | N/A |
| 88010106 | 11 | 15.2 S | 51.3E | 29.6 | N/A | N/A | 25 | 0 | N/A | N/A |



| Tropical Cyclone 07P |  |  | Average | $\frac{00 \mathrm{~h}}{24}$ |  |  |  | WW_ER | 24 WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  | 112 | 198 |  |  |  |  |
|  |  |  | \# Cases | 14 | 12 | 10 |  |  |  |  |
| DTG | W\# | BT LAT | BI LON | POS ER | $24 . E R$ | 48 ER | BT WN |  |  |  |
| 88010712 | 1 | 5.8S | 177.8E | 26.7 | 130.9 | 212.7 | 35 | 0 | 5 | -5 |
| 88010800 | 2 | 6.85 | 176.0E | 18.9 | 56.1 | 138.7 | 35 | 5 | 5 | -25 |
| 88010812 | 3 | 7.75 | 174.2 E | 71.6 | 201.2 | 216.3 | 45 | 5 | 0 | -40 |
| 88010900 | 4 | 9.05 | 172.9E | 54.3 | 205.5 | 399.5 | 50 | 5 | -20 | -60 |
| 88010912 | 5 | 10.05 | 171.1E | 8.4 | 90.8 | 138.5 | 65 | -5 | -40 | -45 |
| 88011000 | 6 | 11.45 | 169.1 E | 11.8 | 68.3 | 41.4 | 90 | -5 | -35 | 0 |
| 88011012 | 7 | 13.05 | 167.1E | 8.4 | 49.4 | 147.7 | 115 | 5 | 5 | 55 |
| 88011100 | 8 | 15.0 S | 165.4 E | 18.9 | 111.5 | 210.5 | 140 | -5 | 30 | 65 |
| 88011112 | 9 | 16.85 | 164.3 E | 12.0 | 168.3 | 306.2 | 130 | 0 | 25 | 45 |
| 88011200 | 10 | 18.5 S | 164.3 E | 23.5 | 137.3 | 170.0 | 115 | 0 | 35 | 45 |
| 88011212 | 11 | 20.15 | 165.2 E | 6.0 | 48.9 | N/A | 85 | 5 | 15 | N/A |
| 88011300 | 12 | 21.75 | 165.8 E | 46.2 | 86.8 | N/A | 60 | 5 | 15 | N/A |
| 88011312 | 13 | 23.45 | 167.2 E | 30.1 | N/A | N/A | 45 | 5 | N/A | N/A |
| 88011400 | 14 | 25.2S | 170.1E | 8.1 | N/A | N/A | 30 | 0 | N/A | N/A |





| Tropical Cyclone 11S |  |  | Average <br> \# Cases | 00h | 48h |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | 25 | 133 | 282 |  |  |  |  |
|  |  |  |  | 16 | 14 | 11 |  |  |  |  |
| DTG | W | BT IAT | BT ION | POS ER | 24.ER | 48 ER | BT WN | WW ER | 24.WE | 48 WE |
| 88020812 | 1 | 15.85 | 101.1E | 63.1 | 302.8 | 568.1 | 30 | 5 | 10 | -5 |
| 88020900 | 2 | 16.0 S | 99.1 E | 50.9 | 194.2 | 439.8 | 35 | 0 | 0 | -35 |
| 88020912 | 3 | 15.9 S | 97.2E | 28.9 | 110.4 | 196.0 | 40 | -5 | -30 | -55 |
| 88021000 | 4 | 15.85 | 95.0 E | 23.9 | 115.5 | 271.0 | 45 | -5 | -30 | -55 |
| 88021012 | 5 | 15.6 S | 93.4E | 16.7 | 156.2 | 322.0 | 60 | -10 | -20 | -10 |
| 88021100 | 6 | 15.75 | 91.9E | 40.9 | 173.4 | 287.4 | 70 | -10 | -35 | -15 |
| 88021112 | 7 | 16.15 | 90.8E | 16.6 | 93.1 | 143.5 | 80 | -20 | -20 | -15 |
| 88021200 | 8 | 17.15 | 90.1E | 8.3 | 123.4 | 379.2 | 90 | -25 | -10 | 0 |
| 88021212 | 9 | 17.85 | 89.3 E | 36.3 | 107.6 | 309.3 | 80 | -15 | 5 | 10 |
| 88021300 | 10 | 18.45 | 88.5E | 24.7 | 36.0 | 37.2 | 70 | -15 | -5 | -10 |
| 88021312 | 11 | 18.65 | 87.8E | 28.4 | 40.2 | N/A | 60 | -10 | -10 | N/A |
| 88021400 | 12 | 19.15 | 86.5E | 16.5 | 198.6 | N/A | 50 | 0 | -15 | N/A |
| 88021412 | 13 | 18.85 | 85. OE | 11.4 | 98.4 | 156.0 | 45 | 10 | 5 | 0 |
| 88021500 | 14 | 18.65 | 83.9E | 17.1 | 114.6 | $\mathrm{N} / \mathrm{A}$ | 45 | 0 | -5 | N/A |
| 88021512 | 15 | 18.25 | 82.6E | 16.6 | N/A | N/A | 40 | -5 | N/A | N/A |
| 88021600 | 16 | 18.25 | 80.9E | 8.3 | N/A | N/A | 35 | -5 | N/A | N/A |


| Tropical | Cyclone 12p |  |  | O0h | 24h | 48h | - |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | Average | 30 | 133 | 261 |  |  |  |  |
|  |  |  | * Cases | 12 | 8 | 7 |  |  |  |  |
| DTG | W ${ }^{\text {H }}$ | BT LAT | BT LON | POS ER | 24 ER | 48 ER | BT WN | WW ER | 24 WE | 48 WE |
| 88022100 | 1 | 13.75 | 158.4E | 53.8 | 255.7 | 419.4 | 35 | 0 | 5 | 5 |
| 88022106 | 2 | 13.95 | 158. OE | 46.6 | 202.1 | 406.4 | 35 | 0 | -5 | -5 |
| 88022118 | 3 | 14.35 | 157.4E | 96.2 | 238.9 | 197.7 | 35 | 0 | 0 | 0 |
| 88022206 | 4 | 14.65 | 156.8 E | 16.7 | 63.4 | 77.4 | 40 | 0 | 0 | 10 |
| 88022218 | 5 | 14.6 S | 156. OE | 52.3 | 615.2 | 179.1 | 40 | 0 | -10 | 5 |
| 88022306 | 6 | 14.6 S | 155.2E | 6.0 | 79.2 | 200.1 | 45 | 0 | 15 | 30 |
| 88022318 | 7 | 14.8 S | 164.5E | 586.0 | 143.0 | 158.3 | 45 | 0 | 15 | 10 |
| 88022406 | 8 | 15.55 | 154.6E | 8.3 | 49.4 | N/A | 40 | 5 | 10 | N/A |
| 88022418 | 9 | 16.6 S | 155.0E | 33.9 | N/A | N/A | 30 | 0 | N/A | N/A |
| 88022806 | 10 | 19.05 | 148.3E | 28.4 | N/A | N/A | 35 | 5 | N/A | N/A |
| 88022818 | 11 | 18.9 S | 147.7E | 6.0 | N/A | N/A | 35 | 0 | N/A | N/A |
| 88022906 | 12 | 19.3S | 147.2E | 13.3 | N/A | N/A | 30 | 0 | N/A | N/A |




| Tropical Cyclone 15P |  |  | Average | $\frac{00 h}{58}$ | $\begin{aligned} & 24 b \\ & 158 \end{aligned}$ | $\frac{48 \mathrm{~h}}{558}$ |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  |  |  |
|  |  |  |  | \# Cases | 8 | 5 | 4 |  |  |  |  |
| DTG | W\# | Bt inat | BT ION | POS ER | $24 . E R$ | 48 ER | BT WN | HW ER | 24_WE | 48 WE |
| 88022818 | 1 | 16.95 | 201.2 E | 56.7 | 95.1 | 349.5 | 25 | 10 | 20 | 20 |
| 88022906 | 2 | 19.25 | 201.1E | 66.1 | 259.8 | 563.9 | 30 | 5 | 0 | 5 |
| 88022918 | 3 | 20.55 | 204.2E | 39.3 | 94.6 | 404.8 | 35 | 10 | -10 | 5 |
| 88030106 | 4 | 22.75 | 207.3E | 20.5 | 227.9 | 913.4 | 40 | 5 | 10 | 15 |
| 88030118 | 5 | 25.45 | 210.3E | 17.3 | 114.5 | N/A | 45 | 5 | 10 | N/A |
| 88030206 | 6 | 28.15 | 212.9 E | 28.8 | N/A | N/A | 35 | 5 | N/A | N/A |
| 88030218 | 7 | 31.2 S | 212.1 E | 57.8 | N/A | N/A | 25 | 15 | N/A | N/A |
| 88030306 | 8 | 34.25 | 209.7E | 177.9 | N/A | N/A | 25 | 5 | N/A | N/A |

Tropical Cyclona 16S

| DTG | W\# | BT LAT |
| :--- | ---: | ---: |
| 88031600 | 1 | 14.6 S |
| 88031612 | 2 | 14.3 S |
| 88031700 | 3 | 14.0 S |
| 88031712 | 4 | 14.5 S |
| 88031800 | 5 | 15.1 S |
| 88031812 | 6 | 15.8 S |
| 88031900 | 7 | 16.3 S |
| 88031912 | 8 | 16.5 S |
| 88032000 | 9 | 16.7 S |
| 88032012 | 10 | 17.9 S |
| 88032100 | 11 | 20.0 S |
| 88032112 | 12 | 21.8 S |
| 88032200 | 13 | 22.6 S |
| 88032212 | 14 | 24.2 S |
| 88032300 | 15 | 27.0 S |
| 88032312 | 16 | 29.7 S |


|  | OOh | $24 h$ | 48 h |
| :--- | ---: | ---: | ---: |
| Average | 31 | 168 | 298 |
| \# Cases | 16 | 14 | 13 |

Tropical Cyclone 175

| DTG | WH | BT LAT |
| :--- | ---: | ---: |
| 88031718 | 1 | 11.05 |
| 88031806 | 2 | $10.1 S$ |
| 88031818 | 3 | $10.2 S$ |
| 88031906 | 4 | $11.8 S$ |
| 88031918 | 5 | $13.1 S$ |
| 88032006 | 6 | $13.7 S$ |
| 88032018 | 7 | $14.4 S$ |

ET LON
50.9 E
51.0 E
52.2 E
53.0 E
54.4 E
55.8 E
56.5 E

| POS ER | 24 ER | 48 EB |
| ---: | ---: | ---: |
| 21.5 | 300.7 | 578.7 |
| 13.3 | 150.6 | 262.3 |
| 16.8 | 199.4 | 283.5 |
| 60.3 | 312.5 | $\mathrm{~N} / \mathrm{A}$ |
| 18.5 | 60.0 | $\mathrm{~N} / \mathrm{A}$ |
| 13.3 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| .0 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
|  |  |  |
|  |  |  |
| 00 h | 24 h | 48 h |
| 19 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 3 | 0 | 0 |


| DTG | W眷 | BT LAT |
| :--- | ---: | ---: |
| 88032706 | 1 | $15.0 S$ |
| 88032718 | 2 | $15.8 S$ |
| 88032806 | 3 | $16.5 S$ |

BT ION
51.4 E
50.4 E
49.5 E

| POS ER | 24 ER | 48 ER |
| ---: | ---: | ---: |
| 18.9 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |
| 26.6 | $\mathrm{~N} / \mathrm{A}$ | $\mathrm{N} / \mathrm{A}$ |

BT
Tropical Cyclone 185

Tropical Cyclone 19P

| DTG | W | BT LAT | BT LON | POS ER | 24 ER | 48 ER | BT WN | HW ER | 24.WE | 48 WE |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88040918 | 1 | 17.75 | 170.2 E | 68.2 | 260.9 | 468.0 | 45 | -10 | -20 | 0 |
| 88041006 | 2 | 18.15 | 170.5E | 18.0 | 133.9 | N/A | 50 | -5 | -5 | N/A |
| 88041018 | 3 | 17.95 | 170.8E | 13.3 | N/A | N/A | 55 | -15 | N/A | N/A |
| 88041106 | 4 | 17.75 | 170.6 E | 24.7 | N/A | N/A | 60 | -20 | N/A | N/A |
| 88041118 | 5 | 18.05 | 170.6 E | 41.4 | 136.2 | N/A | 65 | -25 | -35 | N/A |
| 88041206 | 6 | 18.25 | 170.8 E | . 0 | 85.1 | 178.4 | 70 | -10 | 20 | 10 |
| 88041218 | 7 | 19.15 | 171.4 E | . 0 | 64.1 | N/A | 65 | -15 | -25 | N/A |
| 88041306 | 8 | 20.05 | 171.8 E | 37.6 | 131.2 | N/A | 60 | -10 | -25 | N/A |
| 88041318 | 9 | 21.05 | 172.2E | 8.2 | 39.6 | N/A | 60 | -15 | -25 | N/A |


| 88041406 | 10 | 22.05 | 172.6 E | 12.0 | 136.0 | 229.6 | 60 | -15 | -10 | -10 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 88041418 | 11 | 23.65 | 173.1E | 18.8 | 72.0 | N/A | 55 | -15 | -15 | N/A |
|  |  |  |  |  |  |  |  |  |  |  |
| Tropical | Cyclona | 205 |  | OOh | 24h | 48h |  |  |  |  |
|  |  |  | Average | 18 | 139 | 277 |  |  |  |  |
|  |  |  | \# Cases | 3 | 1 | 1 |  |  |  |  |
| DTG | W\# | BT IAT | BT LON | POS ER | 24.EB | 48 EB | BT WN | WW EB | 24WE | 48 WE |
| 88050918 | 1 | 15.05 | 64.6 E | 24.7 | 139.1 | 276.5 | 40 | -5 | 5 | 5 |
| 88051006 | 2 | 15.2 S | 64.9 E | 13.0 | N/A | N/A | 35 | 5 | N/A | N/A |
| 88051018 | 3 | 15.3S | 65.5 E | 16.7 | N/A | N/A | 35 | -5 | N/A | N/A |
| Tropical | Cyclone | 215 |  | 00h | 24h | 48h |  |  |  |  |
|  |  |  | Average | 32 | 207 | N/A |  |  |  |  |
|  |  |  | \# Cases | 4 | 2 | 0 |  |  |  |  |
| DTG | W\# | BT IAT | BT ION | PQS EB | 24.EB | 48 ER | BT WN | WW ER | 24. FE | 48FE |
| 88051900 | 1 | 12.75 | 96.5E | 42.9 | 193.0 | N/A | 35 | -5 | 15 | $N / A$ |
| 88051912 | 2 | 16.05 | 97.9E | 36.6 | 221.1 | N/A | 35 | 0 | 20 | N/A |
| 88052000 | 3 | 19.2 S | 101.1E | 16.5 | N/A | N/A | 30 | 0 | N/A | N/A |
| 88052012 | 4 | 22.9S | 105. OE | 34.3 | N/A | N/A | 25 | 5 | N/A | N/A |

## APPENDIX I DEFINITIONS

BEST TRACK - A subjectively smoothed path, versus a precise and very erratic fix-to-fix path, used to represent tropical cyclone movement.

CENTER - The vertical axis or core of a tropical cyclone. Usually determined by cloud vorticity patterns, wind and/or pressure distribution.

EPHEMERIS - Position of a body (satellite) in space as a function of time; used for gridding satellite imagery. Since ephemeris gridding is based solely on the predicted position of the satellite, it is susceptible to errors from vehicle wobble, orbital eccentricity and the oblateness of the Earth.

EXPLOSIVE DEEPENING - A decrease in the minimum sea-level pressure of a tropical cyclone of 2.5 $\mathrm{mb} / \mathrm{hr}$ for 12 hours or $5.0 \mathrm{mb} / \mathrm{hr}$ for six hours (Holliday and Thompson, 1979).

EXTRATROPICAL - A term used in wamings and tropical summaries to indicate that a cyclone has lost its "tropical" characteristics. The term implies both poleward displacement from the tropics and the conversion of the cyclone's primary energy source from the release of latent heat of condensation to baroclinic processes. It is important to note that cyclones can become extratropical and still maintain winds of typhoon or storm force.

EYE - The central area of a tropical cyclone when it is more than half surrounded by wall cloud.

FUJIWHARA EFFECT - A binary interaction where tropical cyclones within about $750 \mathrm{~nm}(1389 \mathrm{~km})$ of each other begin to rotate about one another. When tropical cyclones are within about $400 \mathrm{~nm}(741 \mathrm{~km})$ of each other, they may also begin to be drawn closer to one another (Brand, 1970) (Dong and Neumann, 1983).

INTENSITTY - The maximum sustained surface wind speed, typically within one degree of the center of a tropical cyclone.

MAXIMUM SUSTAINED WIND - The highest surface wind speed averaged over a one-minute period of time. (Peak gusts over water average 20 to 25 percent higher than sustained winds.)

RAPID DEEPENING - A decrease in the minimum sea-level pressure of a tropical cyclone of $1.25 \mathrm{mb} / \mathrm{hr}$ for 24-hours (Holliday and Thompson, 1979).

RECURVATURE - The turning of a tropical cyclone from an initial path toward the west and poleward to east and poleward.

SIGNIFICANT TROPICAL CYCLONE - A tropical cyclone becomes "significant" with the issuance of the first numbered warning by the responsible warning agency.

SIZE - The areal extent of a tropical cyclone, usually measured radially outward from the center to the outermost closed isobar.

STRENGTH - The average wind speed of the surrounding low-level wind flow, usually measured within one to three degrees of the center of a tropical cyclone.
SUBTROPICAL CYCLONE - a low pressure system that forms over the ocean in the subtropics and has some characteristics of a tropical circulation, but not a central dense overcast. Although of upper cold low or low-level baroclinic origins, the system can transition to a tropical cyclone.

SUPER TYPHOON - A typhoon with maximum sustained surface winds of $130 \mathrm{kt}(67 \mathrm{~m} / \mathrm{sec})$ or greater.

TROPICAL CYCLONE - A non-frontal, migratory low-pressure system, usually of synoptic scale, originating over tropical or subtropical waters and having a definite organized circulation.

TROPICAL DEPRESSION - A tropical cyclone with maximum sustained surface winds of $33 \mathrm{kt} \mathrm{( } 17 \mathrm{~m} / \mathrm{sec}$ ) or less.

TROPICAL DISTURBANCE - A discrete system of apparently organized convection, generally 100 to 300 nm ( 185 to 556 km ) in diameter, originating in the tropics or subtropics, having a non-frontal, migratory character and having maintained its identity for 12 - to 24 -hours. It may or may not be associated with a detectable perturbation of the wind field. It is the basic generic designation which, in successive stages of development, may be classified as a Tropical Depression, Tropical Storm, Typhoon or Super Typhoon.

TROPICAL STORM - A tropical cyclone with maximum sustained surface winds in the range of 34 to 63 kt ( 17 to $32 \mathrm{~m} / \mathrm{sec}$ ) inclusive.

TROPICAL UPPER-TROPOSPHERIC TROUGH (TUTT) - A dominant climatological system and a daily upper-level synoptic feature of the summer season, over the tropical North Atlantic, North Pacific and South Pacific Oceans (Sadler, 1979).

TYPHOON (HURRICANE) - A tropical cyclone with maximum sustained surface winds of 64 to 129 kt ( 33 to $66 \mathrm{~m} / \mathrm{sec}$ ). West of 180 degrees longitude they are called typhoons and east of 180 degrees longitude hurricanes.

WALL CLOUD - An organized band of cumuliform clouds that immediately surrounds the central area of a tropical cyclone. The wall cloud may entirely enclose or partially surround the center.

## APPENDIX II

## NAMES FOR TROPICAL CYCLONES

| Column 1 | Column 2 | Column 3 | Column 4 |
| :--- | :--- | :--- | :--- |
|  |  |  |  |
| ANDY | ABBY | ALEX | AGNES |
| BRENDA | BEN | BETTY | BILL |
| CECIL | CARMEN | CARY | CLARA |
| DOT | DOM | DINAH | DOYLE |
| ELLIS | ELLEN | ED | ELSIE |
| FAYE | FORREST | FREDA | FABIAN |
| GORDON | GEORGIA | GERALD | GAY |
| HOPE | HERBERT | HOLLY | HAL |
| IRVING | IDA | IAN | IRMA |
| JUDY | JOE | JUNE | JEFF |
| KEN | KIM | KELLY | KIT |
| LOLA | LEX | LYNN | LEE |
| MAC | MARGE | MAURY | MAMIE |
| NANCY | NORRIS | NINA | NELSON |
| OWEN | ORCHID | OGDEN | ODESSA |
| PEGGY | PERCY | PHYLLIS | PAT |
| ROGER | RUTH | ROY | RUBY |
| SARAH | SPERRY | SUSAN | SKIP |
| TIP | THELMA | THAD | TESS |
| VERA | VERNON | VANESSA | VAL |
| WAYNE | WYNNE | WARREN | WINONA |

NOTE: Names are assigned in rotation and alphabetically. When the last name in Column 4 (WINONA) has been used, the sequence will begin again with the first name in Column 1 (ANDY).

SOURCE: CINCPACINST 3140.1S

## APPENDIX III

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## APPENDIX IV

## PAST ANNUAL TROPICAL CYCLONE REPORTS

Copies of the past<br>Annual Tropical Cyclone Reports can be obtained through:<br>National Technical Information Service<br>5285 Port Royal Road<br>Springfield, Virginia 22161

Refer to the following acquisition numbers when ordering:

| YEAR | ACQUISITION NUMBER |
| :---: | :---: |
| 1959 | AD 786147 |
| 1960 | AD 786148 |
| 1961 | AD 786149 |
| 1962 | AD 786128 |
| 1963 | AD 786208 |
| 1964 | AD 786209 |
| 1965 | AD 786210 |
| 1966 | AD 785891 |
| 1967 | AD 785344 |
| 1968 | AD 785251 |
| 1969 | AD 785178 |
| 1970 | AD 785252 |
| 1971 | AD 768333 |
| 1972 | AD 768334 |
| 1973 | AD 777093 |
| 1974 | AD 010271 |
| 1975 | AD A023601 |
| 1976 | AD A038484 |
| 1977 | AD A055512 |
| 1978 | AD A070904 |
| 1979 | AD A082071 |
| 1980 | AD A094668 |
| 1981 | AD A112002 |
| 1982 | AD A124860 |
| 1983 | AD A137836 |
| 1984 | AD A153395 |
| 1985 | AD A168284 |
| 1986 | AD A184082 |
| 1987 | AD A191883 |
|  |  |

UNCIASSIFTED

2. PERSONAL AUTHOR(S)

| 3a. TYPE OF REPORT ANFIUAL | 13b. TIME COVERED FROM TAN 88 TO DEC 88 | $\begin{aligned} & \text { 14. DATE OF REPORT (Year, Month, Day) } \\ & 1988 \end{aligned}$ | $\begin{aligned} & \text { 15. PAGE COUNT } \\ & 216 \text { PLUS } i-x \end{aligned}$ |
| :---: | :---: | :---: | :---: |

6. SUPPLEMENTARY NOTATION

| 7. | COSATI CODES |  |
| :---: | :---: | :---: |
| FIELD | GROUP | SUB-GROUP |
| 04 | 02 |  |
|  |  |  |

18. SUBJECT TERMS (Continue on reverse if necessary and identify by block number) TROPICAL CYCLOHES TROPICAL STORMS TROPICAL DEPRESSIONS TYPHOONS/SUPER TYPHOONS TROPICAL CYCLONE RESEARCH METEOROLOGICAL SATETHTMTE
19. ABSTRACT (Continue on reverse if necessary and identify by block number)

ANNUAL PUBLICATION SUMMARIZING TROPICAL CYCLONE ACTIVITY IN THE WESTERN NORIH PACIFIC, BAY OF BENGAL, ARABIAN SEA, WESTERN SOUTH PACIFIC AND SOUTH INDIAN OCEANS. A BEST TRACK IS PROVIDED FOR EACH SIGNIFICANT IROPICAL CYCLONE. A BRIEF MARRATIVE IS GIVEN FOR AL工 TYPHOONS AND SELECTED TROPICAL CYCLONES IN THE WESTERN NORTH PACIFIC AND NORTH INDIAN OCEANS. ALL RECONNAISSANCE DATA USED TO CONSTRUCT THE BEST TRACKS ARE PROVIDED, UPON REQUEST, ON FLOPPY DISKETTES. FORECAST VERIFICATION DATA AND STATISTICS FOR THE JOIEP TYPHOON WARNING CENTIER (JTWC) ARE SUBMITYED.


BLOCK 18 (CONTINUED)
TROPICAL CYCLONE INTENSITY
TROPICAL CYCLONE BEST TRACK DATA
TROPICAL CYCLONE FORECASTING
TROPICAL CYCLONE RECONNAISSANCE
DYNAMIC TROPICAL CYCLONE MODEL
TYPHOON ANALOG MODEL.
TROPICAL CYCLONE STEFRING MODEL CLIMATOLOGY/PERSISTENCE TECHNIQUES TROPICAL CYCLOIE FIX DATA
DRIFPITNG BUOYS


[^0]:    table 2-1
    POSITION CODE NUNERRSS (PCN)

    PCN METHOD FOR CENTER DETERMINATION/GRIDDING

    EYE/GEOGRAPHY
    EYE/EPHEMERIS
    WELL DEFINED CIRCULATION CENTER/GEOGRAPHY HELL DEFINED CIRCULATION CENTER/EPHEMERIS

    POORLY DEFINED CIRCULATION CENTER/GEOGRAPHY
    6 POORLY DEFINED CIRCULATION CENTER/EPHEMERIS

[^1]:    * THE TECHNIQUE FOR CALCULATING RIGHT-ANGLE ERROR WAS REVISED IN 1981. THEREFORE, A DIRECT COMPARISON IN RIGHT-ANGLE ERROR STATISTICS CANNOT BE MADE BETWEEN ERRORS COMPUTED BEFORE 1981 AND THOSE COMPUTED SINCE 1981.
    ** IN 1986, RIGHT-ANGLE ERROR' WAS REPLACED BY CROSS-TRACK ERROR. (SEE FIGURE 5-1B FOR THE DEFINITION OF CROSS-TRACK ERROR).

