

Did a gamma-ray burst initiate the late Ordovician mass extinction?

A.L. Melott¹, B.S. Lieberman², C.M. Laird¹, L.D. Martin³, M.V. Medvedev¹,
B.C. Thomas¹, J.K. Cannizzo⁴, N. Gehrels⁴ and C.H. Jackman⁵

¹Department of Physics and Astronomy, University of Kansas, Lawrence, KS 66045 USA
e-mail: melott@ku.edu

²Departments of Geology and Ecology and Evolutionary Biology, University of Kansas, Lawrence, KS 66045, USA

³Museum of Natural History, University of Kansas, Lawrence, KS 66045, USA

⁴Laboratory for High Energy Astrophysics, NASA Goddard Space Flight Center, Code 661, Greenbelt, MD 20771, USA

⁵Laboratory for Atmospheres, NASA Goddard Space Flight Center, Code 916, Greenbelt, MD 20771, USA

Abstract: Gamma-ray bursts (GRBs) produce a flux of radiation detectable across the observable Universe. A GRB within our own galaxy could do considerable damage to the Earth's biosphere; rate estimates suggest that a dangerously near GRB should occur on average two or more times per billion years. At least five times in the history of life, the Earth has experienced mass extinctions that eliminated a large percentage of the biota. Many possible causes have been documented, and GRBs may also have contributed. The late Ordovician mass extinction approximately 440 million years ago may be at least partly the result of a GRB. A special feature of GRBs in terms of terrestrial effects is a nearly impulsive energy input of the order of 10 s. Due to expected severe depletion of the ozone layer, intense solar ultraviolet radiation would result from a nearby GRB, and some of the patterns of extinction and survivorship at this time may be attributable to elevated levels of UV radiation reaching the Earth. In addition, a GRB could trigger the global cooling which occurs at the end of the Ordovician period that follows an interval of relatively warm climate. Intense rapid cooling and glaciation at that time, previously identified as the probable cause of this mass extinction, may have resulted from a GRB.

Received 12 January 2004, accepted 12 April 2004

Key words: Population and evolution, mass extinction, gamma-ray burst, Ordovician, ultraviolet ozone.

Introduction

As mass extinctions have become well documented, interest in them has grown, partly out of concern for our current environmental situation. Extraterrestrial causes have been seriously considered in recent years. Building on previous studies, we outline the devastating effects a gamma-ray burst (GRB) would have on the biosphere and present evidence that a GRB could have contributed to the late Ordovician mass extinction (~440 million years ago), which is one of the five great mass extinctions in the history of life. We recognize that the role of a GRB in mediating the extinction is at this time only a 'credible working hypothesis' but one worthy of further consideration.

GRBs (Mészáros 2001) were initially detected as bursts of X-ray and γ radiation, and were later determined to lie at cosmological distances, implying very large energies. They are probably associated with some supernovae, producing narrow beams of radiation during gravitational collapse. Some of these beams are visible if they happen to be pointed at the observer. When localized, GRBs have been found to lie in galaxies. It is consistent with observation that our galaxy has hosted GRBs and a rate can be estimated, as well as the probability that any random locale would be irradiated.

It seems likely that a GRB has affected the Earth (Thorsett 1995; Scalo & Wheeler 2002) and that a GRB would have a substantial effect on living organisms. The consequences of a GRB depend on the flux of photons incident on the Earth, not the intrinsic total energy of the GRB. For most purposes, beaming is irrelevant, as the actual GRB rate inferred from data varies inversely with the beaming angle, leaving the observed rate constant. Due to small number statistics, lack of knowledge of the recent GRB rate is a serious source of uncertainty in estimating the role they may have played in mass extinctions. Most GRBs are at high redshift and the cosmological evolution of the rate is poorly constrained (Weinberg *et al.* 2001). However, even the most conservative rate estimates indicate that a catastrophic event within the geological record is likely.

The local rate of 'dangerous' GRB events

We discuss the threat by estimating the probable interval between 'dangerous' GRB events. Of course, the threat continuously varies with distance. For definiteness, we estimate the rate of events at low redshift (most of the geological record lies within the time corresponding to cosmologically low redshifts). We estimate the distance and power of the

nearest probable event within the disk of our galaxy (which is sufficiently thin to be treated as two-dimensional for this purpose). As we shall see, that a danger exists has a high certainty. For reasons that will emerge from our later atmospheric discussion, we estimate that a typical GRB pointed at the Earth from 3 kpc (or about 10 000 light years) or closer constitutes a serious threat to the biosphere.

There are various ways to estimate the GRB frequency. One approach is to use best fits to an estimate of the GRB rate at low redshift, about $0.44 \text{ Gpc}^{-3} \text{ y}^{-1}$ (Guetta *et al.* 2003). This can be compared with the number density of typical galaxies, about $3 \times 10^{-3} \text{ Mpc}^{-3}$ (using a fit to the luminosity function in Loveday *et al.* (1992), and assuming $h=0.7$). This gives an event rate of $1.5 \times 10^{-7} \text{ gal}^{-1} \text{ y}^{-1}$; since 3 kpc is about one-fifth of the radius of our galaxy, we would expect a dangerous event rate of $6 \times 10^{-9} \text{ y}^{-1}$, in other words about 170 My between events within 3 kpc. This is a crude method, if only because the galaxy does not have a definite boundary.

A cleaner approach, following Scalo and Wheeler (2002) is to scale from blue luminosity density. Blue luminosity is produced by the large, hot stars with short lifetimes that are thought to be the precursors of most GRBs. They present results based on various assumed rates of GRB frequency evolution – the rate of GRBs is thought to have been larger in the past, consistent with scaling with the star formation rate. We cite estimates for the strong evolution case. This produces the smallest number of GRBs within ‘recent’ cosmological time, when the five largest mass extinctions occurred. Current research (e.g. Kewley *et al.* 2002) favours older estimates of strong evolution in the star formation rate, when reddening is corrected. In this case the typical distance to the nearest GRB irradiating a locale within our galaxy would be $1.9 \text{ kpc } t_G^{-1/2}$, where t_G is the time interval in Gy. As we shall see, the greatest threat appears to come from ozone depletion resulting in increased solar UV reaching the Earth’s surface. The ozone shield is thought to have been significant for about 1 Gy, so we shall use this timescale. We expect the nearest burst irradiating the Earth to have been at a distance of about 2 kpc over the last 1 Gy. This is, of course, only an average. This method may bias the result downward somewhat, because there is evidence that GRBs preferentially go off in regions of low galaxy density, rather like the one in which our galaxy is situated (Bornancini *et al.* 2004). Thus, the rate in the Milky Way might be somewhat higher than the global average. This is not expected to be a large effect.

Another method is to use the fit of Guetta *et al.* (2003) to the low redshift effective GRB rate, again $0.44 \text{ Gpc}^{-3} \text{ y}^{-1}$. Using the universal blue luminosity density of $6.0 \times 10^{43} \text{ W Gpc}^{-3}$ (for $h=0.7$; Zucca *et al.* 1997) and the blue luminosity surface density of our galaxy of $8.0 \times 10^{27} \text{ W pc}^{-2}$ (Binney & Merrifield 1998), we estimate a rate of $5.8 \times 10^{-11} \text{ y}^{-1} \text{ kpc}^{-2}$. Using the Poisson process model, the estimated distances to nearest events are only 10% greater than those of Scalo & Wheeler (2002). This is excellent agreement, since their estimate is based on fits to evolution of the star formation

rate and the latter estimate is based on the independent Guetta *et al.* (2003) estimate of the GRB evolution rate.

The latest results (Firmani *et al.* 2004), scaled from the star formation rate in the Milky Way galaxy, suggest a higher frequency, about $1.5 \times 10^{-7} \text{ y}^{-1}$ in our galaxy (pointed at us). Again assuming a 15 kpc nominal radius for our galaxy, the Poisson model would place the typical nearest burst much closer, at about 1 kpc $t_G^{-1/2}$. Firmani *et al.* (2004) represent an extreme assumption, which is, however, not inconsistent with GRB evolution data (e.g. Weinberg *et al.* 2001).

Of course, the Poisson process model ignores local rate variation, such as enhanced star formation in spiral arms. However, given that we are averaging over long times, in which our solar system has migrated in and out of such regions, as well as the possibility of an event which might have occurred within a large fraction of the surface area of this galaxy, the approach should be valid for mean rates. Most GRBs are probably beamed, with γ -rays confined to a narrow cone. However, these estimates give the probability based on the observed flux; if they are beamed, events occur more often, with only a small fraction aimed at a given observer.

GRB have a wide range of duration and ‘isotropic equivalent energy’ – the energy of the GRB if the observed flux is assumed to be radiated equally in all directions. Guetta *et al.* (2003) estimate a typical isotropic equivalent luminosity of $4.4 \times 10^{44} \text{ W}$ at low redshift from an analysis of BATSE data and the observed redshift distribution of events for which this information is available. Lloyd-Rooney *et al.* (2002) deduce a slightly larger number, about $6 \times 10^{44} \text{ W}$, from a statistical analysis of 220 BATSE GRBs with redshifts estimated from a luminosity–variability relation. We use an average power $5 \times 10^{44} \text{ W}$ for the purposes of this discussion.

It can be shown that, due to the shallow slope of the GRB luminosity function, the rarer, more distant, and luminous GRBs offer an additional threat of comparable magnitude. So far we have presented estimates for a local event given a ‘typical GRB’. However, GRBs show large diversity in total isotropic luminosity ranging from about 10^{43} to 10^{47} W . Here we estimate the contribution of more energetic but less frequent events. Yonetoku *et al.* (2003) exploited the correlation (Amati *et al.* 2002) between the isotropic energy and the peak energy of the GRB spectrum to derive the luminosities of 684 GRBs. The derived cumulative luminosity function $N(>L)$, i.e. the number of GRBs with luminosity greater than L , at redshift zero (i.e. the luminosity evolution with redshift being removed), is approximated at luminosities greater than 10^{45} W as

$$N(>L) \sim (L/10^{47} \text{ W})^{-1.4}.$$

To produce the same energy flux at the Earth a more energetic GRB may be seen from a larger distance:

$$F_{\text{obs}} = L/R^2 \sim \text{constant}.$$

This increases the chance probability of such a GRB being within the distance R from us:

$$P = V_{\text{GRB}}/V_{\text{Galaxy}} \sim R^2,$$

where V_{GRB} and V_{Galaxy} are the volume of the Galaxy within the distance R from a GRB and the total volume of the Galaxy, respectively. Note that with R being of order or greater than 3 kpc, the galactic disk may be treated as two-dimensional. Note also that, because of large dust opacity around the Galactic centre, V_{GRB} cannot be greater than about 50% of V_{Galaxy} . Normalizing the rate of dangerous events to R_t , the rate of typical GRBs, we estimate the rate of dangerous events as a function of luminosity to be $\sim R_t (L/10^{45} \text{ W})^{-0.4}$. As one can see, the dependence on the luminosity is very shallow, which means that large but rare events are almost as dangerous as the typical, more frequent and less energetic GRBs. Ultimately, this increases the total rate of potentially dangerous GRBs by a factor of two.

There also exists a category of short-burst GRBs, about which much less is known. They have much harder photon spectra but also a lower rate and lower total energy. We do not include them in our estimate of the hazard to biospheres.

The most probable equivalent distance to the nearest event in the last 1 Gy could range from 0.7 to 2 kpc, depending on which of the event rate estimates we adopt and whether or not we include the effect of more distant powerful GRBs. Two events GRB1 and GRB2 bracket the range. GRB1 and GRB2 both have an equivalent power of $5 \times 10^{44} \text{ W}$ for 10 s. GRB1 is at 2 kpc and GRB2 is at 700 pc. GRB1 uses the low end of the rate estimates; GRB2 uses the high end and includes as a correction the contribution of the more luminous, distant events. From GRB1 the upper atmosphere of the Earth is irradiated by about 10 kW m^{-2} and in GRB2 it is about 80 kW m^{-2} in X-rays and γ -rays.

It has been suggested that copious quantities of ultra high-energy cosmic rays may be emitted by GRBs (e.g. Wick *et al.* 2003; Waxman 2004). If this were true, cosmic rays above 10^{18} eV would travel from our fiducial bursts with so little deflection as to arrive within the beam of the burst photons. However, in the absence of any direct detection of such cosmic rays or the neutrinos that would be expected to accompany them (Kravchenko *et al.* 2003), we will not as yet consider this additional stress on the biosphere.

Atmospheric and biological effects of GRBs

The flux of order 10 kW m^{-2} from GRB1 is at first sight not unusual, since it is about seven times the intensity of sunlight at the radius of the Earth's orbit. However, its effects are far different. Photon energies are typically hundreds of keV, in the X-ray and γ -ray region, rather than eV, the visible light region. The mean free path of such photons in the atmosphere is quite short and most of the energy is dumped in the stratosphere where ozone concentrations are highest (Gehrels *et al.* 2003). Moreover, the effects are quite different: each photon is able to produce many ionizations or dissociations of molecules, about one ionization per 35 eV, as used in Gehrels *et al.* (2003). Smith *et al.* (2003) have computed the propagation of such photons in Earth-like planetary atmospheres. About 0.002 of the energy reaches the surface, primarily in the form of dangerous UVB (280–320 nm) photons. The short burst

UV power of 20 W m^{-2} from GRB1 is about five times the typical UVB flux at the Earth's surface on a bright sunny day, but is not likely to be a major effect. On the other hand, 160 W m^{-2} from GRB2 reaching the ground would likely be quite damaging, even in a short period. So, the danger to one side of the Earth in the initial burst depends critically upon the proximity of the GRB. Most of the burst energy incident upon the atmosphere goes into ionization and dissociation of molecules. As we shall see, this has much more disruptive long-term effects.

If ultra high-energy cosmic rays (mostly protons) accompanied the GRB, they could irradiate the surface with muons, secondaries produced by interaction with the atmosphere, while producing radionuclides by spallation. However, as discussed earlier the cosmic-ray contribution is speculative. Most cosmic rays will scatter and add a contribution comparable to the background from supernovae. Most GRB photons are insufficiently energetic to induce any nuclear reactions. Thus, the instantaneous biotic effects will be UV-dominated and, for our minimal GRB, modest (compared to the long-term effects) and confined to the facing side of the Earth. However, we identify as a major uncertainty the controversy over GRB cosmic-ray production, which could in principle greatly exacerbate the effects (Wick *et al.* 2003; Waxman 2004).

GRBs may be approximately as disruptive as supernovae, on average, or considerably more so, depending on assumptions. The computations of Gehrels *et al.* (2003) show that a supernova at a distance of 8 pc is likely to happen about 1.5 Gy^{-1} , close to the assumed frequency of the GRB events we consider. It would irradiate the atmosphere with about 25 kJ m^{-2} , which is one-fourth that of GRB1. The fluence in excess of 1 MJ m^{-2} from GRB2 is far in excess of any likely result from a supernova. Moreover, the supernova irradiation would be spread out over a period of many months rather than seconds. Any processes that act to restore the former equilibrium, such as the destruction of oxides of nitrogen, will reduce the impact of the supernova, since the synthesis of these compounds is essentially instantaneous. Previous computations (Laird *et al.* 1997) show that the timescale for decay of oxides of nitrogen and the timescale for recovery of the ozone layer, once these destructive catalysts are removed, is of the order of several years. The slow steady irradiation from a supernova must compete with decays, but the GRB irradiation accumulates rapidly. As we now discuss, the NO_x production should be somewhat self-limiting, partially due to recombination of N atoms with NO molecules to form N_2 .

The NO_x (N, NO, NO_2) production is somewhat difficult to estimate. The nitrogen atoms produced through the dissociation of N_2 by the gamma rays and their associated secondary electrons create NO through the reaction $\text{N} + \text{O}_2 \rightarrow \text{NO} + \text{O}$ at a rate of $k_1 = 1.5 \times 10^{-11} \exp(-3600/T)$ and destroy NO through the reaction $\text{N} + \text{NO} \rightarrow \text{N}_2 + \text{O}$ at a rate of $k_2 = 2.1 \times 10^{-11} \exp(100/T)$ (Sander *et al.* 2003). The highest energy gamma rays can penetrate to about the middle of the stratosphere, where they have the most influence on total

ozone. Our typical GRB photons penetrate to near the middle of the stratosphere at 10 hPa (~ 30 km); most of the ozone is near this altitude. With an assumed temperature of 220 K, the timescale for reaction of N with O_2 is about 9 s ($k_1 = 1.2 \times 10^{-18} \text{ cm}^3 \text{ s}^{-1}$) and for reaction of N with NO it is about 25 s ($k_2 = 3.3 \times 10^{-11} \text{ cm}^3 \text{ s}^{-1}$), assuming an NO mixing ratio of 2.5 ppbv. The NO_x will thus increase substantially in the middle stratosphere before being self-limited through the second reaction. If the region is heated by about 10 K, as seems reasonable, the creation of NO will go slightly faster and its destruction slightly slower.

There is an additional possible reaction that may limit NO_x production. This is $N+N \rightarrow N_2 + \gamma$, which could in principle destroy N atoms before they can go on to make NO_x . This reaction has a very limited phase space and is highly unlikely due to the simultaneous requirement of momentum and energy conservation. A more likely reaction is the three-body reaction $N+N+M \rightarrow N_2+M+\gamma$, where M is an arbitrary reactant. However, this is likely to be a small effect and rate limiting for very large NO_x production levels only. The extreme GRB2 produces about 5×10^{13} ionizations cm^{-3} . At 10 hPa, the total number density is about $3 \times 10^{17} \text{ cm}^{-3}$. The N_2 abundance would thus be about $2.3 \times 10^{17} \text{ cm}^{-3}$ and O_2 abundance would be about $6.3 \times 10^{16} \text{ cm}^{-3}$. Assuming that we produce about 1.25 N atoms per ionization, then there would be 6.2×10^{13} N atoms cm^{-3} produced by GRB2. The ratio of N to N_2 would be 2.7×10^{-4} ; the ratio of N to O_2 would be 9.8×10^{-4} ; and the ratio of N to M (total number density) would be 2.1×10^{-4} . Thus, the N abundance compared with O_2 or N_2 is quite small; they would be even smaller for GRB1. It would thus seem logical that there is more probability of N reacting with O_2 than N reacting with N (especially in a three-body reaction). It therefore seems very probable that a GRB will create NO, even though the production is self-limited by the $N+NO \rightarrow N_2+O$ reaction. This rate limitation should only come into play after rather high NO densities are reached.

Rusch *et al.* (1981) discuss the efficiency for production of NO_x during an extreme particle precipitation period, the August 1972 solar proton event, and show that the efficiency for NO_x production is dependent on altitude and the duration of the event. The reaction rates k_1 and k_2 for $N+O_2$ and $N+NO$, respectively, assume that nitrogen atoms are in their ground state, $N(^4S)$. Rusch *et al.* (1981) present model simulations with the branching ration for production of nitrogen atoms in the excited state, $N(^2D)$, from 0.5 to 1.0. They conclude that a branching ratio of 0.8 for production of $N(^2D)$ appears to produce an ozone destruction most compatible with the measurements. The reaction rate of $N(^2D)+O_2 \rightarrow NO+O$ ($k_3 = 5.3 \times 10^{-12} \text{ cm}^3 \text{ s}^{-1}$, from Rees (1989)) is much faster than k_1 , given above. The self-limiting $N(^2D)+NO \rightarrow N_2+O$ ($k_4 = 7 \times 10^{-11} \text{ cm}^3 \text{ s}^{-1}$ from Rees (1989)) reaction rate is similar in value to reaction rate k_2 . This discussion suggests that NO_x will be increased by a GRB, possibly in a substantial way, and will cause a depletion of ozone. It is possible that other reactions need to be considered during the 10 s duration of a GRB, in order to perform a reliable computation of the

ozone depletion. A detailed computation of the ozone impact is outside the scope of this paper and is under careful investigation.

Long-term effects of GRB would spread around the Earth (by prevailing winds and diffusion processes) and include ozone layer depletion, acid rain and global cooling. The effects of a nearly polar burst would never fully affect both hemispheres, because the timescale for diffusion into the opposite hemisphere is of the order of a few years, equalling the decay time of the catalytic products. However, due to the solid angle subtended, a purely polar burst is highly improbable. Bursts as high as 60 degrees latitude that spill significantly into the opposite hemisphere have considerable impact on the opposite hemisphere, extending all the way to the pole in a few months, depending on the time of year. On the other hand, longitudinal diffusion around the Earth takes only a week or two.

The chemical effects result primarily from ionization and dissociation of molecules of N_2 and O_2 , the dominant constituents of the atmosphere. The resulting highly reactive products of N_2 dissociation are various oxides of nitrogen. Nitric acid comparable to or exceeding anthropogenic levels is a possible product (Thorsett 1995). Global cooling may occur (Reid *et al.* 1978) from the absorption of visible light by NO_2 . On the other hand, very substantial ozone depletion will result from NO_x constituents, which catalyze conversion (Gehrels *et al.* 2003) of ozone to oxygen molecules. NO_x species persist for several years in the atmosphere (Jackman *et al.* 2000), so there should be a comparable timescale for ozone depletion. Ozone now absorbs about 98% of biologically damaging solar UV radiation before it reaches the surface, so the depletion of the ozone layer would have serious consequences. A GRB may have paradoxically produced darkened skies and heightened UV radiation.

Our minimal GRB1 event would produce about 6×10^{12} ionizations cm^{-3} , mostly in the stratosphere, using 35 eV per ionization, as in Gehrels *et al.* (2003). GRB2 would produce about 5×10^{13} ionizations cm^{-3} . In comparison, the largest solar proton event in the last 40 years produced of the order of 1.7×10^9 ionizations cm^{-3} , also in the stratosphere (Jackman *et al.* 1995). This event was computed to have produced a maximum total ozone depletion of about 4%. The simulations of Gehrels *et al.* (2003) showed significant ozone depletion for a supernova with fluence four times less than that of GRB1, albeit spread over many months. A linear extrapolation is not valid, but it is possible that a GRB could produce significant ozone depletion. The short timescale and high intensity of the burst is a significant challenge when evaluating the atmospheric changes, which is beyond the scope of this paper. Pending the results of these computations, we regard 3 kpc as a reasonable estimate of the proximity at which a GRB becomes disruptive. Such an event has twice the fluence of the likely nearest supernova and many orders of magnitude more than a solar flare, which produces significant ozone depletion. With the conservative assumptions leading to GRB1, such an event is likely every 500 My.

Even with NO_x-limiting recombination processes, it seems likely that it will have significant damaging effects.

Even modest (10–30%) increases in UV flux, particularly around 300 nm, can be lethal to a variety of organisms (Kiesecker *et al.* 2001; Häder *et al.* 2003; see also Cockell & Blaustein 2000), including phytoplankton, the basis of the marine food chain, as well as oxygen production. UV is attenuated by water, although the precise absorption is heavily dependent upon particulates and dissolved organics (Häder *et al.* 2003). Penetration depths vary from metres to tens of metres. As would be expected, UV effects on micro-organisms have been found to decrease with water depth (Sommer *et al.* 1999).

Given the total number of ionizations expected from our GRB, of the order of $10^{18-19} \text{ cm}^{-2}$, it seems possible that the nitrate deposition could be substantial. From the opacity of NO₂ (a brown gas), there is a possible significant reduction in visible light as a consequence of this event. The reduction in global temperatures resulting from this may have possibly initiated the Ordovician ice age, which came on suddenly in the middle of a period of high climate stability.

Aspects of the late Ordovician mass extinction potentially compatible with a GRB

The late Ordovician is one of the largest mass extinctions in terms of its scale and scope. It appears to comprise two large, abrupt extinction events, separated by 0.5–2 My (Brenchley *et al.* 1994, 1995, 2003; Hallam & Wignall 1997). All major marine invertebrate groups show high rates of extinction during this interval. The late Ordovician mass extinction has been related to a period of rapid cooling followed by a period of rapid global warming (Orth *et al.* 1986; Brenchley *et al.* 1994, 1995, 2003). Brenchley (1984), Brenchley *et al.* (1991, 1994, 1995, 2003), Briggs *et al.* (1988), Owen *et al.* (1991), Fortey & Cocks (2003), Sheehan (2001) and others have convincingly demonstrated how global cooling and warming events could have led to a mass extinction, and we do not dispute the contribution of global cooling to this extinction.

Prior to the late Ordovician cooling, temperatures were relatively warm and it is the suddenness of the climate changes and the elimination of habitats due to sea-level fall that are believed to have precipitated the extinctions (Brenchley *et al.* 1991, 1994, 1995, 2003; Melchin & Mitchell 1991; Owen *et al.* 1991; Sheehan 2001). We emphasize that there is evidence for a possible link between GRBs and global cooling. GRBs produce opaque nitrogen dioxide in the Earth's atmosphere, providing a mechanism for global cooling. Berner (1994), Brenchley *et al.* (1994) and Sheehan (2001) remarked that it is a puzzle that a glaciation developed in the late Ordovician given warm climatic conditions in the middle Ordovician, although Brenchley *et al.* (1995, 2003) did provide an intriguing explanation based on an increase in the rate of organic carbon burial. It is noteworthy that palaeoclimate models (Herrmann & Patzkowsky 2002; Herrmann *et al.* 2003) have shown that late Ordovician glaciation could not have occurred without some external forcing mechanism; one

such impulse could be reduced solar insolation precipitated by a GRB.

The late Ordovician has been studied comprehensively (e.g. Brenchley *et al.* 1991, 1994, 1995, 2003; Melchin & Mitchell 1991; Owen *et al.* 1991; Droser *et al.* 1996; Adrain *et al.* 1998; Anstey *et al.* 2003; Fortey & Cocks 2003) yet generalizations about which species become extinct and why are difficult. (Indeed, this is also true of all the other known mass extinctions.) There may be some patterns of extinction, however, compatible with an initial burst of UV radiation followed by a greatly reduced ozone layer, with subsequent relatively rapid global cooling followed by global warming. The oxygen level of the atmosphere in the Ordovician was not greatly different from current levels (Berner *et al.* 2003) and an ozone shield was in place. Its strong reduction would almost certainly produce catastrophic consequences associated first with more UV radiation reaching the Earth's surface similar to those observed in modern organisms (Kiesecker *et al.* 2001; Häder *et al.* 2003).

Notably, one would predict that planktonic organisms or larvae would more likely be affected than benthic organisms because the former are less shielded from radiation (unless they occupied deep water). Similarly, epifaunal organisms would be more affected than infaunal organisms buried and shielded in sediments. Both of these patterns may exist in the late Ordovician. Furthermore, Brenchley *et al.* (2003) suggest that some extinctions may have begun before the episode of global cooling, indicating the possibility of a precursor event to the glaciation and supporting the notion that global cooling alone may not be sufficient to explain the mass extinction.

One factor that correlates well with an increased likelihood of extinction in trilobites, one of the important groups in the Ordovician (e.g. Briggs *et al.* 1988), is the amount of time a typical organism spent in the water column, especially as a larva. Chatterton & Speyer (1989) quantified the proportion of genera of trilobites with different larval types that became extinct at the end of the Ordovician. Trilobites inferred to have had a planktonic larval phase are proportionately 1.5 times more likely to become extinct than species with a benthic larval phase. In fact, it appears to not only have been whether the larvae were planktonic but the amount of time the larvae spent in the plankton: longer planktonic larval phases, when such can be inferred, are associated with increased extinction probability relative to that of shorter planktonic larval phases. Trilobite species whose larvae floated in the water column as opposed to those that crawled and were partly protected in sediments would be expected to be much more vulnerable to the initial affects of a GRB, including the elimination of the ozone layer leading to exposure to high levels of UV radiation.

The jury is still out on whether organisms with planktonic larvae suffered high extinction rates during any of the four other mass extinctions known to have occurred in the last 500 My. The Cretaceous-Tertiary mass extinction, attributed to bolide impact and when the dinosaurs and numerous marine organisms went extinct, is particularly well characterized. For marine organisms from the Cretaceous period that

are well known and have been intensively studied, the difference between extinction in non-planktonic and planktonic species is not statistically significant (Smith & Jeffery 1998; also see Valentine & Jablonski 1986). Clearly more data are needed on all of the mass extinctions, including the late Ordovician, to pin down this correlation, but the analyses available now are at least suggestive.

The differential extinction not only of species with a planktonic larval type, but also planktonic organisms in general, may hold for a variety of taxa in the late Ordovician, although it is hard to draw precise conclusions (Sheehan 2001). Several groups of marine organisms with a planktonic lifestyle more exposed to UV radiation than groups that lived in the benthos suffered severely during the late Ordovician. These groups include graptolites, acritarchs and cephalopods (McKinney 1985; Colbath 1986; Martin 1988; Melchin & Mitchell 1991; Sheehan 2001). Most trilobites, irrespective of their larval phase, are inferred to have occupied a benthic habit as adults; however, there were some Ordovician species whose adults lived in the plankton, and these species were completely eliminated in the mass extinction (Chatterton & Speyer 1989; Owen *et al.* 1991). Branchley *et al.* (2003) have also shown that typically organisms that dwelled in the plankton were affected before benthic organisms during the mass extinction, and species dwelling in shallow water were more likely to become extinct than species dwelling in deep water (also see Branchley *et al.* 1995; Sheehan 2001). Benthic and deep-water organisms would typically be more shielded from UV radiation than planktonic and shallow-water organisms.

Also interesting is the pattern of extinction in bivalves during the late Ordovician crisis, with burrowing infaunal forms that would be shielded from UV radiation more likely to survive than epifaunal forms living on the surface that were not shielded (Kriz 1984). Finally, the surviving fauna that populated the subsequent Silurian marine world appear to have come from areas of deeper water and/or higher latitude (Melchin & Mitchell 1991; Sheehan 2001).

Conclusions

Radiation from a GRB ionizes and dissociates molecules in the atmosphere, with a burst of UV radiation reaching the ground. Furthermore, this triggers depletion of the ozone layer leading to greatly increased solar UV radiation reaching the surface, whilst also causing other affronts to life, including acid rain possibly followed by global cooling; the result: a one, two punch for life on the planet. Given estimated GRB rates, it is probable that at least one significant event occurred in the last billion years. We present as a credible working hypothesis that a GRB could have contributed to the late Ordovician extinction, first by subjecting the Earth's biota to elevated levels of UV radiation, partially due to the initial burst but primarily by depleting the ozone layer, and then by precipitating relatively rapid global cooling which was followed by rapid global warming. At present we see reasons for associating a GRB with the Ordovician mass extinction

but clearly additional tests are required. Additional atmospheric chemistry modelling is needed to pin down the radiation fluence needed to trigger given ozone depletion, nitric acid rain and global cooling levels. Given the uncertainty in the evolution of the GRB rate, it is possible that GRBs were involved in other mass extinctions or that many much more distant GRBs could have smaller effects (Scalo & Wheeler 2002). A major challenge for astrophysics is to evaluate the likely flux and spectrum of cosmic rays accompanying a GRB.

Acknowledgements

A.L.M. and B.S.L. acknowledge support from the US National Science Foundation; B.S.L. acknowledges support from a Self Faculty Award; B.C.T. acknowledges support from a Self Graduate Fellowship; and A.L.M. and B.C.T. acknowledge support from NASA and the Graduate Research Fund at the University of Kansas. R. Kaesler, D. Smith, J. Scalo, S. Thorsett and A. Miller provided helpful comments on earlier versions of this paper. Referee J. C. Wheeler made a number of extremely helpful suggestions, which improved presentation.

References

- Adrain, J.M., Fortey, R.A. & Westrop, S.R. (1998). Post-Cambrian trilobite diversity and evolutionary faunas. *Science* **280**, 1922–1925.
- Amati, L. *et al.* (2002). Intrinsic spectra and energetics of BeppoSAX gamma-ray bursts with known redshifts. *Astron. Astrophys.* **390**, 81–89.
- Anstey, R.L., Pachut, J.F. & Tuckey, M.E. (2003). Patterns of bryozoan endemism through the Ordovician–Silurian transition. *Paleobiology* **29**, 305–328.
- Berner, R.A. (1994). Geocard II: a revised model of atmospheric CO₂ over Phanerozoic time. *Am. J. Sci.* **294**, 56–91.
- Berner, R.A., Beerling, D.J., Dudley, R., Robinson, J.M. & Wildman, R.A. (2003). Phanerozoic atmospheric oxygen. *Ann. Rev. Earth Planet. Sci.* **31**, 105–134.
- Binney, J. & Merrifield, M. (1998). *Galactic Astronomy*. Princeton University Press, Princeton.
- Bornancini, C.G., Martinez, H.J., Lambas, D.G., LeFloc'h, E., Mirabel, I.F. & Minniti, D. (2004). *Astrophys. J.* in press (astro-ph/0402576).
- Branchley, P.J. (1984). Late Ordovician extinctions and their relationship to the Gondwana glaciation. *Geol. J.* **11**, 291–315.
- Branchley, P.J., Romano, M., Young, T.P. & Storch, P. (1991). Hirnantian glaciomarine diamictites – evidence for the spread of glaciation and its effect on Upper Ordovician faunas. *Geol. Surv. Can. Pap.* **90–9**, 325–336.
- Branchley, P.J., Marshall, J.D., Robertson, D.B.R., Carden, G.A.F., Long, D.G.F., Meidla, T., Hints, L. and Anderson, T.F. (1994). Bathymetric and isotopic evidence for a short-lived late Ordovician glaciation in a greenhouse period. *Geology* **22**, 295–298.
- Branchley, P.J., Carden, G.A.F. & Marshall, J.D. (1995). Environmental changes associated with the ‘first strike’ of the late Ordovician mass extinction. *Modern Geology* **20**, 69–82.
- Branchley, P.J., Carden, G.A., Hints, L., Kaljo, D., Marshall, J.D., Martma, T., Meidla, T. & Nölvak, J. (2003). High-resolution stable isotope stratigraphy of Upper Ordovician sequences: constraints on the timing of bioevents and environmental changes associated with mass extinction and glaciation. *GSA Bulletin* **115**, 89–104.
- Briggs, D.E.G., Fortey, R.A. & Clarkson, E.N.K. (1988). Extinction and the fossil record of the arthropods. In *Extinction and Survival in the Fossil*

- Record (Systematics Association Special Volume 34)*, ed. Larwood, G., pp. 171–209. Clarendon Press, Oxford.
- Chatterton, B.D.E. & Speyer, S.E. (1989). Larval ecology, life history strategies, and patterns of extinction and survivorship among Ordovician trilobites. *Paleobiology* **15**, 118–132.
- Cockell, C.S. & Blaustein, A.R. (2000). ‘Ultraviolet spring’ and the ecological consequences of catastrophic impacts. *Ecology Lett.* **3**, 77–81.
- Colbath, G.K. (1986). Abrupt terminal Ordovician extinction in phytoplankton associations, southern Appalachians. *Geology* **14**, 943–946.
- Collar, J.I. (1996). Biological effects of stellar collapse neutrinos. *Phys. Rev. Lett.* **76**, 999–1002.
- Droser, M.L., Fortey, R.A. & Li, X. (1996). The Ordovician radiation. *American Scientist* **84**, 122–131.
- Firmani, C., Avila-Reese, V., Ghisellini, G. & Tutukov, A.V. (2004). Formation rate, evolving luminosity function, jet structure, and progenitors for long gamma-ray bursts. *Astrophys. J.* submitted (astro-ph/0401580).
- Fortey, R.A. & Cocks, L.R.M. (2003). Palaeontological evidence bearing on global Ordovician–Silurian continental reconstructions. *Earth-Sci. Rev.* **61**, 245–307.
- Gehrels, N., Laird, C.M., Jackman, C.H., Cannizzo, J.K., Mattson, B.J. & Chen, W. (2003). Ozone depletion from nearby supernovae. *Astrophys. J.* **585**, 1169–1176.
- Guetta, D., Piran, T. & Waxman, E. (2003). The luminosity and angular distributions of long GRBs (astro-ph/0311488).
- Häder, D., Kuma, H.D., Smith, R.C. & Worrest, R. (2003). Aquatic ecosystems: effects of solar ultraviolet radiation and interactions with other climatic change factors. *Photochem. Photobiol. Sci.* **2**, 39–50.
- Hallam, A. & Wignall, P. (1997). *Mass Extinctions and their Aftermath*, p. 328. Oxford University Press, Oxford.
- Herrmann, A.D. & Patzkowsky, M.E. (2002). Modeling the response of the Late Ordovician climate system to different forcing factors. *Astrobiology* **2**, 560–561.
- Herrmann, A.D., Patzkowsky, M.E. & Pollard, D. (2003). Oblivious forcing with 8–12 times preindustrial levels of atmospheric pCO₂ during the Late Ordovician glaciation. *Geology* **31**, 485–488.
- Jackman, C.H. *et al.* (1995). Two-dimensional and three-dimensional model simulations, measurements, and interpretation of the influence of the October 1989 solar proton events on the middle atmosphere. *J. Geophys. Research* **100**, 11 641–11 660.
- Jackman, C.H., Fleming, E.L. & Vitt, F.M. (2000). Influence of extremely large solar proton events in a changing stratosphere. *J. Geophys. Research* **104**, 11 659–11 670.
- Kewley, L.J., Geller, M.J. & Jansen, R.A. (2002). O[II] as a star formation rate indicator. *Astron. J.* **127**, 2002–2030.
- Kiesecker, J.M., Blaustein, A.R. & Belden, L.K. (2001). Complex causes of amphibian population declines. *Nature* **410**, 681–684.
- Kravchenko, I. *et al.* (The RICE Collaboration) (2003). Limits on the high-energy electron neutrino flux from the RICE experiment. *Astroparticle Phys.* **20**, 195–213.
- Kriz, J. (1984). Autecology and ecogeny of Silurian Bivalvia. *Palaeontology* **32**, 183–195.
- Laird, C.M., Gehrels, N., Jackman, C.H. & Chen, W. (1997). Modeling supernovae effects on the Earth’s atmosphere: implications for ozone and odd nitrogen levels. *Trans. Amer. Geophys. Union* **79**, 17, April 29, S91.
- Loveday, J., Peterson, B.A., Efstathiou, G. & Maddox, S.J. (1992). The Stromlo-APM Redshift Survey. I – The luminosity function and space density of galaxies. *Astrophys. J.* **390**, 338–344.
- Lloyd-Rooney, N.M., Fryer, C.L. & Ramirez-Ruiz, E. (2002). Cosmological aspects of gamma-ray bursts: luminosity evolution and an estimate of the star formation rate at high redshifts. *Astrophys. J.* **574**, 554–565.
- Martin, F. (1988). Late Ordovician and Early Silurian acritarchs. *Bull. Brit. Mus. Nat. Hist. (Geol.)* **43**, 299–309.
- McKinney, M.L. (1985). Mass extinction patterns of marine invertebrate groups and some implications for a causal phenomenon. *Paleobiology* **11**, 227–233.
- Melchin, M.J. & Mitchell, C.E. (1991). Late Ordovician extinction in the Graptoloidea. *Geol. Surv. Can. Pap.* **90–9**, 143–156.
- Mészáros, P. (2001). Gamma-ray bursts: accumulating afterglow implications, progenitor clues, and prospects. *Science* **291**, 79–84.
- Orth, C.J., Gilmore, J.S., Quintana, L.R. & Sheehan, P.M. (1986). Terminal Ordovician extinction: geochemical analysis of the Ordovician–Silurian boundary, Anticosti Island, Quebec. *Geology* **14**, 433–436.
- Owen, A.W., Harper, D.A.T. & Rong, J.-Y. (1991). Hirnantian trilobites and brachiopods in space and time. *Geol. Surv. Can. Pap.* **90–9**, 179–190.
- Rees, M.H. (1989). *Physics and Chemistry of the Upper Atmosphere*. Cambridge University Press, Cambridge.
- Reid, G.C., McAfee, J.R. & Crutzen, P.J. (1978). Effects of intense stratospheric ionization events. *Nature* **275**, 489–492.
- Rusch, D.W., Gerard, J.-C., Solomon, S., Crutzen, P.J. & Reid, G.C. (1981). The effect of particle precipitation events on the neutral and ion chemistry of the middle atmosphere – I. Odd nitrogen. *Planet. Space Sci.* **29**, 767–774.
- Sander, S.P. *et al.* (2003). Chemical kinetics and photochemical data for use in atmospheric studies. *JPL Publication 02–25*.
- Scalo, J. & Wheeler, J.C. (2002). Astrophysical and astrobiological implications of gamma-ray burst properties. *Astrophys. J.* **566**, 723–737.
- Sheehan, P.M. (2001). The Late Ordovician mass extinction. *Ann Rev. Earth Planet. Sci.* **29**, 331–364.
- Smith, A.B. & Jeffery, C.H. (1998). Selectivity of extinction among sea urchins at the end of the Cretaceous period. *Nature* **392**, 69–71.
- Smith, D.S., Scalo, J. & Wheeler, J.C. (2003). Transport of ionizing radiation in terrestrial-like exoplanet atmospheres. *Icarus* in press (astro-ph/0308311).
- Sommer, R., Cabaj, A., Sandu, T. & Lhotsky, M. (1999). Measurement of UV radiation using suspensions of microorganisms. *J. Photochem. Photobiol. B.* **53**, 1–6.
- Thorsett, S.E. (1995). Terrestrial implications of cosmological gamma-ray burst models. *Astrophys. J.* **444**, L53–L55.
- Valentine, J.W. & Jablonski, D. (1986). Mass extinctions: sensitivity of marine larval types. *Proc. Natl. Acad. Sci., USA* **83**, 6912–6914.
- Waxman, E. (2004). High energy cosmic-rays from gamma-ray burst sources: a stronger case. *Astrophys. J.* **606**, 988–993.
- Weinberg, N., Graziani, C., Lamb, D.Q. & Reichart, D.E. (2001). Determining the gamma-ray burst rate as a function of redshift. In *Gamma-Ray Bursts in the afterglow Era*, Proceedings of the International Workshop held in Rome, eds Costa, E., Frontera, F. & Hjorth, J., p. 252. Heidelberg: Springer, Berlin.
- Wick, S.D., Dermer, C.D. & Atoyan, A. (2003). High-energy cosmic rays from galactic and extragalactic gamma-ray bursts. In *Proc. 2003 Gamma-ray Burst Conf. (Santa Fe, NM)* to appear (astro-ph/0312213).
- Yonetoku, D., Murakami, T., Nakamura, T., Yamazaki, R., Inoue, A.K. & Ioka, K. (2003). GRB formation rates inferred from the spectral peak energy – luminosity relation (astro-ph/0309217).
- Zucca, E. *et al.* (1997). The ESO Slice Project (ESP) galaxy redshift survey. II. The luminosity function and mean galaxy density. *Astron. Astrophys.* **326**, 477–488.