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# Timing of glacial retreat in the Wicklow Mountains, Ireland, conditioned by glacier size and topography

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## Abstract

This paper presents 170 Schmidt Hammer exposure ages from moraine boulders and glacially-sculpted bedrock to reveal the post-Last Glacial Maximum (LGM) history of the Wicklow Mountains, Ireland. These data suggest that large ice masses survived for 4-7 ka after retreat of the Irish Sea Ice Stream and were sustained by summit ice-fields until ~16.6 ka. Post-LGM retreat was dynamic, with re-advance moraines deposited in response to Heinrich Stadial 1. However, these events reflect short-term ice front oscillations ( $\leq 1$  ka) during the long-term retreat phase. Retreat from re-advance positions was synchronous across the range and paced by climate, with time-progressive deglaciation from low to high elevation. In contrast, marked asynchronicity in the timing of Younger Dryas deglaciation is closely linked to snow redistribution and indicates that for small cirque glaciers ( $\leq 1$  km<sup>2</sup>), topography can exert the primary control on glacier survival. This result has important implications for palaeoclimate reconstructions as cirque glacier dynamics may be unrelated to climate. This is further complicated by post-depositional processes which can result in moraine ages (e.g. <sup>10</sup>Be) which post-date retreat. Future palaeoclimate studies should prioritise cirque glaciers where snow contributing areas are small and where post-depositional disturbance is limited (matrix-poor, boulder-rich moraines).

## Keywords

- Schmidt Hammer exposure dating (SHED)
- Wicklow, Ireland

- Glacier chronology
- Topographic controls
- $^{10}\text{Be}$  dating

## Introduction

Recent studies have focused on the retreat dynamics of the Irish Sea Ice Stream (ISIS) (Chiverrell et al., 2013; Smedley et al., 2017a), a major outlet of the British-Irish Ice Sheet (BIIS; Smedley et al., 2017b). The timing of ISIS retreat after the Last Glacial Maximum (LGM; 23.3 – 27.5 ka; Hughes and Gibbard, 2015) has important implications for the dynamics of terrestrial ice masses and, in particular, proximal mountain ice caps (Clark et al., 2012). In the Irish Sea Basin, ice caps centred on the mountains of Wales (Hughes et al., 2016), the Lake District (Wilson et al., 2013) and Wicklow, Ireland (Ballantyne et al., 2006) coalesced with the ISIS during the LGM and may have persisted after ISIS retreat. However, while the deglacial chronologies of the Welsh and Lake District ice caps are constrained by  $^{10}\text{Be}$ , OSL and  $^{14}\text{C}$  ages (Ballantyne et al., 2009; McCarroll et al., 2010; Glasser et al., 2012; Lloyd et al., 2013; Hughes et al., 2016; Smedley et al., 2017b), there is a paucity of geomorphological and geochronological evidence for post-LGM activity in the Wicklow Mountains.  $^{10}\text{Be}$  ages from summits in Wicklow (Ballantyne et al., 2006) and the adjacent Blackstairs Mountains (Ballantyne and Stone, 2015) indicate summit deglaciation soon after the LGM ( $n = 5$ ; 21.0 – 22.9 ka), while a single  $^{36}\text{Cl}$  age from the Mottee Stone (~20 km SE of Lugnaquilla (925m) at 52.886993, -6.208196) indicates separation of mainland ice and the ISIS by  $23.1 \pm 2.2$  ka (Bowen et al., 2002). This timeframe accords with Bayesian modelling of ice stream retreat, with the ISIS retreating ~80 km along the Irish coast during the interval 20.8 – 23.9 ka (Smedley et al., 2017b). Recent research has begun to establish the geomorphological context for deglaciation in the Wicklow Mountains (Knight et al., 2017). However with the exception of isolated  $^{14}\text{C}$  ( $n = 2$ , 11.5 – 11.6 ka; Colhoun and Synge, 1980) and  $^{36}\text{Cl}$  ages ( $n = 1$ ;  $17.9 \pm 1.0$  ka; Bowen et al., 2002) from Lough Nahanagan, the geochronology of the deglaciation of the Wicklow Ice Cap is poorly understood. To address this knowledge gap, this study presents 170 Schmidt Hammer (SH) exposure ages from cirque and valley moraines and from a summit overridden by ice at the LGM. These data provide new insights into the contrasting deglacial chronologies of the ISIS and the Wicklow Ice Cap and the climatic and topographic factors which conditioned post-LGM retreat.

## Methods

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3 To develop a deglacial chronology, sampling was focused on prominent moraines and  
4 boulder accumulations as these are the best geomorphological indicators of the dimensions  
5 of former mountain glaciers (Barr et al., 2017). Key sites along the main SW-NE axis of the  
6 mountain range were targeted for Schmidt Hammer exposure dating (SHED; Tomkins et al.,  
7 2016) including glacially-deposited boulders on prominent cirque moraines (>400 m) at  
8 Kelly's Lough, Lough Nahanagan, Mullaghcleevaun and Upper Lough Bray. These moraines  
9 exhibit good spatial coherence (Kirkbride and Winkler, 2012) as they are generally matrix-  
10 poor, boulder-rich and feature clearly defined moraine crests, although the outer cirque  
11 moraine at Lough Nahanagan is degraded (Colhoun and Synge, 1980). This moraine is ~1 km  
12 in length, broadly convex in cross-profile form and consists of unsorted granite, sand and  
13 gravel deposits with entrained glacially smoothed granite boulders (1 - 4 m diameter;  
14 Colhoun and Synge, 1980). Multiple nested moraines and boulder accumulations are  
15 preserved within the inferred YD glacial limit at each site, which conforms to a pattern of  
16 active oscillatory retreat (Bickerdike et al., 2017). In addition, samples were obtained from  
17 valley moraines (250 – 400 m, c. 2 - 4 km from cirque headwalls) at Carrawaystick Brook,  
18 Upper Glendasan, Lough Brook and Glenmacnass Waterfall and from ice-moulded bedrock  
19 and erratic boulders from the summit of Carrigshouk (571 m), which was overridden by ice at  
20 the LGM (Fig. 1; Table 1). 20 surfaces were sampled at each site (Carrawaystick Brook; n = 10)  
21 and 170 surfaces were sampled in total, comparable to previous applications of SHED in the  
22 Mourne Mountains, Northern Ireland (Barr et al., 2017).  
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30 R-values were generated per surface (Niedzielski et al., 2009). Sampled boulders were of  
31 sufficient size (> 25 kg; Sumner and Nel, 2002; Demirdag et al., 2009) and all sampled  
32 surfaces were free of surface discontinuities (Williams and Robinson, 1983) and lichen  
33 (Matthews and Owen, 2008). *Instrument calibration* (Correction Factor = 1.017) and *age*  
34 *calibration* (Correction Factor = 0.992) were performed using the SHED-Earth online  
35 calculator (<http://shed.earth>) following the recommendations of Dortch et al. (2016). SH  
36 exposure ages and  $1\sigma$  uncertainties were calculated based on the arithmetic mean for each  
37 surface (Mean of 30 R-values) and based on the updated calibration curve of Tomkins et al.  
38 (2017) which includes  $^{10}\text{Be}$  dated surfaces from Blackstairs Mountain, Wexford (n = 2;  
39 Ballantyne and Stone, 2015) and Bloody Foreland, Donegal (n = 6; Ballantyne et al., 2007;  
40 Clark et al., 2009). These data fit the trend established at calibration sites in Scotland and NW  
41 England and indicate that errors in SH exposure age estimates due to climatic variability  
42 appear unlikely (Barr et al., 2017). Calibration site exposure ages are calculated using the  
43 online calculators formerly known as the CRONUS-Earth online calculator  
44 (<http://hess.ess.washington.edu/math/>, Wrapper script 2.3, Main calculator 2.1, constants 2.3,  
45 muons 1.1; Balco et al., 2008) and are based on the time-dependent  $L_m$  scaling (Lal, 1991;  
46 Stone, 2000), the Loch Lomond Production Rate (LLPR; Fabel et al., 2012;  $4.02 \pm 0.18$  atoms  
47  $\text{g}^{-1} \text{a}^{-1}$ ) and assuming  $0 \text{ mm ka}^{-1}$  erosion. The LLPR is constrained by independent  $^{14}\text{C}$  ages  
48 (MacLeod et al., 2011) and is the most widely used local production rate in the British Isles.  
49 However, the results presented here will be subject to recalibration in light of future  
50 refinement of local production rates. For each sampled site (n = 9), probability density  
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estimates (PDEs) were produced and modelled to separate out the highest probability Gaussian (Fig. 2; Dortch et al., 2013) to account for geological uncertainty. This analytical method has been employed in studies using  $^{10}\text{Be}$  (Dortch et al., 2013; Murari et al., 2014) to identify ages that are too young (*moraine degradation*; Heyman et al., 2011) or too old (*inheritance*; Hallet and Putknonen, 1994).

Based on the results of SHED, three dimensional reconstructions of cirque glaciers were generated using the GLaRe tool (Pellitero et al., 2016; Basal shear stress = 100 kPa; Step length = 10 m) and used to estimate palaeo equilibrium-line altitudes (ELAs). Valley glaciers were also reconstructed for individual catchments using this method although ELAs were not calculated for these ice configurations as geochronological data are not available for all glacier outlets. ELAs were estimated using the GIS tool of Pellitero et al. (2015), applying the area-altitude balance ratio method (AABR =  $1.9 \pm 0.81$ ; Rea, 2009). ELAs are controlled by climate (Ohmura et al., 1992; Hughes and Braithwaite, 2008) but are also strongly influenced by non-climatic factors (Table 2), such as the supply of snow and ice from indirect sources (Mitchell, 1996; Kern and Laszlo, 2010). To assess the impact of 'redistributed' snow and ice, combined snow and avalanche contributing areas ( $A_c$ ) were calculated (c.f. Ballantyne, 2007a,b; Barr et al., 2017; Dominant wind direction W/SW = 210 - 300°, Avalanche slopes  $\geq 25^\circ$ ) and compared to total glacier surface areas ( $A_g$ ). The  $A_c/A_g$  ratio is a proxy for the potential contribution of redistributed snow to glacier accumulation.

## Results

Gaussian SH exposure ages (Table 1) are in correct stratigraphic order in individual glacier catchments, are broadly consistent with comparable deglacial chronologies across the British Isles (Clark et al., 2012), and clearly differentiate cirque and valley moraines, with deposition during the Younger Dryas (YD; 11.7 - 12.9 ka) and Oldest Dryas respectively (GS-2.1a; 14.7 - 17.5 ka). Moreover, these datasets are chronologically robust (Kirkbride and Winkler, 2012), with well-dated moraine sequences in Glenmalur, Glendasan and Glenmacnass (Fig. 1), and provide a framework for a wider morphostratigraphic deglacial chronology for the Wicklow Mountains (Knight et al., 2017). At cirque sites, SHED indicates deglaciation by  $12.31 \pm 0.51$  ka at Upper Lough Bray (ULB),  $12.00 \pm 0.44$  ka at Kelly's Lough (KL),  $11.40 \pm 0.13$  ka at Mullaghcleevaun (MC) and  $10.93 \pm 0.26$  ka for the outer moraine at Lough Nahanagan (LN; Colhoun and Syngé, 1980). In contrast, valley moraines were deposited at  $16.46 \pm 0.58$  ka at Glenmacnass Waterfall (GW),  $16.21 \pm 0.60$  ka at Upper Glendasan,  $15.48 \pm 0.35$  ka at Carrawaystick Brook (CB) and  $15.41 \pm 0.30$  ka at Lough Brook (LB). Finally, SHED indicates the emergence of Carrigshouk (CS) by  $16.64 \pm 0.82$  ka. This date provides a minimum age for wider summit deglaciation in the Wicklows Mountains due to its comparatively low elevation (571 m) and central position on the range divide. Reconstructed cirque glaciers (Table 2) range in size from  $0.35 \text{ km}^2$  (ULB) to  $1.10 \text{ km}^2$  (LN) while snow contributing areas ( $A_c$ ) range

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3 from 0.12 km<sup>2</sup> (ULB) to 1.07 km<sup>2</sup> (LN). At Lough Nahanagan and Mullaghcleevaun, extensive  
4 upland plateaus to the west and south (210 - 300°) account for large snow contributing areas  
5 ( $A_c \geq 1$  km<sup>2</sup>). In contrast, restricted upslope areas within the glacier drainage basin likely  
6 limited the potential for significant snow redistribution ( $A_c \leq 0.5$  km<sup>2</sup>) at Kelly's Lough and  
7 Upper Lough Bray. AABR ELAs for cirque glaciers range from 513 m (LN) to 648 m (KL) and  
8 show no clear spatial clustering. Finally, reconstructed valley glaciers range in size from 4.96  
9 km<sup>2</sup> (UGD) to 12.46 km<sup>2</sup> (GW) and demonstrate a progressive reduction in total glacier area  
10 ( $A_g$ ) throughout the period 15.4 - 16.5 ka.  
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## 16 Discussion

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19 Firstly, these data demonstrate that significant ice masses persisted in the Wicklows after  
20 retreat of the ISIS, with large valley glaciers (Length: ~4 km, Area: ~12.5 km<sup>2</sup>) present until  
21 the onset of the Late-glacial (Greenland Interstadial-1; Rasmussen et al., 2014). While lowland  
22 ( $23.1 \pm 2.2$  ka; Bowen et al., 2002) and summit deglaciation ( $n = 3$ ; 21.0 – 21.9 ka; Ballantyne  
23 et al., 2006) was coeval with ISIS retreat (Smedley et al., 2017b), SH exposure ages from the  
24 summit of Carrigshouk (571 m) indicate that summit ice fields were present on the range  
25 divide until  $16.64 \pm 0.82$  ka. However, distal summits were ice free as early as  $21.9 \pm 1.1$  ka  
26 (Djouce Mountain, 725 m),  $21.2 \pm 1.1$  ka (Scarr, 641 m) and  $21.0 \pm 1.1$  ka (Kanturk, 523 m)  
27 and evidence a significant time lag in summit deglaciation (~4.4 ka). Collectively, <sup>10</sup>Be and  
28 SHED ages indicate rapid downwastage of the Wicklow Ice Cap soon after the LGM and a  
29 transition to summit ice fields which sourced discrete outlet glaciers (e.g. Glenmacnass,  
30 Glendasan, Glenmalur; Fig. 1). This chronology is consistent with <sup>10</sup>Be ages from summits in  
31 Wales (Glasser et al., 2012; Hughes et al., 2016), which record downwastage of the Welsh Ice  
32 Cap at 19 - 20 ka ( $\leq 600$  m), while <sup>14</sup>C ages from proximal Welsh lowlands ( $15.82 \pm 0.39$  cal.  
33 ka BP; Lowe, 1981; Reimer et al., 2013), show that Alpine-style valley glaciers were present for  
34 ~4 ka after initial summit emergence (Hughes et al., 2016). These data also accord with a <sup>36</sup>Cl  
35 age from ice-sculpted bedrock in the Scafell massif, which indicates substantial downwastage  
36 of the Lake District ice cap by  $17.3 \pm 1.1$  ka ( $< 750$  m; Ballantyne et al., 2009). Deglaciation of  
37 Carrigshouk at ~16.6 ka marks a shift to topographically confined ice flow, with glaciers  
38 sourced from high elevation cirques, and likely reflects a time-progressive response to  
39 reduced moisture availability and winter aridity during this interval (Kelly et al., 2010).  
40 However, these data indicate that summit ice fields and large valley glaciers persisted in the  
41 Wicklows for 4 - 7 ka after ISIS retreat (Smedley et al., 2017b), and reveal similarities in the  
42 post-LGM histories of the Wicklow, Welsh (Hughes et al., 2016) and Lake District ice caps  
43 (Ballantyne et al., 2009). These data accord with wider evidence for post-LGM disintegration  
44 of the BIIS into component ice caps (Clark et al., 2012).  
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54 Secondly, the geomorphological record indicates that post-LGM retreat was dynamic, with  
55 valley and cirque re-advance moraines deposited during the Oldest Dryas and Younger  
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3 Dryas respectively (Fig. 2A). Re-advance of valley glaciers during the interval 15.4 - 16.5 ka  
4 was coeval with peak ice rafted debris flux (Bard et al., 2000; Eynaud et al., 2009) and reduced  
5 sea surface temperatures (Bard et al., 2000) during Heinrich Stadial 1 (Fig. 2B; HS1) and the  
6 re-advance of the Irish Ice Sheet (IIS) and the ISIS during the Killard Point Stadial (~16 - 17.1  
7 ka; McCabe et al., 2007; Clark et al., 2012). The re-advance of glaciers in Ireland during the  
8 Oldest Dryas was matched further down the North-East Atlantic margin in Spain (Palacios et  
9 al., 2017). These re-advances are linked to North Atlantic climate perturbations (HS1) and  
10 likely reflect short-term oscillations of the ice front ( $\leq 1$  ka) during the long-term post-LGM  
11 retreat phase (~8 ka). Moreover, these chronological data match recent morphostratigraphic  
12 assessments which support a widespread pattern of sustained retreat interrupted by minor  
13 glacier readvance (Knight et al., 2017). Retreat from re-advance positions, estimated at 1.82  
14  $\text{km ka}^{-1}$  for the Glenmacnass glacier, was synchronous across the Wicklow Mountains, as  
15 demonstrated by progressive deglaciation from low to high elevation (Fig. 3A;  $R^2 = 0.9116$ ;  $p$   
16  $= 0.045$ ). These data are indicative of climate-controlled retreat with independent outlet  
17 glaciers responding synchronously to reduced moisture availability (Kelly et al., 2010),  
18 irrespective of contrasting glacier aspects, source areas or glacier extents. Oldest Dryas valley  
19 glaciers were extensive ( $\leq 12.5 \text{ km}^2$ ), sustained by ice fields and prior to ~16.6 ka,  
20 overtopped low-lying summits (~571 m). As a result, the potential for significant  
21 redistribution of snow and avalanche material ( $A_c > A_g$ ) was limited, particularly during  
22 periods of winter aridity (Kelly et al., 2010). Therefore, while topography likely influenced the  
23 retreat pattern in individual valleys, post-LGM retreat (~15 - 17 ka) was primarily paced by  
24 climate.

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33 In contrast, marked asynchronicity in the timing of final YD deglaciation (Fig. 3; 11.4 - 12.3 ka)  
34 is unrelated to cirque elevation (Fig. 3A;  $R^2 < 0.01$ ,  $p = 0.97$ ), palaeo-ELA (Fig. 3E;  $R^2 = 0.04$ ,  $p$   
35  $= 0.81$ ) or site latitude ( $R^2 = 0.10$ ,  $p = 0.69$ ). These data suggest that climate was not the  
36 dominant control on the timing of final YD deglaciation. Instead, glacier retreat was strongly  
37 controlled by local topography and the redistribution of wind-blown snow and avalanche  
38 material (Fig. 3B;  $R^2 > 0.99$ ,  $p < 0.01$ ). Combined snow and avalanche contributing areas ( $A_c$ )  
39 range from just  $0.119 \text{ km}^2$  at Upper Lough Bray to  $1.071 \text{ km}^2$  at Lough Nahanagan. For  
40 glaciers with large  $A_c$  areas, topography may exert the primary control on glacier formation  
41 and survival, and may account for the comparatively late-deglaciation of Lough Nahanagan  
42 and Mullaghcleevaun during the early-Holocene. By comparison, glaciers with small  $A_c$  areas,  
43 where the potential for redistribution of snow and avalanche material is limited, may respond  
44 quasi-synchronously to climate warming. For example, the early deglaciation of Upper Lough  
45 Bray at  $12.31 \pm 0.51$  ka is coeval with a gradual rise in summer air temperatures after ~12.5  
46 ka (Brooks and Birks, 2000) which was likely sufficient to raise the 'climatic' ELA above cirque  
47 elevations ( $A_c/A_g = 0$ ; Barr et al., 2017) and initiate mass wastage. In contrast, abundant snow  
48 redistribution ( $A_c \geq 1$ ), conditioned by existing topographic configurations (extensive upland  
49 plateaus), was likely sufficient to locally suppress the 'local' (non-climatic) ELA and promote  
50 glacier survival at other sites. However, the contribution of redistributed snow to glacier  
51 accumulation almost certainly diminished throughout the YD as summer air temperatures  
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3 increased rapidly towards the onset of the Holocene (Brooks and Birks, 2000), thus limiting  
4 snowpack preservation. These data indicate that while regional climate provides the baseline  
5 conditions for glacier growth and decay, cirque glacier dynamics may primarily reflect the  
6 influence of topography. However, there is a weak correlation between glacier size ( $A_g$ ) and  
7 deglaciation age (Fig. 3C;  $R^2 = 0.77$ ,  $p = 0.12$ ), although the size variation between the  
8 smallest (ULB;  $0.35 \text{ km}^2$ ) and largest reconstructed glacier (LN:  $1.10 \text{ km}^2$ ) is minimal ( $\sim 0.75$   
9  $\text{km}^2$ ). As such, significant within-mountain range variation in glacier response times is not  
10 anticipated (Raper and Braithwaite, 2009). This may account for the weak correlation  
11 between  $A_c/A_g$  ratios and deglaciation ages (Fig. 3D;  $R^2 = 0.58$ ,  $p = 0.24$ ) although the  
12 observable trend demonstrates the probability of early deglaciation for glaciers with small  
13  $A_c/A_g$  ratios. Based on this reasoning, we conclude that for small YD glaciers, local  
14 topoclimatic controls can be more significant than wider regional climate in determining  
15 cirque glacier dynamics, and in particular, the timing of final deglaciation. This result has  
16 important implications for palaeoclimate reconstructions based on dating of cirque moraines  
17 (e.g. using  $^{10}\text{Be}$ ), as cirque glacier ELAs are not representative of the regional climate and  
18 consequently glaciers dynamics are likely to be decoupled from climatic changes occurring  
19 in the North Atlantic region. This phenomenon is observed today in the behaviour of small  
20 glaciers in marginal glaciated settings such as the Italian Alps (Colucci, 2016) and other  
21 Mediterranean mountains (Hughes, 2018). These data from Ireland also show that the impact  
22 of topography on glacier dynamics is most significant when glaciers are small ( $\leq 1 \text{ km}^2$ ),  
23 resulting in clear asynchronicity in deglaciation (Fig. 3A;  $R^2 = < 0.01$ ), and provide further  
24 evidence that the climatic integrity of cirque glaciers may be limited (Kirkbride and Winkler,  
25 2012). In contrast, large glaciers ( $\sim 12.5 \text{ km}^2$ ), with limited potential for snow redistribution,  
26 have been shown to respond synchronously to climate forcing (Fig. 3A;  $R^2 = 0.9116$ ).  
27 Modelling studies have shown that topoclimatic variables (solar radiation/snow  
28 redistribution) can predict the style of deglaciation (moraine distribution) for small Younger  
29 Dryas glaciers (Coleman et al., 2009; Bickerdike et al., 2017). The chronological data  
30 presented here provides new evidence that topographic controls not only influence the style  
31 of deglaciation, but can determine the timing of final deglaciation, with clear within-  
32 mountain range variability.

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43 A further challenge in linking cirque glacier dynamics to climatic fluctuations is the potential  
44 impact of moraine stabilisation (Hallet and Putknonen, 1994). This post-depositional process  
45 can result in moraine ages (e.g.  $^{10}\text{Be}$ ) which post-date glacial retreat. A 1-2 ka early  
46 stabilization period has been recorded for Alpine moraine sequences in Alaska and the Alps  
47 (Briner et al., 2005; Ivy-Ochs et al., 2006, 2008; Dortch et al., 2010a). In these environments,  
48 glaciers can produce distinctive asymmetric ice-contact fans that undergo rapid gullyng and  
49 post-depositional reworking on their ice-proximal slopes (e.g. Hambrey et al., 2012; Lukas et  
50 al., 2012). However, these landforms are topographically and sedimentologically distinct from  
51 the low-relief, topographically concordant valley and cirque landsystems found in the  
52 Wicklow Mountains. Moreover, there has been comparatively little attention to processes of  
53 moraine development and deglaciation dynamics in these environments, with analogue  
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3 studies predominantly focused on 'hummocky moraine' landsystems (e.g. Benn and Lukas,  
4 2006), and a relative paucity of research into moraine processes in smaller cirque type  
5 landsystems, likely under the assumption of rapid stabilisation after deglaciation. Recent  
6 work has highlighted the importance of self-censoring in cirque environments due to  
7 oblitative overlap (Barr and Lovell, 2014) or ice-cored moraine degradation (Crump et al.,  
8 2017; Tonkin et al., 2017), while external-censoring due to slope instability may also provide  
9 a control on moraine stabilisation (Barr and Lovell, 2014). As a result, there is considerable  
10 uncertainty regarding the robustness of chronological datasets for cirque moraine systems  
11 (Kirkbride and Winkler, 2012). To produce better-resolved glacier chronologies, researchers  
12 can account for moraine stabilisation through (1) morphostratigraphic comparison of  
13 moraine sequences, (2) Gaussian separation of exposure ages (Dortch et al., 2013), (3)  
14 assessment against independent geochronological data (e.g.  $^{14}\text{C}$ , OSL) and (4) consideration  
15 of modern process studies and likely modern analogues for moraine assemblages when  
16 assessing site suitability for a give geochronological approach (e.g. Çiner et al. 2015). Based  
17 on these criteria, we infer that moraine stabilisation may be a key post-depositional process  
18 for the outer moraine at Lough Nahanagan. Firstly, this moraine is degraded (Colhoun and  
19 Synge, 1980) and is morphologically distinct from sampled cirque moraines at  
20 Mullaghcleevaun, Kelly's Lough and Upper Lough Bray which are tall (> 3 m), matrix-poor,  
21 boulder-rich, and feature clearly defined moraine crests (Fig. 4). Secondly, Gaussian  
22 separation of SHED data for this site ( $n = 20$ ) reveals a clear 'two-peak' probability density  
23 function (Fig. 2C), with Gaussian ages of  $10.93 \pm 0.26$  ka ( $n = 9$ ) and  $11.38 \pm 0.26$  ka ( $n = 9$ ).  
24 The youngest age post-dates the YD by  $\sim 0.8$  ka and is inconsistent with wider evidence for  
25 deglaciation of the British Isles by the YD/Holocene transition (MacLeod et al., 2011)  
26 although in isolation, this observation is insufficient to reject this age at this stage. In  
27 addition, this method highlights clear outlier ages ( $n = 3$ ; 12.8 - 13.5 ka). These Gaussian  
28 ages can be rejected as they are comprised of fewer than 3 ages (c.f. Fig. 3 in Dortch et al.,  
29 2013). Finally, independent  $^{14}\text{C}$  ages indicate deglaciation during the late YD and early  
30 Holocene (11.5 - 11.6 ka; Colhoun and Synge, 1980). In contrast, a single  $^{36}\text{Cl}$  age suggests  
31 ice free conditions since  $17.9 \pm 1.0$  ka (Bowen et al., 2002), although this age likely reflects  
32 prior exposure (*inheritance*) and is rejected from further analysis. Based on these data, we  
33 conclude that the older Gaussian age of  $11.38 \pm 0.26$  ka is more representative of final  
34 deglaciation as this age is consistent with previous  $^{14}\text{C}$  ages and accounts for both the  
35 distinctive geomorphological assemblage at this site and the clear 'two-peak' distribution of  
36 SHED ages. This conclusion indicates that moraine stabilisation and boulder exhumation may  
37 account for the degraded moraine surface and comparatively 'young' SHED ages. These data  
38 provide further evidence that moraine ages are more likely influenced by post-glacial  
39 instability than prior exposure (Shanahan & Zreda 2000; Putkonen & Swanson 2003; Zech et  
40 al. 2005; Heyman et al. 2011; Applegate et al., 2012). As a result, the dynamics of small cirque  
41 glaciers (< 1 km<sup>2</sup>), as determined by radiometric methods ( $^{10}\text{Be}$ ), may not only primarily  
42 reflect topographic controls, but may be profoundly influenced by post-depositional  
43 processes. The post-depositional evolution of YD moraine systems is largely unexplored at  
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3 present and a clear co-benefit of the SHED approach is the potential insight into these  
4 processes; insight that was not afforded by other geochronological approaches. Future  
5 research should carefully consider landform context (Barr and Lovell, 2014) and prioritise  
6 sampling of cirque environments where snow and avalanche contributing areas ( $A_c$ ) are small  
7 (Warren, 1991; Mills et al., 2012; Barr and Lovell, 2014), where postglacial erosion is limited  
8 and where short transport distances promote the formation of matrix-poor boulder-rich  
9 moraines (Fig. 5; Pallàs et al., 2010). In these environments, snow redistribution is limited and  
10 moraines are more likely to stabilise rapidly after deglaciation. As such, these glaciers may  
11 respond quasi-synchronously to climatic fluctuations and may produce more robust  
12 palaeoclimatic reconstructions.  
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## 19 Conclusions

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21 This study provides the first chronological constraints on post-LGM retreat of the Wicklow  
22 Ice Cap. 170 SH exposure ages from cirque and valley moraines and from a summit  
23 overridden by ice at the LGM demonstrate that significant ice masses persisted in the  
24 Wicklows after retreat of the ISIS and were sustained by summit ice-fields until ~16.6 ka.  
25 Post-LGM retreat was dynamic, with re-advance moraines deposited during the Oldest Dryas  
26 in response to HS1 cooling. However, these events reflect short-term oscillations ( $\leq 1$  ka) of  
27 the ice front during the long-term retreat phase (~8 ka), which was driven by reduced  
28 moisture availability and winter aridity. This chronology is matched by  $^{10}\text{Be}$  and  $^{14}\text{C}$  ages from  
29 the Welsh Ice Cap, with downwastage during the interval 19-20 ka, while alpine valley  
30 glaciers persisted for ~4 ka after initial summit emergence. Retreat from re-advance  
31 positions was paced by climate, with time-progressive deglaciation from low to high  
32 elevation ( $R^2 = 0.9116$ ). In contrast, asynchronicity in final Younger Dryas deglaciation is  
33 accounted for by local topography and the redistribution of snow and avalanche material.  
34 Contrasting synchronicity in the timing of glacial retreat during these periods is conditioned  
35 by glacier size, with small YD glaciers ( $< 1 \text{ km}^2$ ) highly sensitive to local topographic controls.  
36 This result has important implications for palaeoclimate reconstructions based on dating of  
37 cirque moraines (e.g.  $^{10}\text{Be}$ ), as cirque glacier dynamics may be unrelated to climate. This is  
38 further complicated by post-depositional processes which can result in ages which post-date  
39 retreat. As a result, future palaeoclimate reconstructions should prioritise cirque glaciers  
40 where snow and avalanche contributing areas ( $A_c$ ) are small and where the potential for  
41 post-depositional disturbance is limited (matrix-poor, boulder rich moraines).  
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## 39 Figure Captions

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42 Figure 1. Generalised geomorphological map of the Wicklow Mountains. Moraines modified  
43 after Clark et al. (2017).  $^{10}\text{Be}$  ages recalibrated from Ballantyne et al. (2006) using the online  
44 calculators formerly known as the CRONUS-Earth online calculator (Wrapper script 2.3, Main  
45 calculator 2.1, constants 2.3, muons 1.1; Balco et al., 2008) based on the Loch Lomond  
46 Production Rate (Fabel et al., 2012), the time-independent  $L_m$  scaling (Lal, 1991; Stone 2000)  
47 and assuming  $0 \text{ mm ka}^{-1}$  erosion. Ca: Camenabologue; Dj: Djouce Mountain; Ka: Kanturk; Ki:  
48 Kippure; Lq: Lugnaquilla; Mu: Mullaghcleevaun; Sc: Scarr; Tg: Tonelagee. Map projection:  
49 UTM WGS 1984.

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53 Figure 2. Gaussian ages related to deglaciation of the Wicklow Mountains. A: The NGRIP  
54 Oxygen Isotope Curve (Rasmussen et al., 2014) plotted against  $1\sigma$  age boundaries for  
55 sampled cirque and valley sites. The Younger Dryas (YD) and Oldest Dryas (OD) periods are  
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3 marked. B: Ice rafted debris and sea surface temperature records from cores SUBI-18 (Bard et  
4 al., 2000) and MD95-2039 (Eynaud et al., 2009) in the North Atlantic. C: Gaussian models for  
5 sampled cirque and valley sites. For each site, the highest probability Gaussian is considered  
6 the most likely timing of deglaciation as all ages are younger than the Last Glacial Maximum  
7 (Dortch et al., 2013). At Lough Nahanagan, the oldest peak with more than 3 ages is selected  
8 (c.f. Dortch et al., 2013) as this moraine is degraded and morphologically distinct from other  
9 sampled cirque moraines (MC, KL, ULB). Moreover, this estimates matches previous  $^{14}\text{C}$  ages  
10 (Colhoun and Synge, 1980).  
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14 Figure 3. Topographic and climatic controls on the timing of cirque deglaciation in the  
15 Wicklow Mountains. A: Site Elevation; B: Snow contributing area ( $A_c$ ); C: Glacier area ( $A_g$ ); D:  
16  $A_c/A_g$  ratio; E: ELA; F: ELAs and  $A_c/A_g$  ratio plots. These data show that for large valley glaciers,  
17 retreat is paced by climate with progressive deglaciation from low to high elevation (A). In  
18 contrast, marked asynchronicity in the timing of cirque deglaciation (A) is strongly controlled  
19 by snow redistribution (B). This asynchronicity is weakly correlated with glacier size (C) and  
20  $A_c/A_g$  ratios (D) and is unrelated to ELA (E).  
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24 Figure 4. Sampled cirque moraines at Lough Nahanagan (A), Mullaghcleevaun (B), Kelly's  
25 Lough (C) and Upper Lough Bray (D). The outer cirque moraine at Lough Nahanagan is  
26 degraded (Colhoun and Synge, 1980) and is morphologically distinct from other cirque  
27 moraines which are sharp crested, boulder-rich and matrix-poor.  
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30 Figure 5. Matrix-poor, boulder-rich moraine at Mullaghcleevaun which likely stabilised  
31 rapidly after deglaciation.  
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Tables

Table 1. Gaussian ages for cirque, valley and summit sites from the Wicklow Mountains.

Group	Site Name	Site Code	Site Elevation (m)	Site Latitude (°)	Deglaciation Age (ka)	±	Glacier Area (A <sub>g</sub> km <sup>2</sup> )
Cirque	Lough Nahanagan	LN	443	53.034	11.38	0.26	1.10
	Mullaghcleevaun	MC	571	53.089	11.40	0.13	0.61
	Kelly's Lough	KL	585	52.960	12.00	0.44	0.38
	Upper Lough Bray	ULB	466	53.178	12.31	0.51	0.35
Valley	Lough Brook	LB	393	53.070	15.41	0.30	7.63
	Carrawaystick Brook	CB	421	52.964	15.48	0.35	1.83
	Upper Glendasan	UGD	332	53.030	16.21	0.60	4.96
	Glenmacnass Waterfall	GW	268	53.062	16.46	0.58	12.46
Summit	Carrigshouk	CG	571	53.086	16.64	0.82	-

Table 2. Snow contributing areas ( $A_c$ ), glacier areas ( $A_g$ ) and ELAs for cirque moraines.

Site Code	Deglaciation Age (ka)	$\pm$	Snow Contributing Area ( $A_c$ km <sup>2</sup> ) <sup>a</sup>	Glacier Area ( $A_g$ km <sup>2</sup> )	$A_c / A_g$ Ratio	ELA <sup>b</sup>
LN	11.38	0.26	1.07	1.10	0.97	513
MC	11.40	0.13	0.99	0.61	1.62	561
KL	12.00	0.44	0.42	0.38	1.10	648
ULB	12.31	0.51	0.12	0.35	0.34	519

<sup>a</sup> Area within the glacier drainage basin within the 210 – 300° quadrant + all other slopes which overlook the glacier (Gradients > 25°), <sup>b</sup> AABR = 1.9 ± 0.81 (Rea, 2009)

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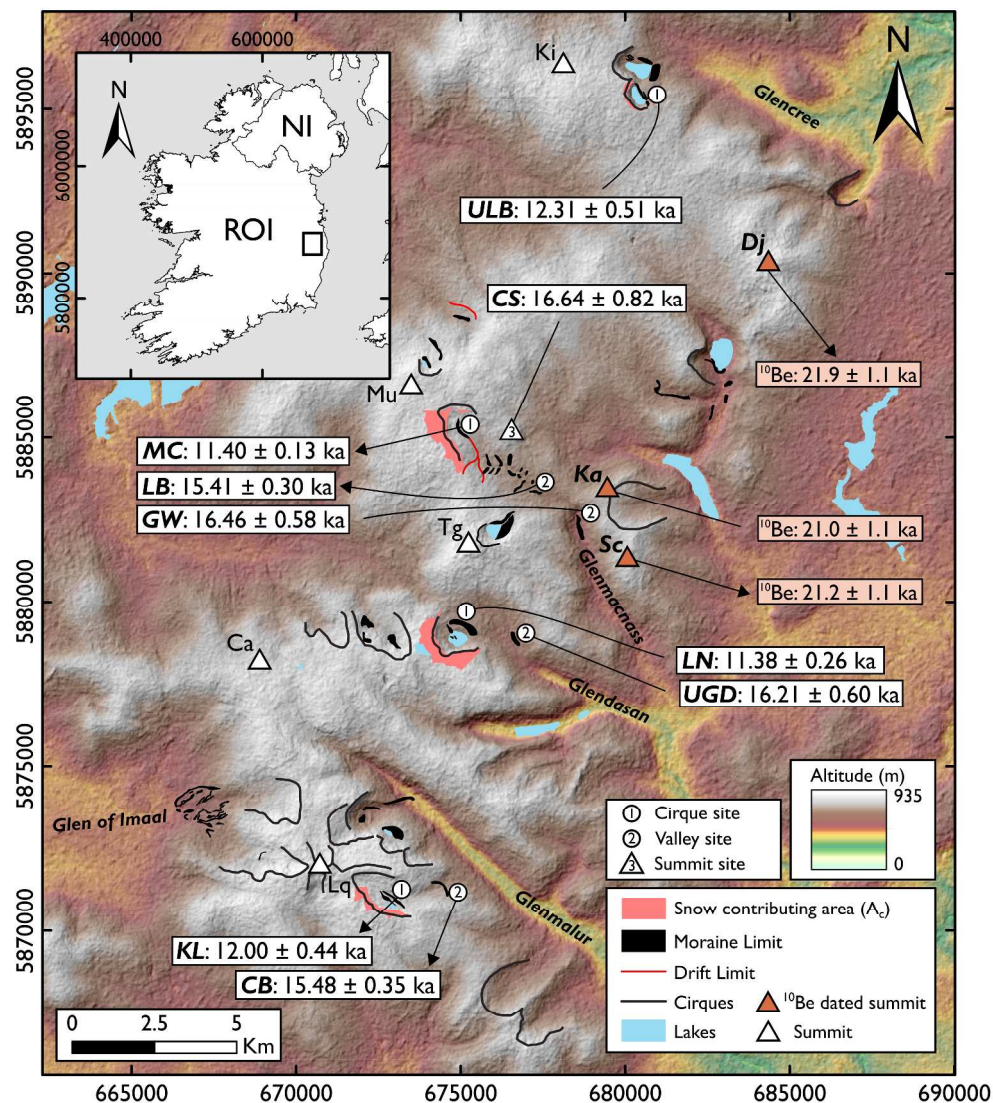


Figure 1. Generalised geomorphological map of the Wicklow Mountains. Moraines modified after Clark et al. (2017).  $^{10}\text{Be}$  ages recalibrated from Ballantyne et al. (2006) using the online calculators formerly known as the CRONUS-Earth online calculator (Wrapper script 2.3, Main calculator 2.1, constants 2.3, muons 1.1; Balco et al., 2008) based on the Loch Lomond Production Rate (Fabel et al., 2012), the time-independent  $L_m$  scaling (Lal, 1991; Stone 2000) and assuming 0 mm  $\text{ka}^{-1}$  erosion. Ca: Camenabologue; Dj: Djouce Mountain; Ka: Kanturk; Ki: Kippure; Lq: Lugnaquilla; Mu: Mullaghcleevaun; Sc: Scarr; Tg: Tonelagee. Map projection: UTM WGS 1984.

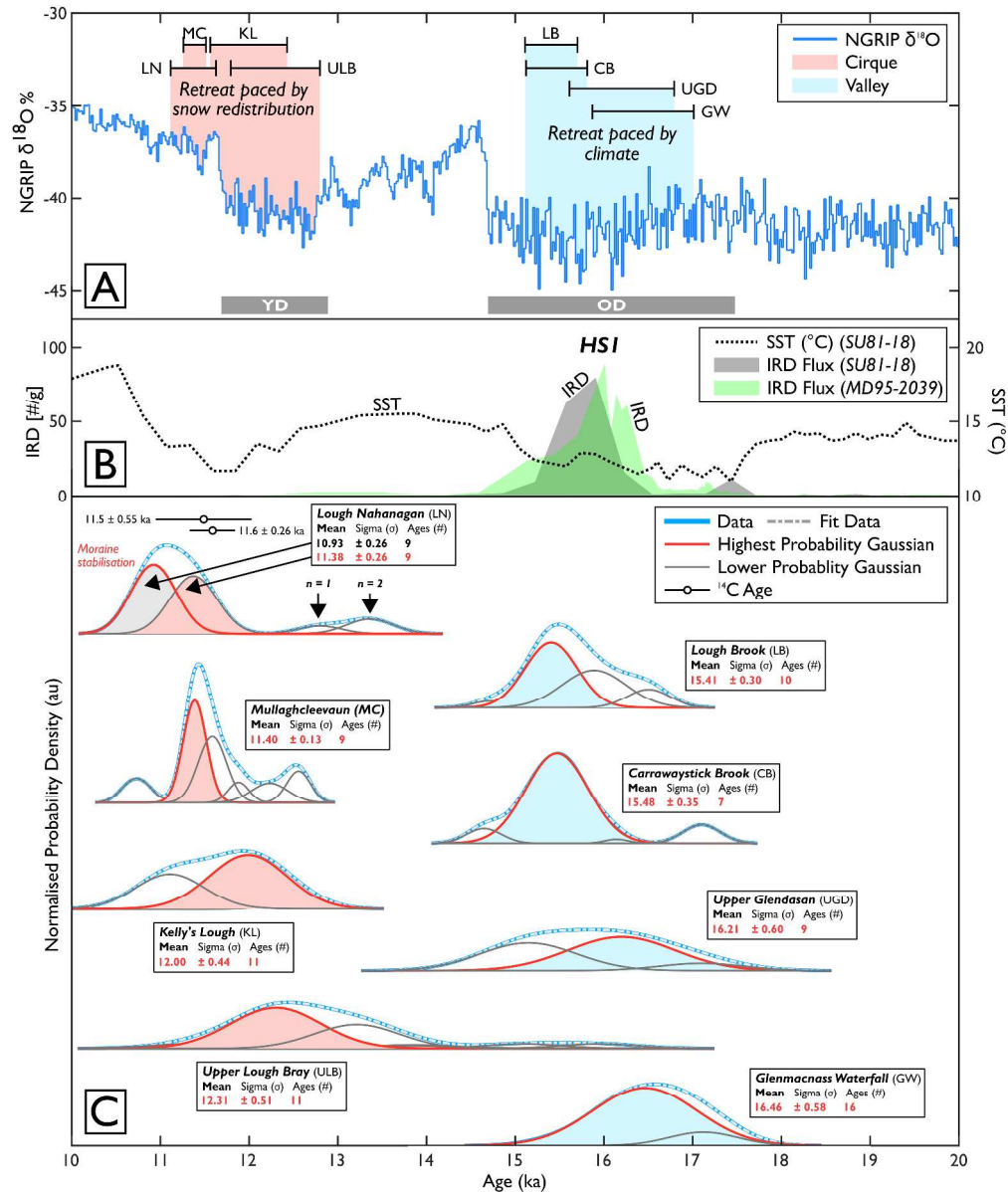


Figure 2. Gaussian ages related to deglaciation of the Wicklow Mountains. A: The NGRIP Oxygen Isotope Curve (Rasmussen et al., 2014) plotted against  $1\sigma$  age boundaries for sampled cirque and valley sites. The Younger Dryas (YD) and Oldest Dryas (OD) periods are marked. B: Ice rafted debris and sea surface temperature records from cores SUB1-18 (Bard et al., 2000) and MD95-2039 (Eynaud et al., 2009) in the North Atlantic. C: Gaussian models for sampled cirque and valley sites. For each site, the highest probability Gaussian is considered the most likely timing of deglaciation as all ages are younger than the Last Glacial Maximum (Dortch et al., 2013). At Lough Nahanagan, the oldest peak with more than 3 ages is selected (c.f. Dortch et al., 2013) as this moraine is degraded and morphologically distinct from other sampled cirque moraines (MC, KL, ULB). Moreover, this estimates matches previous  $^{14}\text{C}$  ages (Colhoun and Synge, 1980).

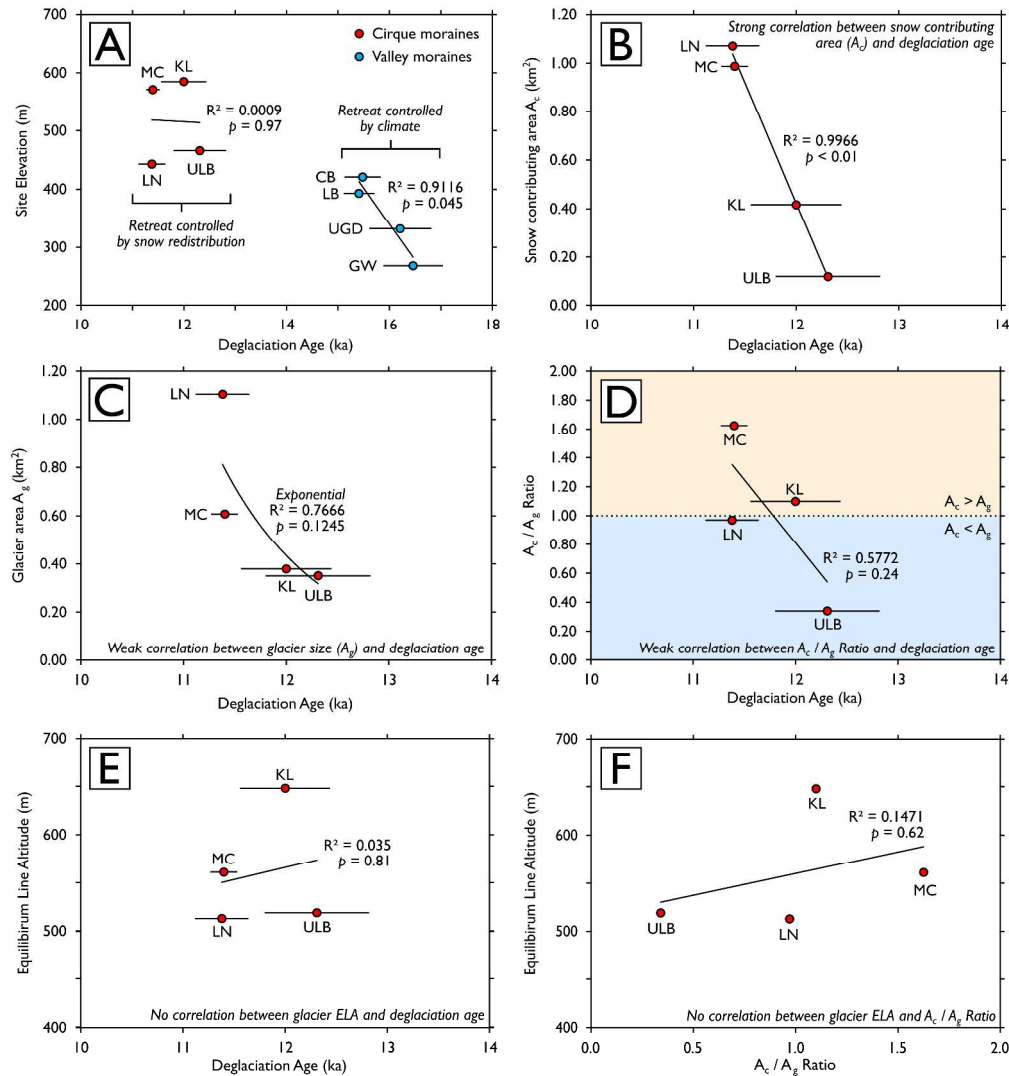


Figure 3. Topographic and climatic controls on the timing of cirque deglaciation in the Wicklow Mountains. A: Site Elevation; B: Snow contributing area ( $A_c$ ); C: Glacier area ( $A_g$ ); D:  $A_c/A_g$  ratio; E: ELA; F: ELAs and  $A_c/A_g$  ratio plots. These data show that for large valley glaciers, retreat is paced by climate with progressive deglaciation from low to high elevation (A). In contrast, marked asynchronicity in the timing of cirque deglaciation (A) is strongly controlled by snow redistribution (B). This asynchronicity is weakly correlated with glacier size (C) and  $A_c/A_g$  ratios (D) and is unrelated to ELA (E).



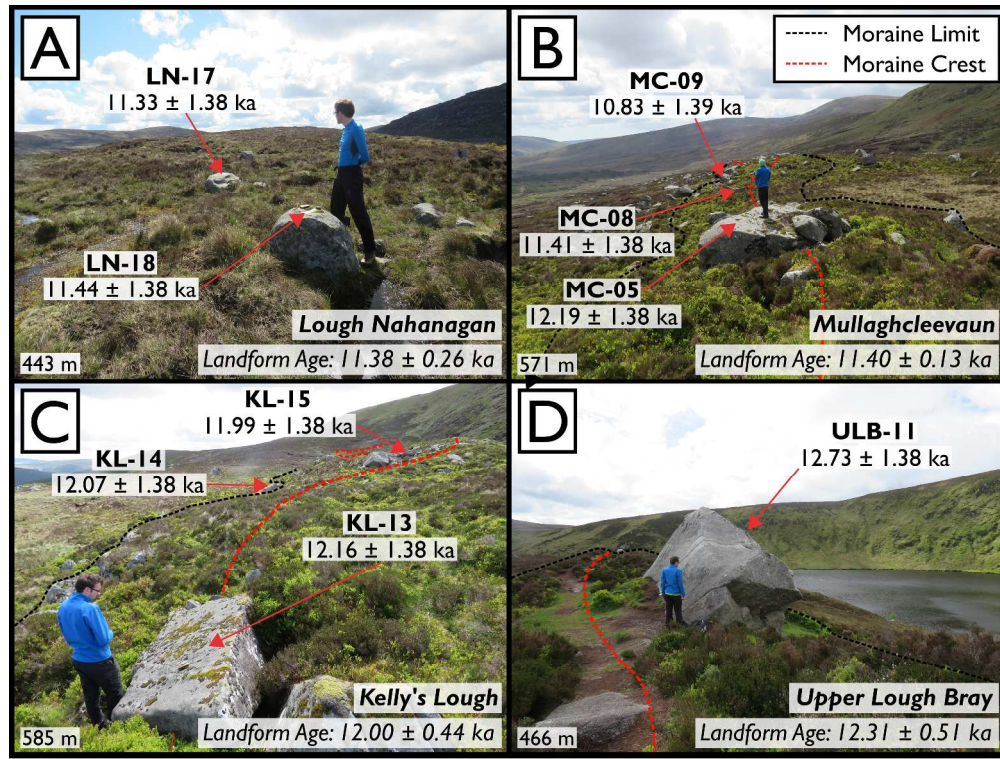
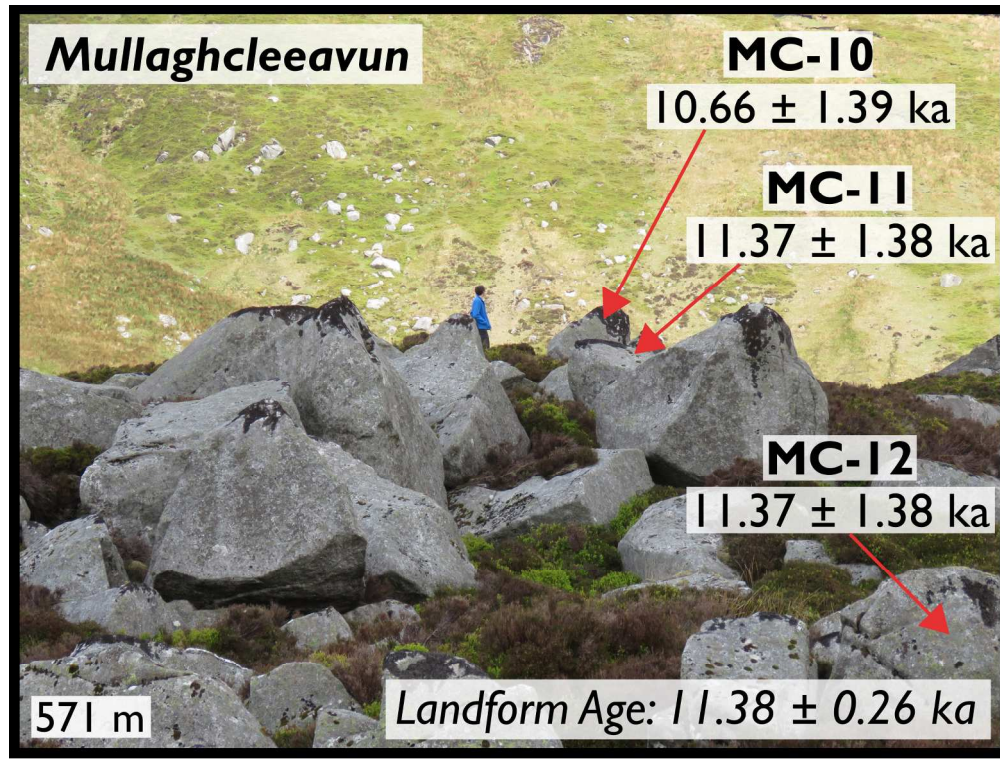


Figure 4. Sampled cirque moraines at Lough Nahanagan (A), Mullaghcleevaun (B), Kelly's Lough (C) and Upper Lough Bray (D). The outer cirque moraine at Lough Nahanagan is degraded (Colhoun and Synge, 1980) and is morphologically distinct from other cirque moraines which are sharp crested, boulder-rich and matrix-poor.



31 Figure 5. Matrix-poor, boulder-rich moraine at Mullaghcleevaun which likely stabilised rapidly after  
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Group	Site Code	Sample Name	Latitude (°)	Longitude (°)	Elevation (m)	Type	Mean R-value	SEM <sup>a</sup>	Age (ka)	1 $\sigma$
Cirque	LN	LN-01	53.03338	-6.3854	445	Boulder	46.10	0.94	11.535	1.382
Cirque	LN	LN-02	53.03352	-6.3858	447	Boulder	46.63	0.92	11.233	1.383
Cirque	LN	LN-03	53.03358	-6.38594	446	Boulder	46.24	0.81	11.46	1.382
Cirque	LN	LN-04	53.03383	-6.3859	446	Boulder	46.91	0.91	11.074	1.384
Cirque	LN	LN-05	53.03407	-6.38591	446	Boulder	47.28	0.98	10.852	1.385
Cirque	LN	LN-06	53.03402	-6.38659	447	Boulder	47.05	1.04	10.971	1.384
Cirque	LN	LN-07	53.03372	-6.3867	447	Boulder	46.42	0.91	11.322	1.383
Cirque	LN	LN-08	53.03381	-6.38705	447	Boulder	42.96	0.64	13.281	1.378
Cirque	LN	LN-09	53.03405	-6.38726	445	Boulder	47.16	0.82	10.892	1.384
Cirque	LN	LN-10	53.03421	-6.38778	444	Boulder	42.63	0.55	13.466	1.377
Cirque	LN	LN-11	53.03463	-6.38824	443	Boulder	46.54	0.87	11.252	1.383
Cirque	LN	LN-12	53.03461	-6.38828	442	Boulder	47.31	0.95	10.816	1.385
Cirque	LN	LN-13	53.03452	-6.38849	443	Boulder	46.18	1.03	11.46	1.382
Cirque	LN	LN-14	53.03478	-6.38881	440	Boulder	47.09	1.05	10.949	1.384
Cirque	LN	LN-15	53.03468	-6.38892	440	Boulder	47.09	0.99	10.949	1.384
Cirque	LN	LN-16	53.03476	-6.39012	439	Boulder	47.23	1.10	10.873	1.384
Cirque	LN	LN-17	53.03477	-6.39019	439	Boulder	46.43	0.94	11.327	1.383
Cirque	LN	LN-18	53.03462	-6.39077	439	Boulder	46.23	0.82	11.441	1.382
Cirque	LN	LN-19	53.03465	-6.39129	438	Boulder	47.01	0.91	11.005	1.384
Cirque	LN	LN-20	53.03481	-6.39209	437	Boulder	43.84	0.82	12.804	1.379
Cirque	MC	MC-01	53.08932	-6.38822	586	Boulder	45.89	0.85	11.636	1.382
Cirque	MC	MC-02	53.08923	-6.38826	586	Boulder	45.50	0.79	11.855	1.381
Cirque	MC	MC-03	53.08913	-6.38826	587	Boulder	46.20	0.76	11.458	1.382
Cirque	MC	MC-04	53.08889	-6.38819	586	Boulder	44.51	0.79	12.414	1.38
Cirque	MC	MC-05	53.08873	-6.38812	583	Boulder	44.92	0.68	12.188	1.38
Cirque	MC	MC-06	53.08862	-6.38807	580	Boulder	46.21	0.91	11.452	1.382
Cirque	MC	MC-07	53.08856	-6.38811	578	Boulder	44.19	0.83	12.599	1.379
Cirque	MC	MC-08	53.08836	-6.38785	575	Boulder	46.28	0.87	11.412	1.383

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4	Cirque	MC	MC-09	53.08817	-6.38766	575	Boulder	47.32	0.90	10.828	1.385
5	Cirque	MC	MC-10	53.08792	-6.38769	575	Boulder	47.62	0.98	10.657	1.385
6	Cirque	MC	MC-11	53.08784	-6.38766	575	Boulder	46.36	0.77	11.367	1.383
7	Cirque	MC	MC-12	53.08763	-6.38737	569	Boulder	46.37	1.09	11.371	1.383
8	Cirque	MC	MC-13	53.08737	-6.38723	568	Boulder	45.41	0.96	11.914	1.381
9	Cirque	MC	MC-14	53.08703	-6.38712	568	Boulder	46.28	0.94	11.418	1.382
10	Cirque	MC	MC-15	53.08699	-6.38697	568	Boulder	46.12	0.97	11.513	1.382
11	Cirque	MC	MC-16	53.08677	-6.38674	566	Boulder	45.85	1.04	11.655	1.382
12	Cirque	MC	MC-17	53.08662	-6.38667	564	Boulder	46.42	0.80	11.341	1.383
13	Cirque	MC	MC-18	53.0861	-6.3857	552	Boulder	46.26	0.87	11.429	1.382
14	Cirque	MC	MC-19	53.0858	-6.38509	543	Boulder	44.24	0.80	12.571	1.379
15	Cirque	MC	MC-20	53.08559	-6.38516	542	Boulder	45.90	0.85	11.63	1.382
16	Cirque	KL	KL-01	52.96091	-6.42741	589	Boulder	47.09	0.69	10.964	1.384
17	Cirque	KL	KL-02	52.9606	-6.42686	586	Boulder	44.40	0.72	12.491	1.379
18	Cirque	KL	KL-03	52.96059	-6.42689	585	Boulder	45.40	0.89	11.918	1.381
19	Cirque	KL	KL-04	52.96058	-6.42695	585	Boulder	45.01	0.59	12.151	1.38
20	Cirque	KL	KL-05	52.96054	-6.42688	585	Boulder	46.57	0.85	11.252	1.383
21	Cirque	KL	KL-06	52.96045	-6.42687	585	Boulder	45.31	0.85	11.961	1.381
22	Cirque	KL	KL-07	52.96037	-6.42682	585	Boulder	46.28	0.66	11.414	1.382
23	Cirque	KL	KL-08	52.96032	-6.42669	585	Boulder	47.08	0.58	10.962	1.384
24	Cirque	KL	KL-09	52.96022	-6.42657	585	Boulder	46.59	0.88	11.24	1.383
25	Cirque	KL	KL-10	52.96015	-6.42655	585	Boulder	44.40	0.83	12.472	1.379
26	Cirque	KL	KL-11	52.96006	-6.42634	585	Boulder	45.67	0.81	11.749	1.381
27	Cirque	KL	KL-12	52.9599	-6.42618	585	Boulder	44.57	0.81	12.368	1.38
28	Cirque	KL	KL-13	52.95984	-6.42605	584	Boulder	44.94	0.81	12.162	1.38
29	Cirque	KL	KL-14	52.95965	-6.42554	582	Boulder	45.11	0.92	12.067	1.38
30	Cirque	KL	KL-15	52.95945	-6.42557	586	Boulder	45.25	0.97	11.99	1.381
31	Cirque	KL	KL-16	52.95947	-6.42537	585	Boulder	47.25	0.99	10.854	1.385
32	Cirque	KL	KL-17	52.95913	-6.42493	586	Boulder	45.56	0.89	11.819	1.381
33	Cirque	KL	KL-18	52.95903	-6.42483	586	Boulder	46.69	0.62	11.176	1.383
34	Cirque	KL	KL-19	52.95883	-6.42416	583	Boulder	45.60	0.74	11.8	1.381
35	Cirque	KL	KL-20	52.95849	-6.42373	583	Boulder	46.30	0.59	11.403	1.383
36	Cirque	ULB	ULB-01	53.18062	-6.30129	464	Boulder	44.30	1.00	12.544	1.379
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4	Cirque	ULB	ULB-02	53.18063	-6.30117	463	Boulder	45.17	1.01	12.043	1.381
5	Cirque	ULB	ULB-03	53.18056	-6.30117	462	Boulder	44.44	1.27	12.459	1.379
6	Cirque	ULB	ULB-04	53.18044	-6.30134	462	Boulder	44.68	1.40	12.327	1.38
7	Cirque	ULB	ULB-05	53.18032	-6.30136	462	Boulder	46.33	1.25	11.38	1.383
8	Cirque	ULB	ULB-06	53.1804	-6.30112	462	Boulder	44.65	1.14	12.334	1.38
9	Cirque	ULB	ULB-07	53.18048	-6.30083	462	Boulder	45.06	1.24	12.109	1.38
10	Cirque	ULB	ULB-08	53.18055	-6.30071	462	Boulder	44.73	1.39	12.294	1.38
11	Cirque	ULB	ULB-09	53.18025	-6.30063	466	Boulder	43.29	1.40	13.116	1.378
12	Cirque	ULB	ULB-10	53.18022	-6.30049	466	Boulder	44.84	1.08	12.234	1.38
13	Cirque	ULB	ULB-11	53.17784	-6.29847	471	Boulder	43.97	1.05	12.728	1.379
14	Cirque	ULB	ULB-12	53.17786	-6.29841	471	Boulder	42.06	0.96	13.813	1.377
15	Cirque	ULB	ULB-13	53.17785	-6.29823	471	Boulder	42.77	0.80	13.409	1.378
16	Cirque	ULB	ULB-14	53.17756	-6.29822	471	Boulder	43.34	0.96	13.086	1.378
17	Cirque	ULB	ULB-15	53.17737	-6.29801	471	Boulder	45.53	0.69	11.846	1.381
18	Cirque	ULB	ULB-16	53.17711	-6.29775	471	Boulder	42.88	0.93	13.345	1.378
19	Cirque	ULB	ULB-17	53.17694	-6.29753	471	Boulder	39.79	0.78	15.1	1.376
20	Cirque	ULB	ULB-18	53.17675	-6.29755	463	Boulder	43.86	0.93	12.79	1.379
21	Cirque	ULB	ULB-19	53.17635	-6.29698	460	Boulder	38.29	0.69	15.951	1.376
22	Cirque	ULB	ULB-20	53.17632	-6.29698	459	Boulder	43.10	0.70	13.231	1.378
23	Cirque	ULB	ULB-20	53.17632	-6.29698	459	Boulder	43.10	0.70	13.231	1.378
24	Cirque	ULB	ULB-20	53.17632	-6.29698	459	Boulder	43.10	0.70	13.231	1.378
25	Valley	LB	LB-01	53.07045	-6.35507	394	Boulder	37.26	1.04	16.538	1.377
26	Valley	LB	LB-02	53.07046	-6.35512	393	Boulder	38.69	0.93	15.726	1.376
27	Valley	LB	LB-03	53.07034	-6.35513	395	Boulder	37.59	0.95	16.345	1.377
28	Valley	LB	LB-04	53.07009	-6.35499	396	Boulder	39.72	1.05	15.137	1.376
29	Valley	LB	LB-05	53.07006	-6.355	396	Boulder	37.93	0.90	16.156	1.377
30	Valley	LB	LB-06	53.07009	-6.35495	396	Boulder	39.43	0.72	15.302	1.376
31	Valley	LB	LB-07	53.07003	-6.35471	396	Boulder	38.44	0.97	15.874	1.376
32	Valley	LB	LB-08	53.07015	-6.3546	396	Boulder	38.71	0.90	15.715	1.376
33	Valley	LB	LB-09	53.07041	-6.35415	390	Boulder	38.44	1.00	15.862	1.376
34	Valley	LB	LB-10	53.07045	-6.35399	390	Boulder	40.34	0.82	14.791	1.376
35	Valley	LB	LB-11	53.07029	-6.35407	392	Boulder	38.82	0.94	15.656	1.376
36	Valley	LB	LB-12	53.0702	-6.35389	393	Boulder	39.35	0.68	15.357	1.376
37	Valley	LB	LB-13	53.07032	-6.35356	393	Boulder	39.02	0.89	15.543	1.376
38	Valley	LB	LB-14	53.07028	-6.35348	394	Boulder	38.06	0.65	16.091	1.376
39	Valley	LB	LB-14	53.07028	-6.35348	394	Boulder	38.06	0.65	16.091	1.376
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Valley	LB	LB-15	53.07027	-6.35349	394	Boulder	37.20	1.11	16.57	1.377
Valley	LB	LB-16	53.07012	-6.35328	394	Boulder	39.07	0.80	15.52	1.376
Valley	LB	LB-17	53.07022	-6.35302	391	Boulder	39.17	0.92	15.454	1.376
Valley	LB	LB-18	53.06963	-6.35252	395	Boulder	39.37	0.93	15.34	1.376
Valley	LB	LB-19	53.06958	-6.35199	389	Boulder	39.14	1.13	15.469	1.376
Valley	LB	LB-20	53.06971	-6.35162	384	Boulder	39.48	0.79	15.281	1.376
Valley	CB	CB-01	52.96432	-6.40324	423	Boulder	39.52	0.66	15.264	1.376
Valley	CB	CB-02	52.96429	-6.40317	424	Boulder	39.79	0.79	15.113	1.376
Valley	CB	CB-03	52.9643	-6.40319	423	Boulder	38.20	0.69	16.021	1.376
Valley	CB	CB-04	52.96428	-6.4031	421	Boulder	39.36	0.64	15.359	1.376
Valley	CB	CB-05	52.96423	-6.40311	421	Boulder	40.53	0.54	14.696	1.376
Valley	CB	CB-06	52.96421	-6.40317	421	Boulder	38.91	0.80	15.616	1.376
Valley	CB	CB-07	52.96425	-6.40302	420	Boulder	38.84	0.79	15.656	1.376
Valley	CB	CB-08	52.96436	-6.4029	420	Boulder	36.25	0.70	17.112	1.378
Valley	CB	CB-09	52.96431	-6.40291	419	Boulder	39.22	0.76	15.431	1.376
Valley	CB	CB-10	52.96418	-6.40294	418	Boulder	38.79	0.90	15.662	1.376
Valley	UGD	UGD-01	53.02901	-6.36507	336	Boulder	38.23	0.93	15.984	1.376
Valley	UGD	UGD-02	53.02903	-6.36502	336	Boulder	37.90	0.91	16.173	1.377
Valley	UGD	UGD-03	53.02916	-6.36542	332	Boulder	39.77	0.90	15.113	1.376
Valley	UGD	UGD-04	53.02936	-6.3655	330	Boulder	39.07	0.90	15.51	1.376
Valley	UGD	UGD-05	53.02939	-6.36557	330	Boulder	38.21	0.96	16.002	1.376
Valley	UGD	UGD-06	53.0296	-6.36602	333	Boulder	37.71	0.73	16.286	1.377
Valley	UGD	UGD-07	53.0297	-6.36598	335	Boulder	39.89	0.97	15.054	1.376
Valley	UGD	UGD-08	53.02973	-6.36654	328	Boulder	40.46	0.97	14.731	1.376
Valley	UGD	UGD-09	53.02976	-6.36657	329	Boulder	37.25	0.91	16.551	1.377
Valley	UGD	UGD-10	53.02974	-6.36667	328	Boulder	36.75	0.68	16.835	1.377
Valley	UGD	UGD-11	53.03012	-6.36654	334	Boulder	36.35	0.84	17.062	1.378
Valley	UGD	UGD-12	53.0302	-6.36666	332	Boulder	36.26	0.78	17.115	1.378
Valley	UGD	UGD-13	53.03029	-6.36672	333	Boulder	39.81	0.71	15.098	1.376
Valley	UGD	UGD-14	53.03049	-6.36659	332	Boulder	38.94	0.82	15.586	1.376
Valley	UGD	UGD-15	53.03056	-6.36662	332	Boulder	37.61	0.99	16.351	1.377
Valley	UGD	UGD-16	53.03062	-6.36666	330	Boulder	39.21	1.03	15.423	1.376
Valley	UGD	UGD-17	53.03066	-6.36667	330	Boulder	39.29	0.96	15.385	1.376



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Valley	UGD	UGD-18	53.03097	-6.367	330	Boulder	37.58	0.93	16.358	1.377
Valley	UGD	UGD-19	53.03113	-6.3668	337	Boulder	39.93	1.12	15.022	1.376
Valley	UGD	UGD-20	53.03128	-6.36709	334	Boulder	38.19	0.83	16.008	1.376
Valley	GW	GW-01	53.06269	-6.33445	275	Boulder	37.29	0.97	16.498	1.377
Valley	GW	GW-02	53.06269	-6.3345	274	Boulder	38.13	0.86	16.021	1.376
Valley	GW	GW-03	53.0626	-6.33461	273	Boulder	38.36	1.09	15.891	1.376
Valley	GW	GW-04	53.06254	-6.33465	269	Boulder	39.13	0.95	15.454	1.376
Valley	GW	GW-05	53.06248	-6.33472	266	Boulder	35.67	0.72	17.422	1.378
Valley	GW	GW-06	53.06242	-6.33468	271	Boulder	36.61	0.83	16.892	1.377
Valley	GW	GW-07	53.06235	-6.33471	270	Boulder	36.35	0.90	17.044	1.378
Valley	GW	GW-08	53.06227	-6.33476	267	Boulder	38.18	1.01	16.002	1.376
Valley	GW	GW-09	53.0622	-6.3347	266	Boulder	35.92	0.81	17.29	1.378
Valley	GW	GW-10	53.06211	-6.33465	268	Boulder	36.22	0.81	17.119	1.378
Valley	GW	GW-11	53.06207	-6.33459	268	Boulder	37.16	1.06	16.589	1.377
Valley	GW	GW-12	53.06197	-6.33464	269	Boulder	36.73	1.00	16.835	1.377
Valley	GW	GW-13	53.06182	-6.3347	269	Boulder	37.43	0.94	16.438	1.377
Valley	GW	GW-14	53.06169	-6.3347	268	Boulder	36.50	0.78	16.968	1.378
Valley	GW	GW-15	53.06165	-6.33473	267	Boulder	37.27	0.86	16.532	1.377
Valley	GW	GW-16	53.06153	-6.33473	268	Boulder	37.14	0.79	16.608	1.377
Valley	GW	GW-17	53.06144	-6.33487	266	Boulder	37.98	0.74	16.135	1.377
Valley	GW	GW-18	53.0614	-6.33493	262	Boulder	36.98	0.79	16.703	1.377
Valley	GW	GW-19	53.06132	-6.33491	261	Boulder	37.72	0.88	16.286	1.377
Valley	GW	GW-20	53.0612	-6.33494	261	Boulder	37.92	0.61	16.173	1.377
Summit	CS	CS-01	53.08641	-6.3626	570	Bedrock	35.45	0.63	17.56	1.379
Summit	CS	CS-02	53.08638	-6.36257	571	Boulder	34.01	0.58	18.389	1.381
Summit	CS	CS-03	53.08649	-6.36254	570	Bedrock	35.62	0.73	17.464	1.378
Summit	CS	CS-04	53.08645	-6.36234	570	Bedrock	39.11	0.82	15.478	1.376
Summit	CS	CS-05	53.08653	-6.36226	570	Bedrock	39.31	0.80	15.367	1.376
Summit	CS	CS-06	53.08663	-6.36216	570	Boulder	37.31	1.00	16.506	1.377
Summit	CS	CS-07	53.0864	-6.36238	570	Bedrock	39.39	0.66	15.321	1.376
Summit	CS	CS-08	53.08634	-6.36248	570	Bedrock	36.21	0.81	17.127	1.378
Summit	CS	CS-09	53.08615	-6.36222	569	Bedrock	39.03	0.56	15.529	1.376
Summit	CS	CS-10	53.08611	-6.36223	569	Bedrock	40.13	0.75	14.905	1.376

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Summit	CS	CS-11	53.08605	-6.36221	567	Bedrock	38.09	0.50	16.059	1.376
Summit	CS	CS-12	53.08595	-6.36209	567	Boulder	37.23	1.02	16.548	1.377
Summit	CS	CS-13	53.08597	-6.36204	568	Bedrock	39.01	0.81	15.548	1.376
Summit	CS	CS-14	53.086	-6.3619	568	Bedrock	39.58	0.70	15.226	1.376
Summit	CS	CS-15	53.08591	-6.3618	568	Bedrock	36.97	0.94	16.703	1.377
Summit	CS	CS-16	53.08586	-6.36144	569	Bedrock	36.74	0.79	16.835	1.377
Summit	CS	CS-17	53.08616	-6.36171	569	Bedrock	36.67	0.71	16.871	1.377
Summit	CS	CS-18	53.08634	-6.36132	569	Boulder	37.58	0.99	16.353	1.377
Summit	CS	CS-19	53.08645	-6.36164	569	Boulder	37.25	0.73	16.544	1.377
Summit	CS	CS-20	53.08644	-6.36157	569	Bedrock	37.66	0.80	16.317	1.377

<sup>a</sup> Standard Error of the Mean